

Opportunities and limits in imaging microorganisms and their activities in soil microhabitats

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1 Opportunities and limits in imaging microorganisms and their

2 activities in soil microhabitats

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19 Abstract

- 20 The soil microhabitat is a heterogeneous and complex environment where local variations can
- 21 modulate phenomena observed at the plot scale. Most of the current methods used to describe
- 22 soil functioning are bulk soil analyses which do not account for fine-scale spatial variability
- and cannot fully account for the processes that occur under the influence of the 3D organisa-
- 24 tion of soil. A good representation of spatial heterogeneities is necessary for the parametrisa-
- 25 tion of new models, which aim to represent pore-scale processes that affect microbial activity.
- 26 The visualization of soil at the scale of the microhabitat can be used to extract descriptors and
- 27 reveal the nature of the relationships between the fine-scale organisation of soil's constituent
- 28 parts and soil functioning.
- 29 However, soil imaging techniques tend to be under-used, possibly due to a lack of awareness
- of the methods or due to a lack of access to the relevant instruments. In recent years, new
- 31 methods have been developed, and continuously improved, offering new possibilities to de-
- 32 cipher and describe soil physical, chemical and biological features of the soil microhabitat in
- 33 evermore exquisite detail.

- 34 This review is structured into several parts in which first imaging methods that are useful for
- describing the distribution of microorganisms and microbial activities, followed by methods
- 36 for characterising the physical organisation of the microhabitat and, finally, methods for char-
- acterising the distribution of soil chemical features, including soil organic matter, are de-
- 38 scribed. Special attention is given to the preparation steps that are required for the proper use
- 39 of the methods, either alone or in combination.

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- 41 Keywords
- 42 Imaging, microhabitat, microscale, microorganisms, soil characterisation, soil organic matter
- 43 Abbreviations
- 44 AFM: Atomic force microscopy
- 45 AHA: L-azidohomoalanine
- 46 BIB: Broad ion beam
- 47 BONCAT: Bioorthogonal non-canonical amino acid tagging
- 48 CTC: 5-cyano-2,3-ditolyl-tetrazolium chloride
- 49 DAPI: Di Aminido Phenyl Indol
- 50 DTAF: Dichlorotriazinylaminofluorescein
- 51 EDX: Energy dispersive X-ray spectroscopy
- 52 EELS: Electron energy loss spectroscopy
- 53 ESEM: Environmental scanning electron microscopy
- 54 FDA: Fluorescein diacetate
- 55 FIB: Focused ion beam
- 56 FISH: Fluorescence in situ hybridization
- 57 FITC: fluorescein isothiocyanate
- 58 FTIR: Fourier transform infrared micro-spectroscopy
- 59 INT: 2-(p-iodophenyl)-3-(p-nitrophenyl)-5-phenyl tetrazolium chloride
- 60 LTSEM: Low temperature scanning electron microscopy
- 61 MALDI: Matrix assisted laser desorption ionisation
- 62 MRI: Magnetic resonance imaging
- 63 NEXAFS: Near edge X-ray fine structure spectroscopy
- 64 NMR: Nuclear magnetic resonance
- 65 PET: Positron emission tomography
- 66 PI: Propium iodide
- 67 SBF: Serial block face

- SEM: Scanning electron microscopy 68
- SHIM: scanning helium ion microscopy 69
- SIM: Structured illumination microscopy 70
- SIMS: Secondary ionization mass spectrometry 71
- 72 SMLM: Single molecule localisation microscopy
- 73 SOM: Soil organic matter
- 74 STED: Stimulated-emission depletion microscopy
- 75 STMX: Scanning transmission X-ray microscopy
- 76 TEM: Transmission electron microscopy
- TOF: Time of flight 77
- 78 TXM: Transmission X-ray microscopy
- 79 VNIR: Visible and near infrared
- 80 XANES: X-ray absorption near edge structure
- X-ray CT: X-ray computed tomography 81
- 82 XRF: -ay fluorescence

83 1. Introduction

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Soils are extremely complex and heterogeneous environments and many properties 85 observed at the profile or at the plot scale are, in fact, determined by microscale conditions and processes (e.g., Falconer et al., 2015; Keiluweit et al., 2017; Steffens et al., 2017). In 86 87 soils, the local environment can differ dramatically across millimetres, or even less, and these 88 variations can control the spatial distribution of microorganisms and their activities (Chenu et 89 al., 2001; Ranjard and Richaume, 2001; Chenu and Stotzky, 2002; Grundmann, 2004; Jasinska et al., 2006; Raynaud and Nunan, 2014; Frey, 2015; Juyal et al., 2020). The majority of 90 the models that are currently used to describe or predict soil functioning are based on bulk soil 91 characteristics, thus implicitly assuming that microscale interactions and processes do not af-92 fect the higher scale properties. This view has been vigorously challenged recently (Baveye et 93 al., 2018). 94 Soils exhibit heterogeneities along spatial, temporal, chemical, physical and biological dimen-95 96 sions (Lehmann et al., 2020). The combinations and interactions among these dimensions 97 mean that soils are made up of a myriad of micro-environments with unique combinations of 98 properties. The microbiological functioning of soils is highly dependent on how microbial cells interact and are affected by the properties of their local environment (Alexander, 1964; 99

Baveye et al., 2018; Chenu and Stotzky, 2002). The access to resources and energy, the availability of water or O₂, the pH that microbial cells are exposed to, the other organisms in proximity for example, can all have dramatic effects on the activity of microbial cells. The environmental properties and the biological neighbourhood that microbial cells experience depend the physical structure of the soil's solid and pore phases across scales, from the micro-environment scale to the scales at which fluxes are regulated. However, we have little or no information on how microbial cells are distributed within this complex and heterogeneous environment, nor do we know much about their interactions with the physical environment or population pressures that they are subjected to. Our understanding of soil microbial functioning is therefore derived from measurements of means and gross trends (Vos et al., 2013) and, in the final analysis, is not built on a solid mechanistic foundation. It is clear that an exhaustive mapping soil microbes would be an arduous and, in view of the temporal variability inherent to soil conditions, futile endeavour. However, a clearer picture of the range of interactions that occur in soil, their prevalence and the effects they can have on microbial functioning would allow us to build models with sounder mechanistic basis. Ultimately, this can only be achieved after studying the spatial relations of soil's constituent parts, as spatial proximity is a strong modulator of the interactions that can occur.

New models, that account for microscale soil functioning, have begun to appear in recent years (Pot et al., 2015; Tecon and Or, 2017; Portell et al., 2018; Wilmoth et al., 2018; Kemgue et al., 2019). The need for data at relevant spatial scales, such as the physical structure and chemical characteristics of soil and the spatial distribution of microorganisms and their activities, for parametrising such models, is paramount if useful information is to be acquired from microscale modelling approaches (Baveye et al., 2018). In particular, visualising soil microorganisms in undisturbed soil samples is essential for understanding how they interact with their local environment, such as the local conditions experienced, their access to trophic resources or their interactions with other microorganisms, and the consequences these interactions have for soil functioning.

For example, it has been suggested that microbial activity hotspots account for a major part of total microbial activity in soils, despite being found in only a small portion of the soil volume (Kuzyakov and Blagodatskaya, 2015). A better understanding of the functioning of such hotspots, in particular of those outside well defined "spheres" such as the rhizosphere or the detritusphere, would contribute greatly to our capacity to predict the response of microbial activity to changing environmental conditions or soil management. Yet, we know close to nothing about such hotspots nor about what causes them to occur. Are they more likely to oc-

cur in the presence of certain types of pore architecture or is it the happy coincidence of complementary microbial taxa being co-located (Kim et al., 2008). To answer these questions, we must first know who is where and how they are organised.

However, visualisation methods of soils at fine scales are challenging, as soils are heterogeneous organo-mineral matrices, composed essentially of very small particles in the submm range that are variably hydrated, and the spatial organisation of which is strongly affected by moisture state. A range of visualisation methods has been used over the last 30 years, but there have been a number of recent methodological developments and a concomitant increase in the number of studies in this area. New methods are now able to provide much more information than before: not only are the resolutions of images higher but it is also possible to obtain spatial information on soil physical, chemical and biological characteristics in three dimensions. Furthermore, image processing and analysis tools have become more efficient, allowing for better correction, segmentation of areas of interest or even predictions of chemical composition (Hapca et al., 2015; Anderson et al., 2020).

Despite the technological progress, imaging methods are still under-utilised in soil sciences because the scientific community is not always aware of them, and due to the fact that they are time consuming, expensive and not always easy to implement. Furthermore, several visualisation methods require access to instruments that are not widely available (e.g. NEX-AFS, nanoSIMS). Many of the imaging methods that are useful for visualising microbial communities in their habitat have extremely small fields of view (e.g. the area covered by a nanoSIMS image is approximately 30 x 30 µm²). The questions then arise, should targetted or random sampling be used (Brus, 2019) and what of the representativeness of the images acquired? Targetted imaging is useful for determining the types of situations that can be encountered, but there is a risk that it will provide a biased view as the areas of interest chosen may not reflect the soil more generally. In order to obtain a representative view, a degree of random sampling is required, though this can be reduced with a judicious use of spatial modelling (Brus, 2019). Whilst sampling considerations are clearly a drawback of these approaches, their exceptional analytical power means that they deserve more attention.

The aim of this review is to give a broad overview of the methods currently available that can be used to obtain information on soil microorganisms and their activities in their micro-environment. We review and present different methods in which (i) microorganisms can be located in their habitats, (ii) different types of microorganisms can be identified, (iii) the characteristics of the immediate environment of microbial communities can be described and

(iv) in situ information on the activities of microorganisms can be acquired. We systematically specify the spatial resolution of the methods, the preparation techniques and quantification possibilities. We address the limits of the different methods, the possibilities of combining them and discuss perspectives in the field.

2. Visualising microorganisms in their habitats

2.1. Localising microorganisms

There are different methods for localising microorganisms in soil samples and for identifying them as bacteria, fungi, archaea etc. (Table 1). Distinguishing microorganisms from soil particles is not always straightforward, but there are a range of useful criteria based on size and shape for facilitating this: rounded shapes, filamentous forms, or the structure of the cell. These may also be combined with general or specific stains in order to increase the contrast between objects of interest and the background.

At the extremes in terms of resolution, stereomicroscopes allow for the observation of fungi (Otten and Gilligan, 1998; Otten et al., 2004; Thompson et al., 2005), identified as such based on criteria such as shape, size or even colour, while bacteria and archea are too small to be visible. Fresh soil samples can be observed directly without any special preparation but thin sections should be prepared if the spatial organisation or spatial relations are of interest. With the greater magnification of light microscopy, both uni- and pluri-cellular organisms can be visualised. Nonetheless, in soil, microorganisms do not show great contrast and it is difficult to identify them precisely based on shape, size and natural colour alone.

If the spatial relations of microbial communities with constituents of their micro-environment is of interest, then the physical structure of samples must be preserved intact, as is done in resin embedded, thin sections of soil (Tippkötter et al., 1986; Tippkötter and Ritz, 1996; Li et al., 2004). In this way, the spatial integrity of samples and the integrity of biological cells are preserved.

Epi-fluorescence microscopy, combined with the use of stains, i.e. fluorochromes, makes it possible to distinguish the targeted organisms from the background and therefore locate and enumerate bacteria (Fig. 1a) (Fisk et al., 1998; Nunan et al., 2001; Juyal et al., 2020) and fungi (Baschien et al., 2001) in 2D. The staining needed to visualise microorganisms can be carried out either before the impregnation, by immersion of the sample in a stain-

ing bath (Nunan et al., 2001), often after a fixation step, or applied after thin section preparation (Juyal et al., 2020). However, only microorganisms situated at the surface of the thin section can be stained and visualised in the latter case. Several fluorochromes specifically stain cell constituents (Table 2). Provided their excitation spectra are not superimposed, and no interferences occur, it is possible to use several staining agents simultaneously (Chen et al., 2007). The main difficulties encountered with the use of fluorochromes are related to unspecific staining, in particular with positively charged fluorochromes such as acridine orange, and to background auto-fluorescence of soil organic particles (Altemüller and Van Vliet-Lanoe, 1990; Li et al., 2004). Figure 1a shows an example of auto-fluorescence (red arrow) with similar shape and size characteristics as bacterial cells, making the identification of bacteria more complicated. In addition, the quality of the staining can be affected by the presence of clay particles, as stains tend to adsorb to clay surfaces resulting in a fluorescence which can hinder the observation of microorganisms (Li et al., 2004). Other factors may also interfere with the staining such as the stain concentration, soil pH and the type of resin used (Altemüller and Van Vliet-Lanoe, 1990; Postma and Altemüller, 1990). It is possible to distinguish microorganisms from organic compounds by collecting visual information on microbes and organic matter auto-fluorescence in different channels and subtracting the signal in one from that in the other (Cardinale, 2014; Schmidt et al., 2018). Differences in signal intensity between dyed and auto-fluorescent objects can also be used to distinguish objects of interest from the background. Finally, new methods based on two-photon excitation fluorescence can be used to take advantage of the native auto-fluorescence of soil and microorganisms to locate fungi and bacteria in soil without using any stain (Lee et al., 2022). Using methods such as those described above, Nunan et al. (2002) measured the spatial organisation of bacteria at the micrometre scale and showed that bacteria were more strongly aggregated in the subsoil than in the topsoil. Juyal et al. (2020) showed that lower soil bulk densities favor the dispersion of inoculated bacteria in soil.

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The resolution limit of light microscopy, imposed by the diffraction of light, is around 200 nm, which allows for the observation of objects between 10⁻³ and 10⁻⁷ meters (Ranjard and Richaume, 2001) in preparations between slide and coverslip or in thin sections after inclusion in a resin.

The resolution of scanning and transmission electron microscopy (SEM & TEM) is much higher (Table 1, Fig. 1b and c). It is possible to reach resolutions of circa 1 nm with SEM (Joy and Pawley, 1992) and 0.05 nm with TEM (Smith, 2008). Preserving the original soil microstructures and the integrity of organisms and organic constituents despite the high

vacuum to which the sample is exposed in the microscope is a challenge and requires the use of specific preparation methods. Samples are generally air-dried before SEM analyses. However, drying strongly affects all hydrated structures and damages biological features (e.g. bacterial cells as well as sheaths of extracellular polysaccharides appear flattened) so that conventional SEM is not recommended (Chenu and Jaunet, 1992). Samples can be observed directly in a moist state without any conducive coating using an Environmental Scanning Electron Microscope (ESEM) (Gleeson et al., 2005; Lin and Cerato, 2014), though it gives poorer quality images in terms of brightness and contrast (Dal Cortivo et al., 2017; Bertola et al., 2019). In order to avoid the damage that is associated with drying, samples can also be observed with a Low Temperature Scanning Electron Microscope (LTSEM) after cryo-fixation (Chenu and Tessier, 1995). To our knowledge, with SEM, image analysis is hardly possible given the 3D surface rendering of the images and the fact that soil microorganisms are only identified based on shape, which is a difficult task. SEM has been used for example to show the presence of fungi in the detritusphere extending between plants cells and adhering soil aggregates (Gaillard et al., 1999) and more recently to study the colonisation of fungi and unicellular organisms involved in the biodegradation of plant residues (Witzgall et al., 2021) or plastic materials in soil (Zumstein et al., 2018).

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Transmission electron microscopy can be used to visualise soil microorganisms in their habitats (Vidal et al., 2016; Watteau and Villemin, 2018). In order to do this, however, samples cannot be simply deposited on TEM grids as a suspension but rather have to be embedded in a resin, from which ultrathin sections are prepared (Watteau et al., 2002, 2012; Elsass et al. 2008). The soil samples are fixed, dehydrated and impregnated with a resin prior to the ultrathin sectionning. These methods have been adapted from biology to soils to account for the sensitivity of soil organic matter to the electron beam (Villemin and Toutain, 1987; Villemin et al., 1995; Elsass et al., 2008). In soil thin sections prepared for TEM, the use of contrasting agents and stains, such as osmium tetroxide (Villemin et al., 1995; Arai et al., 2019), uranium acetate, lead citrate (Foster, 1988; Chenu and Plante, 2006; Elsass et al., 2008) or more specific stains, helps with shape and structure based identification by adding contrast between cells wall structure and the surrounding environment. TEM allows manual quantification if studied structures are recognisable (Chenu and Plante, 2006; Watteau and Villemin, 2018) but the method is most often used qualitatively to characterise proximity and thus potential interactions between microorganisms and the surrounding minerals or aggregates (Fig 1c) (Vidal et al., 2019).

The surface of a soil sample and microbial cells can be visualised with scanning he-

lium ion microscopy (SHIM or HIM) (Qafoku et al., 2019). This relatively new method is still rarely used in soil science, even though the ion beam is less destructive for the surface of the sample than the electron beam of SEM (Bandara et al., 2021; Schmidt et al., 2021), samples do not need coating prior to analysis and resolutions of 5 Å are attainable (Joens et al., 2013), which makes it very promising for studying soils.

Microorganisms can also be identified with Atomic Force Microscopy (AFM) by visualising their topography at very fine scales. This is achieved with a stylus that moves vertically with a vertical resolution of a few Å (Binnig et al., 1986). It requires fewer sample preparation treatments than electron microscopy (Kherlopian et al., 2008); however it requires a flat surface at the base and the scale of the study objects (cell surfaces) is necessarily small (Gaboriaud and Dufrene, 2007). It has therefore not been used on soil samples but rather on model systems, e.g. to study the adhesion of bacteria to clay mineral surfaces with kaolinite, montmorillonite, goethite or hematite particles (Lower et al., 2001; Huang et al., 2015; Qu et al., 2019) and to study the weathering of biotites, chlorites or serpentine minerals by mycorrhizae and other fungi (McMaster, 2012; Li et al., 2021).

Several of the methods presented here only allow the visualisation of soil microorganisms in 2D. The localisation of microbial cells in 3D is possible either (i) by using the depth of field of the microscope (stereomicroscope, SEM, HIM, AFM), (ii) by reconstituting the sample in a non-destructive way using confocal laser scanning microscopy (Li et al., 2004), but due to the opacity of soil this is mainly useful in the analysis of transparent artificial soil models as in Sharma et al. (2020), (iii) by progressively abrading the sample with a focused ion probe (FIB) (Berleman et al., 2016; Vidal et al., 2018), (iv) by reconstitution from thin sections: superimposed serial block face (SBF) sections or broad ion beam (BIB). However, the latter have been mainly used to study geo-materials (Desbois et al., 2010; Houben et al., 2013; Hemes et al., 2015).

The data derived from the imaging of microbial distributions, such as those in Nunan et al. (2001), have been used to develop a statistical model of the spatial distribution of bacteria in soil (Raynaud and Nunan, 2014). Schnepf et al. (2022) have advocated that such an approach should be implemented more widely to analyse the distribution and organisation of microbial communities in the rhizosphere. Indeed, such data are slowly becoming available, e.g. data on the distribution of microbes around root cells (Schmidt et al., 2018), or spatial distributions of inoculated microorganisms in the soil porosity across time (Juyal et al., 2020), and start being used in biogeochemical models (Pagel et al., 2020).

In addition, electron microscopy is one of the few methods with which viruses can be visualised, their size ranging between 30 and 80 nm (Kuzyakov and Mason-Jones, 2018). Viruses are believed to influence the assembly of bacterial communities, even if this is still poorly understood for soils.

2.2. Identifying microorganisms

Prokaryotes can be distinguished from fungi on the basis of shape and size with SEM and fluorescence microscopy and on the basis of the cell wall structure with TEM. However, these tools do not allow to differentiate archaea or actinomycetes from bacteria (Foster, 1993), nor a finer differentiation within bacterial, archeal or fungal groups.

Fluorescence *in situ* hybridization (FISH) can be used for the *in situ* identification of microorganisms. The general principle of FISH is to use an oligonucleotide probe coupled to a fluorescent marker which binds to a specific sequence of RNA or DNA within the microbial cell. It is used in conjunction with epi-fluorescence microscopy (Bandara et al., 2021) or with confocal microscopy (Muggia et al., 2013). A database of probes, called probeBase, can be used to search for, and identify, relevant probes (Greuter et al., 2016). More or less specific probes can be used to target microorganisms at different taxonomic levels: archaea, crenarachaea, bacteria and fungi, bacterial phyla, or identifying gram positive bacteria (Baschien et al., 2001; Eickhorst and Tippkötter, 2008a, 2008b; Zarda et al., 1997; Kobabe et al., 2004). The FISH methods can also be used to target certain functional groups of microorganisms, e.g. those involved in denitrification (Pratscher et al., 2009; Hoshino and Schramm, 2010), nitrification and methane oxidation (Torsvik and Øvreås, 2002).

The main limitations of FISH are the percentage of successfully detected cells among the total targeted cells (Bouvier and Del Giorgio, 2003), as well as the high levels of background noise due to non-specific fixation of the stains on soil particles. It explains why FISH has been used so infrequently on soil samples. Different probes have been developed, making it possible (i) to amplify the signal, e.g. catalysed reporter deposition FISH or CARD-FISH (Kubota, 2013; Schmidt and Eickhorst, 2014; Juyal et al., 2018; Schmidt et al., 2018; Bandara et al., 2021), (ii) to increase the resolution by allowing observations *via* SEM using GOLD-FISH (Kenzaka et al., 2005; Schmidt et al., 2012) (Fig. 1d, e, f and g), and (iii) to limit unspecific staining or at least detect it . NON338 or NONEUB probe complementary to the

EUB338 probe and serving as a negative control for non-specific labelling) (Eickhorst and Tippkötter, 2008a).

3. Characteristics of microbial microhabitats

Understanding the drivers of microbial activity requires that the soil habitat be well described. The soil matrix determines the physical accessibility of microorganisms to resources and water, the local conditions that control microbial activities and also regulates trophic relations (predation, competition etc.). Unravelling the distribution of microorganisms and resources, as well as their probabilities of encounter, will likely improve our understanding of soil microbial activity.

3.1. Physical characteristics of microhabitats

Soil structure affects the spatial distribution of microorganisms and trophic resources and consequently the access that microorganisms have to substrate and controls the transport of oxygen and water, key factors for microbial growth and activity (Schlüter et al., 2020). Soil structure can be heterogeneous at the microscale and must therefore be studied at this scale. The most frequently used method to describe the 3D structure of a soil is X-ray computed tomography (X-ray CT) or X-ray micro computed tomography (µ X-ray CT) (Fig. 2). It has been used to study the heterogeneity of soil organisation (Elyeznasni et al., 2012; Hapca et al., 2015; Lucas et al., 2020) and its relation with soil biotic organisation and functioning (Helliwell et al., 2014; Kravchenko et al., 2019; Schlüter et al., 2019a, 2019b; Rohe et al., 2021). Samples used in these studies vary in size from cm, with a voxel resolution of tens of micrometres (Elyeznasni et al., 2012; Hapca et al., 2015; Rohe et al., 2021), to mm sized aggregates, where resolutions can be in the micrometre range.

The variation in absorption of the X-rays by different soil materials makes it possible to obtain images, in Hounsfield units (represented in gray levels), that reveal the different phases of the soil. As it is based on differences in density, it is mainly used to distinguish solids from voids, but the water phase can also be detected if located in pores that are sufficiently large relative to the image resolution) (Landis and Keane, 2010; Tippkötter et al., 2009) (Table 3).

Schlüter et al. (2019b) incubated sand-based microcosms in which they placed microbial hotspots either randomly or in dense layers and found that both the spatial distribution of the hotspots and the water saturation critically affected NO and N_2O emission rates. Determining air connectivity, pores tortuosity and diffusion lengths with X-ray μ CT, they concluded that local oxygen supply was the driving variable, paving the way to use soil structural attributes to predict denitrification via parametrized models.

Light and transmission electron microscopy may also be used to describe the structure of a soil, but these methods provide far less complete information, since they are in two dimensions. However, they can be adapted when better resolutions are needed (Table 3). For example, 3D acquisitions can be done on small samples such a soil aggregates using FIB-SEM (e.g., Vidal et al. 2018). Microscale biogeochemical models with an explicit representation of soil structure use X-ray μ CT images as input information, but are limited by the lack of information on sub-resolution pores in which many microorganisms reside (Pot et al., 2021). 3D high resolution methods such as FIB-SEM might help to overcome these limitations.

3.2. Chemical characteristics of microhabitats

3.2.1. Localisation of organic matter in soil

Being the trophic resource of heterotrophic microorganisms, localizing and characterizing in-situ soil organic matter is crucial. At the moment, even in microscale biogeochemical models, SOM spatial distribution is not described from soil imaging, but is prescribed assuming either homogeneous or heterogeneous distributions (Pot et al., 2021). Organic matter can be observed in soil at different resolutions (Fig. 1 and 2); first with optical microscopy on the basis of its shape and colour. It is possible to visualise organic matter after staining, as in a recent study of Merino et al. (2021), who used confocal laser micro-

With electron microscopy, organic matter can be identified using shape criteria, in the case of SEM and TEM, and by electron density in the case of TEM. It is also possible to stain SOM using heavy metals in order to amplify the contrast with the rest of the matrix (Foster, 1988; Elsass et al., 2008; Chenu et al., 2015) (Fig. 2). The use of electron microscopy coupled with energy dispersive X-ray spectroscopy (EDX), enabled by the genesis of X-ray photons

scopy to measure the decomposition of lignin that was stained with safranin-O.

following the interaction of electrons with atoms, has been used to determine the elementary composition of OM (Table 3) (e.g., Chenu and Plante, 2006; Hapca et al., 2015). Similarly, electron energy loss spectroscopy (EELS) can provide information on the elementary composition and the chemical bonds of OM (Watteau et al., 1996, 2002; Watteau and Villemin, 2018) allowing, for example, an elemental (C and N) characterisation of the polyphenolic substances during root senescence and biodegradation (Watteau et al., 2002). More recently it was used to characterize contrasting N and oxidized C contents between organo-organic and organo-mineral interfaces in a soil and suggested different organic matter stabilization processes at these interfaces, at the nanometre scale (Possinger et al., 2020).

Scanning transmission X-ray microscopy (STXM) provides an image of thin objects by transmission, as does TEM, but in this case from X-rays. The distribution of elements can be mapped based on the absorption of X-rays at different energy levels (Table 3). The samples must be very thin, but the method has the advantage that ultrathin sections of hydrated materials can be used (Solomon et al., 2012). Information on functional groups can be derived from X-ray absorption near edge structure (XANES) by coupling the STXM with Near-edge X-ray fine structure spectroscopy (NEXAFS) (Keiluweit et al., 2012; Remusat et al., 2012) at nanometre resolutions (Table 3, Fig. 3). This method was used by Lehmann et al. (2005) to map the distribution of different forms of organic carbon at the nanometer scale in soil micro-aggregates. Using ultrathin sections of rapidly frozen samples obviates the need for inclusion in a resin (Lehmann et al., 2008; Solomon et al., 2012). Otherwise, the resin used must be distinguishable from the mapped elements and for distribution maps of carbon must be carbon-free, such as polymerized elemental sulphur (Lehmann et al., 2005) or with a differentiable signature (Vidal et al., 2018; Vergara Sosa et al., 2021).

Fourier transform infrared micro-spectroscopy (micro-FTIR) can be used to characterise, quantify and locate organic molecules based on the bonds of functional groups (Singh and Gräfe, 2010). However, the resolution is generally low (Table 3). Several infrared spectro-microscopy methods using different wavelengths such as the visible and near infrared (VNIR) have enabled, using predictive models, the acquisition of images with a 53 µm resolution on dried soil, with detailed estimations of the organic carbon spatial distribution in a soil as a function of depth (Fig. 4) (Steffens and Buddenbaum, 2013; Hobley et al., 2018). This method requires almost no prior preparation except drying samples and ultrathin sectioning to 200nm using a cryo microtome (Hernandez-Soriano et al., 2018; Weng et al., 2021). Other micro-spectroscopy methods are still little used in soil science, such as Raman spectrometry with which organic matter can be characterised based on vibrational and rotational mo-

lecular analyses (Bandara et al., 2021; Lee et al., 2022; Musat et al., 2012).

Spectro-microscopy based on synchrotron radiation (SR), which is brighter and, above all, better focused than other sources of radiation, gives a finer OM identification than that obtained by spectroscopy used with conventional sources of radiations (Singh and Gräfe, 2010; Hota, 2021). It also provides information on the location and nature of the OM contained in a sample at finer scales than with conventional radiations. These methods are, for the most part, non-destructive and with higher resolutions than the same methods used with conventional sources of radiations (Weng et al., 2021), which makes it possible to obtain information on various molecular species of OM (depending of their absorption characteristics at a given wavelength) at the scale of microbial habitat (Singh and Gräfe, 2010; Hernandez-Soriano et al., 2018). Depending the photon energies used for image acquisition, some methods can be more harmful than others. SR based X-ray spectromicroscopy can harm and kill microorganisms, because of the high levels of electron volts applied to the sample. Lower energy levels are used with SR-based infrared spectromicroscopy, meaning that measurement can be repeated more easily and time sequences can be recorded (Holman and Martin, 2006).

Elementary maps of samples can be established at high resolutions using Secondary Ion Mass Spectrometry (Bandara et al., 2021; Remusat et al., 2012; Schurig et al., 2015). There are several types of Secondary Ion Mass Spectrometry (SIMS): static SIMS or Time of flight SIMS (TOF-SIMS) and dynamic SIMS, CAMECA or nanoSIMS (Myrold et al., 2011). A primary ion beam (usually Cs⁺ for samples with organic matter, or O⁻) is used to sputter the sample surface and release secondary ions, which are collected, separated and analysed. These secondary ions are characterised using a mass spectrometer for dynamic SIMS and a time-offlight mass spectrometer for static SIMS (Myrold et al., 2011). Scanning the sample gives access to elementary and isotopic map of the soil sample at very fine resolutions, i.e., less than 100 nm (Herrmann et al., 2007b). However, as with many recent techniques, SIMS suffer from heavy technical constraints, particularly in sample preparation: the samples must be dry, stable, conductive, flat and resistant to a very high vacuum (Herrmann et al., 2007b). Hence, samples are usually embedded in a carbon containing epoxy resin which contributes to a background signal that has to be removed (Mueller et al., 2013). In addition, the method is relatively destructive since the ion beam sputters the sample's surface. Quantitative elemental analyses are complicated to obtain because adequate standards are needed (Mueller et al., 2013). Finally, the observed field with nanoSIMS is very small (from 5x5 to 50x50 µm),

which makes it very difficult to identify the areas of interest (Herrmann et al., 2007b). This suggests that when the spatial scales targeted are between 100 µm and 1 cm, TOF-SIMS may be more appropriate (Myrold et al., 2011; Bandara et al., 2021), although it is rarely utilised for soils due to a lower sputtering rate of the surface samples (impacting a thinner part of the sample surface) and is therefore more exposed to problems with surface contamination. Nevertheless, the method is still worthwhile as it is one of the only methods that can be used to trace isotopic elements, such as ¹³C or ¹⁵N labelling, in a spatially explicit manner.

Recently, laser ablation-isotope ratio mass spectrometry (LA-IRMS) has been used to characterize natural ¹³C/¹²C abundance of soil organic matter, which was found to be heterogeneous, at a 10μm spatial resolution, in soil aggregates of the rhizosphere (Rodionov et al., 2019).

Finally, Matrix assisted laser desorption ionisation (MALDI), with which various analytes (sugars, lipids, amino acids, metabolites etc.) may be mapped on roots after their extraction from the soil, is worthy of consideration, as shown by Rudolph-Mohr et al. (2015).

Is it possible to localize organic matter un 3-D soil samples ? X-ray CT is appropriate as samples are subjected to an X-ray beam from different angles and this is followed by a 3D reconstruction of the object (Roose et al., 2016). With X-ray CT, as with transmission microscopy, SOM is distinguished on a density basis but this remains difficult given the density similarity between SOM and other soil constituents (Roose et al., 2016). A number of heavy metal stains have been tested in order to accentuate the contrast (Peth et al., 2014; Van Loo et al., 2014; Maenhout et al., 2021). So far, only osmium (Peth et al., 2014; Zheng et al., 2020; Maenhout et al., 2021) and iodine (Lammel et al., 2019) stains have proved to specifically stain organic matter (though Schlüter et al. (2022) suggest that Osmium binds to some minerals), to provide a detectable staining and to diffuse through the soil matrix. Osmium was successfully used to map soil organic matter (Rawlins et al., 2016).

Micro-spectroscopy, μ X-ray fluorescence (μ XRF) tomography, where X-ray induce the reemission of a X-ray fluorescent radiation from the sample, is a promising method which provides information about the chemistry of 2D (Schlüter et al., 2022) or 3D samples. Its use is very rare in 3D, in particular because improvements still have to be made to circumvent problems of fluorescence attenuation in soils (Feng et al., 2021; Hapca et al., 2015; Roose et al., 2016). Fluorescent light can be absorbed by the sample, in particular in the case of thick samples, thus attenuating the fluorescent signal. It has already be used in 3D to analyse diatom cells at the micrometre scale, or plant seedling (Pushie et al., 2014). Attempts have been

made to use the method with soil aggregates (Antipova et al., 2018).

Another option to go from 2D to 3D images is to use one of the methods described above and to carry out 3D reconstructions using statistical approaches. For example, Hapca et al. (2015) obtained several elementary 2D maps of a soil in SEM-EDX and extrapolated the chemical characteristics after combination with 3D images from X-ray tomography (Fig. 5).

3.2.2. Oxygen and CO₂ distributions in soil

As oxygen consumption and carbon dioxide emissions are directly related to the respiratory activity of soil organisms (Kuzyakov and Blagodatskaya, 2015), they are often monitored quantitatively in the atmosphere around the soil sample but are rarely monitored at the microhabitat scale. Methods are being developed to visualise the distribution of partial gas pressures in soils using probes, called needle-type or planar optodes (Santner et al., 2015) (Table 4). The former are needles, often a set, pinched in the soil with which gas partial pressures are measured (Elberling et al., 2011); it can bring additional image information if combined with other methods like X-ray computed tomography (Rohe et al., 2021). There are different types of planar optodes, but they generally involve a combination of gels and optical sensors (Pedersen et al., 2015). They are based on the use of fluorochromes that are sensitive to the presence of a gas, which acts as an exciter or an extinguisher of fluorescence, causing variations in brightness which is recorded by a camera (Rudolph et al., 2012; Pedersen et al., 2015). These fluorescent chemical probes are very sensitive and the reactions are reversible (Rudolph et al., 2012). The associated resolutions vary depending on the cameras used and the experimental device (Santner et al., 2015), but resolutions of a few tens of µm have already been reached (Larsen et al., 2011; Pedersen et al., 2015).

Most of these probes allow for the simultaneous visualisation of several parameters, such as oxygen, carbon dioxide, pH (Cf paragraphe 3.2.4.) and temperature (Borisov et al., 2011). However, this technique still requires some improvement. Indeed, measuring simultaneously different parameters can result in interferences which affect the quality of the results (Borisov et al., 2011; Pedersen et al., 2015). The dynamics of oxygen and water were studied non-invasively using this approach, thus providing new understanding of the activation of root systems regarding root respiration in a rhizosphere soil (Rudolph et al., 2012), and in a soil treated with pesticides (Rudolph-Mohr et al., 2015) with respective resolutions of 0.21 mm and 50 μ m per pixel (Fig. 6).

3.2.3. Water distribution in soil

Water is a major factor driving microbial activity in soil and can be visualised with a few methods (Table 4). Magnetic resonance imaging (MRI) can be used to visualise water dynamics in soil (Roose et al., 2016), but rarely is. It has been tested at the rhizosphere scale (0.6 mm resolution) by Pohlmeier et al. (2008), who observed changes in soil water content as a

consequence of water uptake by roots. This method has the advantage of offering 3D images, which few other methods currently offer.

Neutron computed tomography is based on the absorption of neutrons by the sample, providing three-dimensional images similar to X-ray CT images (Koliji et al., 2010). It has been used by Carminati et al. (2007) to study water flows in soil. Furthermore, neutron radiography, in which neutrons transmitted through a sample are analysed, also gives information on the distribution of water, but in two dimensions and only at resolution of a few tens of μ m (Carminati et al., 2010; Rudolph et al., 2012) (Fig. 7). Recently, fast neutron tomography has been used to visualise and monitor in 3D the root uptake of water, as 3D images can be acquired in few seconds (Tötzke et al., 2021), and determine which roots are preferentially at the origin of water uptake in soil.

The resolution of these methods explains why they are mainly used at the rhizosphere scale and are not adapted for studies at the microbial scale. Finally, X-ray CT, delivering greyscale images that are characteristic of the different phases of soil, including water, can also provide information on the spatial distribution of water. A segmentation treatment makes it possible to distinguish water from soil, in 3D and with a resolution which may be less than a mm (Landis and Keane, 2010). As with the mapping of organic matter, a number of studies have used contrast agents (heavy elements dissolved in water like CdSO₄, KI or AgNO₃) in order to increase the water signal (Van Loo et al., 2014). Others have attempted to combine images of a dry and wet soil to subtract the "dry soil" signal from the "wet soil" signal in order to isolate the "water" signal (Tracy et al., 2015). However, this approach is not suitable for clay soils, because of the risk of shrinking and swelling of clays with soil moisture changes (Baveye et al., 2018).

3.2.4. Other soil chemical characteristics

Planar optodes, described in section 3.2.2., also allow for the visualisation of other molecules such as NH_4^+ or PO_4^{3-} (Pedersen et al., 2015), and for the mapping of pH and redox potential (Eh) (Pedersen et al., 2015; Roose et al., 2016). Reagents sensitive to acid-base variations have, for example, been used by Rudolph-Mohr et al. (2015) to describe pH variations within a soil after the addition of pesticides.

4. In situ information on the physiological state and activities of microorganisms

Visualizing the microstructure of soil, localising microorganisms in the pore network, localising organic constituents which are potential substrates sets the scene for microbial activity. This also needs to be assessed directly and a range of methods have been developed to gain information on the *in situ* activities of soil microorganisms. Here, we focus on those that are compatible with visual observations of soils. The *in situ* activities of soil microorganisms at the micro-scale can be studied using methods that (i) differentiate the physiological state of microbial cells, using markers, (ii) demonstrate *in situ* substrate uptake, using tracers, and (iii) visualise potential activities of extracellular enzymes.

4.1. Physiological state of soil microorganisms

Microorganisms can be present in different physiological states in soil. While dead microorganisms are in an irreversible state in which no growth, cell elongation, nor protein synthesis can take place, active microorganisms are, as defined by Blagodatskaya and Kuzyakov (2013), « the portion of total microbial biomass that i) is involved in current utilization of substrates, ii) readily responds to substrate input e.g., by respiration, producing enzymes, or iii) is growing and reproducing ». The dormant state of soil microorganisms is that of microbial cells exhibiting strongly reduced physiological activity, e.g., resting cells forming spores or cysts. Dormant microorganisms may switch more or less rapidly from inactive/dormant state to activity (Blagodatskaya and Kuzyakov (2013).

These different physiological states can be directly or indirectly observed *in situ* by microscopy using fluorescent markers targeting microbial nucleic acids or proteins. Specific markers, such as Fluorescein Diacetate (FDA), 2-(p-iodophenyl)-3-(p-nitrophenyl)-5-phenyl tetrazolium chloride (INT), 5-cyano-2,3-ditolyl-tetrazolium chloride (CTC), make the direct visualisation of active microorganisms possible because they are subjected to enzymatic oxidation or hydrolysis by active cells (Table 5).

Recently, a promising method for identifying the active fraction of microorganisms *in situ*, called bioorthogonal non-canonical amino acid tagging (BONCAT), has been introduced to soil science (Couradeau et al., 2019). The method is based upon the incorporation *of* an ad-

ded L-methionine analog, L-azidohomoalanine (AHA) combined with a chemically reactive azide group, during protein synthesis. A biorthogonal azide-alkyne click reaction is then used to tag the molecule with a fluorophore which renders the active fraction of the microbiome detectable (Hatzenpichler et al., 2014). Although this method holds much promise, it has not yet been used on undisturbed samples. Nevertheless, BONCAT has already shown that a surprisingly high proportion of cells extracted from a soil at two depths was active (25-70%) Couradeau et al. (2019).

While its reliability is sometimes questioned (Shi et al., 2007), propidium iodide (PI) allows the visualisation of dead microorganisms because damaged cells-walls are PI permeable, contrary to living cells (Table 5). These stains may be combined with fluorochromes such as DAPI or Calcofluor White (Table 2), to both locate the total microorganisms and infer their physiological state. Dormant microorganisms are estimated by deducting active and dead microorganisms from the total microorganisms (Maraha et al., 2004).

These methods generally allow for a good estimate of dead, active and dormant microorganisms (Blagodatskaya and Kuzyakov, 2013) but present some limitations. For example, intact dead cells are not labelled by PI (Maraha et al., 2004). These labelling techniques also present limits, in particular within a soil matrix, in which, in addition to unspecific staining (Li et al., 2004), accessibility may be limited for the marker to reach its target.

4.2. Enzymatic activities

Finally, a third way to characterise the activity of microorganisms *in situ* is to visualise their production of extracellular enzymes. Zymography has been developed quite recently (Pedersen et al., 2015) and is used by only a limited number of soil research teams (Spohn and Kuzyakov, 2013; Spohn et al., 2013; Spohn and Kuzyakov, 2014; Razavi et al., 2016; Sanaullah et al., 2016; Ma et al., 2017; Guber et al., 2018; Bilyera et al., 2020). Zymography produces images of the spatial distribution of enzymes on the surface of a soil sample. A gel or membrane containing a substrate that changes colour when it comes into contact and reacts with a specific enzyme is placed on the surface of a soil sample. Colour zones signal the presence of the targeted enzyme in the sample (Guber et al., 2018; Razavi et al., 2019). The resolution of zymography (~ tens of μm and often used at the mm scale) is often lower than the

techniques described in the previous two paragraphs. Nevertheless, research is under way to adapt zymography to the micro-scale environment and resolutions (< 100µm), by preparing thin sections from resin impregnated samples and analysing these with epifluorescence microscopy (Ghaderi et al., 2020). Recently, zymography has been adapted to also visualise oxido-reductases (Khosrozadeh et al., 2022), and time lapse imaging has been implemented allowed to better describe enzyme activities by accounting diffusion losses and the kinetics of signa development (Guber et al., 2021). It has, for example, allowed the visualisation of the distribution of phosphatases at the rhizobox scale (Spohn and Kuzyakov, 2013; Spohn et al., 2013; Razavi et al., 2016; Liu et al., 2017) (Fig. 9). Heitkötter and Marschner (2018) demonstrated that microbial hotspots of activity, revealed by zymography, represented less than 3% of cores surface area, but after spraying glucose onto the surfaces of interest, the enzymatic activity dramatically increased outside of these initial hotspots, demonstrating that apparently dormant areas of soil are easily stimulated.

There are few examples of correlative imaging using zymography with other methods, even if it has been used with X-ray tomography and 14 C imaging (Kravchenko et al., 2019; Becker and Holz, 2021). However, this technique has enabled the identification of hotspots of activity of numerous enzymes (β -glucosidase, α -glucosidase, xylanase, phosphatase, chitinase, peroxidase etc.) (Kuzyakov and Blagodatskaya, 2015; Heitkötter and Marschner, 2018; Razavi et al., 2019).

4.3. Assimilation of substrates

Physiologically active heterotrophic microorganisms assimilate organic substrates and, therefore, isotopically labelled organic substrates can be used to label and visualise active microorganisms within the soil architecture. It requires a combination of methods for detecting both the microorganisms and the isotopically labelled constituents, the superposition of which can reveal active cells, as described below.

Radioisotope labelled organic substrates in combination with autoradiography have been used to measure the spatial distribution of substrate assimilation activities. When placed in contact with an emulsion or a photographic film, the distribution of radioactive source is recorded, such as zones of root exudation (Holz et al., 2019), rhizosphere hotspots (Becker and Holz, 2021) and assimilation zones (Torsvik and Øvreås, 2002) at scales ranging from cmmm. It has allowed, for example, the visualisation of the transfer of ¹⁴C photosynthates from

Pinus roots to mycorrhizae in rhizoboxes with ¹⁴C (Leake et al., 2001), ³²P (Lindahl et al., 1999) and ³³P (Wu, 2014), and to locate methanotrophs assimilating ¹⁴C-CH₄ according to soil depth and aggregates size in an afforestation chronosequence on subalpine pasture (Karbin et al., 2017). Autoradiography can be applied at much smaller scales when combined with electron microscopy, i.e. micro-autoradiography, but this has seldom been used with soils (Lee et al., 1999; Rogers et al., 2007). Radioisotope based methods are very sensitive, meaning that low substrate concentrations can be used, but they demand specific safety procedures, unlike stable isotopes.

Similar information can be obtained using substrates labelled with stable isotopes and SIMS. Cliff et al. (2002) visualised the assimilation of ¹⁵N and ¹³C by bacteria and fungi grown in a model soil system made of kaolinite with TOF-SIMS. SIMS techniques have been used to monitor the fate of labelled organic root exudates, thus highlighting microbial hotspots in the rhizosphere and shedding light on rhizosphere functioning through the visualisation of the transfer of organic carbon from the plant to rhizosphere microorganisms (e. g. Vidal et al., 2018). They have also been used to visualise microorganisms that have assimilated N after ¹⁵N nitrogen fertilizers addition to soil, as shown in Fig. 8 (Herrmann et al., 2007b).

No technique can visualise microorganisms and their activities in 3D at present. However, positron emission tomography (PET) is a promising approach, as it can locate isotopes in 3D in a sample. Garbout et al. (2012) used this method to visualise the assimilation of ¹¹C-CO₂ by a plant. However, as ¹¹C is very unstable (half-life of 20.4 min), the observation time was short. Although commonly used in the medical sciences field, this technique, which does not require any particular sample preparation, is nearly un-used in soil science. In addition, the resolutions obtained so far are very low, i.e. in the order of a mm, which mean that this methodology not really suitable for investigating microorganisms in soils.

The methods presented so far target assimilation sites. In order to establish links between the spatial distribution of microorganisms and the distribution of their activities in the soil structure, several methods need to be combined. In some studies microorganisms were located with FISH labels and their activities with micro-autoradiography FISH (MAR-FISH) (Lee et al., 1999; Ouverney and Fuhrman, 1999; Torsvik and Øvreås, 2002; Nubel et al., 2002; Musat et al., 2012). In others, such as in Schlüter et al. (2019a), fluorescence micro-

scopy and electron microscopy were used to locate microorganisms in a soil structure that was visualized with computed microtomography. In this study nanoSIMS was also used to identify microorganisms that had assimilated the organic substrate. Similarly, the use of FISH followed by the use of nanoSIMS (FISH-SIMS) made visualising microorganisms that had assimilated stable isotopes possible (Musat et al., 2012; Mueller et al., 2013; Schurig et al., 2015). The FISH methods can be adapted using very electronegative halogen markers (I, Br) (HISH-SIMS) which can be detected in SIMS directly (Li et al., 2008; Musat et al., 2012; Mueller et al., 2013). However, to our knowledge, this technique has not yet been applied to soil samples. Finally, the cross-use of SIMS with electron microscopy is frequent (Watteau and Villemin, 2018). It facilitates the prior identification of areas of interest and makes it possible to obtain information related to microhabitat (Mueller et al., 2013).

5. Methodological challenges and solutions

All observation methods have different advantages and disadvantages and the choice of one rather than another will depend on the scientific question to be answered. Different critical points must be considered before choosing the most relevant method: the sample preparation, possible combination of different methods, existing imaging treatments and representativeness of these images with regards to the scientific question.

5.1. Sample preparation methods

The sample preparation impose a number of constraints, which need to be taken into account when choosing a method of analysis. First, observation methods often, though not systematically, require the production of thin sections and/or impregnation for the soil structure to remain undisturbed. The procedure used must be adapted to the sample and the objective of the study, and all samples and controls should be prepared in equivalent conditions and technical controls to assess the quality of preparations.

5.1.1. Fixation and dehydratation

When fragile objects need to be kept in their original state or microorganisms activity needs to be stopped, a fixation can be performed, by cross linking proteins. Fixatives often are toxic compounds, which require to take precautions to ensure user safety. Glutaraldehyde solutions are frequently used for this purpose (Tippkötter et al., 1986; Altemüller and Van Vliet-Lanoe, 1990; Nunan et al., 2001; Elsass et al., 2008; Vidal et al., 2018) but avoided with FISH because glutaraldehyde impedes hybridization of nucleic acids (Solovei, 2010), making formaldehyde the fixative of choice with FISH (Schmidt et al., 2012; Schmidt and Eickhorst, 2014). Fixatives can also have substantial influence on the samples' chemistry, which is susceptible to modify the information acquired with Raman spectra for example (Bandara et al., 2021). Thus, control samples are needed to assess their effect. Cryo-fixation, using liquid nitrogen, propane or ethane for rapid freezing without ice crystal formation is another fixation option. Also, rewetting soil samples with fixatives should be performed as carefully as possible as it may displace the microorganisms and modify particles organisation.

Dehydration, which is essential for some methods (TEM, SEM and nanoSIMS for example) can alter soil structure (Tippkötter and Ritz, 1996) because of clay shrinkage. Different protocols exist, such as air drying (Gutiérrez Castorena et al., 2016), freeze-drying (Tippkötter et al., 1986) or dehydration with water-ethanol or water-acetone gradients (Tippkötter et al., 1986; Nunan et al., 2001; Elsass et al., 2008; Mueller et al., 2012). Water-acetone exchanges have proven their efficiency in limiting impacts on soil structure (Altemüller and Van Vliet-Lanoe, 1990). However, biological material must be fixed prior dehydration with water-acetone gradients in order to avoid cell damage (Elsass et al. 2008). Water-ethanol exchanges are believed to minimize cells damage because ethanol is a weaker solvent than acetone (Bandara et al., 2021). Following cryo-fixation, freeze-substitution allows dehydration by substituting frozen water with a solvent at sub-zero temperatures. Supercritical drying is also efficient for drying soil containing natural "gels" (allophanic soils) (Woignier et al., 2005, 2008; Calvelo Pereira et al., 2019).

5.1.2. Staining

Stains (fluorochromes or contrast agents) used in particular with epi-fluorescence, confocal microscopy and electron microscopy for observing microorganisms may have limited efficiency. Many stains, and particularly in undisturbed soil, may have difficulty penetrating the sample and reach their target (Peth et al., 2014). Testing this by quantifying the gradi-

ents of the stain in soil aggregates as performed by Zheng et al. (2020) is hence very useful. In addition, non-specific staining, as well as autofluorescence can also occur, depending on soil characteristics. Therefore, controls should be considered whenever possible with unstained thin sections or samples. Stains can be applied on the surface of previously prepared thin sections (Elsass et al., 2008; Gutiérrez Castorena et al., 2016; Juyal et al., 2019; Schlüter et al., 2019a), which is useful when used in combination with surface imagery methods.

5.1.3. Embedding

The embedding step can present some difficulties with ensuring that the resins saturate the entire porosity and with avoiding the presence of air bubbles, which lead to fragile samples for subsequent processing (cutting, polishing). The relatively high viscosity is at cause here. These problems are particularly true for large samples, which require long periods under high vacuum conditions. A solution is to perform multiple additions of resins at different dilutions with acetone making it less viscous, thus ensuring that the porosity is better filled (Nunan et al., 2001; Elsass et al., 2008; Mueller et al., 2012; Vidal et al., 2018).

In addition, the choice of embedding resin should suit the study objectives. For example, nanoSIMS requires the use of special resins such as Aradite-502 (Mueller et al., 2012; Vidal et al., 2018) for their resistance to the applied pressures (Herrmann et al., 2007a; Mueller et al., 2012). Any study targeting the localisation and characterisation of soil carbon requires the use of resins that can be differentiated from the sample (Mueller et al., 2012; Vidal et al., 2018; Vergara Sosa et al., 2021). For the specific case of nanoSIMS, even though most samples are embedded, direct deposition of samples is also possible as long as they are very flat (< 1μ m topography for natural abundance and 30 μ m for stable isotopes enriched samples) (Mueller et al., 2013).

5.2. Image acquisition and processing

When acquiring images, it is important to note that very often the operator will have an impact on the quality of the final image. Indeed, parameters such as time of exposure or aperture, that are generally set manually, are very important and should be taken into account. Regardless of the choices made when establishing a protocol, the different steps should be systematically recorded as, failing that, variations in choices from one operator to another and

from a day to another may have an impact on the results.

 The image processing steps are also of fundamental importance for the extraction of relevant information. It allows analysing the image and thus distinguishing objects, enumerating them, but also extracting data about their size and shape and characterizing their spatial distribution. The images obtained with the different methods present characteristics which differ according to the modes of acquisition (size and depth fields, resolutions, voxels or pixels, number of channels, bit depth) and which require different image processing and analysis strategies. Starting from the raw image, the identification of different soil constituents can require significant expertise for choosing the optimal processing procedure. Indeed, observation methods based on shapes and colours for example can generate errors and also lead to different conclusions from one observer to another (Baveye et al., 2010). Replications of observations at different dates and by several observers may be useful (Kleber et al., 2003; Chenu and Plante, 2006), but it is costly and time consuming.

Image processing often requires segmentation, correction and filtering steps, amongst others (Schlüter et al., 2014; Roose et al., 2016; Withers et al., 2021; Jeckel and Drescher, 2021). To avoid operator effects, image analyses should be automated as much as possible. The choice of a segmentation threshold is subjective and depends on the observer. Therefore, most studies recommend the use of a fully automatic thresholding. However, thresholds must be chosen correctly depending on the targeted object characteristics and the type of sample (Iassonov et al., 2009; Hapca et al., 2013; Bilyera et al., 2020; Pot et al., 2020). The same is true for all stages of image processing (Kaestner et al., 2008; Houston et al., 2013) and this should always be carefully described in any scientific publication. There are more and more methodological publications describing specific methods to perform total or locally adaptative methods (Sauzet et al., 2017; Gao et al., 2019; Bilyera et al., 2020). Machine learning and especially supervised pixel classification have been increasingly found to provide reliable classification of different minerals and organic matter in nanoSIMS images (Steffens et al., 2017; Vidal et al., 2021), in SEM microscopy of shales (Wu et al., 2019) and in hyperspectral microscopy where bacteria were classified ex situ (Liu et al., 2021). We can thus expect that it will help future research to extract more easily the images interest objects.

All samples, including controls, should be acquired in the exact same conditions with the same parameters and the image processing should follow the same steps, otherwise they cannot be considered equivalent and compared.

5.3. Image representativeness

As the representativeness of images is linked to their size, the question of the representativeness of the images arises when the imaging resolution is high. The study areas must be chosen with care and otherwise it can lead to focusing on exceptional areas, which can result in incorrect conclusions being drawn. Then it may be difficult to link observation results to the overall functioning of the soil. Microbial densities are much higher in hot-spots than un bulk, which has consequences on the choice of imaging methods because the necessary field to describe processes inside these environments has to be considered. NanoSIMS is adapted to the study of the rhizosphere and detritusphere as the probability of encountering microorganisms that have assimilated isotopically-labelled resource is relatively high, contrary to the bulk soil (Védère, 2020).

Furthermore, it should not be forgotten that many of the techniques available provide a 2D image of a 3D object. For example, the organisation of porosity in 2D images is not representative of the 3D pore network. To overcome this, correlative imaging can help to reach a more complete understanding of the spatial organisation of soil constituents soil functioning, as was done by Schlüter et al. (2019a).

5.4. Combination of observation methods

5.4.1. Compatibility of the different methods

Reviewing the literature shows that there is no single method to study both the spatial and temporal fate of soil microorganisms and their activities in their microhabitats. This is why understanding the functioning of the soil as a whole requires a combination of different methods and correlative-imaging. Recent publications demonstrate the great potential of this approach. The first step is to check that the methods are compatible. When the analyses can be carried out on the same samples, then the question of the order of the methods to be used must be considered.

Indeed, the methods previously described require specific preparations and are more or less destructive, such as SEM or SIMS, which require the samples to be covered with a gold layer or which sputter the sample, respectively. These methods should therefore be implemented in the final stages of observation. Other methods such as tomography or zymography are advantageous since, not requiring specific prior soil preparations, they allow for

multiple observations of the same object without such concerns. Moreover, X-ray computed tomography has negligible effects on root growth (Zappala et al., 2013), archaea and bacteria and their functioning in soil (Bouckaert et al., 2013; Schmidt et al., 2015).

 Altogether, we suggest to favour first the methods which do not require any particular preparation (X-ray or neutron tomography for example), secondly the observations using flat probes or gels (zymography or planar optodes) and requiring only a cut in the material, then the methods requiring the preparation of thin sections and finally all the methods whose operation consists in bombarding the surface of the object and which can irreversibly damage the surface of the objects (SEM, nanoSIMS). Other methods, such as Raman scanning can be harmful for the resin (Bandara et al., 2021). All these particularities should be considered in order to define the best pathway to do correlative imaging. For example, Bandara et al. (2021) successfully applied a thin section protocol in order to combine and correlate images from nothing less than six different imaging methods by using LR white resin. Otherwise, the methods can be applied to different samples prepared under the same conditions, but the number of samples to be prepared can quickly become huge.

A challenge to overcome with correlative imaging is related to the superimposition of two images of the same exact zone acquired with different techniques. Solutions are to use a single same sample holder including a coordinate system (Bandara et al., 2021) or to target landmarks directly present in the samples (Juyal et al., 2020; Schlüter et al., 2019a) to find back the region of interest. For surface imaging methods, it is also possible to artificially mark the sample surface to create references helping at finding previously targeted region of interest using the electron microscopy beam (Bandara et al., 2021). Moreover, the difficulty increases when one wants to combine optical techniques, which imply a field depth, and surface techniques which have none. Indeed, some features below the sample surface may be visualized with field depth techniques but not with surface techniques. For example, if the zone of interest is not on the sample surface, although visible by light microscopy, then, it will not be possible to reach it with surface techniques such as scanning electron microscopy or nanoSIMS (except if the object is very close to the surface where you can expect to reach it by an abrasion of surface). This limitation makes the correlative imaging not trivial. However, powerful tools, such as correlia, a pluggin developed for imageJ, help correlating images from different methods (Bandara et al., 2021; Rohde et al., 2020).

5.4.2. Upscaling issues

Focusing on microorganisms in their habitat results in considering a very small field of observation compared to the global functioning of the soil and raises the question of the upscaling. How can we account for these observations on larger scales? Are microscale observations representative of the phenomena described on a larger scale? Some studies recommend the creation of mosaics of high-resolution images in order to obtain a wide image of the samples. Others advice to work with an average volume which consists of averaging microscopic description into a Representative Elementary Volume (Baveye et al., 2018) or a homogenisation which assumes that structure is sufficiently periodic to be considered as composed of repeated units for variables as porous structure (Roose et al., 2016). The use of these options is still under discussion considering the trade-off between their potential and the considerable loss of data that they can involve (Baveye et al., 2018).

The combination of observation methods can also address this problem with complementary observations at different spatial scales. Such studies have already been carried out, such as that of Schlüter et al. (2019), who used correlative imaging with X-ray tomography at a centimetre scale, optical microscopy and fluorescence at a mm scale, then electron microscopy on the scale of a few hundred microns and finally nanoSIMS on the tens of microns scale in order to describe the microhabitat in a decaying leaf detritusphere combining structural, geochemical and biological data. They showed that if bacteria were mainly present in pores < 10 μm, they were preferentially concentrated near macropores and organic matter. Juyal et al. (2019, 2020) also combined observations at mm (X-ray tomography, light and fluorescence microscopy) and micrometric (light and fluorescence microscopy) scales to study the localisation and spread of microorganisms in the soil structure and in particular as a function of porosity indicators. Keiluweit et al. (2012) characterised the soil-microorganisms interaction by combining nanoSIMS and NEXAFS with STMX allowing them to propose a conceptual model of the fate and transformation of fungal cell wall compounds in soil and its relation with mineral particles. Bandara et al. (2021) acquired images allowing description of the rhizosphere using six different imaging methods, i.e., light microscopy, epifluorescence microscopy, HIM microscopy, Tof-SIMS, nanoSIMS, SEM-BSE and -EDX and confocal Raman spectroscopy.

2D (SEM-EDX) and 3D (X-ray CT) approaches may be combined to build 3D chemical maps of soil samples based on a statistical approach such as that proposed by Hapca et al. (2015). Similarly, Anderson et al. (2020) combined 2D (FIB-SEM) with 3D images (TXM,

transmission X-ray microscopy) using machine learning on sediment rock materials. Another approach is to merge 2D images acquired at 3 different scales to create one final image with a better resolution and multiscale porosity information (Karsanina et al., 2018). The development of computer calculation capacity also allows thinking that reconstruction from 2D to 3D images will be facilitated in the future.

5.5. Perspectives

5.5.1. Methodological developments. Methods of the future.

There are a number of potentially useful methods that are not or rarely used in soil science for a variety reasons. μ X-ray fluorescence CT for example (Fig 2) is seldom used because of necessary developments and improvements regarding the issue of energy attenuation (Bleuet et al., 2010; Hapca et al., 2015; Roose et al., 2016). Nuclear magnetic resonance (NMR) is little compatible with soil samples that contain paramagnetic particles (Schmidt et al., 1997; Baveye et al., 2018). Neutron radiography and positron emission tomography (PET) scanning do not yet have sufficient spatial resolutions for the study of soils at the scale of microorganisms. A number of methods therefore still have technical limitations for applying them to the soil environment and to the microorganisms' spatial scales.

Other methods, however, are underused, even though they would be appropriate for investigating soil microorganisms in their habitats. Super-resolution methods which give access to unprecedented resolutions in optical microscopy could be used, e.g. to study microorganisms and soil particles interactions. Super resolution techniques such as STED (stimulated-emission depletion microscopy), SIM (structured illumination microscopy) or SMLM (single molecule localisation microscopy) have resolutions down to a few tens of nanometers (Turkowyd et al., 2016), but have not yet been used to visualise soil microorganisms. With STED, the sample is scanned using two lasers. The first one stimulates the emission of fluorescence subjected to diffraction on a targeted zone and the second, in the form of a "donut", uses a de-excitation beam and comes to switch off a part of the emission of the fluorochrome leaving only its central emission source. This process counterbalances the effects of diffraction and means that resolutions of between 20-70 nm can be achieved. As this method is based on activation and successive repeated extinction of fluorochromes, all fluorochromes must be photo-stable (Turkowyd et al., 2016).

The SIM method uses a mask containing a pattern of lines regularly spaced with a frequency, orientation and phase known, rotating above the sample to achieve a controlled illumination of the fluorochrome. The exploitation of the "moiré" effect obtained by this approach is used to reconstruct the structure of the observed object at a resolution of approximately 80-100 nm. One advantage is that it does not require any staining (Turkowyd et al., 2016).

Finally, the SMLM (single molecule localisation microscopy) methods use the blinking of certain fluorochromes to isolate them one by one via their isolated emissions, allowing very small resolutions to be obtained between 10 and 50 nm. Here the fluorochromes must be photo-convertible and photo-stable (Turkowyd et al., 2016).

These last methods are very sophisticated and the density of the staining, the type and the size of the fluorochrome used strongly impact the final result (Huang et al., 2009). The SIM and SMLM methods also require a time-consuming post-processing of images that hinders their use. These methods are used in cell biology and chemistry but not yet in soil science.

Similarly, FIB, BIB and SBF are rarely used while they could provide much 3D information on objects not accessible to the Xray CT resolutions. Other methods of the future are to use transparent soil media to overcome soil's opacity and thereafter use 3D imaging methods not previously suitable to study undisturbed soil. Sharma et al.(2020) tested different artificial transparent soils to obtain 3D images using CLSM and Raman microscopy. They recorded microorganisms position and carbon uptake and were able to observe a higher activity of bacteria at the vicinity of dead fungal hyphae following a drying/ rewetting cycle. Yang et al. (2021) by using such transparent model of soil could observe under confocal microscopy the dynamics of organic matter sorption and protection by clay and the effect of microbial enzyme activities on this protection.

5.5.2. Dynamic observations

While the importance of hot moments is recognized for soil biogeochemical processes, still few visualisation studies address temporal dynamics. The first reason is that visualisation methods for soil microorganisms in their microhabitat are time consuming and expensive. Further, if the visualisation methods are destructive, it requires to prepare and process a large number of replicate samples. Nevertheless, a few methods allow non-destructive time monit-

oring visualisation of soil. X-ray tomography, neutron radiography and zymography allow repeated observations on the same sample at different dates. However, the scanning time of these methods is sometimes too long for dynamics to be captured. For example the direct monitoring of water flow through the porosity remains difficult to study because CT tomography could not scan fast enough to capture the water flow (Baveye et al., 2018). Neutron tomography has improved enough to allow fast imaging and monitoring of water uptake in soil with 1 minutes iterations as recently demonstrated by Tötzke et al. (2021) and high speed synchrotron-based X-ray computed tomography allowed few seconds iteration scanning but it is currently not used in soil sciences (Berg et al., 2013).

5.5.3. Precious experimental data for modelling

It is now acknowledged that the spatial distribution of microorganisms, trophic resources, air and water in the soil architecture at the microscale largely determine the biogeochemical fluxes at the macroscale. A new generation of biogeochemical models is emerging that are based on an explicit description of soil structure and water distribution at the microscale (Pot et al., 2015; Kemgue et al., 2019; Ruiz et al., 2020; Pot et al., 2021), in order to more accurately forecast the dynamics of organic matter in soil under a wide range of climate and management conditions (Monga et al., 2008; Portell et al., 2018; Ruiz et al., 2020). In these microscale models, the distribution of microorganisms and trophic resources are up to now ascribed a priori (Pot et al. 2021), and not based on experimental data. However experimental data, such as the distribution of microorganisms (Raynaud and Nunan, 2014; Juyal et al., 2019; Schlüter et al., 2019a) or organic matter (Peth et al., 2014; Quigley et al., 2018; Schlüter et al., 2022), are necessary input data for calibrating the models and assessing their performance. The many technological advances and developments in the field should allow for rapid progress in this area.

5.5.4. Imaging soils at the microscale for soil microbial ecology

Despite the numerous constraints and difficulties, these methods, when combined, offer major perspectives for characterizing the location of total and active microorganisms in their environment and better understanding soil functioning. Considering the scant available knowledge on soil microorganisms distribution and activities in their microhabitats, there is an open field for microbial ecology research using imaging at the microscale. We can imagine different workflows providing valuable data and an overall better understanding of soil functioning at the microscale and will present two exemples.

A first example concerns the detritusphere. In the detritusphere, soil moisture controls the transfers of organic solutes, enzymes and microorganisms, the accessibility of microorganisms to their substrates and modulates biodegradation (Védère et al., 2020). X-ray computed tomography can allow to describe the organisation of the porosity in the vicinity of labelled plant residues and the porosity directly connected to it (distances to the residues, connectivity, tortuosity, direction...). Following the pores description, thin sections can be produced on the same samples and the spatial distribution of microorganisms in the vicinity of previously targeted porosity and at increasing distances from the residues could be assessed using epifluorescence microscopy. Finally, nanoSIMS can allow to observe C and N transfers from the residue to the soil and to locate microbial assimilation hotspots at this interface. Combining several imaging methods can address questions such as the influence of plant residues on soil porosity, whether soil moisture modifies the spatial distribution of microorganisms decomposing plant residues and the interplay between mineralisation and stabilization of carbon from these plant residues.

A second example addresses the possible influence of biochar addition of rhizosphere functioning. Biochars are increasingly used to improve soil properties like water retention or cation exchange capacity of soil. They present a high porosity that can be used by microorganisms as habitats. Rhizosphere development and functioning is influenced by biochar addition (Atkinson et al., 2010). Unfortunately no clear description of microorganisms spatial distribution relative biochar particle exists. X-ray computed tomography can describe soil porosity inside the rhizosphere artificially produced in a rhizobox. Then, the surface of the box could be exposed to direct zymography in order to give information about enzymes diffusion in the porosity. Once those measurements being done, thin sections of soil at the vicinity of root and biochar particles could be prepared to localize microorganisms using epifluorescence microscopy. Such a workflow could address questions like: Is rhizosphere porosity influenced by biochar addition in soil? Do biochar particles provide new habitats for microorganisms and modify their spatial distribution? Are enzyme activities affected by biochar in the rhizosphere?

6. Conclusion

Imaging techniques have developed significantly in recent years, have become increasingly accessible and are attracting unprecedented attention in the soil's scientific community. Microorganisms can be located and identified at fine scales with more accuracy and confidence using modern visualisation methods. The mapping of certain microbial activities, such as enzyme activities and substrate assimilation, in the soil structure is also possible. The improvement of existing techniques has made detailed descriptions, at better resolutions, possible and the development of new technologies opens new horizons. However, technical development is still necessary, particularly to reduce analytical time and costs and to optimise the combination of different methods. Moreover, imaging techniques often need a high-level expertise to be used properly and these are rarely present at a single location. Therefore, a better understanding of soil microhabitats will involve multi-disciplinarity and collaborative studies. The information obtained on the spatio-temporal evolution of microorganisms and their activities in the soil structure should make it possible to improve knowledge and lead to a better understanding of soil functioning.

Conflict of interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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