

Boulder fall activity in the Jezero crater, Mars

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12	Key Points:
13 14	• We report on previously unrecognized boulder fall activity in Jezero crater – landing site of Mars 2020 Perseverance rover
15 16	• Hazardous boulders bigger (length >2.7 m) than the rover have recently fallen on the front of the western delta deposits within the crater
17 18	• Recently fallen boulders may have freshly exposed surfaces, which can be ideal targets for astrobiological investigations by the rover
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25 Abstract

Jezero crater is the landing site for the Mars 2020 Perseverance rover. We report 63 boulder fall 26 27 tracks within the crater using High Resolution Imaging Science Experiment (HiRISE) images. The boulder tracks have both fresh and faded morphologies similar to those reported elsewhere on 28 Mars, but reported for the first time in Jezero crater. We combine observations from 16 boulder-29 tracks on the western delta deposit with 47 in the surrounding regions to infer possible process(es) 30 31 of boulder destabilization, which can be tested with rover observations. This newly-found hazard should be taken into account for rover operations. Boulders associated with tracks are geologically 32 "recent falls", so it is possible that the surfaces of these boulders may provide an opportunity to 33 sample material less exposed to radiation than other rocks at the martian surface and could be ideal 34 35 targets to analyze for organics.

36 Plain language summary

We have found tracks left by falling boulders within Jezero crater on Mars, which is the landing 37 site of the Mars 2020 Perseverance rover. The boulder tracks are observed on the slopes of the 38 delta deposits that the rover will explore, on the wall of Jezero crater and on the wall of a ~ 2 km 39 40 diameter crater on the floor of Jezero crater. Boulders of a few meters in size have fallen down moderately steep slopes, moving distances from a few tens of meters to nearly a kilometer. The 41 42 boulder tracks appear both fresh and degraded. The positions of the boulders and the appearance of the tracks have not been observed to change, but the simple existence of the tracks suggests that 43 44 the boulder falls are recent and could be ongoing. Our results reveal a new location on Mars where boulder fall activity is observed and indicate a new hazard to be taken into account by the rover 45 team. These recently fallen boulders may also have surfaces less exposed to radiation and could 46

be ideal candidates for obtaining samples that could be analyzed for signs of ancient life. Suchsigns usually degrade rapidly with exposure to radiation.

49 **1 Introduction**

Boulder falls are energetic events and are known to occur on Mars (e.g. Roberts et al., 2012; 50 51 Brown and Roberts, 2019; Kumar et al., 2019). Hence, understanding if such a process could 52 occur within the landing site of a roving mission on Mars is a key factor when planning operations to ensure the safety of the rover (e.g. Golombek and Rapp, 1997; Golombek et al., 2003, 2008, 53 54 2012). Recent boulder falls are characterized by a track marked in the regolith in the wake of the 55 boulder. The tracks are continuous and/or discontinuous depressions of a similar width to that of 56 the boulder with slightly raised edges and are produced by rolling, bouncing and rotation of the 57 boulders (van der Bogert and Plescia, 2014). Over time, the boulder tracks may be erased by the deposition of dust (Beyer, 2012), infilling by aeolian deposits, and/or aeolian abrasion. 58

Isolated boulders are often key science targets for roving missions, as they are more easily accessible than outcrops and they can reveal information both on the mineralogy and geochemistry of the surface (e.g. Morris et al., 2000; Hamilton and Ruff, 2012; Squyres et al., 2006; Stolper et al., 2013; Pajola et al., 2017; Wiens et al., 2020). Boulders with tracks have the additional advantage that they can be traced back to the specific source region and hence enable a roving mission to sample rocks from the locations that would generally be difficult to reach (Schmitt and Cernan, 1973; Senthil Kumar et al., 2016).

The landing site of the Mars 2020 Perseverance rover (Jezero crater; centered at 18.42°N, 77.67°
E) is a proposed paleolake with two inlet channels feeding the crater from the north and west
(Figure 1a; Fassett and Head, 2005). The approximate position of the landing ellipse of the rover

69 and the potential traversable routes of the rover to regions of interest within the landing site 70 surrounds the partially eroded delta deposits (also known as western delta deposit) at the mouth of the western inlet channel (Figure 1b; Fassett and Head, 2005; Goudge et al., 2015, 2017; 71 Farley, 2018; Mangold et al., 2020; Williams et al., 2020). Presence of aeolian landforms such 72 as yardangs and transverse aeolian ridges (TARs) has been reported over the volcanic deposits on 73 the front of the delta deposits (Chojnacki et al., 2018; Day and Dorn, 2019), and aeolian erosion 74 has been suggested to have extensively eroded the delta deposits (Fassett and Head, 2005; Schon 75 et al., 2012). In addition, the front of \sim 50-60 m high western delta deposit hosts numerous boulders 76 77 on its top surface, over its talus slope and at the base of the talus slope. Numerous boulders are found scattered across the wall of Jezero crater. If a boulder has detached from the face of the delta 78 front or the crater wall in the very recent past at geological time scales, it is likely that the boulder 79 would have produced a track as it rolled/bounced over the slope materials under the force of 80 gravity. To our knowledge, boulder fall activity within Jezero crater has never been examined in 81 detail. In this study we used 13 Mars Reconnaissance Orbiter (MRO) HiRISE images (0.25 or 0.5 82 m/pixel) and 4 HiRISE stereo-pair derived digital terrain models (DTM) (1 m grid spacing) 83 (McEwen et al., 2007; Kirk et al., 2008) (Table S1 and S2, Figure S7) to search for boulder fall 84 tracks, and analyze the morphology of identified boulder tracks. We examined all slopes for 85 boulder tracks, including the flanks of delta deposits, Jezero crater wall, the walls of any other 86 craters on the floor of Jezero, on the flanks of mounds adjacent to the Jezero crater wall, and on 87 the flanks of any other topographic or hill-like features on the Jezero crater floor. To test whether 88 boulder falls are currently active, we examined images with a long time separation (Table S3) to 89 seek signs of change in the boulder position and the boulder track morphology (for example, fading 90 91 of boulder tracks). We used the empirical shadow angle method (Lied, 1977; Evans and Hungr,

92 1993) to analyze the runout distance of boulder falls on the delta deposits. Shadow angle is the 93 angle between the horizontal line from the talus apex and the line that connects the talus slope apex 94 with the most distant boulder fall (Duszynski et al., 2015). Please see the supporting information 95 for details on the data and methods. Finally, we discuss the possible process of boulder 96 destabilization and explore the potential scientific implications for the Mars 2020 Perseverance 97 rover.

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99 **2 Results**

We have identified 63 boulder tracks within Jezero crater (Dataset S1). 16 boulder tracks are on 100 the front of western delta deposits and the remaining 47 boulder tracks are in other regions of the 101 crater (43 boulder tracks on/near the wall of Jezero crater and 4 boulder tracks on the wall of a ~2 102 km crater on the floor of Jezero crater) (Figure 1b-1d). We did not identify any convincing 103 changes in the position of the boulders and tracks have not faded over the ~12 years of available 104 images (Figure S1) (Table S3). We found that the tracks reported in this study do not 105 appear/disappear with respect to the change in the illumination conditions (such as incidence angle 106 107 and subsolar azimuth) (Table S1).

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Figure 1. (a) Mars Orbiter Laser Altimeter (MOLA) and High Resolution Stereo Camera (HRSC) 114 blend color elevation map over the Context (CTX) camera mosaic of the region around Jezero 115 crater (Dickson et al., 2018). The location of Jezero crater (star) is shown on the global view of 116 117 Mars (top-left) using Viking color imagery. The locations of panels b-d are shown by black boxes. The Mars 2020 Perseverance rover landing ellipse is in yellow. (b) Boulder fall tracks identified 118 on the western delta deposits and western wall. (c) Additional boulder tracks identified on the 119 western wall. (d) Boulder tracks identified on the wall of a ~2 km crater on the floor of Jezero 120 121 crater. Background is CTX mosaic in panels b-d. CTX image mosaic credits: NASA/JPL/MSSS/Caltech Bruce Murray Lab. 122

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124 **2.1 Boulder falls on the western delta deposits**

We have identified 14 faded and 2 fresh boulder tracks on the front of western delta deposits (Figures 1b, 2a). The morphology of faded boulder tracks varies widely between the tracks (Figure 2b-d). For five of the faded boulder tracks, depressions are still identifiable, but for the 9 other faded boulder tracks only the track margin is discernible. We interpret this difference in boulder track morphology to be related to the extent to which they are mantled by dust, and therefore could give us an indication of the relative ages of the tracks. Fresh boulder tracks are characterized by distinct depressions (Figure 2e).

132 In general, all the boulder tracks end on the talus slope with an average length of 31 m (Figures 2b-c, 2e, 2g). The boulders associated with the tracks measure <4 m in length (mean: 1.7 m, 133 134 standard deviation: 0.8 m) (Figure 2f). Widths of boulder tracks vary from 1 m to 3.3 m (mean: 135 1.8 m, standard deviation: 0.7 m) (Figure 2f). The average slope between the start and end of the tracks ranged from 16° to 29° with most between 17° and 24° (mean: 22°, standard deviation: 3.8°) 136 137 (Figure 2h). The longest observed boulder track (67 m) is associated with a big boulder (length: 3.7 m) that has travelled beyond the base of talus slope, and the track superimposes aeolian 138 landforms (e.g. TARs/dunes) (Figure 2d). Compared to the other fallen boulders around the delta 139

140	deposits, this big boulder is double their average length (1.7 m). In comparison to the 63 boulder
141	falls reported in this study, this big boulder is nearly 3 times their average length (1.3 m). Both
142	longer (44 m to 67 m) and shorter (16 m to 30 m) boulder tracks have an average slope between
143	16° and 21°. The average slope of the boulder source regions varies from 5° to 47° (mean: 23°,
144	standard deviation: 13.6°) (Figure 2h).
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Figure 2. (a) Overview of the front of western delta deposits within Jezero crater on which fresh 164 (green) and faded (red) boulder fall tracks are found. Locations of panels 2b-e are labeled. (b-d) 165 Detailed view of faded boulder tracks (arrows). In Figure 2d, a custom stretch was applied to 166 167 enhance contrast of the boulder track. (e) Detailed view of a fresh boulder track (arrow). Scale is same in panels b-e. HiRISE Image ID (a-e): PSP 002387 1985, credits: NASA/JPL/University of 168 Arizona. (f-h) Morphometric characteristics of boulder fall on the front of western delta deposits. 169 Histograms of (f) boulder length and boulder track width, (g) boulder track length, (h) boulder 170 171 track slope and boulder source slope. The measured values plotted in f-h are rounded to 1 decimal 172 place.

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174 **2.2 Boulder falls in the other regions of the crater**

40 boulder tracks out of the 47 observed tracks (Figures 1b-d and 3a-b; 36 tracks on/near the 175 Jezero crater wall and 4 tracks on the wall of a ~2 km crater on the Jezero crater floor) are classified 176 as faded tracks (Figures 3f-g and 3i). The remaining 7 boulder tracks (on/near the Jezero crater 177 wall) have a fresh morphology (Figures 3c-e and 3h). The faded tracks lack identifiable 178 depressions, and appear infilled or mantled by dust, whereas the fresh tracks are characterized by 179 distinct depressions. The boulders observed here (mean size: 1.2 m, standard deviation: 0.6 m) 180 have formed tracks of length and width in the range 33 m - 900 m (mean: 190 m, standard 181 deviation: 155 m) and 0.5 - 2.6 m (mean: 0.9 m, standard deviation: 0.5 m), respectively. The 182 slope of the boulder source regions is slightly steeper (mean: 24°, standard deviation: 4°) than the 183 184 slope of the boulder tracks (mean: 20°, standard deviation: 2.1°).

We find that the orientation of the slopes on which the boulder tracks are found are preferentially:
(1) southeast (Figure 3m), (2) east-southeast (Figure 3n), and (3) west-southwest/east-northeast
(Figure 3o).





Figure 3. (a-b) Overview of the fresh (green) and faded (red) boulder tracks newly identified
on/near the wall of Jezero crater. HiRISE Image ID (a): ESP_037119_1985. HiRISE Image ID
(b): PSP_001820_1985. Locations of panels 4c-i are shown. (c-e): Detailed views of example

faded boulder tracks (arrows). (f-i): Detailed views of example fresh boulder tracks (arrows). Scale 192 193 is same in panels c-i. HiRISE Image credits: NASA/JPL/University of Arizona. (j-l) Morphometric characteristics of the boulder falls on/near the wall of Jezero crater. Histograms of (j) boulder 194 195 length and boulder track width, (k) boulder track length, (l) boulder track slope and boulder source slope. The measured values plotted in j-l are rounded to 1 decimal place. Orientation of slopes on 196 which the boulder fall tracks are observed on (m) delta deposits, (n) crater wall, and (o) the walls 197 of 6 small craters (diameter: $\sim 1.5 - 5$ km) located within two to three crater radii of Jezero crater 198 199 (Figure S2, S8).

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201 **2.3 Boulder runout distance and shadow angle**

Runout characteristics of the boulders fallen around the delta deposits is inferred from the ratio 202 between the horizontal distance travelled by the boulder (L) and the vertical fall height (H) of the 203 boulder. The L/H ratios for boulder falls on the delta deposits is greater than ~1.6 with a mean 204 value of 2.5 (Figure 4a). The correlation coefficient for best-fit is 0.75 using a linear regression 205 function. We find that on the delta deposits, both smaller ($\leq 1-1.2$ m) and bigger (3.7 m) boulders 206 207 have L/H ratio \geq 3, suggesting that boulder size is not the primary driver for runout distance on the delta deposits. The shadow angle is estimated as the arctangent of the relationship between the 208 vertical drop of the boulder (elevation difference between the apex of talus slope and final 209 elevation of the fallen boulder) and the horizontal distance travelled by the boulder (distance 210 between the apex of the talus slope and final location of the fallen boulder). The shadow angle 211 ranges between 16.2° and 27.1° for boulder falls on the delta deposits (Figure S3). On the front of 212 delta deposits, the increase in runout distance (67 m) and reduction in minimum shadow angle 213 (16.2°) resulted in boulder runout beyond the base of talus slope. This is consistent with the 214 215 observational evidence that the runout of the boulders on the talus slope was not influenced by the aeolian landforms on the front of delta deposits. The mobility of the boulders that we have 216 calculated (using the shadow angle) is within the range of those calculated for terrestrial rock falls 217 (Evans and Hungr, 1993; Copons et al., 2009). Hence, there cannot be a significant acceleration 218

- 219 (or addition of energy) above that of simply gravity acting alone. No trend is identifiable in the
- 220 plot of boulder fall size (m³) versus tangent of the shadow angle (**Figure 4c**).



Figure 4. (a) Plot of horizontal distance (L) versus vertical drop (H) with fitted linear regression lines for delta deposits ($R^2 = 0.75$). (b) Plot of boulder volume (volume of individual boulders in m³) versus tangent of the shadow angle for boulders with tracks sourced at the delta deposits.

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226 **3 Discussion**

We report previously undetected boulder fall activity at the landing site of Mars 2020 Perseverance 227 rover. The average slopes (22° for the delta front and 20° for the other regions of the crater) onto 228 229 which the boulders have fallen are lower than the angle of repose (30°) for dry particles on Mars (Kleinhans et al., 2011; Atwood-Stone and McEwen, 2013). Some (3 of the 16 boulders with 230 tracks on the western delta deposits) of the boulders that have moved (length >2.7 m) are bigger 231 232 than the Mars 2020 Perseverance rover (width 2.7 m, and height 2.2 m) (Figure S4). The proposed traverse of the Perseverance rover is along the front of delta deposits; therefore, it is important to 233 understand the processes of boulder destabilization and whether it could still be ongoing today. 234 which we will discuss further below. 235

236 **3.1 Process of boulder destabilization**

237 Evidence of frost has been reported at low latitudes on Mars (15-30°) (Schorghofer and Edgett, 2006; Carozzo et al., 2009; Vincendon et al., 2010a, 2010b) and at the Opportunity rover landing 238 site (at 1.95° S) (Landis, 2007). If seasonal frosts also occur at Jezero crater, the rocks therein 239 240 could be broken down by the stress produced due to volume change of frost condensed in cracks. Seasonal and diurnal temperature differences at latitudes poleward of 20-30° are estimated to 241 develop tensile stress >2-3 MPa in ice cemented ground (Mellon, 1997). However, whether thin 242 frosts in fractured rocks at equatorial latitudes can induce such thermal contraction cracks is 243 unknown but is a factor to consider for breaking down and destabilizing boulders in Jezero crater. 244 Solar insolation induced thermal stress has been proposed as a mechanism for rock breakdown on 245

the surface of Mars (Eppes et al., 2015), and has been shown to lead to rock falls with preferential orientations (Tesson et al., 2020). However, we find that the boulder tracks do not exhibit any similarity in the preferred orientations (Figure 3m-o). So although thermal stresses probably contribute to rock breakdown, they do not seem to be the principal driver for destabilization.

250 Wind could scour beneath boulders causing them to fall (e.g., Malin and Edgett, 2001). Aeolian bedforms indicate modern easterly winds in the interior of Jezero crater (Chojnacki et al., 2018; 251 Day and Dorn, 2019) and Jezero crater is located in a high sand flux region of Mars (Chojnacki 252 et al., 2018). Previous studies have implicated aeolian erosion for the erosion of delta deposit 253 within the crater (Fassett and Head, 2005; Schon et al., 2012; Day and Dorn, 2019) and a less 254 erosion-resistant rock layer is known to be removed by aeolian scour and induce scarp retreat in 255 Jezero crater (Williams et al., 2020). Hence, aeolian processes seem a likely driver for boulder 256 destabilization in Jezero crater. 257

258 Lithological differences can favor boulder destabilization. Mass wasting (in form of boulder falls 259 and movement of the boulders) of the sulfate-rich sediments in the equatorial region of Mars is 3-4 orders faster than the average erosion rate on the martian surface (Thomas et al., 2019). 260 261 Mineralogical mapping of the delta front where we have observed boulder falls indicates the presence of Fe/Mg smectites (Goudge et al., 2017), but fresh boulder tracks are not confined to 262 this unit (Figure 1b), and various minerals have been reported in these regions (Goudge et al., 263 2015, 2017). Hence, we do not find a specific role for lithology in explaining the boulder falls in 264 Jezero crater. 265

Boulders can be mobilized downslope by the ground shaking caused by meteorite impacts or 266 seismic shaking. We found no fresh impact craters within two to three crater radii of Jezero crater. 267 268 We found boulder fall tracks within Jezero crater as well as on the wall of 6 small craters distributed within a radial distance of 60-70 km around the Jezero crater (two to three crater radii) 269 (Figure S2, S8). We have not observed any recent impact crater within this 60-70 km radius around 270 271 Jezero crater to account for impact causing shaking in the study region. Previous work has found 272 that boulder falls resulting from seismic shaking can be observed over a distance of several 273 kilometers from visible surface faulting (for example in Cerberus Fossae) (Roberts et al., 2012). 274 Jezero crater is close to a group of large, concentric grabens, called Nili Fossae, although there 275 have been no reports of these faults being active on geologically recent timescales. A detailed 276 examination of the HiRISE images of the Nili Fossae would be needed to understand if recent 277 mobilization of boulders has occurred systematically along the strike of the faults in the region, 278 which is beyond the scope of this paper. Hence, we maintain that seismic shaking is a plausible hypothesis for boulder mobilization at this location, but have no corroborating evidence. A 279 280 previous study reported a paleomarsquake of moment magnitude 'M_w' up to 7.3-7.8 from boulder track data in Grjota Valles (**Brown and Roberts, 2018**), but InSight (Interior exploration using Seismic Investigations, Geodesy and Heat Transport) instruments have only observed marsquakes, ' M_w ' <4 (**Banerdt et al., 2020**). Therefore, detailed modeling would be needed to understand the proximity and magnitude of seismic events to mobilize boulders as large as 3.7 m at Jezero crater.

The data presented here are not sufficient to determine the recurrence interval of boulder falls at 285 Jezero crater; hence we cannot assess the likelihood of boulder falls happening during the Mars 286 2020 mission. We have observed fresh and faded tracks in close proximity (Figures 1b-1d, 2, 3), 287 suggesting that neighboring tracks have different ages (as they should share similar environments), 288 and the processes destabilizing the boulders within Jezero crater has been active over a prolonged 289 period. The evidence of fresh tracks on the western delta deposits, on the western wall of crater, 290 and on the wall of small craters in the vicinity of Jezero crater suggests that the boulder 291 destabilization process is not local and was last active relatively recently. If the boulders continued 292 to move once the track has formed, then this motion has a magnitude less than a meter in ~ 12 293 294 years.

3.2 Other Implications for the Mars 2020 mission

While the process for boulder destabilization is still unclear, the hypotheses we present could be tested by the Mars 2020 Perseverance rover, scheduled to arrive in Jezero crater in early 2021. For instance, the mast-mounted camera system (MastCam-Z) has wavelength filters that could detect frost deposits (**Rice et al., 2020**), while the Radar Imager for Mars' subsurface eXperiment (RIMFAX) would be able to identify any subsurface ice (**Russell et al., 2020**). The onboard cameras can be used to monitor aeolian transport and abrasion and look for clues of thermal breakdown. The role of wind or frost could also be tested by the Mars Environmental Dynamics Analyzer (MEDA), which will monitor wind speed and directions, surface temperature and
 measure local humidity (de la Torre Juárez et al., 2020).

305 Our observations open the door for in-situ analyses of boulders whose source areas would ordinarily not be accessible directly by the rover. The tracks allow their provenance to be better 306 ascertained than boulders without tracks. Some of the boulders at the foot of the delta are present 307 on slopes as low as 4-15° (Figure S5), thus accessible by a rover, and have been detached from 308 source regions where Fe/Mg smectites are thought to be present (Goudge et al., 2017). Boulders 309 present on the foot of crater wall at slope as low as 20° (Figures 1c, 3a-b and Figure S6) could 310 also provide access to the crater rim material, thus to ancient pieces of the crust (e.g. Goudge et 311 al., 2017). In addition boulders with tracks are geologically "recent falls", so these recently fallen 312 313 boulders may provide an opportunity to drill fresh exposures which might be less exposed to radiation than other rocks at the martian surface and hence could present ideal targets to analyze 314 315 for organics (Cockell et al., 2016).

316 4 Conclusions

We demonstrate a previously unknown boulder fall activity on the talus slope of the western delta 317 318 deposits and on/near the wall of Jezero crater. The relative paucity of monitoring data (time step 319 of before-and-after images limited to \sim 12 years) does not allow us to conclude whether boulder 320 fall and boulder track modification are actively occurring at the present-day or not. Nevertheless, 321 our observations suggest boulders with visible tracks to have detached during recent geological 322 time scales. Surfaces of these recently fallen boulders therefore present a set of potential targets for measurement of organic molecules by the Mars 2020 Perseverance rover. With the available 323 data we cannot determine which process or processes may have been responsible for the boulder 324

falls in and around Jezero crater, but propose hypotheses that can be tested by the rover (such as:thermal contraction, aeolian erosion, seismic shaking).

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