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Springer Handbook of Glass

1. Silicate glasses

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Summary:

Silicate glasses are important cultural, societal and geological materials. Geologic glasses testify for the igneous activity of the Earth and, for instance, represented important source of tools and ornamental objects during the Paleolithic. Nowadays, silicate glasses are used to build technical materials, such as smartphone screens or glass matrix for stabilizing hazardous radioactive wastes. Therefore, silicate glasses are central to the history of the Earth and of the humanity. The compositional landscape of natural and industrial silicate glasses is vast, with various elements that all influence differently the glass properties and structure. The SiO_4 tetrahedral framework, backbone of silicate glasses, is variously influenced by the introduction of network modifier metal cations or network former aluminium cations. Industrial and geologic silicate glasses further contain multivalent elements (e.g., $\text{Fe}^{2+/3+}$), rare-earth elements, and volatile elements (H, C, S, Cl, F, I) that play different roles on the glass structure and properties. This chapter proposes to review the link between the structure, the properties and the chemical composition of silicate glasses.

1.1 Silicate glasses: Historical and industrial importance

The origin of the first silicate glass is a vast question. The first glass tools actually were manufactured by breakage and date approximately of 10000 BCE, as shown by obsidian arrow heads of this age found in Turkey for example. Manufacture of silicate glasses dates back to the earliest antiquity; it follows the control of fire, and is synchronous with the Bronze Age which itself varies geographically. The first objects made by melting date of ~6000 BCE. Hazard was probably the first factor in this melting manufacturing. Henrivaux [1] proposed that the appearance of glass should be contemporary with the manufacture of bricks cooked by fire and of certain ceramics. Indeed, a plausible hypothesis is that brick and ceramic manufacturers were the first glass manufacturers because various materials subjected to the action of an intense fire may present the first symptoms of vitrification in the form of a veneer or surface glaze. This hypothesis may explain the Mesopotamian glazes observed on several ceramics and located at Suse between 7000 and 5000 BCE. However, this hypothesis contradicts with the story reported by Pline the Elder (23-79 BC). Indeed, in his Natural History [2], Pline reported that it would be Phoenician merchants who, cooking their food on the banks of the river Belus in pots supported by blocks of Natron, would have seen an unknown substance flow:

“The story is, that a ship, laden with nitre, being moored upon this spot, the merchants, while preparing their repast upon the sea-shore, finding no stones at hand for supporting their cauldrons, employed for the purpose some lumps of nitre which they had taken from the vessel. Upon its being subjected to the action of the fire, in combination with the sand of the sea-shore, they beheld transparent streams flowing forth of a liquid hitherto unknown: this, it is said, was the origin of glass”

This observation is unlikely when one knows that temperatures close to $\sim 1000^{\circ}\text{C}$ are necessary to observe the flow of silicate melts, conditions difficult to achieve using campfires. Pliny's account thus probably is a legend, and current historical knowledges suggest that the formation of glass as a byproduct of metallurgic and ceramic activities is more plausible [e.g., see the review of 3]].

Regardless of the origin of its discovery, glass was considered as precious during the antiquity. This is shown, for instance, by the Hebrew term that designates glass, “כֹּזְחִית” (pronounced “zkhourhit”), which comes from the Hebrew word “זָכָה” that means pure. Glass thus remained an expensive material during antiquity. As such, it was reserved for noble objects as jewelry and amulets (Fig. 1), competing with precious metals such as gold or silver for such uses.

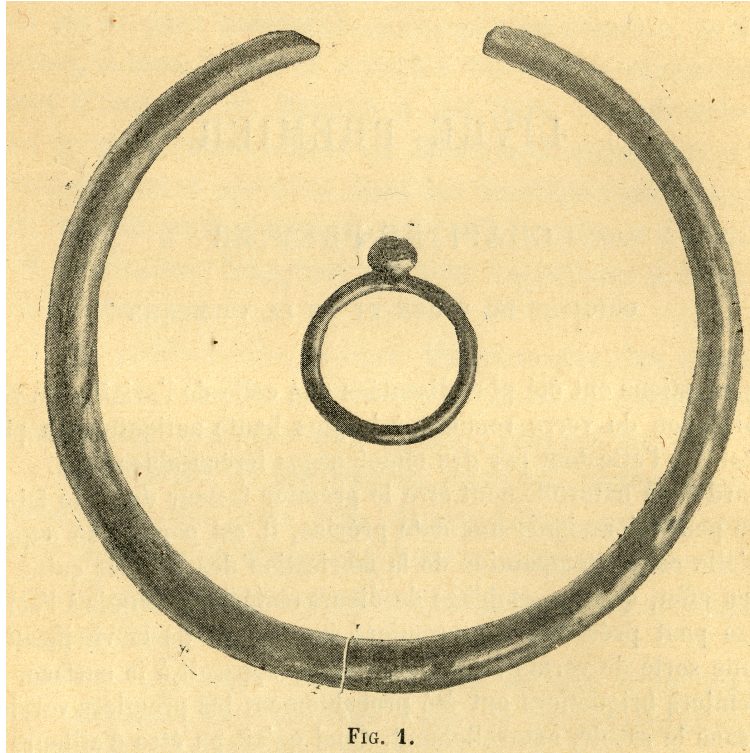


Figure 1: Ring and earring reproduction. Image reproduced from Henrivaux, 1897, “Le Verre et le cristal” Vicq-Dunaud Editors, Paris [1].

From 5000 to 3000 BC, objects partly made of glass are characterized by a faience coating placed on the surface of a ceramic material. Then, as the furnaces of the age of bronze and copper appeared, the work of the glass became easier, allowing the manufacture of colored glass bead for instance. Core-molding and cast glass methods, using molds for shaping glass objects, were subsequently developed in the ~ 1500-500 BCE period in Mesopotamia and Egypt. Some sources attribute the invention of glass-blowing to Syrian glassmakers around 300 BCE [e.g., 4], but this disagrees with the representation of glass-blowing in Egyptian tombs [5], as for instance in a tomb from the fifth dynasty at Sakkarah or the painting at Beni-Hassan (Fig. 2) that dates from the reign of Userhessen I. (2758-2714 BCE).

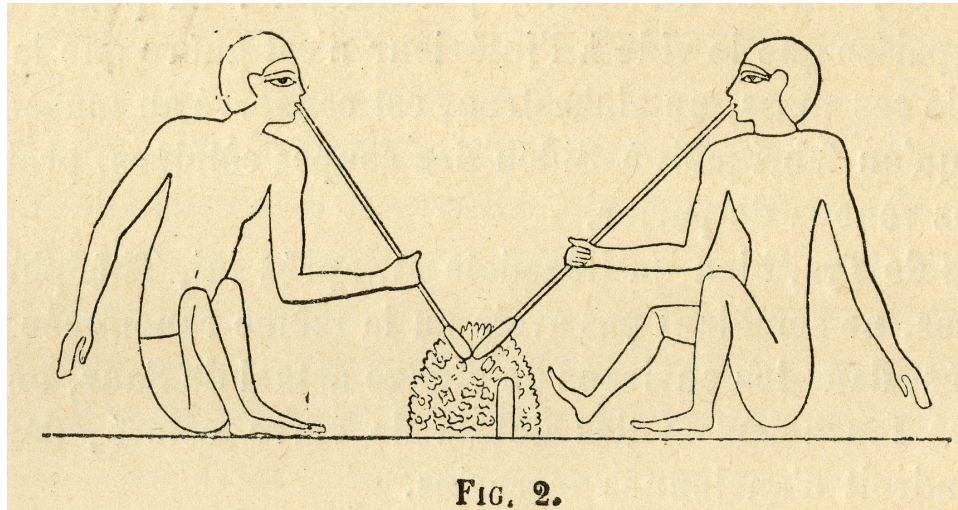


Figure 2: Representation of Egyptian glassworkers blowing glasses on brazier at Beni-Hassan. Reproduced from Henrivaux (1897) “Le Verre et le cristal” Vicq-Dunaud Editors, Paris [1].

Historical glass compositions are based on the use of silica, lime and soda. The compositions of historical glasses remain in a tight chemical domain. They are still close to those of present glasses (Fig. 3), illustrating that the optimization of glassmaking processes was achieved early in the history of human glass manufacture. An interesting fact is that early glass makers already aimed at controlling the color of translucent or slightly opaque glasses through the addition of metals such as copper (red and blue-red), iron (black, brown and green), antimony (yellow), cobalt (blue) or tin (white). Such coloring substances were already identified and known before the age of glass manufacturing, as several of them were used as pigments [6]. Early glass manufacturers probably had access to raw materials presenting variable quality. Production was therefore subject to many hazards, against which they could only contrast empirical knowledge.

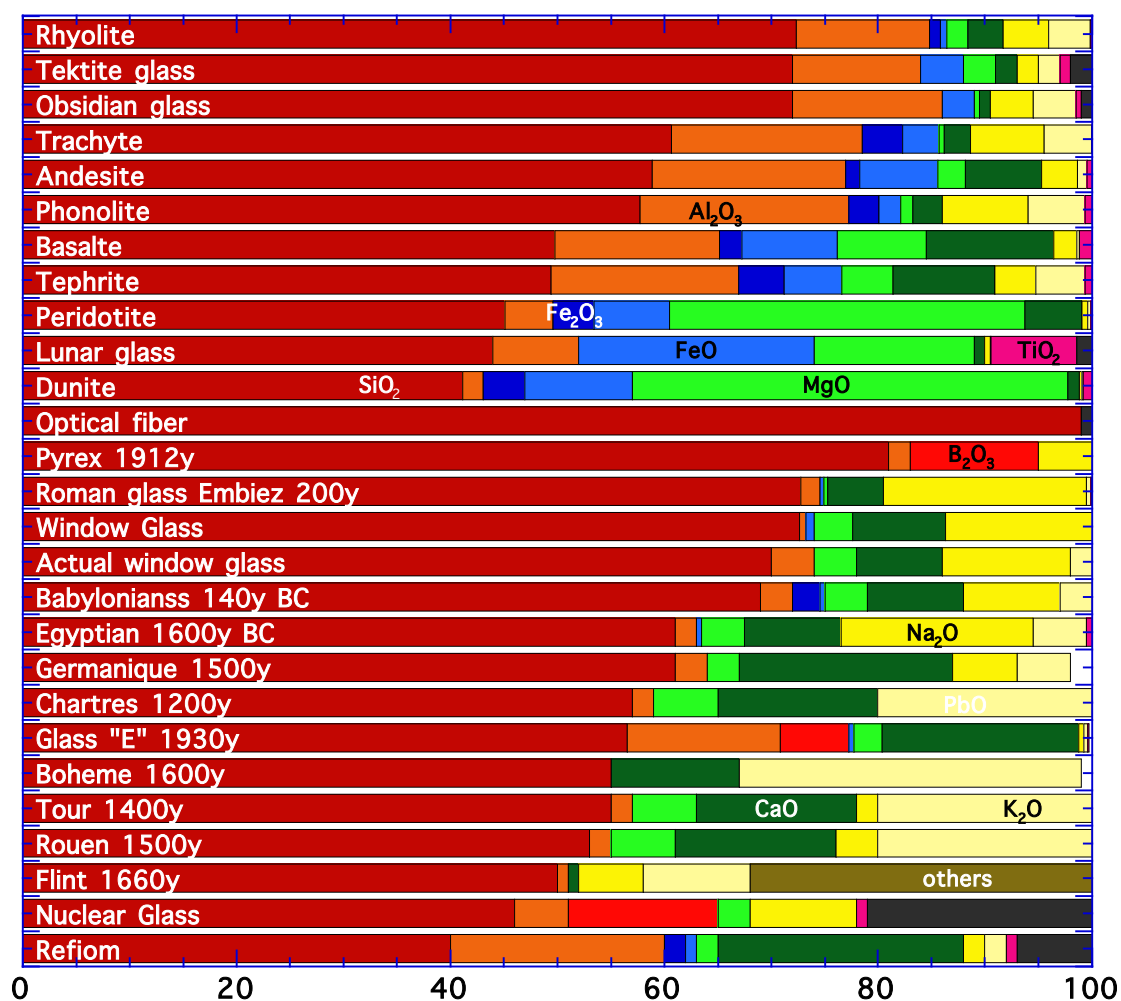


Figure 3: Examples of chemical composition of various geologic (lunar glass, dunite, peridotite, tephrite, basalt, phonolite, andesite, trachyte, rhyolite, tektite, obsidian), industrial (Optical fiber, Window glass, Glass E, Nuclear Glass, Refiom, Pyrex 1912, Actual window glass) and historic (Roman glass Embiez 200y, Flint, Egyptian, Babylonians, Tour, Chartres, Rouen, Germanique, Boheme) glasses [7–10].

Around the 15th century BCE, improvements of the furnaces allowed higher working temperatures. As a result, fusion quality improved, leading to a technological transformation as the first translucent glasses were produced. Glass objects presenting a large variety of shape appeared, such as bottles and containers used to preserve ointments, perfumes, food or beverages, for instance. For instance, to form small containers, core-molding technics were used: a rather viscous silicate melt was spread on a core, made for example of a mixture of clay and manure; the core was then removed from the inside of the new piece. Other shaping methods were developed, and, from that time, glasses were decorated. In addition, technological glasses appeared, as for instance *millefiori* glasses (Fig. 4a) composed of many glass filaments embedded in a transparent glass, forming a unique piece. Thin glaze films covering various

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objects were also formed using silicate melts rich in lead oxide and, hence, presenting low viscosity at working temperature. Towards the 4th century BCE, this new process improved greatly potteries, making them waterproof. From its Mesopotamian-Egyptian origin, the art of glass spread to neighboring countries. Pearls made at the end of the second millennium were found from Italy to Central Asia and China (Fig. 4b). Glass containers were found in Greece, Italy and China, respectively dating from the 13th, 8th and 5th centuries BCE. The availability and quality of the raw materials may have determined the abundance and interest in producing glass at such time. Indeed, glass transparency requires using raw materials with high purity during the glass-forming process. In particular, the starting materials should be very poor in iron oxide, one of the main glass colorant that is abundant on Earth surface. As such pure raw materials were easily found in Palestina and Egypt, for instance, this may explain the abundance of glasses made in these areas.

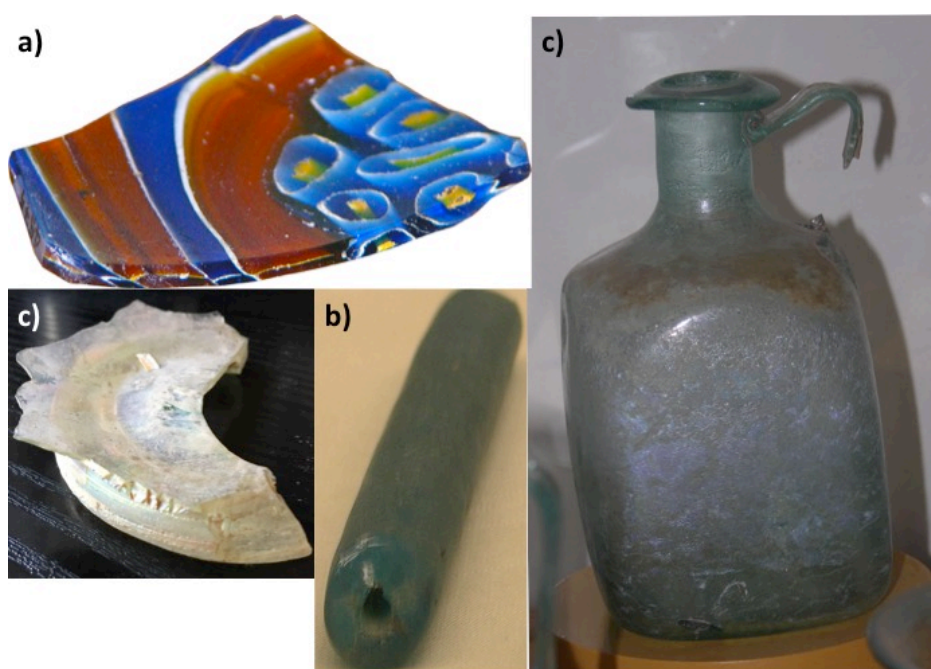


Figure 4: a) *millefiori* glasses, b) pearls, c) fragment of Roman glass and d) bottle from the Roman empire. Pictures Daniel R. Neuville.

The use of glass blowing during the Roman empire (Fig. 2) gave birth to a veritable industrial revolution. This process allowed fast, easy and thus cheap shaping of the glass. The size and shape of the pieces were also less limited than with previous methods. Whether intended for utilitarian or decorative

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purposes, glass was able to find its place in all homes under the Roman empire. Even the deceased benefited of such technological evolution as their ashes were preserved in glass urns. Glasses dating from the Roman empire cover a small chemical composition domain. This suggests a form of globalization of the glass market at that time, with the production of glass concentrated in various factories located along the syro-palestinian costline [11]. This concentration probably resulted from the abundance of local deposits of pure sand and natron, giving this region a great competitive advantage. From there, glass was exported in the form either of finished products or of glass ingots, which could be re-melted in small furnaces at their destination to shape the final objects. For instance, the Ouest-Embiez ship wreck found in the Mediterranean Sea testifies for the occurrence of this trade [11].

The decline of the Roman Empire followed by the Arab conquests probably led glass trading routes to be cut off [12]. As a result, starting at the beginning of the Middle Ages (5th-15th century BCE), glasses had to be produced locally in Western Europe, with the available resources on the spot. The local cullet was used together with ashes to help melting local sands in order to produce glass. Using ashes helped compensating the absence of natron. Figure 5 represents the composition of ashes obtained from different plant and tree species. In Europe, ash from beech, oak and fern were mostly used in glass making processes. Such ashes are rich in potassium (Fig. 5), such that potash-lime glasses were produced in northern Europe. Unfortunately, potassium is not an element ensuring the stability of the glass in regards of alteration, explaining why potash-lime glasses from the early Middle Ages are not well preserved. Soda-lime glasses were also produced in regions where the local availability of saltwort was good (Fig. 5). This induced a typical geographic repartition of soda-lime and potash-lime compositions in Western and North-Eastern Europe at this time. This changed during the 11th century, as saltwort becomes to be imported from Spain for producing glass.

In parallel to the events in Europe, the development of glass in the East took a totally different path after the fall of the Roman empire. In the Byzantine empire, glassmaking processes used during the Roman period survived, such that the glass production that characterized the Roman glasses was conserved. The rise of the Muslim Arabs saw the development of new glasses. The Islamic glassmakers used their knowledge of the Roman and Byzantine glassmaking technics and their own experience to produce high quality glasses. In particular, this led to the introduction of the *Islamic soda-lime glass*, introduced between the 8th and 10th century CE. After the first crusades in the 11th century, glassmaking centers progressively shifted from the Islamic glassmakers to the Venice (Italy) center.

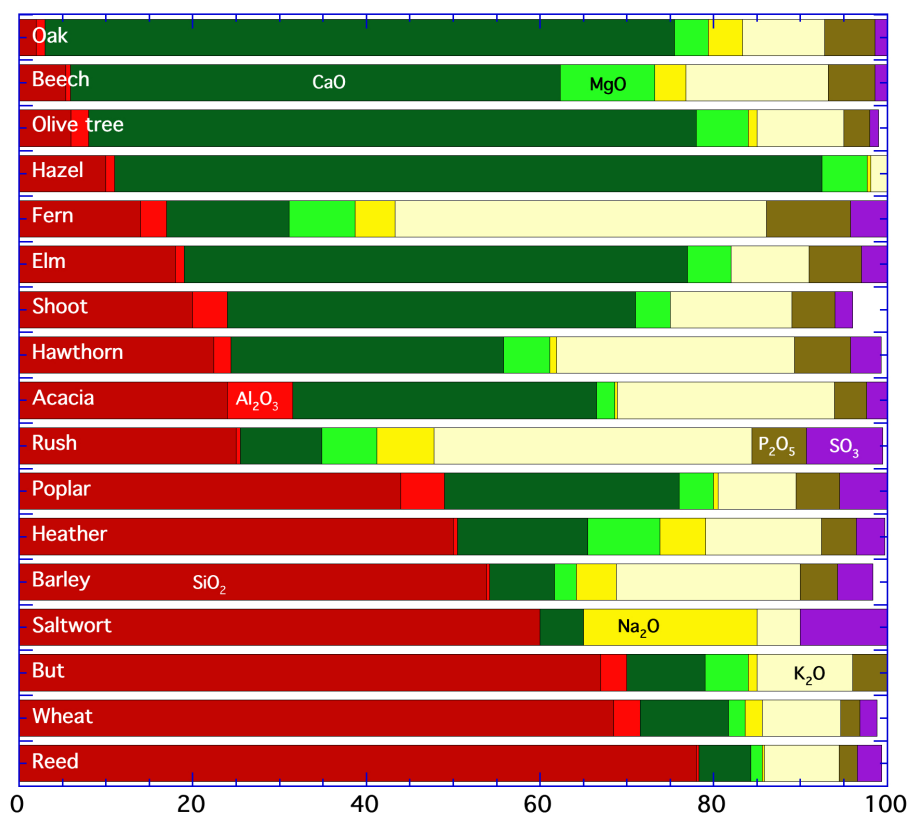


Figure 5: Chemical composition of ashes generated by burning different plant and trees. Data from Vernioles [13].

Indeed, the glassmaking knowledges acquired during the prosperous periods of the Roman empire never really died, and new contacts with the Byzantine empire following the first crusades allowed the transmission of new knowledges from the Islamic to the Venetian glassmakers. In addition, starting from the 11th century, Venice was becoming one of the most important center of trade, commerce and arts in the Europe and Mediterranean regions. As a result of the combination of those factors, the Italian glassmaking industry started to flourish again from around 1200 CE. In 1291, the glassmaking factories moved to the Island of Murano, to keep the Venice city safe from fires. This further helped keeping the secrets of the Venetian glassmaking processes safe. By the 15th century, Murano became one of the most important glassmaker in Europe. Venetian glassmakers continued to improve the quality of the glasses they produced, this leading to the production of the *cristallo* for instance, a transparent, shiny glass prepared from soda ash and quartz pebbles from Ticino (Italy). Using this glass composition that is characterized by a long working time, the Murano glassmakers manufactured thin and relatively light glass windows and mirrors that were exported all over Europe.

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While the Venetian glasses were of high quality, they also were expensive. This actually pushed the birth of a French glassmaking company that is nowadays one of the largest glass and material producer in the world. Indeed, when the French king Louis XIV decided to build the Versailles castle, the price asked by the Murano glassmakers for the required large quantities of glasses was particularly depressing. This pushed Louis XIV to grant a royal privilege in 1665 to a new glass factory, which finally settled in 1693 in the village of Saint-Gobain. In 1695, it became the French Royal Glass Factory, famous for the production of flat glass through cooling and flattening the melt directly on a wood table of several meters. This process was notably depicted in the painting “The visit of the duchess of Berry at Saint-Gobain in 1822”.

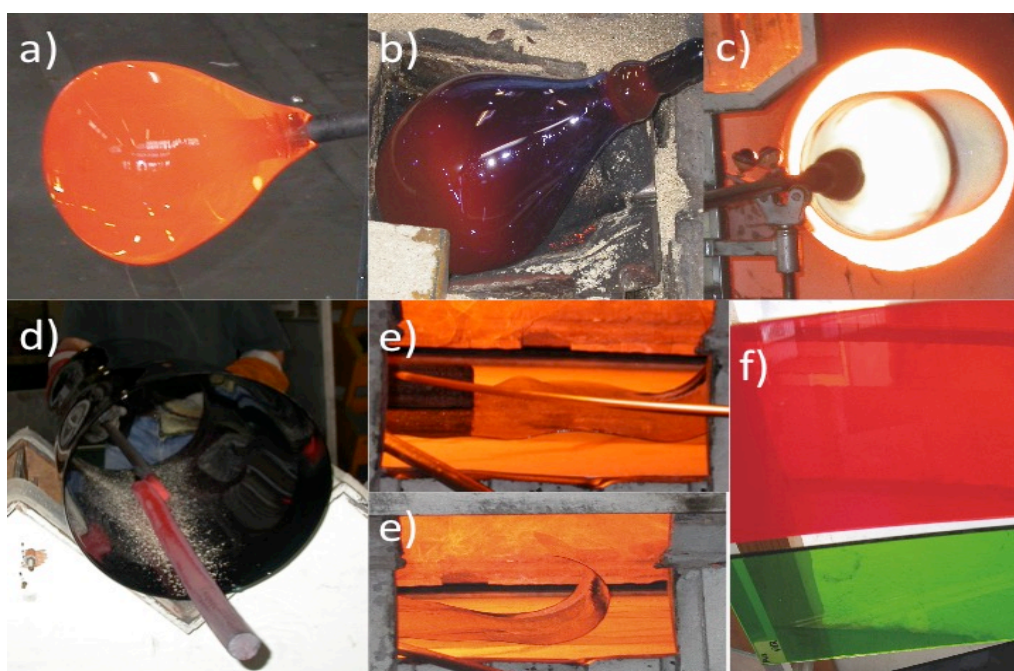


Figure 6: Formation of flat glass using the cylindrical sleeve technique: a) b) and c) a cylinder of glass is formed by blowing, d) each end is removed and one of the cylinder side is cut, e) the cylinder is then softened under fire to obtain a flat window, and f) final product can be colored or uncolored. Pictures Daniel R. Neuville, from the Saint-Just Factory – Saint-Gobain Compagny.

The process cited thereabove was a late evolution of two technics used to produce flat glass, in high demand for the construction of windows and mirrors for religious and royal edifices in Western Europe since the Middle Ages. Indeed, the first productions of flat glass were performed using either the *crown* or the *cylindrical sleeve* processes. The *crown* process consisted in obtaining a cylinder of flat glass from a vase, initially formed by blowing. This cylinder of flat glass can then be cut in any wanted shapes. The *cylindrical sleeve* process used a different approach, forming a glass sheet from a glass cylinder (Fig. 6).

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Glasses formed through those processes were expansive, such that their use was reserved to religious or important edifices until the end of the 17th century.



Figure 7: Representation of the three different furnaces needed to make glass and proposed by Agricola. This representation is extracted from “Dans l’art de la verrerie” from Neri, Merret and Kunckel (1752) [6].

While the brazier represented in Figure 2 is one of the first representations of a glass furnace, the one developed since the Middle Ages were highly sophisticated. An example of a furnace used at this time is given by Agricola in *De Re Metallica* in 1556 and later reported by Neri et al. [6]. Figure 7 shows the representation displayed in Neri et al. [6], in which we distinguish three different furnaces: -i) the first called *calcaria* was used to prepare oxide and carbonate powders from raw materials (shells, chalk, marls,

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wood, sea salt, vegetable or animal ash...); - ii) the second furnace was typically a pot furnace used to melt the materials generated with the first furnace, in order to obtain from 100 to 300 kg of glass (a bath furnace could also be used for generating large glass volume of several cubic meters); -iii) a third furnace was used for annealing the glass pieces generated by blowing the silicate melt from the second furnace.

At the end of the Middle Ages and the beginning of the Renaissance, academic societies and the renaissance *savants* started to make further important progress in knowledge. Such time saw the development of new glasses, and glassmaking technics. For example, the Prince Rupert started to look at some spectacular and curious properties of glass, presented to the Royal Society of London in 1661. Indeed, dropping melts at high temperature in water, Prince Rupert noticed the formation of glass drops very resistant to violent shocks, up to a point after which they exploded suddenly and turned into fine powder. Spectacularly, these first tempered glasses, well known as the *Prince Rupert's drops*, illustrated the role of internal stresses that accumulate in the absence of annealing (Fig. 8). George Ravenscroft (1632-1683) invented the *crystal* glass through the addition of an important quantity of lead oxide in the composition. From this invention, patented in 1673 by Ravenscroft, lead-based glass spread among Europe, representing an important turnover as it allowed improving the appearance of glass as well as lowering the melting and working temperatures. In parallel, the naturalist and physicist René-Antoine Ferchault de Réaumur (1683-1757) observed the partial crystallization of a glass and created the first glass-ceramic in 1727. While such aspect was considered to be a defect at this time, it is a particular focus in glass and materials sciences nowadays because glass-ceramic materials open new doors to the manufactures of new strategic materials [e.g., see 14].



Figure 8: Picture of two glass drops observed under polarized and analyzed light. The sample at the left is unrelaxed, while that at the right was annealed, removing any intern stress. Picture courtesy of Prof. D. de Ligny.

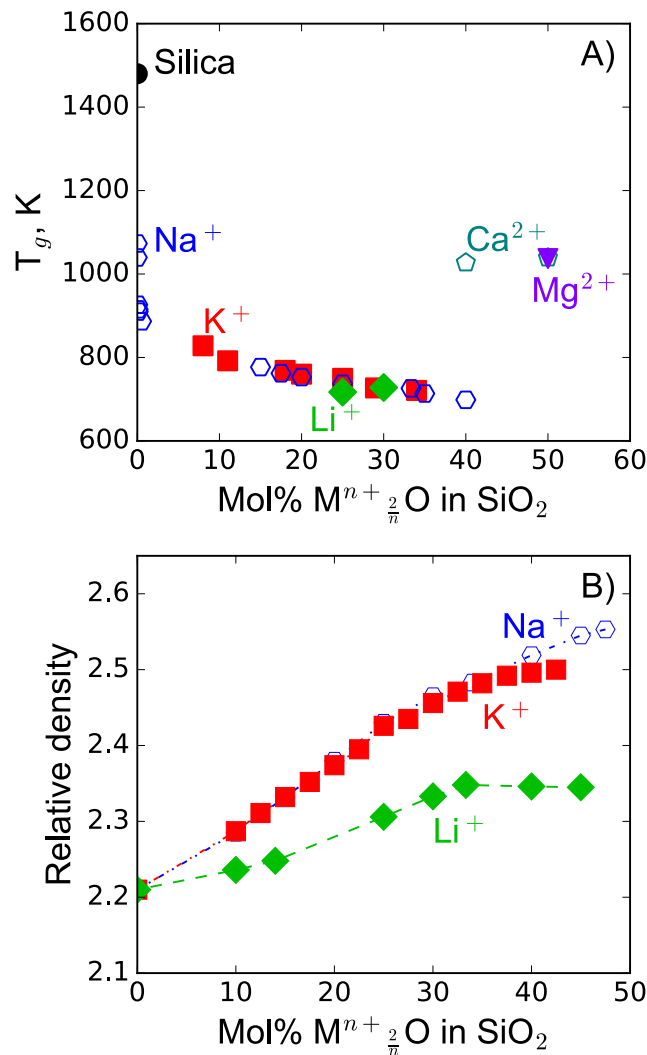
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The 19th and 20th century saw a rapid evolution of the glass market, as glass was needed as a technological material for various applications. For instance, in 1893, the *Duran* boro-silicate glass, was developed by Otto Schott, a German glass maker. This glass presented a low thermal expansion, such that the Corning Glass Works company used it to form the first light bulb for Thomas Edison. At the turn of the 20th century, other technical problems appeared with the development of railroad, automobiles and other mechanical transportation means. For instance, the glass globes of signal lanterns on United States of America railroads were sensible to extreme temperature changes (e.g., heavy rain on hot glass), leading them to shatter. This problem was solved by the development of the Nonex and the famous Pyrex[®] glasses, respectively in 1908 and 1915 by the Corning Glass Works company. Those boro-silicate glasses present a very low thermal expansion, making them particularly suitable for uses in environments with extreme temperature changes. The E glass, a calcium boro-silicate, was subsequently developed for electrical insulators has thus been found to have excellent mechanical properties. From the middle of the 20th century, it was this glass that was produced in the form of fibres to reinforce composite materials. In parallel of the discovery of such new glass compositions, the fast spreading of mechanical glassmaking processes allowed lowering the price of glass, making it suitable for its use in many different areas such as food storage, building or glass lenses for instance. Following World War II, many new glass families were made and developed, and the invention of the *float glass process* by Sir Alastair Pilkington and Kenneth Bickerstaff between 1953 and 1957 further revolutionized the glass production processes. Nowadays, silicate glasses still represent the largest quantity of human-made glasses. They are used in many aspects of everyday life, e.g. fiber optics, cellphone and computer screens, car windshields. However, the importance of silicate glasses is becoming smaller, thanks to the development of new and exciting glasses, such as, for example, chalcogenide, metallic or fluoride glasses. Such new glasses have a high technological importance, with critical applications as for instance for infrared cameras, photonics or data storage.

From the previous discussion, silicate glasses used in many technical and technological applications through historical times as well as natural geologic glasses present a wide range of chemical composition (Fig. 3). Such large variations in chemical composition are associated with large differences in term of glass properties, e.g., glass transition temperature T_g , density, hardness, chemical durability. For example, the T_g and relative density of simple binary silicate glasses $M^{n+}_{2/n}O-SiO_2$ are reported as a function of the molar percentage of $M^{n+}_{2/n}O$ in them in Fig. 9 (M=Na, K, Mg, Ca); along the $M^{n+}_{2/n}O-SiO_2$ binary systems, T_g vary non-linearly of hundreds of degrees. In addition, glasses composed of a given amount of SiO_2 but different metal cations M present different T_g (Fig. 9a). Density changes are closer to a linear

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trend that vary quite strongly with composition (Fig. 9b). Such changes in glass properties translate changes in the glass disordered molecular structure upon variations of the glass chemical composition. Indeed, while glasses do not present long-range order arrangement as crystals do, they still present specific molecular arrangements at short and intermediate range order. For instance, the environment at short range order of Si^{4+} cations is well defined as being a SiO_4 tetrahedral units in silicate glasses. At a scale of a few SiO_4 tetrahedral units, arrangements of such tetrahedral species in rings and cages is also reported, as it will be presented in the following sections. However, one should remember that even such molecular arrangements are disordered in glasses: for instance, they can present a broad range of bond angle and bond lengths, as shown by the board distribution of the Si-O-Si bond angle in the silica SiO_2 glass [15–19].



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Figure 9: A) Glass transition temperature T_g and B) relative density of binary silicate glasses along the $M^{n+}_{2/n}O-SiO_2$ join. T_g of Na_2O-SiO_2 and K_2O-SiO_2 glasses compiled by [20]; T_g of Li_2O-SiO_2 glasses: [21]; T_g of $MgO-SiO_2$ and of $CaO-SiO_2$ glasses: [22, 23]. Density data are from [21, 23–25].

In the present chapter, the structure and properties of silicate glasses will be reviewed. We define silicate glasses as glasses containing dominantly SiO_2 as the glass network former oxide component that allows their vitrification. Glasses built upon the presence of other network formers in their structure will be discussed in the following chapters (see chapters XXXXX). Silicate glasses usually contain various elements such as, for example, aluminum, iron, alkali and alkaline-earth metal cations. In a first time, the structure and properties of the silica glass will be briefly reviewed in section 1.2. Then, we will add new levels of complexity by looking at the structure and properties of Al-free and Al-bearing silicate glasses in sections 1.3 and 1.4. Section 1.5 will review multivalent and rare-earth elements in silicate glasses, and redox effects will be discussed, while sec. 1.6 gives a brief overview of volatile elements dissolved glasses.

1.2 Silica glass

1.2.1 Structural concepts

Silicate glasses inherit their names from the basic oxide component that allow their formation: silica oxide or SiO_2 . Interestingly, the possibility to quench the simple SiO_2 melt into glass indicates that Si^{4+} is a *network former* cation, similar to other cations such as Ge^{3+} or B^{3+} for instance (see chapters XXX for further information). Network former elements in oxide glasses participate in building an interconnected molecular network, forming the backbone of glassy materials. Regarding the pure silica glass, it is composed of Si^{4+} cations and O^{2-} anions arranged in SiO_4 tetrahedral units with Si^{4+} at the center and O^{2-} at the summit. Those SiO_4 units are interconnected by their apical bridging oxygen atoms, forming a tridimensional tetrahedral network.

Because of its simple chemistry, one may think that the structure of the silica glass is simple. However, it actually exhibits such a level of complexity that there is no definitive consensus about the three-dimensional arrangements of SiO_4 tetrahedral units at medium-range order. An early description of the silica glass structure was provided by Zachariasen [26], who proposed the Random Network (RN) model to describe the structure of oxide AX_2 glasses (with A the network former cations and X the anions).

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Following this model, Si and O structural sites are energetically equivalent, and the Si and O atoms randomly distribute between those sites following 4 rules that govern the formation of AX_2 glasses:

- 1) the system contains enough X anions to ensure the formation of triangle or tetrahedra of X anions around the central network former A cations;
- 2) those triangle/tetrahedra share the apical X anions at their summit;
- 3) a given X anion only link 2 A cations;
- 4) the triangle/tetrahedra only share their summit, and not their faces or edges.

In the silica glass, this model predicts the formation of a three-dimensional disordered network composed of rings and cages of SiO_4 interconnected tetrahedral (Fig. 10). First X-ray diffraction data of the silica glass in 1934 corroborated such description [27], and the latest observations of thin 2-dimensional films of silica deposited on graphene further support such view [28] (Fig. 10).

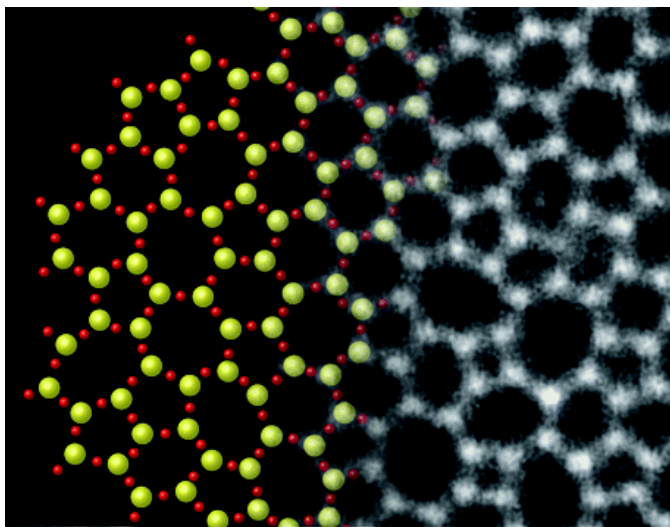


Figure 10: Image linking the transition between the theoretical 2D view the structure of the silica glass as drawn by Zachariasen [26] in 1932 (on the left) and the experimental observation by Huang et al. [28] in 2012 of a thin 2D film of amorphous silica deposited on graphene (on the right), made with using annular dark-field scanning transmission electron microscopy (ADF-STEM). Reprinted with permission from Huang, P.Y., Kurasch, S., Srivastava, A., Skakalova, V., Kotakoski, J., Krashenninnikov, A.V., Hovden, R., Mao, Q., Meyer, J.C., Smet, J., Muller, D.A., Kaiser, U., 2012. Direct Imaging of a Two-Dimensional Silica Glass on Graphene. Nano Letters 12, pages 1081–1086, doi:10.1021/nl204423x. Copyright 2012 American Chemical Society.

Pre-1990 analysis of X-Ray diffraction data [e.g., 29] and Raman spectra [e.g., 30] of the SiO_2 glass as well as early structural calculations [e.g., 31] supported the idea that the structure of the SiO_2 glass at medium range order was composed of tetrahedral rings, but no consensus existed about the number of tetrahedral units in the rings. In the Raman spectrum of SiO_2 , the asymmetric R band centred near $\sim 460 \text{ cm}^{-1}$, which extends from ~ 200 to $\sim 650 \text{ cm}^{-1}$, is actually assigned to Si-O-Si and O-Si-O bending vibrations in such SiO_4 rings (Fig. 11).

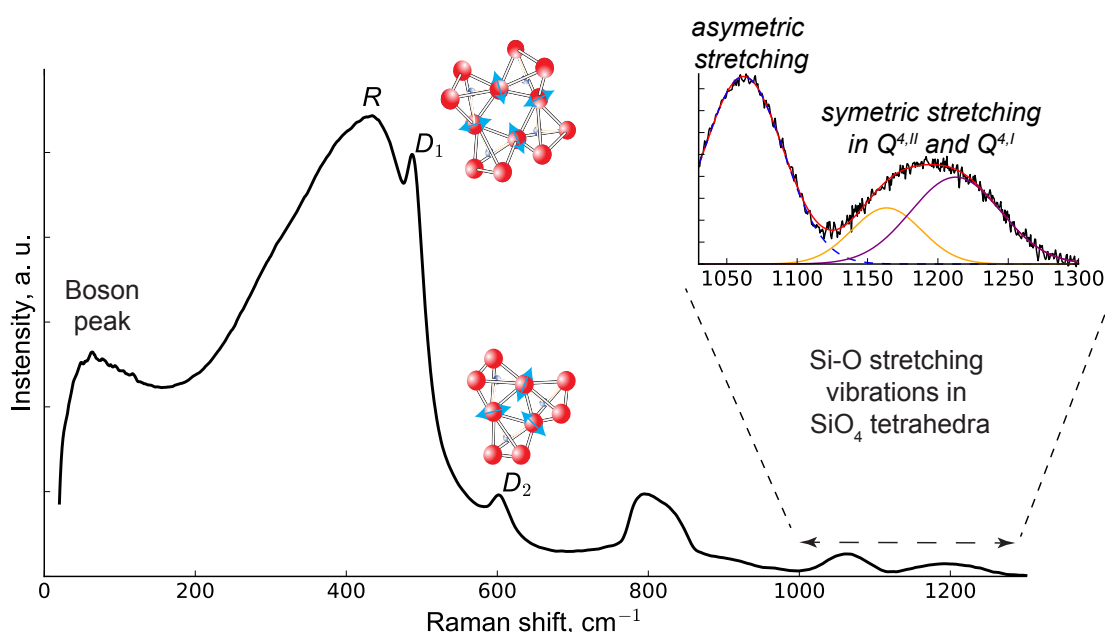


Figure 11: Raman spectrum of the Lybian glass, a glass found in the Lybian desert that probably results from a meteoritic impact. This glass presents a composition of 99% silica and 1% impurities (Al_2O_3 , alkalis...). The Raman spectrum of this glass is very similar to that of the pure SiO_2 glass. It displays: the Boson peak at $\sim 65 \text{ cm}^{-1}$, that can be assigned to collective transverse acoustic vibrational modes [32] promoted in silicate glasses by cooperative inter-tetrahedral librations [33–35]; the R band, an asymmetric band with an intensity maxima near 440 cm^{-1} that may be assigned to O movements along the Si-O-Si intertetrahedral bonds of tetrahedral rings with 5, 6 or more tetrahedral [36–48]; the D_1 and D_2 sharp peaks, assigned to breathing vibrations of four and three membered rings, respectively [39, 41–43, 46, 48–51]; an asymmetric band near 800 cm^{-1} , assigned to asymmetric vibrations of SiO_4 units with both O and Si movements [43, 52–54]; a peak near 1065 cm^{-1} , assigned to Si-O asymmetric stretching vibrations in SiO_4 units [25, 55–57]; and a broad band between 1100 and 1250 cm^{-1} that is assigned to Si-O symmetric stretching vibrations in SiO_4 units [36–38, 43].

At higher frequency, peaks at $\sim 1060\text{ cm}^{-1}$ and $\sim 1195\text{ cm}^{-1}$ are observed, and are assigned to Si-BO asymmetric and symmetric stretching modes of SiO_4 units, respectively [30, 55, 56, 58, 59]. Peak-fitting the $\sim 1180\text{ cm}^{-1}$ band with Gaussian components, Seifert et al. [30] noticed that two peaks were necessary because this $\sim 1180\text{ cm}^{-1}$ band is asymmetric (Insert in Fig. 11). They interpreted this splitting of the Si-O symmetric stretching in Q^4 units in two components as indicating an asymmetry of the Q^4 environments in the silica glass. This could be related to a slight asymmetry of the T-O-T inter-tetrahedral angle in SiO_2 glass (Fig. 12), as suggested by results from molecular dynamic simulations [18, 60–62], ^{29}Si NMR spectroscopy [19] and X-Ray absorption spectroscopy [15, 29, 63]. Seifert et al. [30] assigned this asymmetry to different puckering extents of 6-membered rings in the structure of the silica glass that yield T-O-T angle differences of $\sim 5\text{--}10^\circ$. In line with such idea, Henderson et al. [29] discussed the X-Ray Diffraction (XRD) data for SiO_2 as suggesting the coexistence of cristobalite-like and trydimite-like 6-membered rings in silica. Such structural models fall in the category of the crystallite network model of the SiO_2 glass, discussed by Wright [64].

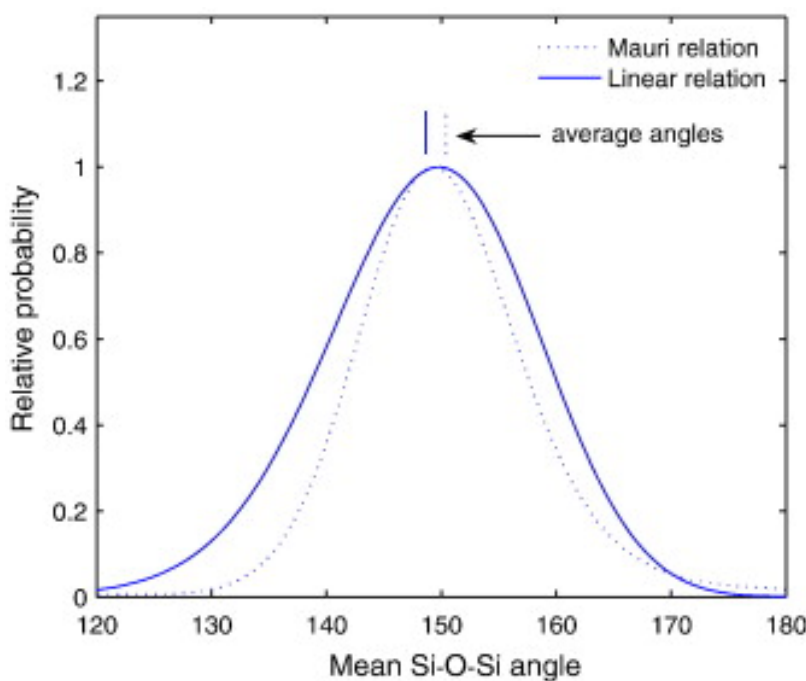


Figure 12: Distribution of the mean Si-O-Si bond angle in silica derived from ^{29}Si NMR spectroscopy, using either a linear relationship between the ^{29}Si chemical shifts and Si-O-Si bond angles measured in SiO_2 polymorphs and zeolites (plain curve), or a correlation determined by density functional theory by Mauri et al [65] (dotted curve). Reprinted from Malfait [19], Copyright (2008), with permission from Elsevier.

The idea of different 6-membered rings coexisting in the structure of the silica glass is still presently plausible, as no consensus has been reached on the structure of the silica glass. For instance, Huang et al. [66] recently followed such idea, showing with the help of Molecular Dynamic simulations that the mixing in different quantities of two types of 6-membered rings, presenting shapes close to those found in α -cristobalite and β -cristobalite, may account for variations of the SiO_2 bulk modulus and density with pressure. However, most Monte Carlo and Molecular Dynamic simulations [e.g., 18, 47, 51, 60, 67] indicate that, at medium range order, the structure of the silica glass is composed of a distribution of rings with different numbers of SiO_4 units, with 5-, 6- and 7-membered rings tending to be dominant (Table I). Therefore, as rings with different number of tetrahedral units present different T-O-T angles [42, 60], Si-O stretching frequency in SiO_4 units should vary as a function of the ring statistic in the SiO_2 glass. For instance, signals assigned to Si-O stretching in SiO_4 units in the main 6- and 7-membered rings will give signals at different frequency because they present different Si-O-Si bond angles and Si-O bond lengths [53, 54, 60]. Therefore, a repartition of SiO_4 units in mostly 5-, 6- and 7-membered rings, as suggested by results from simulations, may actually explain the asymmetry observed in the Raman data. In all cases, the asymmetry of the R band further agrees with considering the structure of the SiO_2 glass as containing a distribution of rings with 3 to more than 8 tetrahedral units. In particular, two sharp peaks at ~ 590 and $\sim 606 \text{ cm}^{-1}$ in Raman spectra of SiO_2 glasses (Fig. 11) can be assigned to minor fractions of 4- and 3-membered rings in the glass structure, respectively [39, 41, 42, 45–49, 51, 52, 68]. Such observation thus support results from simulations that tends to predict low concentrations of such rings in the glass structure [e.g., 46–48, 51]. The vibrational decoupling of such rings from the remaining part of the glass network allow them to contribute significantly to the glass Raman spectra [49]. Overall, the picture of the structure of the silica glass still is incomplete, but advances in simulations and spectroscopic technics may bring further determining pieces of information in a near future.

	Value	References
Si-O bond distance	1.59 – 1.61 Å	[18, 67]
Si-O-Si bond	146° – 151°	[17–19]
% of 3-membered rings	0.0 – 0.7 %	[18, 47, 51, 60]
% of 4-membered rings	0.0 – 7.7 %	
% of 5-membered rings	17.0 – 23.0 %	
% of 6-membered rings	35.0 – 42.0 %	
% of 7-membered rings	17.0 – 31.5	
% ≥ 8-membered rings	11.0 – 17.0	

Table 1: Examples of structural characteristics of the silica glass, obtained either experimentally or from molecular dynamic simulations. Rather than proposing single values, a range of values is reported in the table.

1.2.2 Properties of Silica

The precedent section depicted the structure of the silica glass as consisting of a tri-dimensional arrangement of SiO₄ tetrahedral units connected by their summit and forming ring structures in the glass. Such interconnected, polymerized structure yields the highest known glass T_g , equal to 1480 K according to the calorimetric measurements of Richet and Bottinga [69] (Fig. 13).

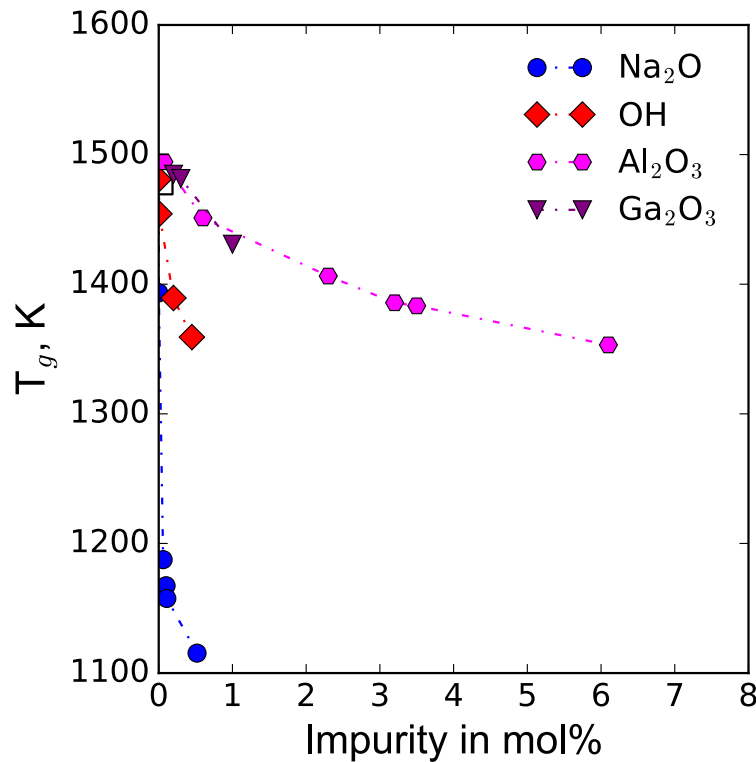


Figure 13: Glass transition temperature T_g of the silica glass as a function of the molar concentrations in Na_2O (blue circles), OH (red diamonds), Al_2O_3 (pink hexagons) or Ga_2O_3 impurities (purple inverted triangles), predicted from the viscosity data of Leko et al. [70] assuming that $\log(\eta) = 12 \text{ Pa s}$ at $T = T_g$. For comparison, the calorimetric T_g of Richet and Bottinga [69] is also reported as a black empty square.

Silica actually is the strongest known liquid, in the strong-fragile denomination of liquids that was introduced by Angell [71, 72]. A strong liquid is characterized by a linear or nearly-linear relationship between the logarithm of its viscosity and the T_g/T ratio, while fragile liquids present strong departure from this linearity. The fragility m of liquids, equal to the slope of the $\log(\eta)$ vs T_g/T curve at T_g , allows one to quantify such deviations. m tends toward ~ 20 for strong liquids such as SiO_2 , and to much higher values for fragile liquids, such as B_2O_3 for instance. The unique interconnexion of the silica tetrahedral network makes the silica glass properties very sensitive to the presence of impurities. For instance, the addition of OH or Na_2O in the silica glass produces a dramatic reduction of hundreds of degrees of the silica glass transition temperature T_g (Fig. 13). Another example is the thermal diffusivity of the silica glass, which varies between 0.795 and $0.870 \text{ mm}^2 \text{ s}^{-1}$ at 298K depending on the impurities present in the glass [73; see Table 2]. Because of such sensitivity, four different types of silica glasses are usually distinguished depending on the level of impurities they contain [74]:

- Type I: glasses produced from natural quartz by electrical fusion under vacuum or under an inert gas atmosphere; they contain about 5 ppm or less of OH groups, but relatively high contents of metal impurities such as Al (30-100 ppm) and Na (~ 4ppm). Commercial names are Infrasil, IR-Vitreosil, G. E. 105, 201, 204;
- Type II: glasses produced from quartz crystal powder by flame fusion (Verneuille-process), using an oxygen-hydrogen flame. They present OH contents ranging between 150 and 400 ppm, but much less metal impurities than type I glasses. Commercial names are Herasil, Homosil, Optosil, O.G. Vitreosil, G.E. 104;
- Type III: synthetic silica glasses produced by hydrolyzation of SiCl_4 . They are free from metallic impurities, but high contents of OH (~ 1000 ppm) and Cl (~ 100 ppm). Commercial names are Suprasil, Spectrosil, Corning 7940;
- Type IV: synthetic silica glasses produced from SiCl_4 in a water vapour-free plasma flame. These glasses, similar to type III glasses, contain ~ 0.4 ppm OH and ~ 200 ppm Cl. Commercial names are Suprasil W, Spectrosil W, Corning 7943.

Another parameter needs to be taken into account when looking at the link between the structure and the properties of the silica glass: the fictive temperature T_f . Indeed, T_g is usually defined as the temperature corresponding to the glass transition upon cooling the melt at a given laboratory cooling rate of ~ 10°/sec. However, cooling the glass much faster or slower than the latter value yields different temperatures for the glass transition. Those different values are referred to as the fictive temperature, T_f . As silica is extremely viscous even at high temperatures, changes of hundreds of degrees of T_f can be obtained easily by varying the cooling rate from hundreds of degrees per seconds to ~ 1° per second. This leads to small variations in the glass density or refractive index, for instance (Table 2). Such variations in the glass density are linked to changes in the silica glass structure: glasses presenting higher T_f also present high concentrations of three- and four-membered rings, as shown by the increasing signals of those rings in the Raman spectra of fast-quenched glasses [75]. However, such causal relationships between T_f , glass structure and properties is complex and depend on the glass composition. For instance, the analysis of a glass with a 75%-25% mixture of SiO_2 - GeO_2 revealed an opposite correlation between the glass T_f , density and the ring statistic [75].

	Value	Remarks	References
Density	2.2002-2.2060 g cm ⁻³	depends on T_f	[74]
Molar Volume	27.24-27.31 cm ³ mol ⁻¹	depends on T_f	Density conversion
Thermal diffusivity D	0.795 – 0.870 mm ² s ⁻¹	at 298 K	[73]
Refractive Index n	1.4583-1.4589	depends on T_f	[74]
Calorimetric T_g	1480 °C	-	[69]
Configurational entropy	5.1 ±2 J mol ⁻¹ K ⁻¹	from calorimetric data	[69]

Table 2: Examples of properties of the silica glass.

In addition to T_f , the densification of the silica glass also modifies the ring statistics, and, hence, its elastic properties. Indeed, upon densification, the frequency of the R band significantly increases, pointing to a puckering of the rings in the network without significant changes in the Si coordination or Si-O bond distance, at least up to pressure of ~ 8 GPa [76, 77]. Such changes are reversible up to 8 GPa. After this pressure, irreversible changes are observed and reflect a change in ring statistics [76]. Compared to the ambient pressure SiO₂ glass, permanently densified glasses typically exhibit higher concentrations of four and particularly three membered rings [78]. The inter-tetrahedral angles in those rings is unaffected by the densification, contrary to the Si-O-Si angles of large rings that significantly decreases of $\sim 6^\circ$ with a 20% densification [79]. Molecular dynamic simulations from Huang et al. [66] further suggest that an increase of dense, α -cristobalite-like 6-membered rings in SiO₂ upon densification may explain the densification mechanism of amorphous SiO₂. The authors of this study indicate that the anomalous elastic property of amorphous SiO₂, i.e. its decreasing elastic modulus upon compression, actually is explained by the conversion of β -cristobalite-like 6-membered rings in α -cristobalite-like 6-membered rings upon compression. Their simulation further shows that the introductions of elements such as rare gas or alkali oxides in the SiO₂ 3D tetrahedral structure further prevent this conversion process, and help reducing the elastic anomaly of the SiO₂ glass.

1.3 Aluminium-free Silicate Glasses

In silicate glasses, other elements are frequently present and play a role different to that of Si⁴⁺. Amongst the major cations present in industrial and geologic silicate glasses, we can cite Al³⁺, Fe³⁺, Ti⁴⁺, Li⁺, Na⁺,

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K^+ , Fe^{2+} , Ca^{2+} and Mg^{2+} (Fig. 1). At the present time, we will focus on a brief description of the Al-free silicate glasses.

As visible in Fig. 13, the properties of the silica glass, such as, for instance, its T_g , can be greatly affected by addition of small fractions of impurities, such as for instance of Na^+ metal cations. The latter actually are considered as *network modifier* cations, which break Si-O-Si bonds, and form ionic bonds with the newly transformed non-bridging apical oxygens. Such process leads to the dramatic depletion in the glass T_g observed in Fig. 13. The following section will describe in more detail such processes at the molecular scale, and their effect on the glass properties.

1.3.1 Network modifier elements and Q^n units

Alkali (Li, Na, K) and alkaline-earth (Ca, Mg) metal cations can be present together with SiO_2 in glasses, forming *sensu stricto* silicate glasses (aluminum free). However, such cations present ionic field strengths ($IFS = Z/r^2$, with Z the cation electronic charge and r its ionic radius) too low to be able to enter the glass structure in tetrahedral units as Si^{4+} does. Indeed, according to Dietzel [80], cations presenting large IFS tend to enter in tetrahedral coordination and act as network former cations. For instance, the IFS of Si^{4+} is equal to 59.1 \AA^{-2} . Other network former elements can present lower ionic field strength, *e.g.* Ge^{4+} in CN 4 presents an IFS of 26.3 \AA^{-2} , but such values still are high enough such that they play network former roles. On the other hand, cation with IFS lower than $\sim 8 \text{ \AA}^{-2}$ do not act as network formers in glasses. Instead, they play either a *network modifier* role, breaking Si-O-Si bonds, or a *charge compensator* role, ensuring electrical neutrality in the vicinity of trivalent network former cations such as Al^{3+} (see sec. 1.4; Al^{3+} IFS = 26.3 \AA^{-2} in CN 4). In Al-free silicate glasses, introduction of network modifiers thus results in glass depolymerisation: network modifiers break Si-O-Si bonds, and form weak bonds with the non-bridging oxygen (NBO) anions at the summit of the SiO_4 tetrahedral units.

As a result of the process previously described, SiO_4 tetrahedral units in alkali and alkaline-earth silicate glasses carry different numbers of NBOs and BOs, and, hence, the local environment of Si^{4+} atoms vary. For instance, the Si-BO distances are shorter than the Si-NBO distances, as revealed by molecular dynamic simulations from Ispas et al. [81]]. This study provides values of 1.63-1.64 \AA for Si-BO distances and 1.58-1.60 \AA for Si-NBO distances in the $Li_2Si_4O_9$ glasses, depending on the potential used for the calculations (empirical or Car-Parrinello *ab initio*). Such difference in the local environment of Si^{4+} results in variations of the frequency of Si-O stretching in different tetrahedral units [82, 58]] as well

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as of the ^{29}Si NMR frequency [83, 84]. Those variations can be detected with using Raman and ^{29}Si NMR spectroscopy, respectively. The study of the Raman and ^{29}Si NMR spectra thus allows observing how the glass polymerization changes with its composition [43, 58, 82, 84–88].

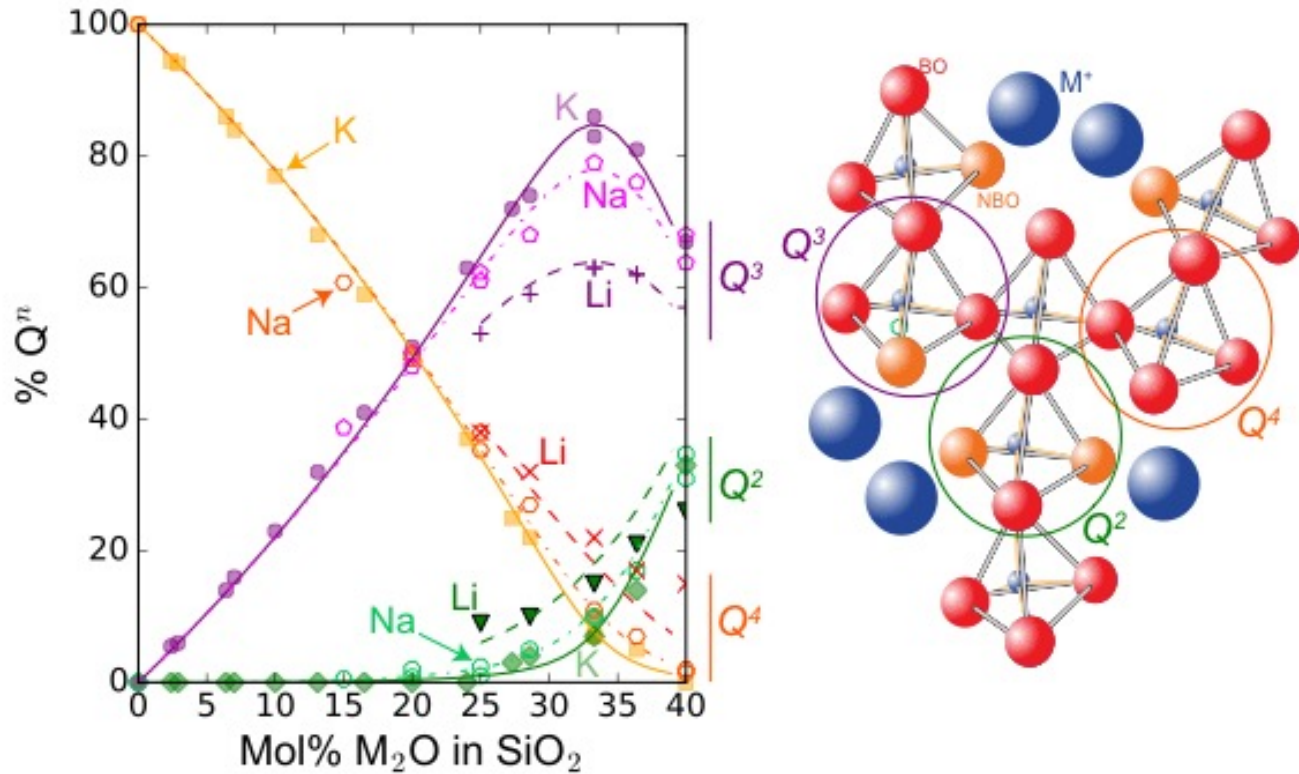


Figure 14: Changes in the fractions of Q^n units measured by ^{29}Si NMR spectroscopy as a function of the alkali metal content of binary $M_2\text{O-SiO}_2$ glasses ($M = \text{Li, Na, K}$). Alkali content is restricted to the 0-40 mol% range, but can reach higher values for $\text{Na}_2\text{O-SiO}_2$ glasses, such that Q^1 and Q^0 units also has been detected in Na_2O -rich silicate glasses. On the right is shown a schematic view of the structure of an alkali silicate glasses, in which the Q^2 , Q^3 and Q^4 units are distinguished; NBO and BO refer to non-bridging and bridging oxygen atoms, respectively, and M^+ refers to alkali network modifier cations. Si^{4+} network former cations are in purple at the center of the tetrahedra. Curves are calculated from the equilibrium constant of the $2Q^3 = Q^2 + Q^4$ relationship, with values of 0.008, 0.02 and 0.08 for the K, Na and Li silicate compositions. Data are from [84, 88–90].

Figure 14 illustrates, for instance, how the fractions of tetrahedral units carrying different numbers of BO and NBOs at their summit vary along the $\text{Li}_2\text{O-SiO}_2$, $\text{Na}_2\text{O-SiO}_2$ and $\text{K}_2\text{O-SiO}_2$ binary joins, from 0 to 40% $M_2\text{O}$ ($M = \text{Li, Na, K}$). As illustrated in this figure, a useful way to distinguish tetrahedral units

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containing different numbers of BO and NBO is the Q^n notation [85], where n refers to the number of BO carried by the SiO_4 tetrahedron Q . For example, the silica glass can be considered as composed of (ideally) only Q^4 units. Contrary to that, a sodium trisilicate NaSi_3O_7 glass (75 mol% SiO_2) presents a distribution of Q^n units with 48% Q^4 , 50% Q^3 and 2% Q^2 [84]. Increasing the concentration in network modifier cations promotes the existence of Q^{n-1} units at the expense of Q^n units in the glasses (Fig. 14), this effect being proportional to the fraction of added network modifiers. In addition, this effect is not ideal and depends on the ionic field strength of the network modifier cation that is present in the glass (Fig. 14). For instance, lithium silicate glasses are richer in Q^2 and Q^4 units than sodium silicate glasses. This translates a tendency of Li^+ network modifier cations to cluster in the vicinity of Q^2 units in the network, leading to local enrichment in Q^4 units by compensation. This clustering may be related to the presence of network modifier percolation channels in the glass structure, as discussed in section 1.3.3. In all cases, increasing the metal cation ionic field strength in silicate glasses shifts the equilibrium:



to the right-end side. One should note that, for a given composition, the equilibrium described by eq. (1) may also depend on the glass fictive temperature, which varies with the glass cooling rate. Indeed, higher temperature promote the shift of the reaction described by eq. (1) to the right, as shown by data from Raman spectroscopy and ^{29}Si NMR spectroscopy of alkali silicate melts [e.g., 91–94].

The knowledge of the Q^n distribution in glasses allows evaluating their ratio of NBO to tetrahedral units, NBO/T, which is a bulk measure of the glass polymerization. The NBO/T can be calculated from knowledge of the Q^n units distribution as:

$$\text{NBO/T} = 4 \times Q^0 + 3 \times (Q^1) + 2 \times (Q^2) + 1 \times (Q^3), \quad (2)$$

or from the knowledge of the molar proportions of the different network formers T, network modifier M^{i+} and oxygen O ions, as [95, 96]:

$$\frac{\text{NBO}}{T} = \frac{2O-4T}{T} = \sum_{j=1}^j \frac{iM_j^{i+}}{T}. \quad (3)$$

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The glass chemical composition thus defines its NBO/T as well as the distribution of its Q^n species. Different models have been proposed to calculate the latter. The binary distribution model, suggested by Dupree et al.[97], predicts that only two species are present in the glass at given SiO_2 content, such that the Q^n units in the glass are distributed as binary combinations of Q^4 - Q^3 , Q^3 - Q^2 , Q^2 - Q^1 and Q^1 - Q^0 as observed in silicate minerals. Contrary to that, the random bonding model predicts that the concentration of a Q^n unit can be calculated from a random distribution of the NBO and BO between the units. Figure 15 shows a representation of the two models.

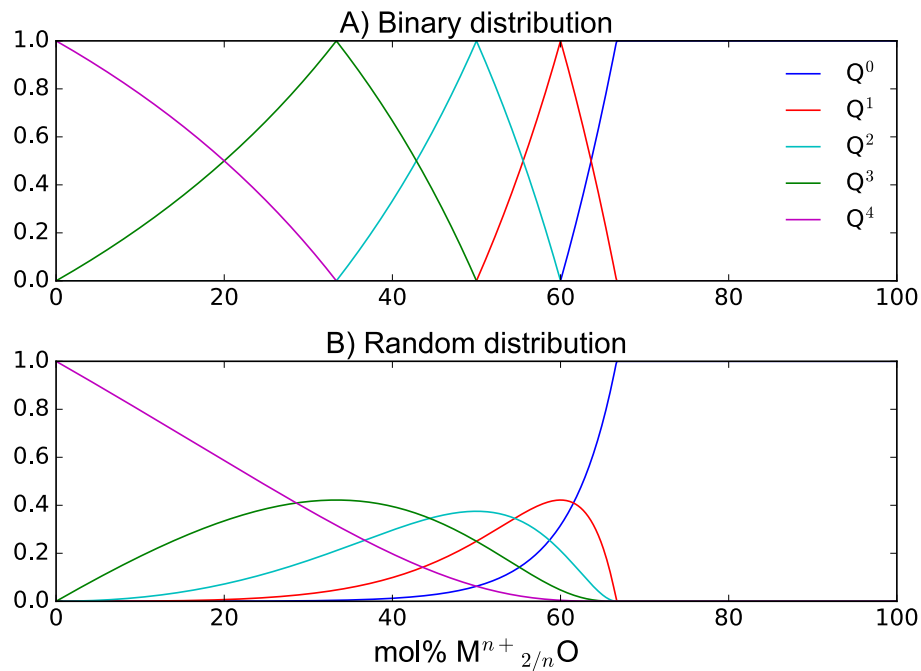


Figure 15: the binary (A) and random bonding (B) models predicting the fractions of the different Q^n units as a function of the mol% of $M^{n+}_{2/n}O$ in SiO_2 (M a network modifier element, such as Li, Na, K, Ca or Mg).

It is clear from the comparison of Figs. 14 and 15 that the reality lies between those two extreme visions. Neither the random bonding model nor the binary distribution model describe the measured distribution of Q^n units observed in silicate glasses. Furthermore, those models do not account for the temperature dependence of the Q^n unit speciation. Indeed, in glasses the Q^n unit distribution is frozen-in, but as highlighted previously, it changes with T in the molten state, such that glasses with different T_g s will present slight variations in their Q^n unit distribution [91–93]. Using the thermodynamic theory of affinity, the model of associated solutions was proposed to solve such issues [98–100]. This model is based on

the idea that melts are media in which chemical reactions, characterized by their standard Gibbs free energies, proceed in the direction of the equilibrium state. The latter is reached when the total Gibbs free energy of the system reaches a minimum. More recently, Gaddam et al. [101] proposed a new model based on statistical mechanics. This model assumes grand canonical ensemble of silicate units, i.e. the Q^n are considered as a statistical ensemble in thermodynamic equilibrium with a reservoir from which they exchange energy and network modifiers. The first test performed with this model, which takes into account the composition and temperature in the calculation of the Q^n unit fractions, are promising and may offer a new way to calculate the distribution of Q^n units in silicate melts. In particular, this model allows calculating the proportions of Q^n_{ijkl} units, where the $ijkl$ subscripts represent the n number of the adjacent tetrahedral units to the central Q unit. For instance, a Q^4 units bonded to 2 other Q^4 units and 2 Q^2 units will be described as Q^4_{2244} . A Q^3 unit bonded to two other Q^3 unit and one Q^4 unit will be written Q^3_{334} , the fourth subscript being omitted because of the NBO carried by the Q^3 unit. Such level of refinement has been reach by using ^{29}Si double quantum NMR spectroscopy, which allows discriminating the different Q^n units bonded to the central Q^n unit (Fig. 16).

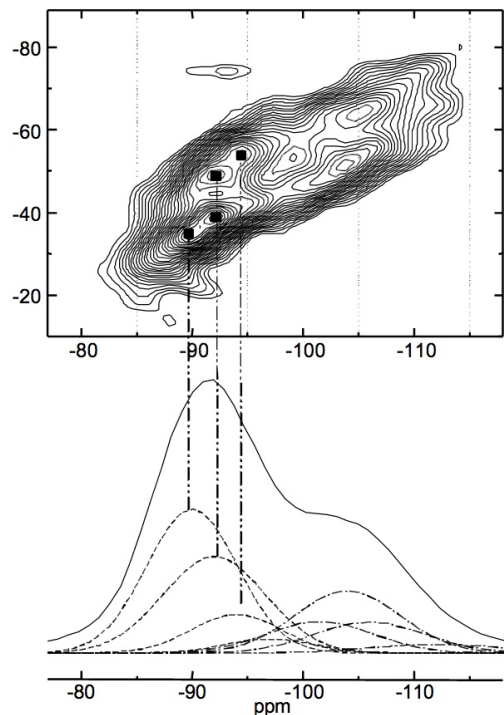


Figure 16: Top: Enlarged ^{29}Si double quantum (DQ) NMR spectrum of a $\text{Na}_2\text{Si}_3\text{O}_7$ glass. The marked peaks with chemical shifts of 89, 92 and 94 ppm on the horizontal axis belong to Q^3_{333} , Q^3_{334} , Q^3_{344} units. These values can be used to fit the 1D ^{29}Si NMR spectrum of the glass with Gaussian peaks, as shown at the bottom. Hatched lines are used for the Q^3_{jkl} lines whereas long hatched lines with dots are used for

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the corresponding Q^4_{jklm} resonances in the linedeconvolution. Reprinted from Olivier et al. [102], Copyright (2001), with permission from Elsevier.

In addition to the statistical mechanic model of Gaddam et al. [101], the random connectivity model of Machacek et al. [103] also allows predicting the different fractions of Q^n_{ijkl} units in the glass network. It assumes that the tetrahedra are linked fully on a random base. The distribution of Q^n_{ijkl} units are calculated using a multinomial distribution. This model can be seen as an extension of the random bonding model presented previously. It thus faces the same limitations, particularly in regards of the effect of temperature as well as the random character of tetrahedral bonding.

1.3.2 Q^n unit distribution: links with glass properties

In the introduction, we showed that the density of an alkali silicate glass increases upon addition of alkali oxide while its T_g dramatically decreases. Such changes actually can be correlated with the changes in the glass structure. In order to calculate the density of silicate glass, Doweidar [24, 104, 105] assigned different partial molar volumes to the different Q^n units. His calculations allow to calculate the density of silicate glass, but, maybe more importantly, give some insights about the volume properties of Q^n units. Indeed, he reported that, while the volume of Q^4 units seem to be constant and equal to that in pure SiO_2 , the volumes of Q^3 , Q^2 and Q^1 units depend on the network modifier present in the glass network (Fig. 17).

First, for a given network modifier, the partial molar volume V_m of the Q^n units increases in the order $V_m(Q^4) < V_m(Q^3) < V_m(Q^2) < V_m(Q^1) < V_m(Q^0)$. This trend is explained by the fact that M-O distances are longer than Si-O distances (Table 3), such that Q^n species surrounded by more network modifiers have higher partial V_m . Turning to the linear correlation between the ionic radius of the network modifiers and the volume of the Q^n units, such volume effect is explained by the difference in the M-O bond length as the ionic radius of the M cation increases. Indeed, results from experiments and molecular dynamic simulations show that the ionic radius and M-O bond length are positively correlated (Table 3).

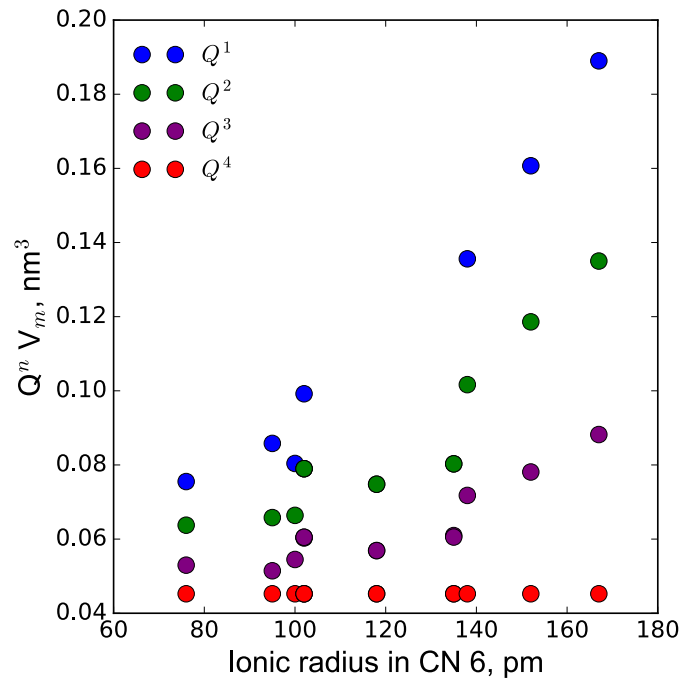


Figure 17: Partial molar volume of Q^n units V_u estimated by Doweidar [105, 106] as a function of the ionic radius of metal network modifier cations (Li, Na, K, Rb, Cs, Ca, Sr, Ba, Cd) in silicate glasses. Ionic radius is reported for a constant CN of 6, using values from Shannon [107]. Variations in the true CN of the modifiers may explain some of the scattering in this relationship [e.g., 108–111].

Table 3 show a comparison of Si-O and M-O distances obtained for alkali tetrasilicate glasses by experimental methods (X-ray and neutron diffraction experiments) and Car-Parrinello molecular dynamic simulations. We note that values from simulation and experiments are in good agreement. Both indicate that Si-NBO distances are shorter than Si-BO distances. Furthermore, in addition to being affected by the ionic radius of the M element, the M-O distance changes as a function of the O atom: M-NBO distances appear to be significantly shorter than M-BO distances.

In addition to affecting the glass density and partial molar volume, the large changes in the polymerization of the glass structure that accompanies the changes in the glass chemistry strongly affect the glass thermodynamic and rheological properties. However, those changes are different. For instance, the glass heat capacity is an additive function of the glass composition, such that it can be calculated using a set of partial molar volumes of the glass oxide components, as provided by Richet [112]. In general, the heat capacity of silicate glasses tends to the Dulong and Petit limit ($=3R$, with R the perfect

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gas constant) as T tends to T_g , such that the glass heat capacity at T_g can be calculated using this limit. However, this relationship should not be extended to other type of glasses, because, for instance, the glass heat capacity of B_2O_3 at T_g only reach values $\sim 0.6 \text{ } 3R$ [see 112 and references therein].

Glass	Si-O	Si-NBO	Si-BO	M-O	M-NBO	M-BO	Method	References
$Li_2Si_4O_9$	1.60(-)	-	-	1.96(-)	-	-	Exp.	[113]
$Li_2Si_4O_9$	1.63(-)	1.58(-)	1.64(-)	1.92(-)	1.87(-)	-	MD.	[81]
$Na_2Si_4O_9$	1.625(11)	1.589(8)	1.65(8)	2.45(40)	2.29(24)	2.52(25)	Exp.	[114]
$Na_2Si_4O_9$	1.63(-)	1.58(-)	1.65(-)	2.28(-)	-	-	MD.	[115]
$K_2Si_4O_9$	-	-	-	2.77	-	-	MD.	[116]

Table 3: Examples M-O bond distances in tetrasilicate glasses reported from molecular dynamic simulations (MD.) and experimental (Exp.) studies.

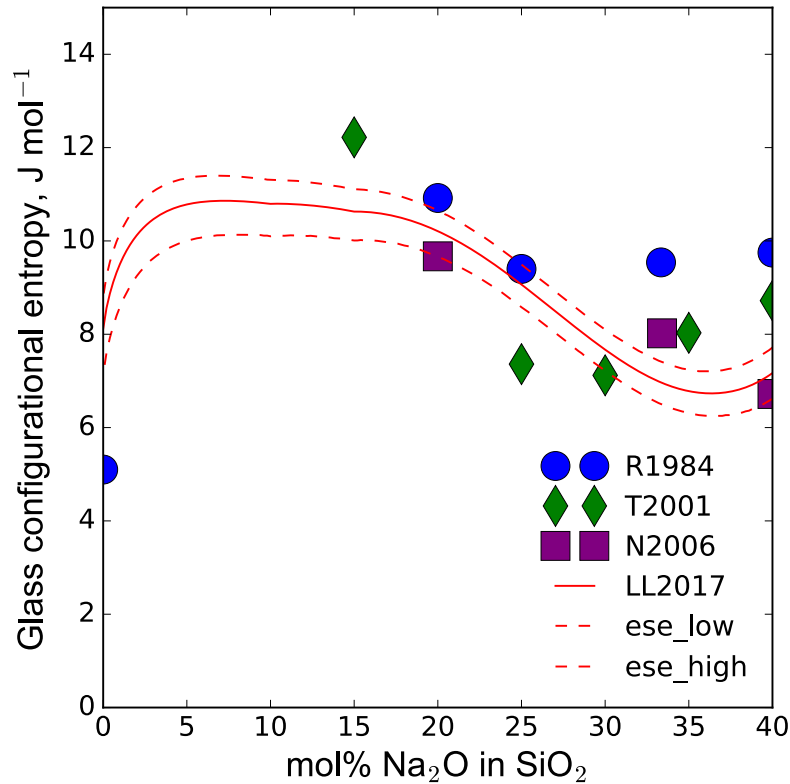


Figure 18: Glass configurational entropy, equal to the melt configurational entropy at T_g , as a function of the mol% of Na_2O added in SiO_2 . Those values were estimated from the fit of the melt viscosity data using the Adam and Gibbs theory of viscous flow (see [chapter XXX](#)), except the value of $5.1 \pm 2 \text{ J mol}^{-1}$

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for SiO₂ that was estimated by [69] using the complete thermodynamic cycle. Errors bars on the points are not reported for clarity; they vary between 0.5 to 2 J mol⁻¹. The red plain line (LL2017 in legend) and its associated errors represented by dashed lines (95% confidence interval) represent the values calculated from the model of Le Losq and Neuville [20]. Other data come from individual fits of the viscosity data by [117; R1984 in legend], [[118; T2001 in legend] and N2006: [[23; N2006 in legend].

While the glass heat capacity is an additive function of the glass composition, the glass configurational entropy at T_g is not. It presents complex and non-linear variations as a function of the glass composition (Fig. 18). Values reported in Fig. 18 were calculated using viscosity data, through the Adam and Gibbs theory that relates the viscous flow of silicate melts to their thermodynamic properties (see [chapter XXX](#)). Values calculated by different authors scatter significantly, showing the sensitivity of the calculation of the glass configurational entropy S^{conf} using melt viscosity data to 1) the quality of the viscosity data and 2) the fitting protocols performed by the different authors. Le Losq and Neuville[[20] recently proposed a simple model to calculate S^{conf} from the distribution of the Q^n species in sodium and potassium silicate melts. This model deviates from an idea originally proposed by Mysen [119], which proposed that the configurational heat capacity of silicate melts C_p^{conf} (equal to the difference between the melt C_p at T and the C_p glass at T_g) can be calculated from partial molar heat capacity values $C_p^{Q^n}$ assigned to different Q^n units:

$$C_p^{conf} = \sum_{n=0}^4 x_{Q^n} C_p^{Q^n}, \quad (4)$$

with x_{Q^n} the fractions of the Q^n units. Figure 19 shows that such model allows reproducing fairly well the C_p^{conf} values derived from calorimetric measurements on alkali silicate glasses and melts.

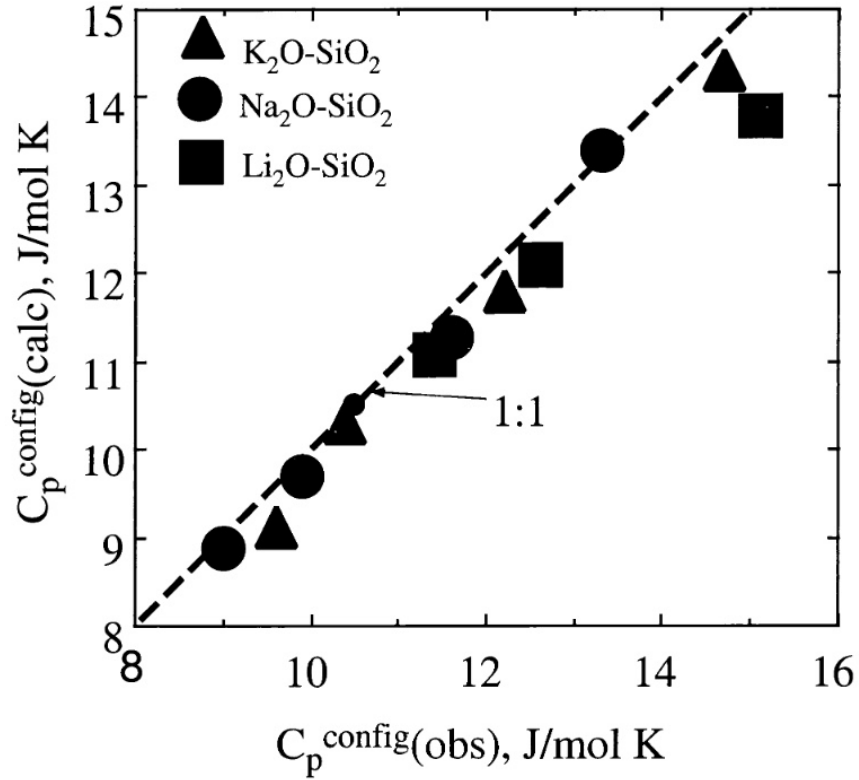


Figure 19: Comparison of the C_p^{conf} values calculated by Mysen [94] using the model resumed in eq. (2) at 1400°C and values derived from measurements. Reprinted from Mysen [94], Copyright (1999), with permission from Elsevier.

Following the idea of Mysen [119], Le Losq and Neuville [20] assigned S^{conf} partial molar values to the different Q^2 , Q^3 and Q^4 units in the melt, allowing one to calculate the topological contribution to the configurational entropy of the glass, i.e. the entropy arising from the disorder in the tetrahedral network in term of Si-O bond distances and Si-O-Si bond length. Indeed, the configurational entropy of glasses can be divided in topological S^{topo} and chemical S^{mix} contributions as:

$$S^{conf} = S^{topo} + S^{mix} . \quad (5)$$

S^{topo} originates from the disorder in the glass network, in term of Si-O bond distances and Si-O-Si bond length for instance. S^{mix} represents the excess of entropy generated by the mixing of different elements in similar structural positions. Le Losq and Neuville [[20] thus calculated S^{topo} as:

$$S^{topo} = \sum_{n=2}^4 x_{Q^n} S_{Q^n}^{conf} , \quad (6)$$

and S^{mix} as:

$$S_{Si}^{mix} = -x_{Si}/x_O * 2R * \sum_{n=2}^4 x_{Q^n} \ln(x_{Q^n}), \quad (7)$$

with R the perfect gaz constant, and x_{Si} and x_O the atomic fractions of Si^{4+} and O^{2-} . In eq. (6), different $S_{Q^n}^{conf}$ values are used for potassium and sodium silicate melts as the coupling between the M alkali modifier and the Q^n tetrahedral species varies with the ionic radius of M. The factor 2 in eq. (7) arises when working using an O_2 base to express the thermodynamic values (i.e. the structural formula is expressed using an O_2 base). This simple model allows one to calculate the configurational entropy of alkali silicate glasses with a good accuracy, falling in between the individual values reported by previous studies (Fig. 18). Adopting a thermodynamic and structural point of view, this model further provides pieces of information regarding the origin of the evolution of the silicate glass T_g along the binary M_2O - SiO_2 join (Fig. 20).

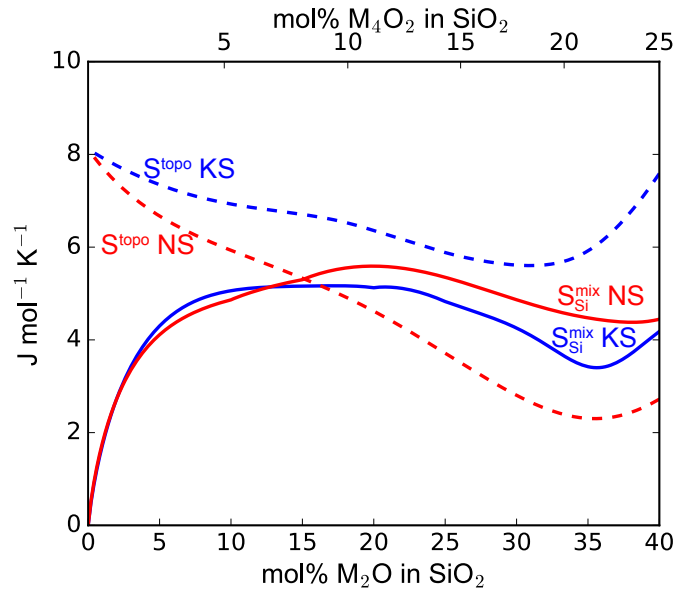


Figure 20: Variation of the topological and mixing contributions to the entropy of potassium and sodium silicate melts as predicted by equation 3 and 4, respectively, as a function of the concentration in alkali oxide in the melt. Values from [20].

Indeed, the model predicts that the apparition of Q^3 species upon addition of a few mol% alkali oxide in SiO_2 leads to a large, dramatic increase in the chemical contribution to the glass configurational entropy because Si can mix between the existing Q^4 and the newly formed Q^3 units. In other term, this illustrates

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that the introduction of NBO in the glass network allows new configurational environments for Si, and lowers the melt viscosity and glass T_g . The model further predicts that the glass topological entropy varies nearly linearly (Fig. 20), being driven by the changes in the Q^n unit distribution. This model remains to be further tested by new research, but together with the works of Mysen [[94, 119] and Doweidar [24, 104, 105], it illustrates that the Q^n unit distribution can be a useful tool for modelling the properties of silicate glasses.

1.3.3 Structure at medium range order and the Mixed Alkali Effect

While at short range order, the structure of silicate glasses can be described as a mixture of interconnected Q^n units with interstitial network modifier elements forming ionic bonds with NBOs as discussed in the previous section, such vision is difficult to extrapolate at the medium range order. Indeed, the network modifier cations have high coordination numbers, and are surrounded by both BO and NBOs from several Q^n units [e.g., 120, 121]. For instance, in aluminum-free silicate glasses, Ca^{2+} and Mg^{2+} coordinence numbers are of $\sim 4\text{--}5$ [109, 122–124], whereas that of Na^+ is ~ 6 [110, 111, 125]. The network modifier cations thus have defined local environment, such that the RN model (see section 1.1) cannot account for their distribution in silicate glasses because it assumes their random repartition in the free volume in the silicate network [126]. An alternative model is thus needed to describe the arrangement of the different network former and modifiers cations in the glass structure.

Two models were proposed to account for the environment of network modifiers in silicate glasses. From the interpretation of X-ray absorption results [126–129], Greaves and co. proposed the Modified Random Network (MRN) model. It depicts the glass structure as a tri-dimensional arrangement of dynamic channels of network modifier metal cations percolating through a disrupted silicate tetrahedral network. Such channels are dynamic features in the molten state, and become frozen-in in the glasses below T_g . However, the interpretations of ^{23}Na and ^{17}O NMR data on sodium silicate and alumino-silicate glasses lead Lee and co. [130–132] to propose an alternative model: the Perturbed Cation Distribution (PCD) model. The PCD model assumes a relatively homogeneous distribution of metal cations in the glass network, neither random nor forming clusters. It describes network modifiers as surrounded by both NBO and BO anions, as indicated by NMR data [e.g., 132]. Whether one adopts the MRN or the PCD model to describe the glass structure, the adopted model needs to explain an effect that greatly affect the glass properties: the Mixed Modifier Effect. This effect is also called the Mixed Alkali Effect (MAE) or

the Mixed Alkaline-earth Effect (MME) in the literature. To avoid confusion, we chose to use the MME denomination in the present handbook.

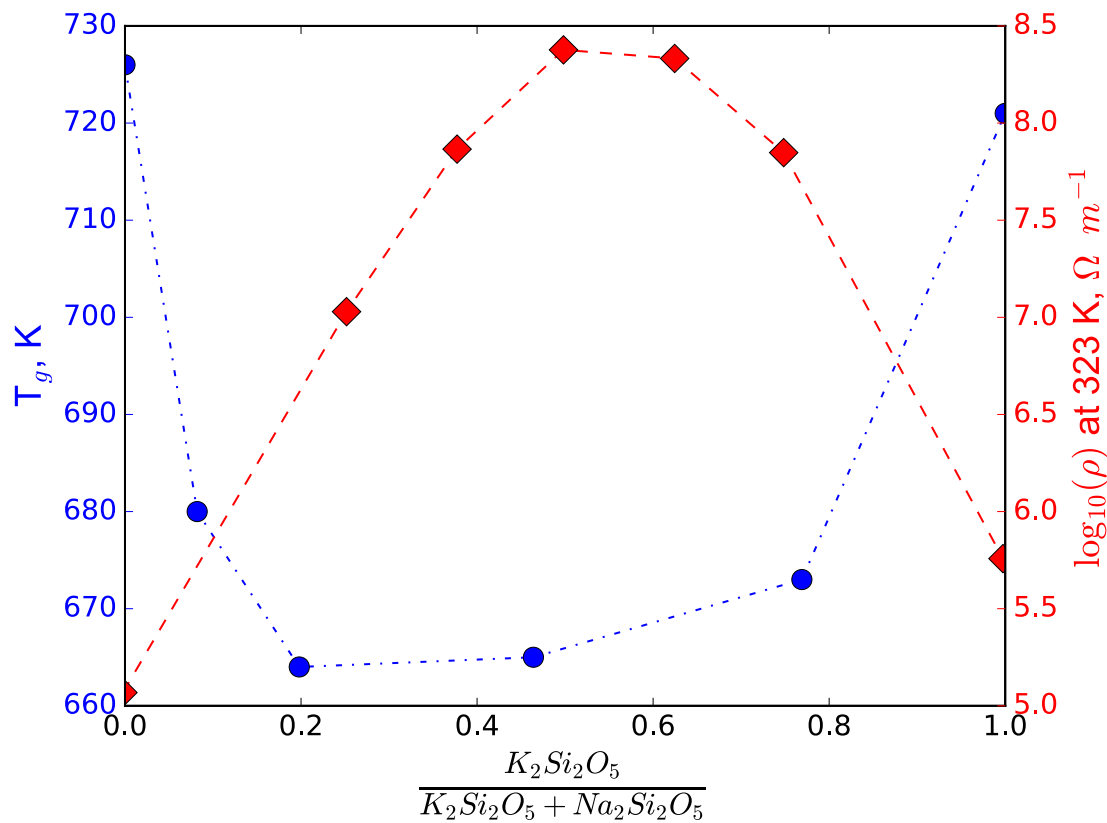


Figure 21: Glass transition temperature T_g and logarithm of the electrical resistivity at 373K of $K_2Si_2O_5$ - $Na_2Si_2O_5$ mixed glasses, as a function of the fraction of K in the glasses. The Mixed Modifier Effect (MME) refers to the non-linear and large deviations of many glasses properties when mixing two network modifier elements in the glasses. Data from [133–135].

The MME refers to the strong and non-linear changes in silicate glass properties that can occur upon mixing two metal cations [e.g., see 136, 137 and references cited therein]. Figure 21 illustrates this effect on the T_g and the electric resistivity of glasses that are mixtures of $K_2Si_2O_5$ and $Na_2Si_2O_5$. Upon mixing Na^+ and K^+ in the glasses, a large depression in the glass T_g is observed while the electric resistivity increases of several orders of magnitude. The MME usually results in maxima or minima of the glass properties when the ratio $A/(A+B)$ between the mixed A and B cations reach 0.5. This effect has been extensively documented for Na and K in silicate glasses [e.g., see 136, 137 and references therein], but it also occurs in silicate glasses when mixing different alkaline-earth cations such as, for instance, Ca and Mg [22, 138, 139]]. In Al-free silicate glasses, mixing alkali modifier cations seems to lead to larger

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deviations to linearity than mixing alkaline-earth modifier cations [139]. The MME also occurs in Al-bearing silicate glasses [140–143], but recent data suggest that it may be more complex than the MME in Al-free silicate glasses [57], such that it will be treated separately in the case of Al-bearing glasses in sec. 1.4.5. In all cases, the MME affects various glass properties, such as, for instance, the glass T_g , Vickers hardness, compressibility, electrical resistivity or ionic conductivity (*e.g.*, Fig. 21).

The source of the MME has been an extensive subject of study, and no real consensus about its origin has yet been reached because of the complexity of ionic diffusion in glasses. The MRN model of the glass structure was initially accounting for this effect by assuming that the diffusions of two different modifier cations were affected by the presence of each other in the channels [128]. Greaves and Ngai [121] further refined the MRN model to explain the MAE. They proposed that the alkali modifiers diffuse in the channels through a mechanism of cooperative intrachannel hopping. According to them, the increase in electrical resistivity observed upon mixing modifier cations (Fig. 21) reflects radical changes and differences in the diffusion coefficients of the modifier cations. Such cation diffusivity can be calculated as a function of the R_{M-M} inter-cationic distances, the hopping attempt frequency ν_o , the diffusion frequency factor D_o , the activation enthalpy W , the Boltzmann constant k and the temperature T as [129]:

$$D_0 = R_{M-M}^2 \nu_o / 6 , \quad (eqDoGN)$$

$$D = D_0 e^{-W/kT} . \quad (eqDGN)$$

Figure 22 shows the results of such calculation: the diffusivity of Na and K are calculated at 25, 200 and 350°C in mixed Na-K trisilicate glasses; this allows the calculation of the electrical conductivity. The agreement between the model and experimental data is also shown, and the model reproduces fairly well the predicted deviations. However, looking in details, it predicts too strong depletion of the electrical conductivity at $K/(K+Na) \sim 0.5$.

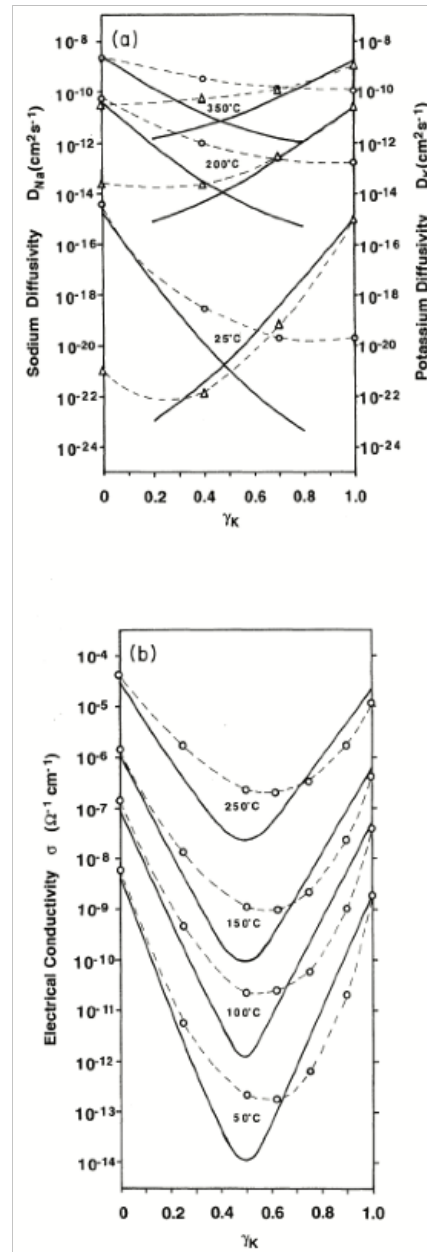


Figure 22: A) Isothermal diffusion coefficients and B) electrical conductivities for the mixed alkali glasses $0.25 [K\gamma KNa1-\gamma K]_2O - 0.75 SiO_2$. Points are experimental data, dashed lines are interpolations of those data, and the plain lines are the prediction by the model proposed by Greaves and Ngai [129]. Reprinted figure with permission from G.N. Greaves and K.L. Ngai, *Physical Review B*. 52, 6358–6380, 1995, doi:10.1103/PhysRevB.52.6358. Copyright (1995) by the American Physical Society

In all cases, the existence of ionic diffusive pathway for modifier cations, as depicted by the MRN model, has been supported by results from atomic force microscopy [e.g., 144], molecular dynamic simulation [e.g., 145, 146], NMR [e.g., 102] and inelastic neutron scattering [e.g., 147–150] experiments. In the

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framework of the Adam and Gibbs [151] theory of viscous flow, which describes liquid relaxation, and hence, viscous flow as a cooperative re-arrangement mechanism of sub-systems in the liquid, the minima in T_g (and, hence, viscosity) observed upon mixing Na and K in silicate melts is explained by an excess of configurational entropy at T_g that originates from a random mixing of pairs of alkali metals [117], as shown by ^{17}O Dynamic Angle Spinning NMR experiments [152]. This interpretation supports the cooperative inter-diffusion and mixing of metal cations in ionic channels percolating through a disrupted silicate tetrahedral network. It is interesting to note that the MAE affects properties that are related to the diffusivity of the network modifiers, but it affects much less properties related to the silicate tetrahedral network. Indeed, mixing alkali elements in silicate glasses yields to nearly-linear variations in the fraction of the Q^n species [see 20 and references cited therein] as well as in that of NBO in the network [152]. Similarly, molar volume or density of silicate glasses are only slightly affected by the MAE [136, 137].

As highlighted previously, not all data support the clustering of modifiers in percolation channels. Lee and co. [130–132] observed that modifier cations are interacting with both BO and NBOs, such that they should be homogeneously distributed in the glass network. Recently, Angeli et al. [153] also interpreted ^{17}O MQMAS data as showing a uniform distribution of Na in sodium silicate glasses, because the two-dimensional projection of the ^{17}O NMR chemical shift and quadrupolar coupling constant distributions for BO oxygens in the $\text{Na}_2\text{Si}_2\text{O}_5$ and the SiO_2 glasses are not overlapping. Therefore, ^{23}Na and ^{17}O NMR results on silicate and aluminosilicate glasses seem to support the alternative Perturbed Cation Distribution (PCD) model, which neither implies a random distribution of the modifiers nor their segregation in specific regions of the network. The PCD model particularly predicts that if different network modifier cations are present in the glass network, they present various degree of ordering between them. The model thus introduces a parameter Q_m that describes the degree of intermixing between different network modifier cation: $Q_m = 1$ for a chemical order favoring dissimilar pairs, $Q_m = 0$ for a random mixing and $Q_m = -1$ for clustering of similar cations [120, 131]. For instance, as indicated earlier, 2D Dynamic Angle Spinning NMR data support a random mixing of Na and K in silicate glasses [152], resulting in an excess of entropy that explains the observed minima in the glass transition temperature during the Na-K mixing (e.g., Fig. 2I). Contrary to that, ^{17}O MAS and 3QMAS NMR data on Ca and Na silicate glasses indicate that those two cations mix in dissimilar pairs [120], this explaining why the random mixing model does not explain the viscosity variations upon mixing Na and Ca in silicate melts [23]. Similarly, Ba shows a stronger affinity for NBO than Na in Ba-Na silicate melts, while in Ba-Mg silicate melts, a Ba and Mg are not randomly ordered around the NBOs [121]. Overall, those results

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lead Lee [131] to suggest that most results from NMR spectroscopy point to Q_m variations in silicate glasses between 0 and 1, and that values < 0 , corresponding to clustering, are unlikely.

While ^{23}Na and ^{17}O NMR from Lee and co. [120, 121, 130–132] or Angeli et al. [153] support the PCD model for accounting of the metal cation distribution in silicate glasses, it must be said that other ^{29}Si NMR data point in favour of the MRN model. Results of Olivier et al. [102] from ^{29}Si Double Quantum NMR experiments and MD simulations indicate that Q^4 units are connected to both Q^4 and Q^3 units, supporting the existence of silica-rich domains as depicted in the MRN model. Performing ^{29}Si and ^{17}O experiments on potassium silicate glasses with SiO_2 contents higher than 75 mol%, Sen and Youngman [88] identified two different type of Q^4 units in silica-rich potassium silicate glasses: some Q^4 are connected to Q^3 units (called Q^{4-3} units by Sen and Youngman) while others are only connected to Q^4 units (called Q^{4-4} units by Sen and Youngman). They proposed that Q^3 and Q^4 units were actually mixing randomly, but this naturally yields to a clustering of Q^3 units in a Q^4 rich network. According to those authors, increasing the mol% K_2O content about a particular threshold, equal to 7.5 mol% K_2O , yields to a three-dimensional percolation of those clusters. This ultimately forms the ionic percolation channels depicted by the MRN model. Sen and Youngman further indicate that such interpretation would explain why results from molecular dynamic simulation from [154] indicated that the MAE was not occurring a very low alkali concentration.

The presence of silica-rich domain in silicate glasses as well as the dependence of the MRN on the modifier nature also is supported by Raman spectroscopy data on alkali silicate glasses. For instance, Matson et al. [155] observed that the alkali ionic radius was influencing the Raman signals in the 200-700 cm^{-1} portion of Raman spectra of alkali silicate glasses. In particular, they observed that the 200-500 cm^{-1} portion of the Raman spectra of Li silicate glasses was similar to that of the SiO_2 glass, as illustrated in figure 23.

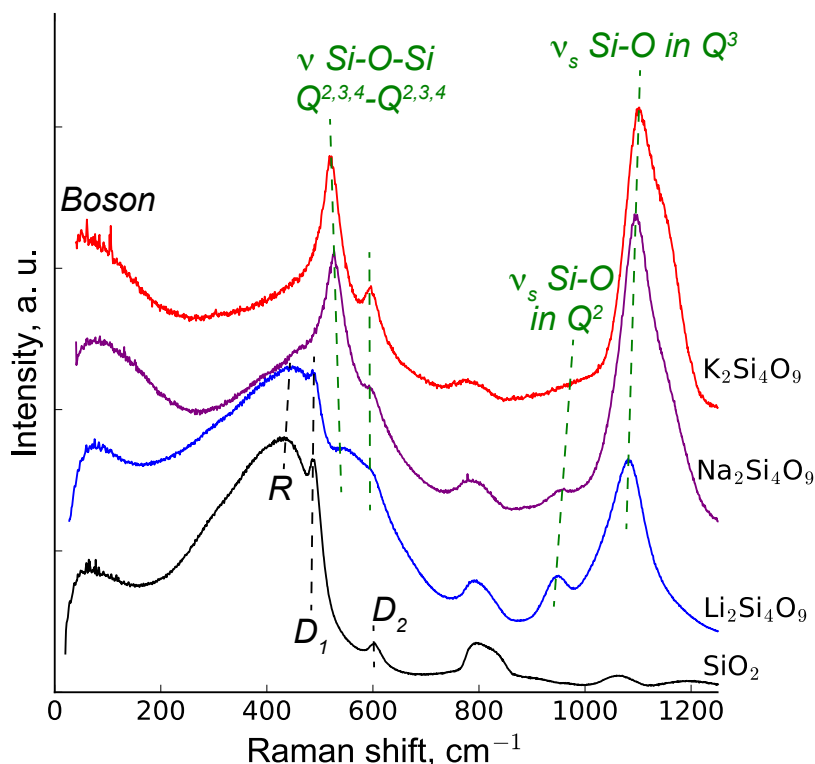


Figure 23: Raman spectra of silica and alkali tetrasilicate glasses. The Boson peak at $\sim 60\text{--}80\text{ cm}^{-1}$ can be assigned to collective transverse acoustic vibrational modes [32] promoted in silicate glasses by cooperative inter-tetrahedral librations [33–35]. In the Raman spectra of the SiO_2 glass, the peaks between 200 and 700 cm^{-1} , with in particular the R, D_1 and D_2 signals, are assigned to vibrations of rings of Q^4 tetrahedral units, see section 1.3.1 and figure 11 for details. In sodium and potassium silicate, new peaks at ~ 520 and $\sim 595\text{ cm}^{-1}$ appear; they do not result from the shift of the pre-existing peaks in SiO_2 , but are assigned to Si-O-Si vibrations involving Q^3 and Q^2 units [43] (those signals are annotated ν_s Si-O-Si $Q^{2,3,4}\text{--}Q^{2,3,4}$ in the figure). The change in the $200\text{--}700\text{ cm}^{-1}$ Raman signal occurring upon addition of Na or K in SiO_2 thus indicate a breaking of the three-dimensional tetrahedral network in those glasses. However, we observe that the $200\text{--}500\text{ cm}^{-1}$ signals from the $\text{Li}_2\text{Si}_4\text{O}_9$ glass are very similar to those in SiO_2 . As this glass contains more Q^2 and Q^4 units than the $\text{Na}_2\text{Si}_4\text{O}_9$ or $\text{K}_2\text{Si}_4\text{O}_9$ glasses (Fig 14), this suggests that it exists silica-like molecular complexes containing Q^4 units in this glass, and, hence, support the modified random network of the structure of silicate glasses. The peaks between 900 and 1300 cm^{-1} are assigned to Si-O vibrations in Q^n units [43, 58, 82, 155]. In particular, the peaks near 950 and 1100 cm^{-1} are assigned to Si-O symmetric stretching vibrations in Q^2 and Q^3 units, respectively. Spectra are from [20, 75, 156].

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Indeed, signals in the $200\text{--}700\text{ cm}^{-1}$ portion of the Raman spectra of silicate glasses are assigned to breathing vibrations of the Si-O-Si bonds [36, 38, 39, 41, 43, 44, 53, 54, 58]. The Raman spectra of LS4 presents a maximum near 460 cm^{-1} and a small sharp peak at 490 cm^{-1} . Those signals are very similar to those found in the SiO_2 Raman spectra, where they are assigned to breathing vibrations of rings with 4 or more tetrahedral units (Fig. 11 and sec. 1.2.1). Contrary to that, the Raman spectra of the sodium and potassium tetrasilicate glasses do not present such signals, but two new peaks centred at 518 and 595 cm^{-1} , assigned to Si-O-Si vibrations involving Q^3 and Q^2 units [43, 155, 157–159]. According to Matson et al[155], the similarity between the Raman spectra of the SiO_2 and $\text{Li}_2\text{Si}_4\text{O}_9$ glasses indicate the presence of silica-rich regions in the Li silicate glass. Therefore, a dissociation of Q^3 species into Q^2 and Q^4 species higher in the Li silicate than in the Na and K silicate glasses (Figure 14) may actually suggest a segregation of Li around Q^2 units, and of Q^4 units in silica-like regions as depicted by the MRN model. For lithium silicate glasses, such interpretation is supported by the $^{29}\text{Si}\{^7\text{Li}\}$ rotational echo double resonance spectroscopy results from Voigt et al[160]]. Indeed, almost independently of the silica content, the Li^7 dipolar field measured at the Q^3 sites is stronger than that measured at the Q^4 sites in Li silicate glasses. This supports the clustering of Li in the network of silicate glasses. To conclude, it should be noted that the precedent interpretation yields to an interesting fact. In alkali silicate glasses, increasing the IFS of network modifiers seems to lead the alkalis to cluster around Q^{n-1} units at the expense of Q^n units, and, hence, affect the distribution of the Q^n units in the glass (Fig. 14). As a result, the Q^n distribution actually tends toward that predicted by a random mixing of NBO and BO amongst Q^n units (Fig. 15), while one may have been tempted to draw the opposite conclusion. This arises from the fact that randomness is not correlated with homogeneity, but, on the contrary, rather may lead to heterogeneity as proposed by Sen and Youngman [88].

1.4 Aluminum in silicate glasses

Addition of aluminum in silicate glasses is known to improve their mechanical properties, such as, for instance, their chemical durability or their hardness. Aluminum is thus vastly used in technical glasses in the industry to confer interesting properties to the final products. For instance, at the present era of smartphones and tablets, one well known example of an alumino-silicate glass is the Gorilla[®] Glass from the Corning[®] company. However, the interest in alumino-silicate glasses is not limited to the industrial domain. Indeed, alumino-silicate glasses are common products of geological processes, deriving from the rapid quench of magmatic liquids. For instance, they can be entrapped in crystals in deep magma

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chambers, and, once retrieved at the surface, they provide fundamental information about geological processes on Earth [e.g., 161–163] or the Moon [164, 165]. On a fundamental note, if we assume the presence of a magmatic ocean at the surface of the Earth after its formation 4.567 Ga ago [see 166 and references cited therein], it then must be said that the properties of the aluminosilicate melt composing the magma ocean drove the segregation of Earth iron core and the formation of Earth mantle and crust. This illustrates the importance of aluminium in silicate melts, which, hence, deserves a particular attention.

1.4.1 Local environment of aluminum

As previously highlighted, numerous studies have documented the effect of the addition of aluminum in silicate composition on the glass properties. For instance, adding Al in a NaSi_3O_7 glass produces a large, non-linear increase in its glass transition temperature (Fig. 24). Such variation points to the fact that Al^{3+} enters in the structure of silicate glasses as a network former element, mostly in tetrahedral coordination [83, 167–170].

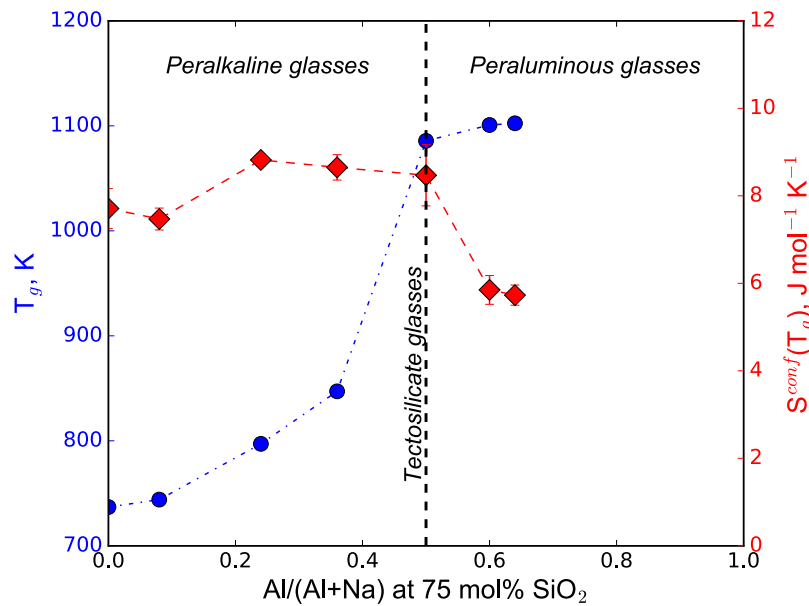


Figure 24: Glass transition temperature T_g and configurational entropy at T_g of aluminosilicate glasses containing 75 mol% SiO_2 and different $\text{Al}/(\text{Al}+\text{Na})$ ratios. Peralkaline glasses present an $\text{Al}/(\text{Al}+\text{M})$ ratio < 0.5 ($\text{M}=\text{Na}^+, \text{K}^+, \text{Mg}^{2+}, \text{Ca}^{2+}$, etc.), tectosilicate glasses a ratio of 0.5 ($[\text{Al}^{3+}] = [\text{M}^{+,2+}]$), and peraluminous glasses present an $\text{Al}/(\text{Al}+\text{M})$ ratio > 0.5 . Data from [25].

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Results from ^{27}Al NMR spectroscopy show that Al also can be present in 5- and 6-fold coordination, in low concentrations that depend on the glass composition [25, 124, 171–175]. Indeed, the ratio between Al and the alkali/alkaline-earth metal cations $\text{M}^{+,2+}$ in the glasses, Al/M, is of fundamental importance in determining the role of the metal cation as well as the coordination and environment of aluminum. When $\text{Al/M} < 1$, the composition is peralkaline (Fig. 24). In such case, one part of the M cations compensates the electrical charge deficits around AlO_4^- units, due to the fact that Al^{3+} substitutes for Si^{4+} in the tetrahedral units. The remaining part of M cations acts as network modifiers. When $\text{Al/M} = 1$, the glass has a tectosilicate composition. The amount of aluminum and metal cations is the same, such that all metal cations should compensate for the charge deficit of AlO_4^- units and the glass should be fully polymerized, i.e. ideally constituted of only Q^4 units. In reality, as ideality and glasses are antagonistic concepts, a few percent of NBO can be present in tectosilicate glasses [176–178] (sec. 1.4.4). When $\text{Al/M} > 1$, the composition is peraluminous: there is an excess of aluminum in the network compared to charge-balancing metal cations.

From ^{27}Al NMR data [25, 171–174], a small fraction of Al is present in 5 ($^{[5]}\text{Al}$ or AlO_5) and 6 ($^{[6]}\text{Al}$ or AlO_6) fold coordination in tectosilicate and peraluminous glasses; the higher the Al/M, the higher the fractions of $^{[5]}\text{Al}$ and possibly of $^{[6]}\text{Al}$. Figure 25 illustrates this dependence in the $\text{CaO-Al}_2\text{O}_3\text{-SiO}_2$ diagram. In this ternary diagram, a significant fraction of Al already is in CN 5 in the peralkaline domain. At all silica contents, increasing the Al content up to reaching the peraluminous domain leads to a very important increase of the $^{[5]}\text{Al}$ fraction in the glasses. In peraluminous Ca-bearing glasses, Neuville et al. [174] also reported a small fraction ($< 3\%$) of $^{[6]}\text{Al}$ at very high Al/Ca ratios.

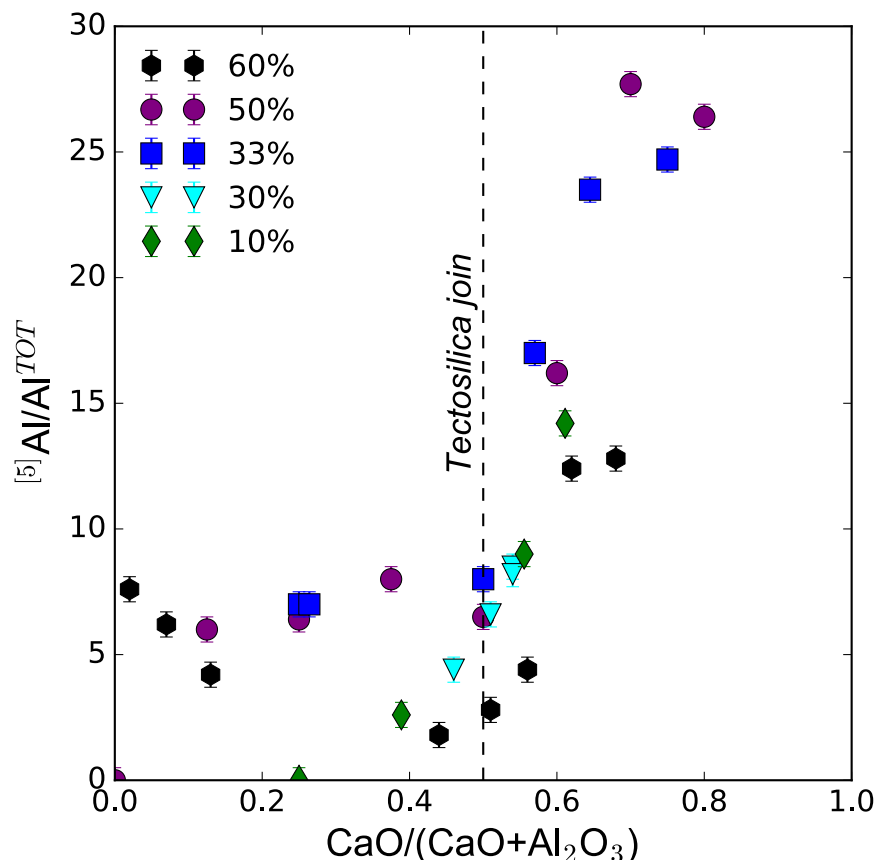


Figure 25: fraction of $^{[5]}\text{Al}$ (five-fold coordinated Al) represented as a function of the ratio $\text{CaO}/(\text{CaO}+\text{Al}_2\text{O}_3)$ in silicate glasses. Data acquired through ^{27}Al NMR spectroscopy at different silica concentration are represented by black hexagons (60 % SiO_2), purple circles (50%), blue squares (33%), inverted cyan triangles (30%) and green diamonds (10%). Data at 50, 33 and 10% are from Neuville et al. [174] and data at 30 and 60 % are from Thompson and Stebbins [178].

The fractions of $^{[4]}\text{Al}$, $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ not only depends on the Al/M ratio, but also on the ionic field strength of the M cation [179–182]. Indeed, Ca aluminosilicate glasses already present a few percents of $^{[5]}\text{Al}$ in peralkaline glasses [174, 183], while Na aluminosilicate glasses don't [25]. In general, alkaline-earth aluminosilicate glasses present higher fractions of $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ than alkali aluminosilicate melts. Differences are also observed between different alkaline-earth aluminosilicate glasses: for instance, Sr, Ca and Mg tectosilicate glasses with 50 mol% SiO_2 present $^{[5]}\text{Al}$ fractions of 0.02, 0.07 and 0.12, respectively [see also results from 182, 184]. In Y and Lu tectosilicate glasses with the same SiO_2 content (50 mol%), $^{[5]}\text{Al}$ fractions reach ~ 21 and 26 %, respectively [185]. Figure 26 illustrates the correlation between the fraction of $^{[5]}\text{Al}$ and the ionic field strength of metal cations: increasing the IFS of the metal cations in aluminosilicate glasses seems to promote $^{[5]}\text{Al}$ at the expense of $^{[4]}\text{Al}$.

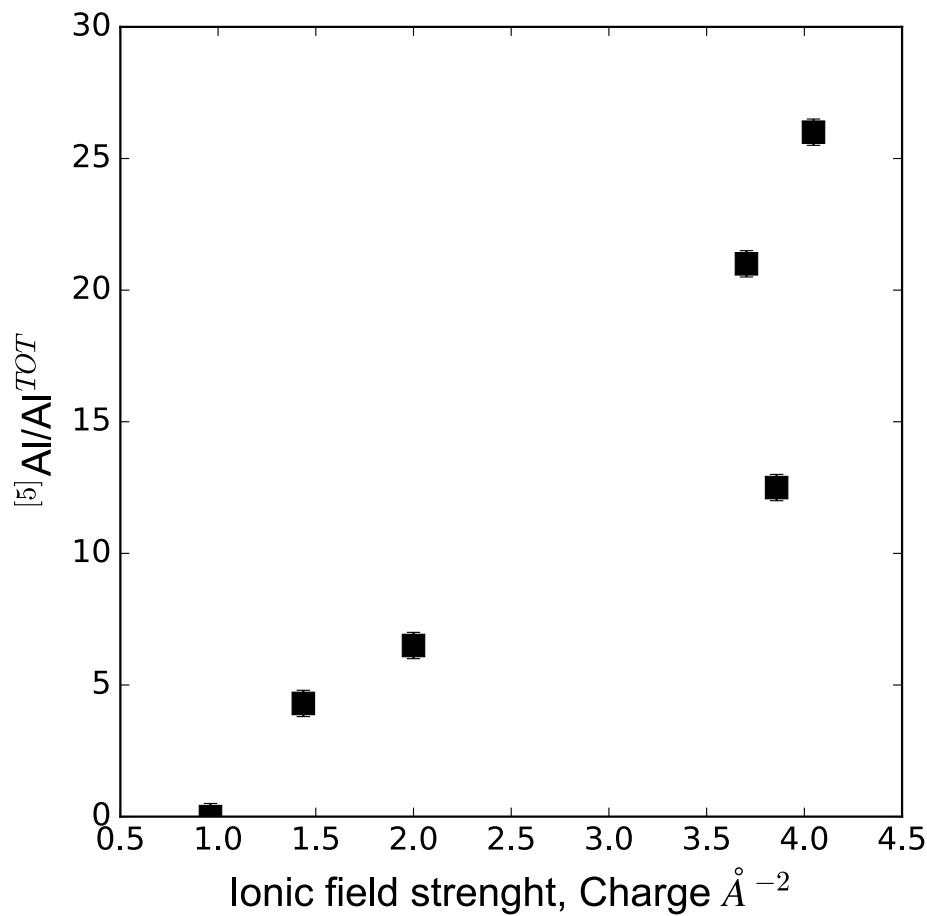


Figure 26: Fraction of $^{[5]}\text{Al}$ detected by ^{27}Al NMR spectroscopy in tectosilicate glasses with 50 mol% SiO_2 as a function of the ionic field strength of the non-network former cations in the glass (i.e., Sr, Ca, Mg, Na, Y and Lu). For ensuring consistency, and because the CN of metal cation in tectosilicate glasses is not perfectly known, the cation ionic field strength has been calculated at a constant CN of 6 for all cations. Therefore, scattering in this figure may result from variations in the CN of the different cations. Ionic radii have are from Shannon [107]; data are from [174, 184–186]

In addition to the effect of the glass chemical composition, the conditions of the glass formation also influence the fractions of $^{[4]}\text{Al}$, $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ in the glass. Increasing the parental melt temperature and pressure leads to increase the coordination of Al in the retrieved glasses, particularly for tectosilicate and peraluminous compositions [179, 187, 175, 188–190, 25]. As a result, the glass quench rate and quench pressure affects the glass fractions of $^{[4]}\text{Al}$, $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$. Glasses formed at high pressure (several GPa) will contain large amounts of $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$, with several percent to tens of percent increases in their proportions depending on the pressure in the 1-10 GPa range [[179, 180, 190, 191].

The effect of temperature on the fraction of highly coordinated Al polyhedral units in the glass is more complex. Indeed, different trends have been reported as a function of the glass composition. In pure alkali or alkaline-earth alumino-silicate glasses, increasing the quench rate and, hence, glass T_f generally leads to a potential increase of a few percent the fraction of $^{[5]}\text{Al}$, as supported by ^{27}Al NMR data from Stebbins et al. [192] and Thompson and Stebbins [193] on Ca alumino-silicate glasses. In agreement with that, Allwardt et al. [187] reported increasing $^{[5]}\text{Al}$ fraction in a $\text{Na}_3\text{AlSi}_7\text{O}_{17}$ glass (synthesized at 10 GPa) with increasing the glass T_f . However, they also reported a decreasing $^{[5]}\text{Al}$ fraction in a $\text{NaAlSi}_3\text{O}_8$ glass (also synthesized at 10 GPa) with increasing the glass T_f . Such observation is intriguing since Le Losq et al. [25] measured *in situ* higher fractions of $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ with increasing temperature in $\text{NaAlSi}_3\text{O}_8$ and other Na alumino-silicate melts (peralkaline and peraluminous), using *in situ* laser-heated ^{27}Al NMR spectroscopy (Fig. 27). Therefore, an additional effect (potentially of pressure) may have affected the temperature effect in the observations of Allwardt et al. [187] on the $\text{NaAlSi}_3\text{O}_8$ glass. In a Na-K alumino-silicate glass synthesized at a pressure of 3.5 GPa, Malfait et al. [190] also reported negligible changes in the Al speciation with changing the glass T_f .

Therefore, from this review, increasing T_f in alkali and alkaline-earth alumino-silicate glasses tend to increase the fraction of highly coordinated Al species. This effect is small compared to that of pressure: it is only of a few % for quench rate variations of several orders of magnitude [e.g., 193]. Furthermore, it also probably depends on the melt composition. The latter dependence is illustrated in Fig. 27, which shows ^{27}Al NMR spectra acquired on Na alumino-silicate glasses at room temperature and acquired on their corresponding melts at high temperature [25]. While ^{27}Al NMR spectra from peralkaline glasses and melts ($\text{Na}/(\text{Na}+\text{Al}) < 0.5$) show a peak at a comparable chemical shift that can be assigned to $^{[4]}\text{Al}$, the tectosilicate and peraluminous glasses and melts ^{27}Al NMR spectra present significant differences: they are broader and, for the peraluminous composition, shifted to lower ^{27}Al chemical shifts. Those differences are assigned to the presence of $^{[5]}\text{Al}$ and even $^{[6]}\text{Al}$ in those tectosilicate and peraluminous composition.

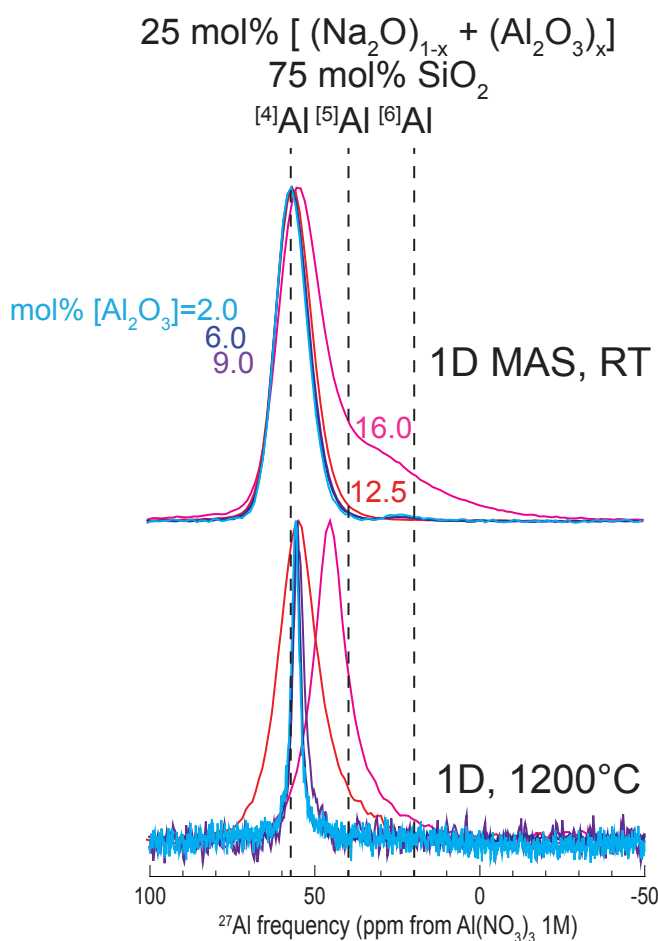


Figure 27: ^{27}Al NMR spectra acquired at room temperature and at 1200°C on sodium aluminosilicate glasses and melts presenting different Na/Al ratios. The peralkaline glasses and melts show no significant difference, as Al is at $\sim 99\%$ in CN4. In tectosilicate and peraluminous compositions, the broadening of the ^{27}Al NMR signal indicate the existence of Al in 5- and even 6-fold coordination at high temperature.

From observations shown in Fig. 27, effects of melt composition, discussed previously in the case of glasses, seem to be enhanced in the molten state. As a result, depending on the glass composition, changing the quench rate leads to different outcomes in term of the Al coordination number. Data from boro-alumino-silicate composition further corroborate this idea. Indeed, while both Morin et al. [182] and Wu and Stebbins [194] reported decreasing coordination of Al with increasing the T_f of Na, Ba, Ca, La and Y aluminoborosilicate, Kiczinski et al. [195] reported an inverse trend in the multi-component boro-alumino-silicate E glass¹, used as the reinforcing phase in fiberglass. Morin et al. [182] explained

¹ Glass E composition in mol%: 57.3% SiO_2 , 6.0% B_2O_3 , 8.7% Al_2O_3 , 24.9% CaO , 0.9% MgO , 0.8% Na_2O , 0.8% F_2 .

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their observations by a coupling between B and Al species in the melt, and suggested that the reverse trend reported by Kiczinski et al. [195] may arise from an increase in association of F with Al at higher temperatures. Therefore, it appears that coupling of Al with other network former elements such as boron or with volatile elements may have an influence on the final coordination of Al in the glass, and its dependence on the glass T_f .

1.4.2 Aluminum coordination and glass properties: examples

The changes in the aluminum coordination number are affecting various properties of the glass. In particular, increasing the aluminum coordination is an important component in the densification mechanism of aluminosilicate glasses [179]. Indeed, there is a clear trend between the change in the glass molar volume with pressure and the average Al coordination (Fig. 28). The scatter visible in Fig. 28 indicates that other mechanisms occur, such as changes in the mean inter-tetrahedral angle and compression of the sites of modifier and charge-compensator cations [179, 180, 191, 196]]. The melt composition may be an important control in determining which mechanism prevails [180, 191].

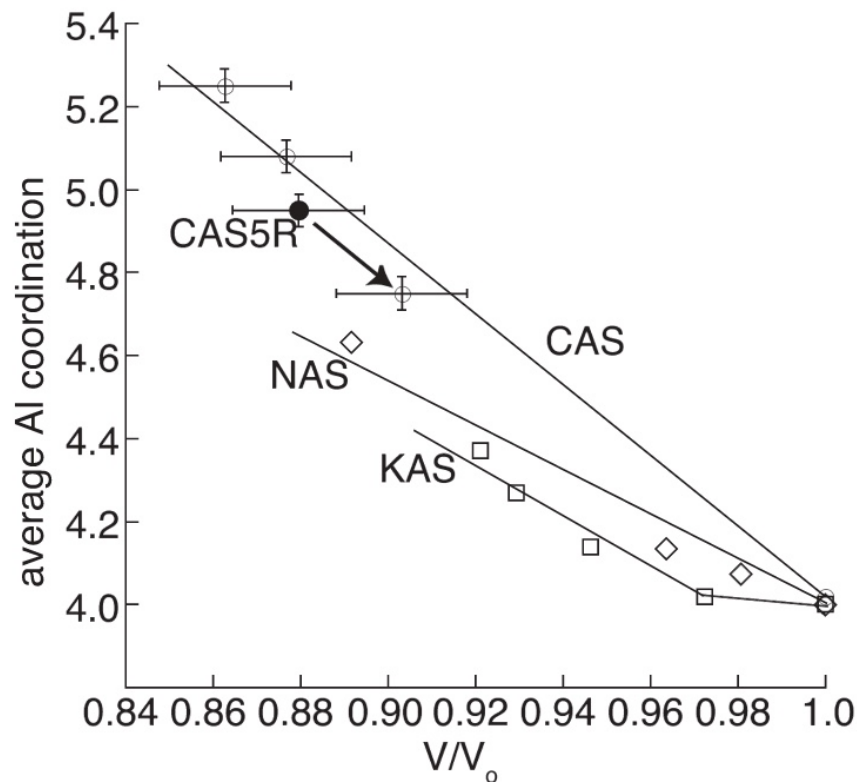


Figure 28: Average Al coordination in $\text{Ca}_3\text{Al}_2\text{Si}_2\text{O}_{18}$ (CAS & CAS5R), $\text{Na}_3\text{AlSi}_3\text{O}_9$ (NAS) and $\text{K}_3\text{AlSi}_3\text{O}_9$ (KAS) glasses reported as a function of their relative molar volume, calculated as the ratio between their

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molar volume at high pressure (3 – 10 GPa) and that at room pressure. Reproduced from Allwardt JR (2005) Aluminum coordination and the densification of high-pressure aluminosilicate glasses. American Mineralogist 90:1218–1222. doi: 10.2138/am.2005.1836, Copyright (2005), with permission from the Mineralogical Society of America.

In addition to the role of highly-coordinated Al species in densified glasses, their presence affect the mechanical properties of the glasses. By combining molecular dynamic simulations and MAS NMR spectroscopy, Iftekhar et al. [185] showed that, in $\text{Re}_2\text{O}_3\text{-Al}_2\text{O}_3\text{-SiO}_2$ (Re = Y, Lu) glasses, there is a direct correlation ($R=0.9$) between the Vickers hardness and the mean coordination number of Al (Fig. 29).

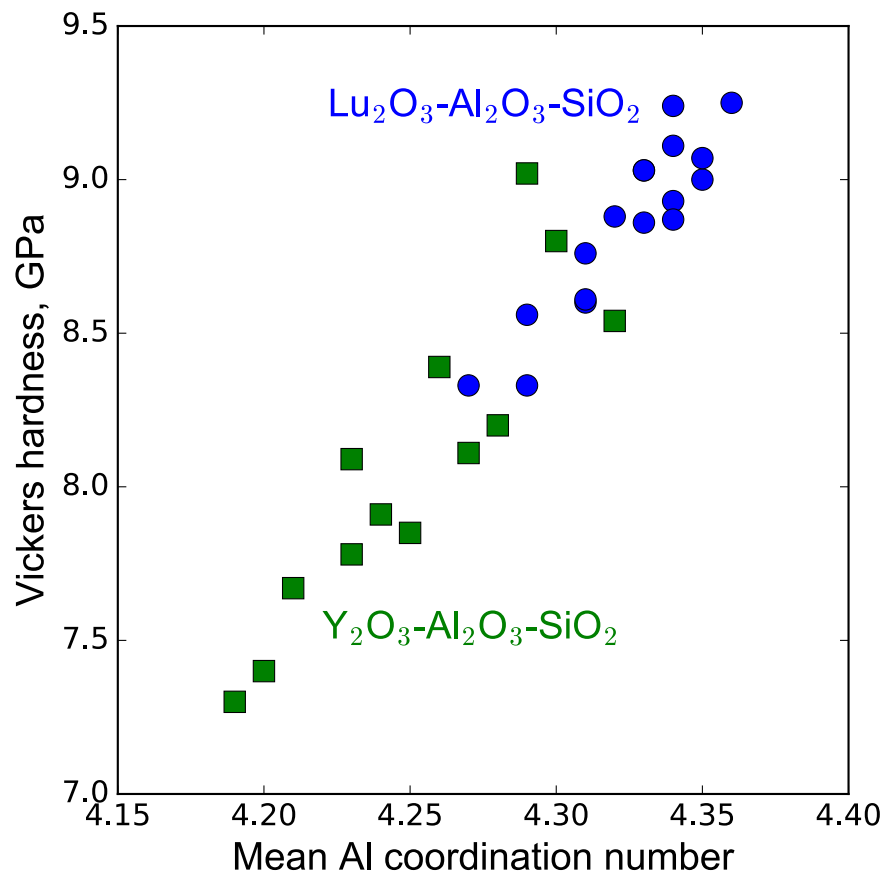


Figure 29: Vickers hardness versus average Al coordination in $\text{Re}_2\text{O}_3\text{-Al}_2\text{O}_3\text{-SiO}_2$ (Re = Y, Lu) glasses with various compositions ($0.99 \leq \text{Al/Lu} \leq 2.4$ and $\sim 49 \leq \text{mol\% SiO}_2 \leq \sim 65$ for $\text{Lu}_2\text{O}_3\text{-Al}_2\text{O}_3\text{-SiO}_2$ glasses; $0.72 \leq \text{Al/Y} \leq 1.77$ and $\sim 41 \leq \text{mol\% SiO}_2 \leq \sim 68$ for $\text{Y}_2\text{O}_3\text{-Al}_2\text{O}_3\text{-SiO}_2$ glasses). Data from [185].

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Iftekhar et al. [185] interpreted such relationship as showing that $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ form strong $^{[5]}\text{Al-O-T}$ and $^{[6]}\text{Al-O-T}$ bonds in the network, enhancing its connectivity, and, hence, its structural strengths. The increase in connectivity following the apparition of $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ in the glass structure is further corroborated by the increase in the T_g of glasses with increasing the fraction of aluminium in peraluminous glasses (Fig. 24; [25]). The variation of the viscosity of Na peraluminous silicate melts further support such idea. Indeed, close to T_g , the viscosity of Na alumino-silicate melts increases with increasing the $\text{Al}/(\text{Al}+\text{Na}^+)$ ratio in the peraluminous domain at iso-temperature [[25, 169], while at high temperature it decreases [168, 197]. This effect occurs because of the decoupling between the alumino-silicate and metal cation sub-lattices relaxation at T_g [198]: Na^+ cations diffuse rapidly in a frozen-in polyhedral alumino-silicate network at, or close to, T_g . In such conditions, $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ participate in ordering the network in glasses, as shown by the large decrease in the glasses configurational entropy at T_g with increasing $\text{Al}/(\text{Al}+\text{Na})$ from 0.5 to 0.7 in Na alumino-silicate compositions (Fig. 24). At high temperature, relaxation of all cations tends toward the same values. In such conditions, the presence of $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ in the melts seems to favor the diffusion of oxygen and, hence, to enhance viscous flow [25, 199].

1.4.3 Aluminum, glass polymerization, and the Al avoidance rule

Sections 1.4.1 and 1.4.2 illustrated that aluminum enters preferentially in CN 4 in glasses, the proportion of CN 5 and 6 being small and dependent on the glass pressure synthesis and composition, and, in a minor extent, T_f . The overall effect of adding aluminium in silicate glasses thus results in increasing their polymerization, as Al enters as a network former into the glass ionic framework. In particular, Al enters in Q^4 units in the glasses with a tendency to avoid formation of Al-O-Al bonds (the so-called Al-avoidance rule or Lowenstein rule [200]) that depends on the melt chemical composition [131, 201, 202]. According to Mysen et al. [203], Al in Q^4 units represent at least 70% of the total aluminum in the peralkaline domain in Na alumino-silicate glasses. A resulting effect of aluminum preference to reside in Q^4 units in glasses is to shift the equilibrium $2 Q^3 \rightleftharpoons Q^2 + Q^4$ to the right end side at constant NBO/T values with increasing the ratio $\text{Al}/(\text{Al}+\text{M}^{n+}_{2/n})$ in peralkaline glasses [203]. The least depolymerized units in the glasses present the lowest Al/Si ratio, i.e. Al avoids entering in depolymerized units.

The “preference” of Al for polymerized environments probably results from the minimization of the energy deficit of Al polyhedral units: carrying NBO will lead to a very strong electrical charge deficit for AlO_z units ($z = 4, 5, 6$). The Al-avoidance effect further results of such concept: Al-O-Al bonds are

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avoided as they are energetically not favorable. However, Al-avoidance is not perfectly respected in alumino-silicate glasses, as shown by results from ^{29}Si Magic Angle Spinning (MAS) and ^{17}O triple quantum MAS (3QMAS) spectroscopy on Ca, Li and Na alumino-silicate glasses [[132, 201, 202, 204] as well as by results from MD simulations [e.g., 205, 206]. For instance, results for ^{29}Si NMR spectroscopy show that the degree of Al avoidance in Na alumino-silicate melts is higher than that in Ca alumino-silicate glasses [201]. A useful parameter for quantifying such degree of Al avoidance is the Q parameter: $Q = 1$ in case of perfect Al avoidance (no Al-O-Al bonds), and 0 in case of a purely random Al-Si distribution [131, 201]. In alkaline-earth and alkali tectosilicate melts, Q values range from 0.65 to 0.99; Figure 30 show how Q values decrease with increasing the ionic field strength of the charge compensating metal cation in the glass framework [201, 204].

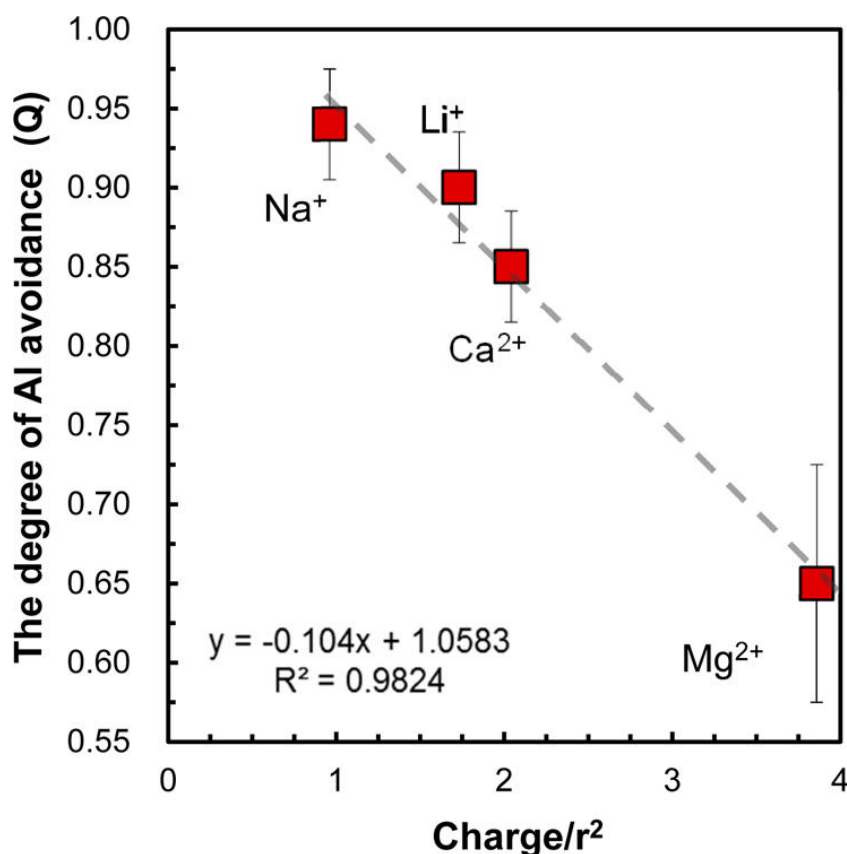


Figure 30: Relationship between the degree of Al avoidance in alumino-silicate melts and the ionic field strength of the metal cation. Reprinted with permission from Lee SK, Kim H-I, Kim EJ, Mun KY, Ryu S (2016) Extent of Disorder in Magnesium Aluminosilicate Glasses: Insights from ^{27}Al and ^{17}O NMR. *J Phys Chem C* 120:737–749. doi: 10.1021/acs.jpcc.5b10799 Copyright (2016) American Chemical Society.

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It thus seems that the IFS of metal cations influences linearly the degree of Al avoidance in aluminosilicate glasses. However, Lee [207] warrant that increasing the cation IFS at values higher than that of Mg^{2+} may lead to deviation from the linear trend observed in Fig. 30, because of the formation of highly-coordinated $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ species and possibly of Al-NBO bonds in such systems.

The degree of Al avoidance can further be related to the thermodynamic properties of the glass, and, in particular, to their configurational entropy. Indeed, the Si-Al ordering implies a non-random process for the integration of Al in Q^4 units. The analysis of the Raman spectra of tectosilicate glasses reveal the presence of two bands at $\sim 1140\text{-}1170$ and $1180\text{-}1210\text{ cm}^{-1}$ (Fig. 31), which are assigned to Si-O stretching in Q^4 units [30, 57, 58, 208].

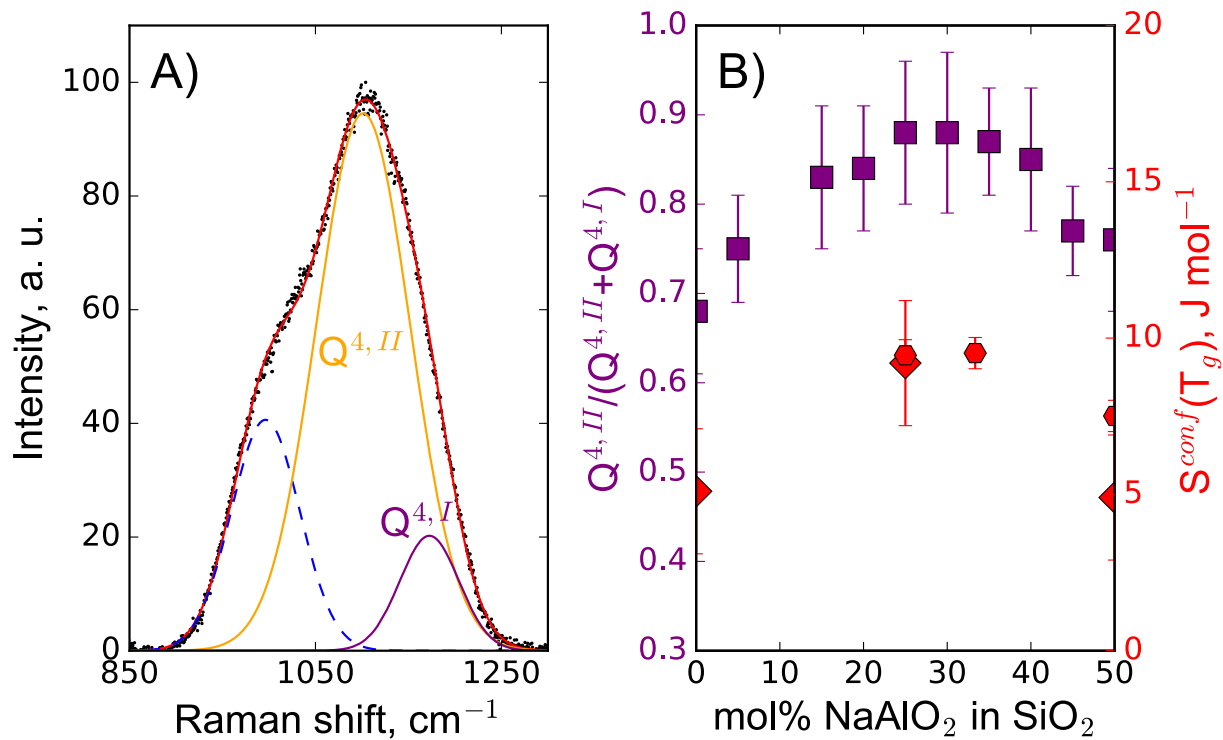


Figure 31: A) 850-1300 cm^{-1} portion of the Raman spectrum of the $\text{NaAlSi}_3\text{O}_8$ glass, presenting an Al/Si ratio of 0.5. The dotted blue peak is assigned to T-BO asymmetric stretching in TO_4 units, while the peaks in purple ($Q^{4,I}$) and orange ($Q^{4,II}$) are assigned to T-O symmetric stretching in TO_4 units [25, 57 and references therein]. In B), the ratio of the area of the $Q^{4,I}$ Raman peak over the sum of the areas of the $Q^{4,I}$ and $Q^{4,II}$ Raman peaks (purple squares) is represented as a function of the mol% NaAlO_2 added in SiO_2 . It presents variations very similar to those of the glass configurational entropy at T_g , $S^{conf}(T_g)$ (Red diamonds [69] and hexagons [209]).

Similar to the observations made in the SiO₂ glass (Sec. 1.2.1), the splitting of the Si-O stretching signal in Q^4 units in two bands indicate a slight asymmetry of the distribution of T-O-T angles in tectosilicate glasses, in agreement with results from molecular dynamic simulations on aluminosilicate glasses and melts [205, 206, 210]. Neuville and Mysen [208] noticed a similarity between the variations of the entropy of configuration at T_g , $S^{conf}(T_g)$, of the glasses between the NaAlSiO₄ and the SiO₂ composition (Fig. 31b), with the formation of a convex curvature. Such curvature in the variations of the glass $S^{conf}(T_g)$ is similar to the one that can be generated by mixing two cations in similar environments. The data of Neuville and Mysen [208] actually provide further information about the mechanism of Si substitution by Al in the glasses network. Indeed, the increase of the area ratio between the two peaks assigned to Si-O stretching in Q^4 units, $Q^{4,II}/(Q^{4,II}+Q^{4,I})$, with increasing NaAlO₂ in SiO₂ below 25 mol% suggest that Al enters primarily in an environment associated with the $Q^{4,II}$ signal. As the Raman shift of the $Q^{4,II}$ peak is lower than that of the $Q^{4,I}$ peak (Fig. 31), we expect the T-O-T bond angle of the environment associated with this signal to be lower than that of the environment associated with the $Q^{4,I}$ peak[53, 54]. Therefore, Al tends to enter in Q^4 units with “small” T-O-T angles in a first time, up to an Al/Si ratio of 0.5, after which Al substitutes for Si in the second environment. This comportment may relate to the length of Al-O bonds, which are longer than Si-O bonds (Table 4). As there is an inverse correlation between T-O bond distances and T-O-T bond angles [211], this will explain the affinity of Al for the $Q^{4,II}$ site.

System	Si-O	Al-O	M-O	Methods	Reference	
Li ₂ O-Al ₂ O ₃ -SiO ₂	1.59-1.61*	1.71-1.73*	2.05*	Mod.	[212]	
Na ₂ O-Al ₂ O ₃ -SiO ₂	1.59-1.61	1.72-1.78	2.4-2.6	Exp., Mod.	[205, 211, 213]	
K ₂ O-Al ₂ O ₃ -SiO ₂	-	-	3.00-3.06	Exp.	[214]	
CaO-Al ₂ O ₃ -SiO ₂	1.60-1.63	1.72-1.77	2.32-2.45	Exp., Mod.	[206, 215–217]	
MgO-Al ₂ O ₃ -SiO ₂	1.61-1.62	1.75-1.77	1.98-2.05	[218]	[207]	

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*Table 4: Examples of ranges of bond lengths for common metal cations in alumino-silicate melts, reported in the literature by experimental (Exp.) and modeling (Mod.) studies of alkali and alkaline-earth alumino-silicate glasses. *values derived from figures.*

Other pieces of explanation may account for the variations of $S^{conf}(T_g)$ along the binary join, as proposed by Toplis et al. [209] or Lee and Stebbins [201]. The former indicated that the variations of the glass $S^{conf}(T_g)$ along the $\text{SiO}_2\text{-NaAlO}_2$ join originate from variations of the topology of the glass network as well as from chemical contributions from cation mixing, the former being dominant and approximately equal to $\sim 5.1 \text{ J mol}^{-1} \text{ K}^{-1}$ while the later accounts for the concave shape of the entropy variations. Lee and Stebbins [201]] further modelled the chemical variation using a mixing of the different Si-O-Si, Si-O-Al and Al-O-Al bonds in the network. From this model, in case of a perfect Si-Al ordering (perfect Al avoidance, $Q = 1$), $S^{conf}(T_g)$ between the SiO_2 and NaAlSiO_4 compositions is expected to show a maximum at the $\text{NaAlSi}_3\text{O}_8$ composition with an Al/Si = 0.5 (25 mol% NaAlO_2) [201, 209]. The Q values in Na tectosilicate melts are between 0.93-0.99, implying strong to nearly perfect Al avoidance [201]. Therefore, such model accounts for the observations made in Fig. 31.

1.4.4 NBO, tectosilicate glasses and triclusters

The study of tectosilicate glasses has been of particular interest in the literature as it brings information on the comportment of metal cations that act as charge compensators of the various Al polyhedral units. While tectosilicate glasses are ideally fully polymerized, there is actually a non-negligible number of NBO oxygens in their structure, subsequently called *non-stoichiometric NBO*. Stebbins and Xu [176] were among the first to report a direct experimental evidence of this, as they estimated using ^{17}O MAS NMR spectroscopy that $\sim 5\%$ of non-stoichiometric NBO are present in a calcium tectosilicate glass with $\sim 50 \text{ mol\% SiO}_2$. Other experimental studies further reported fractions of a few percents of non-stoichiometric NBO in tectosilicate glasses, as listed in Table 5.

Compensating cation	Si/(Si+Al)	% non-stoichiometric NBO	Remark	Ref.
Ca	0	5(-)	-	[219]
Ca	0.3	5.9(4)	-	[178]
Ca	0.33	5(1)	-	[220]
Ca	0.43	8(-)		
Ca	0.5	5(1.0)	-	[176]
Ca	0.5	7.6(1.9)-9.4(2.4)	for 2 different T_f	[192]
Ca	0.6	2.7(4)	-	[178]
Ca	0.75	1(-)	-	[221]
Ba	0.3	5.0(7)	-	[181]
K	0.6	0.9(4)-1.6(4)	-	[178]
K	0.8	n.d.	-	[178]

Table 5: Proportion of NBO in percent in tectosilicate glasses, measured by various NMR spectroscopy technics (principally ^{17}O NMR).

At fixed SiO_2 content, the variations of the fraction of non-stoichiometric NBO with the $\text{M}^{\text{n}+}_{2/\text{n}}\text{O}/(\text{M}^{\text{n}+}_{2/\text{n}}\text{O}+\text{Al}_2\text{O}_3)$ ratio describe linear trends (Fig. 32). The disappearance of non-stoichiometric NBOs occur in the peraluminous domain, for $\text{M}^{\text{n}+}_{2/\text{n}}\text{O}/(\text{M}^{\text{n}+}_{2/\text{n}}\text{O}+\text{Al}_2\text{O}_3) < 0.5$ (reciprocally $\text{M}^{\text{n}+}_{2/\text{n}}\text{O}/(\text{M}^{\text{n}+}_{2/\text{n}}\text{O}+\text{Al}_2\text{O}_3) > 0.5$).

While up to 9 percent of non-stoichiometric NBO are reported in Ca tectosilicate glasses, the fractions reported in K tectosilicate glasses are lower (Fig. 32). An effect of the ionic field strength of the modifier cation thus appears possible: reducing the cation ionic field strength apparently results in a reduced content of non-stoichiometric NBO in the glass[178]. However, very few direct data are available in other systems (*e.g.*, Na-bearing glasses) to test this statement, such that it should be considered with caution. Indirect evidences from Raman spectroscopy and viscosity measurements suggested the presence of non-stoichiometric NBOs in Na and Mg tectosilicate glasses [177, 197], but do not provide any quantification. The molecular dynamic simulation results from Xiang et al. [205] on Na aluminosilicate glasses with 60 mol% SiO_2 indicate that, at such silica concentration, the tectosilicate glass may

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contain 4-5 % non-stoichiometric NBO, depending on the potential used during the calculations. Such value is an estimation to consider with caution, because of the high fictive temperature of simulated glasses that may affect the modelled structure [222]. In addition to a potential effect of the metal cation ionic field strength, the existing data on the Ca tectosilicate glasses suggest that the non-stoichiometric NBO content also may vary with the Si/(Si+Al) ratio of the glass [178]. Indeed, it is comprised between 5 and 8 % for $0 < \text{Si}/(\text{Si}+\text{Al}) < 0.5$, and then decreases down to the percent level at $\text{Si}/(\text{Si}+\text{Al}) > 0.5$ (Table 5).

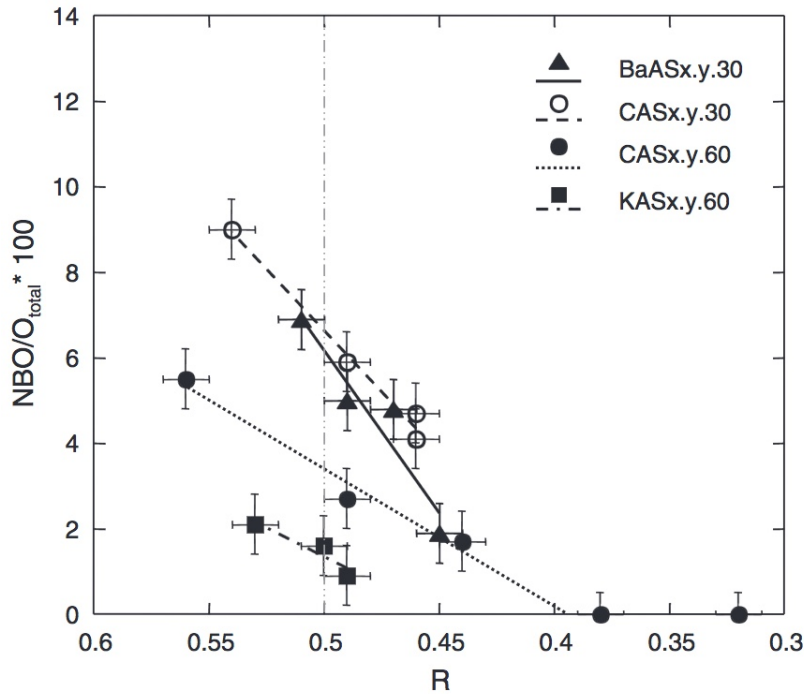
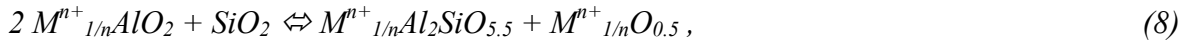


Figure 32: Proportion of non-stoichiometric non-bridging oxygens NBO relatively to the total proportion of oxygens in the network of tectosilicate glasses as a function of the ratio R of the concentrations of metal cation and aluminum $M^{n+}_{2/n}O/(M^{n+}_{2/n}O+Al_2O_3)$. BaASx.y.30: barium tectosilicate glasses with 30 mol% SiO_2 ; CASx.y.30 and CASx.y.60: calcium tectosilicate glasses with 30 and 60 mol% SiO_2 , respectively; KASx.y.60: potassium tectosilicate glasses with 60 mol% SiO_2 . Lines fits are shown as guides to the eye. Reprinted from Thompson LM, Stebbins JF (2012) Non-stoichiometric non-bridging oxygens and five-coordinated aluminum in alkaline earth aluminosilicate glasses: Effect of modifier cation size. *Journal of Non-Crystalline Solids* 358:1783–1789. doi: 10.1016/j.jnoncrsol.2012.05.022. Copyright (2012), with permission from Elsevier.

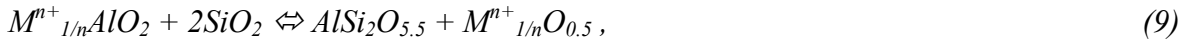
In addition of $^{[5]}Al$ and non-stoichiometric NBO, the study of tectosilicate glasses and melts further brought to light the possibility of the existence of tri-coordinated oxygen atoms in such compositions.

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Lacy [223] originally proposed the idea of the presence of triclusters in alumino-silicate melts. This idea was pursued by Toplis et al. [197], to explain why, at fixed SiO₂ concentration, the isothermal viscosity of Na alumino-silicate melts with varying Al/Na ratios form a maximum centered in the peraluminous domain at liquidus temperatures, and not, as expected, on the tectosilicate join. Following Lacy's idea, Toplis et al. [197] proposed that such oxygen triclusters was linking either 2 SiO₄ tetrahedra to one AlO₄ tetrahedron, or two AlO₄ tetrahedra with one SiO₄ tetrahedron. Such linkage will allow the presence of non-stoichiometric NBO in tectosilicate and peraluminous glasses following the reactions [197]:



or



with $M^{n+}_{1/n}Al_2SiO_{5.5}$ and $AlSi_2O_{5.5}$ the triclusters. However, as highlighted in the previous paragraph, the presence of non-stoichiometric NBOs occur simultaneously with the presence of excess ^[5]Al in the structure of the melts. Stebbins et al. [192] highlighted that the formation of ^[5]Al and NBO could be related through some mechanism implying both species. Indeed, the fall of the non-stoichiometric NBO content in the peraluminous region seem to relate to the rise of the ^[5]Al content in the glasses (see figures 32 and 25, sec. 1.4.1). Stebbins et al. [192] further noted that increasing the glass fictive temperature was leading to increase both the NBO and ^[5]Al contents, an observation confirmed by the recent ¹⁷O data on Ca tectosilicate glasses of Thompson and Stebbins [193]. The latter reported increases of up to 1.7% of the NBO fractions in Ca tectosilicate glasses with increasing their *T_f* of ~ 100-140°. However, they found that such increases depend on the glass composition in a complex way. In general, the links between the formation of non-stoichiometric NBO and ^[5]Al in glasses are not fully clear, leading Thompson and Stebbins [193] to propose that several mechanisms probably are playing at the same time.

The presence of oxygen triclusters in glasses thus cannot be excluded. However, while molecular dynamic simulations also suggest their presence in the network of alumino-silicate glasses [e.g., 205, 206], the only direct experimental observation of such oxygen triclusters have been made on a Ca aluminate glass by Iuga et al. [219] using heteronuclear high-resolution NMR spectroscopy. We finally should note that, while oxygen triclusters have been proposed as linking three Al- and Si-bearing

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tetrahedral units, it also is possible that $^{[5]}\text{Al}$ and potentially $^{[6]}\text{Al}$ polyhedra carry three-coordinated oxygen anions, $^{[3]}\text{O}$, at some of their summit. Indeed, Le Losq et al. [25] noticed that $^{[3]}\text{O}$ are present at the summits of $^{[5]}\text{Al}$ and $^{[6]}\text{Al}$ species in Sillimanite or in other minerals like Low or High Albite for instance [see 224 for instance]], while $^{[4]}\text{Al}$ and $^{[4]}\text{Si}$ species are linked by two-fold coordinated oxygen anions ($^{[2]}\text{O}$) in such minerals. As a result, those authors proposed that $^{[3]}\text{O}$ may be present in the structure of tectosilicate and peraluminous glasses, but they connect highly coordinated Al species to other tetrahedral units rather than only tetrahedral units. This explanation will allow understanding why, for instance, both the viscosity of the $\text{NaAlSi}_3\text{O}_8$ melt and its NBO content decreases with increasing pressure [225, 226]. Indeed, such paradox can be explained assuming that the formation of $^{[5]}\text{Al}$ apporportion the existing NBOs, but also favor the diffusivity of O species in the melt such that it lowers its viscosity. However, this proposition does not explain the positive correlation between the T_f of Ca tectosilicate glasses and their non-stoichiometric NBO and $^{[5]}\text{Al}$ contents, reported by Stebbins et al. [192] and Thompson and Stebbins [193]. Indeed, assuming $^{[5]}\text{Al}$ apporportion NBOs in the glass network, a decreasing fraction of non-stoichiometric NBO might be expected upon increasing the glass T_f and $^{[5]}\text{Al}$ fraction. Such contradiction illustrates that the links between glass composition, structure, $^{[5]}\text{Al}$, non-stoichiometric NBO and possibly $^{[3]}\text{O}$ (either as triclusters or bonded to $^{[5]}\text{Al}$) are not yet understood, and that several mechanisms that depend on melt composition and temperature probably occur simultaneously.

1.4.5 Medium range order structure in alumino-silicate glasses and the Mixed Alkali Effect

From the above description of the environment of aluminum at short-range order in alumino-silicate glasses, the structure at medium-range order of such glasses is complex as it needs to accommodate a few particularities. In particular, models for describing the structure at medium range order of alumino-silicate glasses need to account for the intimate relationship between AlO_4 units (and possibly AlO_5 and AlO_6 units) and the charge-compensator metal cations. Furthermore, the models should account for the fact that the distribution of Al and Si in Q^n units presents a certain degree of randomness, depending on the metal cation ionic radius. As aluminum presence shift the equilibrium (6) to the right [203], its presence in silicate melts potentially promotes regions rich in inter-connected Q^4 units and regions rich in Q^2 units surrounded by network modifier cations. Therefore, it is tempting to extend the MRN model of Greaves [126, 127] to the case of charge-compensating cations in alumino-silicate glasses.

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Following this line of thinking, Greaves and Ngai [129] proposed the Charge-Compensated Random Network (CCRN) model. Applied to tectosilicate glasses, it depicts their structure as composed of segregated regions rich in compensating metallic and aluminum ions. Indeed, the specific coordination of charge-compensator cations may lead to a distribution of those elements that is not random in the alumino-silicate network. Such idea may explain the weak and particular mixed alkali effects (MAE) observed in alkali tectosilicate glasses (Fig. 33). Indeed, the mixture of the $\text{NaAlSi}_3\text{O}_8$ and KAlSi_3O_8 glasses leads to deviations from linearity in the glasses T_g that are not explained by a random mixing of Na and K in the alumino-silicate network [57]. T_g variations suggest that Na and K have specific and different environments in alumino-silicate glasses. In particular, they probably present different coordination numbers, as observed in different alumino-silicate glasses and minerals: values between 6 and 8 have been reported for Na [109–111], whereas values between 9 and 12 have been reported for K^+ [214, 224, 227]. Translating such CN values in ionic radius, ionic radius values range between 102 and 118 pm for Na^+ and 155 and 164 pm for K^+ , using the Shannon [107] ionic radius dataset.

We should note that the ionic radius of K^+ even exceed that of O^{2-} , of ~ 135 pm [107]. Indeed, O^{2-} actually is one of the largest ion in usual silicate glasses, as the ionic radius of the common Ca^{2+} , Mg^{2+} , Na^+ or Li^+ cations are much smaller than that of O^{2-} . Le Losq and Neuville [57] suggested that such large K^+ ionic radius and CN in alumino-silicate compositions may promote the segregation of K^+ ions in ionic percolation channels, with “walls” rich in AlO_4 units. Such hypothesis may explain the variations of the T_g and configuration entropy of Na-K tectosilicate glasses reported in Fig. 33 because it implies a non-random mixing of Na and K in the glass structure. In addition, it implies that, compared to Na^+ , less configurations are available to accommodate K^+ in the aluminosilicate network, explaining the higher T_g and lower configurational entropy of the potassic tectosilicate glasses.

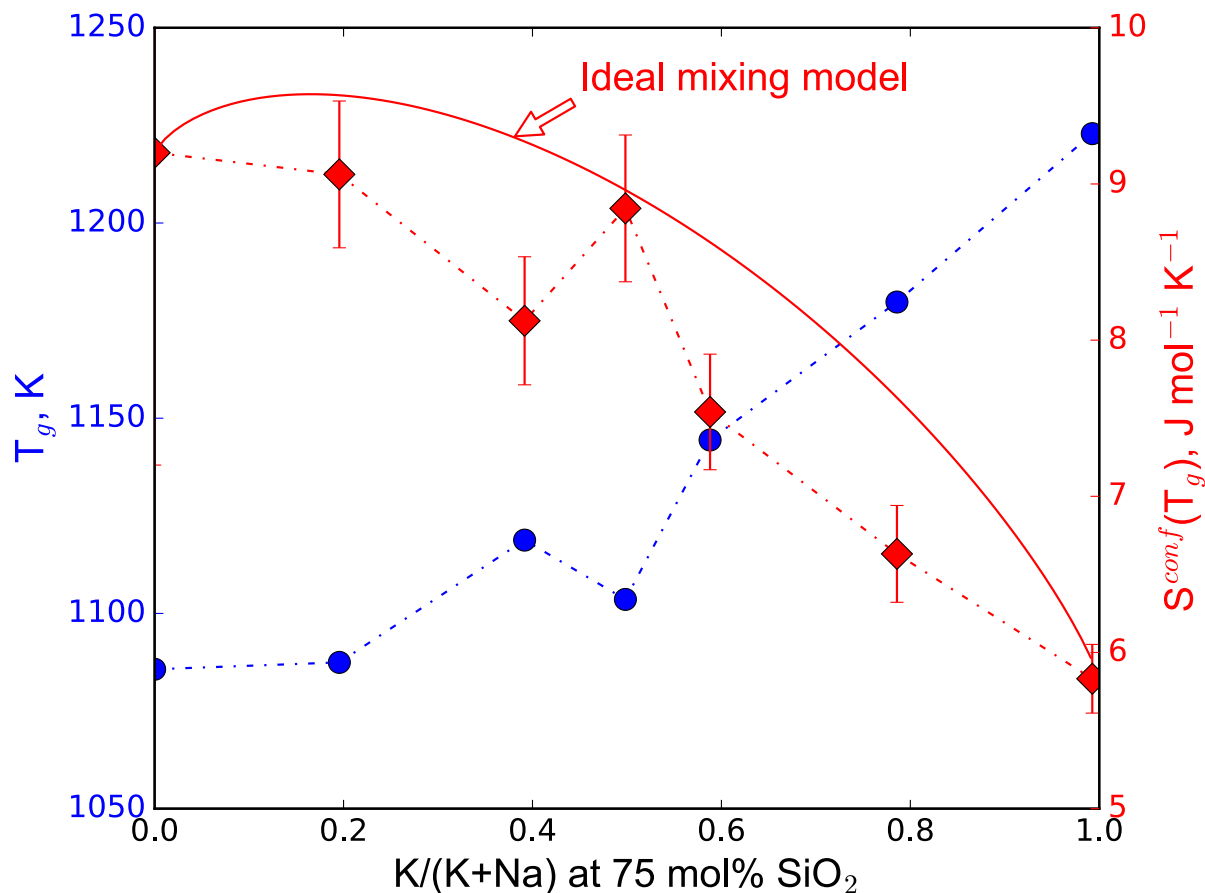


Figure 33: Viscous glass transition temperature T_g and configurational entropy at the glass transition temperature $S^{conf}(T_g)$ of $NaAlSi_3O_8$ - $KAlSi_3O_8$ glasses as a function of the glass $K/(K+Na)$ ratio. $S^{conf}(T_g)$ values have been determined performing a global fit of the viscosity data in the binary $NaAlSi_3O_8$ - $KAlSi_3O_8$ system from [57, 228], with using the framework of the Adam and Gibbs equation linking viscosity, heat capacity and entropy (see [chapter XXX](#) and [117]). A common high temperature viscosity limit was used during the fit. Dotted lines are guides for the eyes. The plain line illustrates the entropy variations expected between the end-member in the case of an ideal mixing of the Na and K cations in the aluminosilicate framework [see 22, 117, 140].

This hypothesis of different local environments and segregation of the charge-compensator Na^+ and K^+ cations in aluminosilicate glasses is corroborated by results from Raman spectroscopy of the Na-K tectosilicate glasses. Indeed, Le Losq and Neuville [57] reported that, upon substituting Na by K in $NaAlSi_3O_8$ - $KAlSi_3O_8$ glasses, the peak assigned to T-O stretching around $\sim 1100\text{ cm}^{-1}$ was shifting to a higher frequency. In addition, the intensities of two peaks at $\sim 570\text{ cm}^{-1}$ and $\sim 590\text{ cm}^{-1}$, assigned to breathing vibrations in three-membered rings containing different amounts of Al [57, 229], change with

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the glasses K/(K+Na) ratio. In particular, the intensity of the 590 cm^{-1} peak increases with substitution of Na by K, pointing to an increase of the fraction of three-membered rings containing only Si atoms. Following this observation, Le Losq and Neuville [57] suggested that, while K cations enter in the glass network, they distort the tetrahedral cages that accommodate them, leading to a general increase of the mean T-O-T angle. Si atoms segregates in denser regions, represented by the Si-bearing three-membered rings. Overall, the variations suggest a segregation of the glass network in Si-rich and Al-K rich regions in tectosilicate glasses. This hypothesis implies that the mixing of the charge compensator Na^+ and K^+ is not ideal, as they occupy different structural positions, such that the variations of the glass T_g and configurational entropy at T_g upon Na-K mixing reported in Fig. 33 are closer to linear trends obtained by a mechanical mixing of two Na-Al-Si-O and K-Al-Si-O sub-lattices than to non-linear trends resulting from an ideal Na-K mixing in a single Si-Al-O sub-lattice.

From the above discussion, the CCRN model could explain the mixing of Na and K as charge-compensator in alkali silicate melts, with a non-ideal mixing of the two alkali cations in different structural environment. However, the absence of an “ideal” MAE in tectosilicate glasses is not general. Indeed, Hummer and Arndt [140] indicated that the variations of viscosity and T_g in mixed $\text{Ca}_2\text{Al}_2\text{Si}_3\text{O}_8$ - $\text{NaAlSi}_3\text{O}_8$ melts can be modelled using an ideal mixing entropic model, implying a random mixing of Na and Ca in the melt structure. In peralkaline melts, a MAE also is observed when mixing Ca and Mg for instance. Indeed, the configurational entropy at T_g of mixed $\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ - $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ tectosilicate glasses follow variations that can be modelled with assuming an excess of entropy generated by a random mixing of pairs of Ca and Mg cations in the glasses [22]]. Similarly, a MAE affects the T_g , the Vickers microhardness as well as the configurational heat capacity $C_p^{\text{conf}}(T_g)$ values ($C_p^{\text{conf}}(T_g) = C_p^{\text{liquid}}(T_g) - C_p^{\text{glass}}(T_g)$) of Na-Ca-Mg peralkaline glasses containing 60 mol% SiO_2 and 16 mol% Al_2O_3 has recently been reported [142]. Those data show the occurrence of a continuous MAE, as observed when mixing Ca and Mg or Na and K in silicate glasses (Fig. 34).

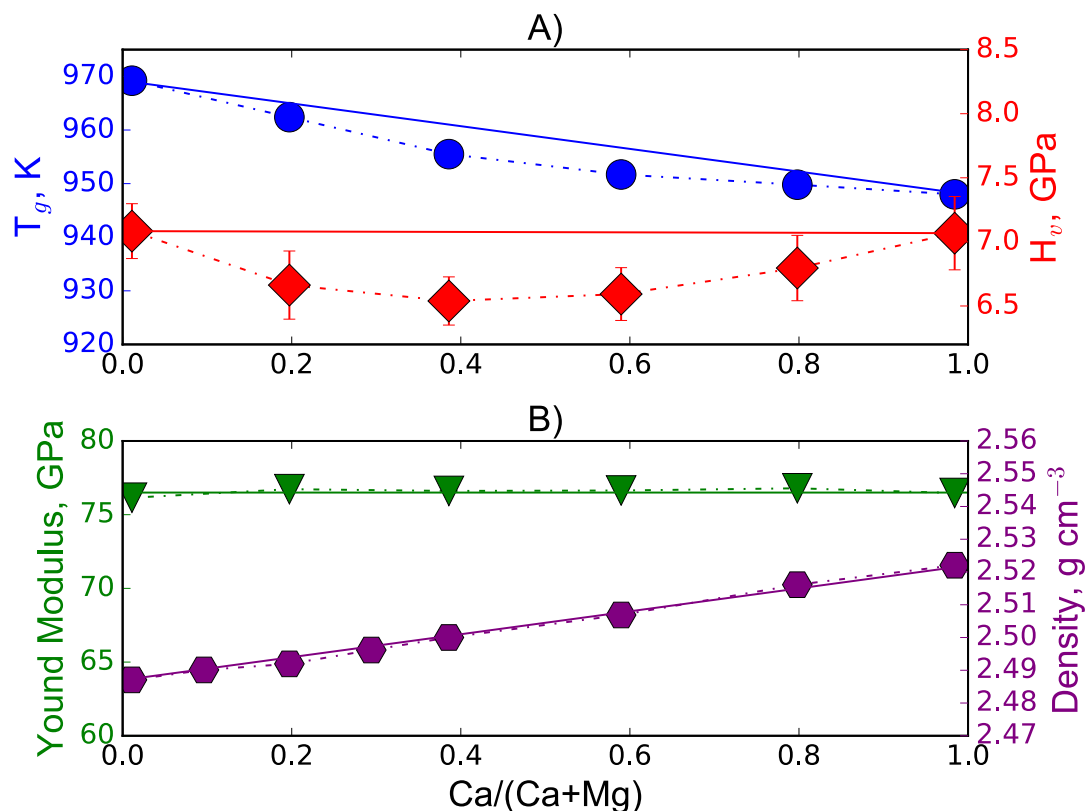


Figure 34: A) Glass transition temperature T_g and Vickers hardness H_v as well as B) Young modulus and density of glasses as a function of the Ca/(Ca+Mg) ratio of glasses composed of 60 mol% SiO_2 , 16 mol% Al_2O_3 , 15.8 mol% Na_2O and the remaining proportion as $\text{CaO}+\text{MgO}$. While glass T_g and H_v values show significant deviations to linearity, marking the occurrence of a Mixed Alkaline-earth Effect, the glass Young modulus and density vary linearly as a function of Ca/(Ca+Mg). Data from [142].

It is interesting to note that the MAE affects properties that depends on the diffusivity of the metal cations, but not properties that relate on the structure of the alumino-silicate network in itself, such as, for instance, elastic properties or glass density (Fig. 34). For instance, Figure 34 shows that, while the Young modulus of Ca-Mg-Na aluminosilicate glasses vary linearly upon Ca-Mg mixing, a MAE affects the Vickers hardness. Kjeldsen et al. [142] indicate that, as the Vickers hardness actually is related to the plastic flow of the glass, the occurrence of the MAE is not surprising. On the other hand, glass density and Young modulus, not affected by MAE, are related to the structural arrangements in the glass. This can be explained by the fact that structural characteristics of the glass network tend to vary linearly upon mixing metal cations between endmember compositions. For instance, data in Al-free silicate melts point to negligible or linear variations of the Q^n unit distribution upon mixing metal cations in the glasses. Indeed, Hater et al. [230] did not observe systematic change of the Q^n unit distribution upon Rb-Na

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mixing in alkali trisilicate glasses using ^{29}Si and ^{23}Na Magic Angle Spinning (MAS) NMR spectroscopy. ^{23}Na , ^7Li , and ^{29}Si NMR data from $(\text{Na,Li})_2\text{Si}_2\text{O}_5$ glasses also suggest an intimate, uniform mixing of the Li and Na alkali metal cations without significant changes in the silicate tetrahedral network [231]. On the other hand, using ^6Li , ^7Li and ^{29}Si static NMR spectroscopy, Bray et al. [232] observed a linear change of the fraction of Q^3 units upon mixing Li and K in disilicate glasses. The ^{29}Si static NMR data from Emerson and Bray [233] further suggest a slight departure from linearity of the Q^3 unit fraction when mixing $\text{Na}_2\text{Si}_2\text{O}_5$ and $\text{Cs}_2\text{Si}_2\text{O}_5$ glasses. Such interpretation may be coherent with ^{17}O Dynamic Angle Spinning NMR data from mixed $\text{Na}_2\text{Si}_2\text{O}_5$ - $\text{K}_2\text{Si}_2\text{O}_5$ glasses [152], showing slightly non-linear changes of the fractions of Bridging (BO) and Non-Bridging (NBO) oxygen anions with changes in the glasses K/(K+Na) ratios. Recent analysis of Raman spectra of mixed $(\text{K,Na})\text{Si}_3\text{O}_7$ and $(\text{K,Na})_2\text{Si}_2\text{O}_9$ glasses by Le Losq and Neuville[20] further indicate that the Q^n distribution of mixed Na-K silicate glasses vary almost linearly upon Na-K mixing; when combined with the data from [152], it appears that slight deviation to linearity appear at high alkali concentration but do not exceed a few percent[20]]. Therefore, almost linear variations of the glass Q^n unit distribution are suggested by NMR and Raman spectroscopy results. This may explain why, for instance, the glass density tends to vary linearly upon metal mixing in silicate and alumino-silicate melts (Fig. 34, see also [137] for instance). Indeed, as presented in section 1.3.1, it is possible to relate the volume of Q^n units to the density of the glasses, such that nearly linear variations of the Q^n unit fractions upon mixing metal cations in silicate and alumino-silicate glasses should result in linear variations of the glass density, as observed (Fig. 34).

The contrast between the variations of T_g upon mixing Na and K in tectosilicate glasses (Fig. 33) and Ca and Mg in Na-Ca-Mg peralkaline glasses (Fig. 34) directly points to the complexity of the MAE. Charge compensators with very dissimilar ionic radius, like Na and K (Table 6), appear to mix non-randomly in the glass structure, while cations with similar ionic radius, like Ca and Mg, appear to mix more randomly. However, even the apparent ideal mixing reported in Ca-Mg alumino-silicate melts, observed from a thermodynamic point of view [22], may not necessarily be perfect. Indeed, recent data from ^{17}O MAS and 3QMAS NMR spectroscopy indicate that, in mixed $\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ - $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ glasses, the mixing of Ca and Mg around NBO is not perfectly random [234]. A slight excess of Mg-NBO pairing is observed, implying that extra Ca must be associated with BO from Al-O-Si bonds according to Kelsey et al. [234]. The latter authors indicate that the $S^{\text{conf}}(T_g)$ values of mixed $\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ - $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ glasses, calculated from the viscosity of the melts [22], are actually slightly lower than the values modelled using an ideal, random mixing of Ca and Mg in the melts structure. As a result, Kelsey et al. [234] indicate that some ordering might take place. In addition, they discuss the fact that the simple

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random mixing model of the $S^{conf}(T_g)$ of $\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ - $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ may be too simple, as it considers the mixing of the glasses as a binary system. Further complexity actually may arise from a change of the Al avoidance or an increasing fraction of $^{[5]}\text{Al}$ in the glasses when substituting Ca by Mg.

Cation	Coordination number	Corresponding Ionic radius from Shannon [107], Å	Method	Reference
Na	6-8	1.02-1.18	Exp., Mod.	[110, 111, 125]
K	9-11	1.55-1.64*	Exp.	[214]
Ca	5-7	1.00*-1.06	Exp., Mod.	[206, 215–217]
Mg	~ 5	0.66	[218]	[213]

*Table 6: Examples of ionic radius reported in experimental (Exp.) and/or modeling (Mod.) studies of the structure of Na, K, Ca and Mg alumino-silicate glasses. *: no value for this CN is provided by Shannon [107], such that the value for the closest CN is reported hereby.*

Therefore, while the data from [22] show that considering Ca and Mg as randomly mixing in the structure of alumino-silicate glasses yields a good approximation for modelling the properties of the Ca-Mg alumino-silicate melts and glasses, NMR data invite to refine such approach. In particular, future models may have to account for the influence of the metal IFS on preferences for NBO/BO bonding and for Al charge compensation, as well as on $^{[5]}\text{Al}$ generation and on the Al avoidance extent. Furthermore, another level of complexity appears when considering that, in the structure of peralkaline glasses, network modifiers (NM) and charge compensators (CC) mix between themselves (NM-NM or CC-CC mixing) and potentially between each-other (NM-CC mixing). According to Greaves and Ngai [129], the structure of such peralkaline compositions should be described as a mix of the CCRN and MRN model. Charge compensator and network modifier metal cations thus will be represented as residing in different type of channels. While the local environment of charge compensator cations will be rich in Al^{3+} ions, that of the network modifier cations will be rich in NBOs. Other representations of the structure of alumino-silicate melts and the repartition of metal cations in it come from the perturbed distribution model, which depicts the network modifiers and charge compensators as randomly distributed nor segregating in clusters in the network, as suggested by ^{17}O NMR data that show that Na^+ interacts with BO from Si-O-

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Si, Si-O-Al and Al-O-Al bonds in Na tectosilicate glasses [130]. To this date, the exact representation of the structure of alumino-silicate glasses and of the repartition of metal cations in them remains an open topic of research.

1.4.6 Elastic properties and density of alumino-silicate glasses

The macroscopic elastic moduli of structural glasses are generally controlled by the glass atomic density and the mean inter-atomic forces. Different metal cations will affect differently the latter properties, such that the nature of metal cations in glasses directly correlate with the glass Shear, Bulk and Young's moduli (Fig. 35).

Therefore, elastic properties of glasses are closely related to the cations present in the glass. On the other hand, the glass Poisson ratio ν is differently affected by the glass chemical composition. Weigel et al. [235] showed that the Poisson ratios of $M^{n+}_{1/n}\text{AlSiO}_4$ glasses ($m=\text{Li, Na, K, Ca, Mg, Sr, Ba, Zn}$) showed no correlation with the glass molar volume, but were divided in 2 global mean values, depending on the charge of the metal cation present in the glass. Indeed, while the Poisson ratio of alkali-bearing glasses was ~ 0.234 , that of alkaline-earth bearing glasses was ~ 0.265 . Such difference arises from the different structure of alkali and alkaline-earth tectosilicate glasses, as shown by the strong difference of their Raman spectra [e.g., 30]. Such differences in glass structure implies a difference in the atomic packing density C_g of the glasses, which, in turn, affects the glass Poisson ratio [236 and references therein]. In this regard, the lower Poisson ratio of alkali $M^{n+}_{1/n}\text{AlSiO}_4$ glasses compared to alkaline-earth $M^{n+}_{1/n}\text{AlSiO}_4$ glasses reported by Weigel et al. [235] highlight a difference in the 3D arrangement of the glass structure. As a result, 3D arrangements at medium-range order in alumino-silicate glasses seem influenced by the electric charge of the charge compensator metal cations.

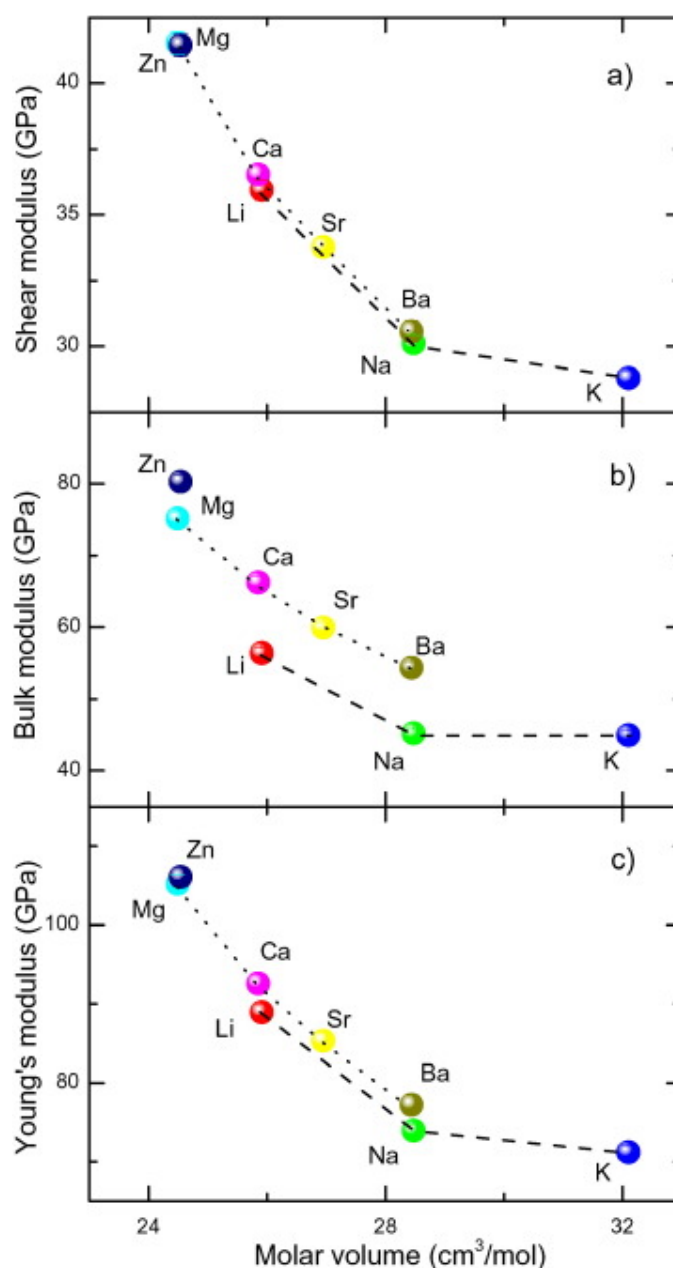
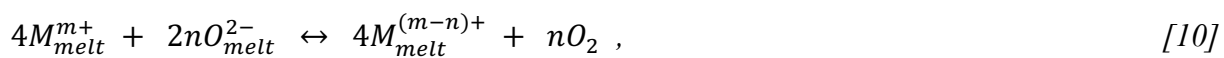


Figure 35: Elastic moduli of $M^{n+}_{1/n}\text{AlSiO}_4$ glasses ($m=\text{Li, Na, K, Ca, Mg, Sr, Ba, Zn}$) as a function of the glasses' molar volume. Reprinted from Weigel C, Le Losq C, Vialla R, Dupas C, Clément S, Neuville DR, Rufflé B (2016) Elastic moduli of $X\text{AlSiO}_4$ aluminosilicate glasses: effects of charge-balancing cations. *Journal of Non-Crystalline Solids* 447:267–272. doi: 10.1016/j.jnoncrysol.2016.06.023. Copyright (2016), with permission from Elsevier.

1.5 Multivalent elements in silicate glasses

All natural glasses and most of manmade glasses contain several multivalent elements, such as, for instance, Fe, Cr, V, Cr, and Eu, in different concentrations. Many properties of silicate glasses and melts depend on the nature, the concentration and the oxidation state of such multivalent elements. For instance, the redox states and the coordination of multivalent elements present in glasses determine their color and optical properties [e.g., 237; see **chapter XX** for further details]. Several authors recently reported the influence of the bulk composition on the photoluminescence tunability of rare-earth elements [238–240], showing that by adjusting the optical basicity of the host glassm it is possible to control both Eu redox and the glass optical properties. Another example is the viscosity of iron-bearing silicate melts, which is controlled by the oxidation state of Fe at given temperature [e.g., 241–243]. In addition to such effects, the oxidation state of multivalent elements further represents a probe for studying the conditions of glass formation in geologic systems. Indeed, the redox state of multivalent elements reflects the temperature and oxygen fugacity of magmatic sources [244], such that it can be used to retrieve the temperature of ancient geologic systems [e.g., 162]. Multivalent element oxidation state further affects the partitioning of elements and mass transfers processes within the Earth [245, 241, 246, 242, 243]. As a result, understanding the behavior of such multivalent elements in silicate melts is particularly important in Materials and Earth sciences.

Reduction-Oxidation (redox) reactions involve the transfer of electrons between chemical species forming the redox couple of a given element, as for instance Fe^{2+} and Fe^{3+} . In silicate melts, oxygen is generally the electron carrier, such that the redox state (ratio of the different species) of a given element is driven by changes in the oxygen chemical potential. In Earth and Material sciences, oxygen fugacity (f_{O_2}) is conventionally used to describe the oxygen chemical potential. Redox equilibria in glass-forming melts have been exhaustively discussed by the early studies of Schreiber [247–253]. In brief, the redox equilibrium of a multivalent element can be described as



with n representing the number of electron transferred. This relationship is controlled by several factors, such as temperature, oxygen fugacity, and bulk chemical composition.

1.5.1 Temperature, Oxygen Fugacity, and Pressure effects

Increasing temperature shifts the RedOx equilibrium (10) toward reduced species, as it favors the endothermic reaction according to the Van't Hoff law [250]. The log of the redox equilibria $M^{(m-n)}/M^{m+}$ vary linearly with reciprocal temperature [e.g., 254–257], with a slope that is equal to the enthalpy of reduction, ΔH , via an approximation of the Van't Hoff relation [250, 254]:

$$\log \left(\frac{[M^{(m-n)}]}{[M^{m+}]} \right) = - \frac{\Delta H}{2.303RT} + b . \quad (11)$$

ΔH is always endothermique, leading to the reduction of the glass-forming silicate melt as temperature increases [250].

The effect of oxygen fugacity is opposite to that of T: increasing f_{O_2} favours the stabilisation of the most oxidized specie. This can be expressed for multivalent elements dilute in silicate melts by manipulating equation (10) to obtain:

$$-\log(f_{O_2}) = \frac{4}{n} \log \left(\frac{[M^{(m-n)}]}{[M^{m+}]} \right) + k , \quad (12)$$

with n the number of exchanged electrons. For a given composition, and at constant temperature, the relationship described by eq. (12) should yield a straight line with slope equal to $4/n$. This has been experimentally corroborated for many redox couples [e.g., 244, 250, 258, 259]. Tacking the redox couple Fe^{2+}/Fe^{3+} as an example (fundamental for both Earth and Material science problems), studies showed that the slope of the $\log(f_{O_2})$ vs $\log \left(\frac{[Fe^{2+}]}{[Fe^{3+}]} \right)$ relationship was close to the ideal value of 4 (or 0.25 if using the $\log \left(\frac{[Fe^{3+}]}{[Fe^{2+}]} \right)$ vs $\log(f_{O_2})$ relationship) [244, 259, 250]. However, the melt composition influences slightly this value, such that it may not be ideal in all cases. For example, Mysen et al. [259] reported a slight decrease of the slope according to the ionic potential of the alkaline-earth cation in alkaline-earth silicate glasses. This trend is corroborated by the recent result of Cicconi et al. [260], who observed that the mean ionic field strength of the alkali metal cations influence the slope of the $\log(f_{O_2})$ vs $\log \left(\frac{[Fe^{2+}]}{[Fe^{3+}]} \right)$ relationship in multicomponent glasses.

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While the effects of oxygen fugacity and temperature on the redox state of multivalent elements are reasonably understood and characterized, that of pressure has been the subject of only few studies and remains very elusively known. In general, existing studies indicate that higher pressures promote reduced species in many compounds. For instance, reduction of Fe^{3+} , Cu^{2+} and Mn^{3+} occurs at high pressure in solid compounds [261–263]. Experimental data confirmed that such behavior occurs for Fe in silicate melts [264, 265], as the volume change accompanying the reduction of Fe^{3+} in Fe^{2+} favors the reduction reaction [e.g., see 257]. Such changes are small: the $\text{Fe}^{3+}/\text{Fe}_{\text{TOT}}$ ratio only changes of a few % per GPa, at most, when P varies between 0 to 7 GPa [264, 265]. At higher pressure, changes in Fe CN can be expected and may affect the pressure effect because of the influence of Fe CN on the volume change associate with the Fe^{3+} - Fe^{2+} redox reaction. While data on the effect of pressure on the Fe redox state are scarce, those existing for other systems are even less abundant (if simply non-existent). The data of Burnham and Berry [266] indicate that the Ce redox state may not be influenced by pressure, but those authors indicated that pressure may also favor stabilization of Ce^{4+} . Data for other elements in silicate melts are virtually inexistent.

1.5.2 Compositional effects

The chemical composition of the initial melt heavily influences the redox equilibrium, as shown by existing experimental studies devoted to Fe [267–269], and to a lesser extent to Ti [247], Cr [247, 270], Mo [271], V [272], Sn [273], Ce and Eu [247, 266, 274, 275]. The optical basicity [276–280] of the glass generally correlates well with the redox state of multivalent elements in the melt equilibrated at given T- f_{O_2} conditions [e.g., 252, 266]. The general glass optical basicity can be calculated from the different basicity of metal cation and network former oxides [e.g., 280]. Increasing the melt basicity results in a shift of the redox equilibrium (eq. 10) toward the reduced species (Fig. 36). In silicate glasses, metal cation present higher optical basicity than network former cations, such that increasing the fraction of metal cations promotes reduction of multivalent species. This relationship induces a general correlation of the redox equilibrium with the glass polymerization, i.e. the equilibrium constant of eq. 10 is correlated with the glass NBO/T. The common network former cations present close optical basicity in silicate glasses [e.g., $\text{Al}_2\text{O}_3 = 0.59$, $\text{SiO}_2 = 0.48$, $\text{TiO}_2 = 0.58$, 279]]. As a result, their influences on the Fe redox state are analogous, either in simplified quaternary system $\text{CaO-MgO-Al}_2\text{O}_3\text{-SiO}_2$ [269] as well as in natural melts [281, 282]. As metal cation can present large differences in optical basicity [e.g., from 0.39 for H_2O to 1.40 for K_2O , 279], they influence differently the equilibrium constant of eq. 10. A consequence of this is that the glass NBO/T may not be the best parameter to truly capture the control of

the melt chemical composition on the redox behavior of multivalent elements because it does not distinguish between different metal cations.

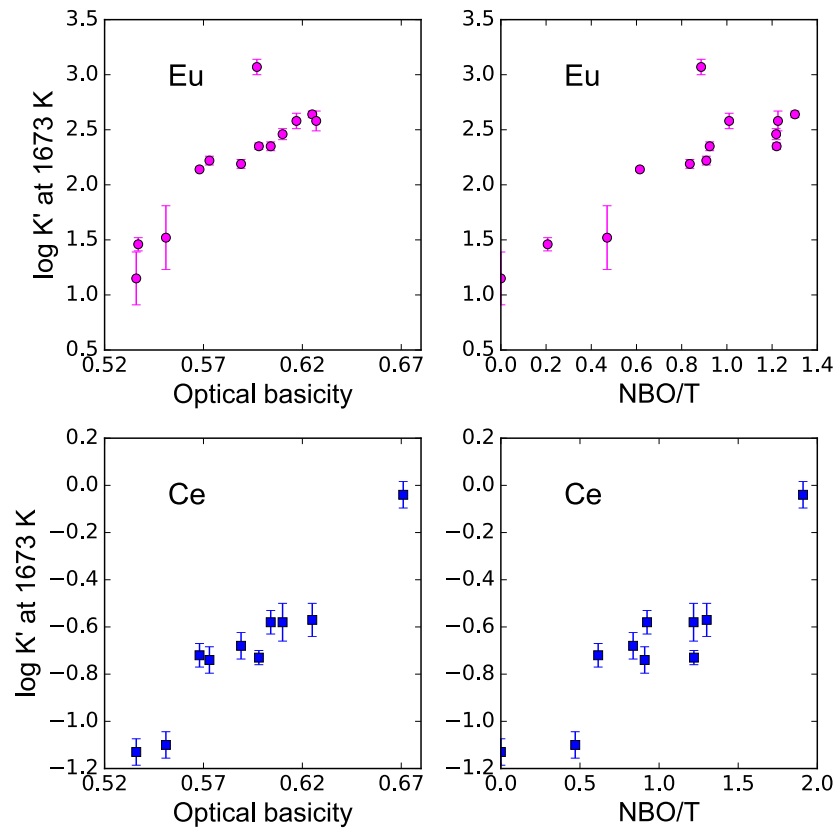


Figure 36: logarithm of the equilibrium constant of eq. (12) at 1673 for Eu and Ce in alumino-silicate melts as a function of the optical basicity or of the NBO/T. Note the outlier in the Eu plots: this data was acquired on a glass, contrary to other points that were acquired on melts. As a result, Eu oxidation state in the glass does not represent that in the melt. Burnham and Berry [266] assign this difference to potential interactions between Eu and Fe upon cooling. Data from [266, 283].

The links between the chemical composition, the glass optical basicity and polymerization, and the redox state of multivalent elements has been studied well for iron-bearing compositions, because of the interests for iron in both Earth and Material sciences. In general, increasing the content of metal cations as well as the glass Al/(Al+Si) ratio result in reducing Fe at given T- f_{O_2} conditions [252, 267, 269]. As different metal cation present different optical basicity, complex effects between the metal cation composition of the glass and its iron redox state further have been reported. For instance, Dickenson and Hess [268] observed that the Fe^{2+}/Fe_{TOT} ratio weakly increases with increasing the glass $K_2O/(K_2O+Al_2O_3)$ ratio in peralkaline melts, whereas the opposite is observed in peraluminous compositions. In Ca aluminosilicate

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glasses, Mysen et al. [269] reported a general decrease of the $\text{Fe}^{2+}/\text{Fe}_{\text{TOT}}$ ratio with the glass $\text{M}/(\text{M}+\text{Al})$ ratio.

From the analysis of 75 peralkaline $\text{K}_2\text{O}-\text{FeO}-\text{SiO}_2$ glasses, Tangeman et al. [284] reported an increase of $\text{Fe}^{2+}/\text{Fe}_{\text{TOT}}$ ratio with increasing $[\text{K}_2\text{O}]$. In peralkaline multicomponent glasses, Vetere et al. (2014) reported that both high NBO/T and high $\text{Na}/(\text{Na}+\text{K})$ stabilize Fe^{3+} in the melt. A stabilization of Fe^{3+} in CN 4 by addition of K_2O was also suggested by several authors [268, 281, 282, 285]. For basaltic and other complex melts, the oxidizing power of metal cations depend on their ionic field strength [e.g., 268, 282, 286]. In addition to the effect of the metal cation ionic field strength, its role in the glass network must be taken into account to explain the Fe redox state variations observed depending on the glass composition. As previously highlighted, glass optical basicity and NBO/T are parameters too simplistic to precisely depict the relationship between glass chemistry and the oxidation state of multivalent elements.

The case of multivalent elements can be further complicated by a potential change of their role in the glass structure as a function of their redox. Indeed, while Fe^{2+} is commonly considered as a network modifier element, Fe^{3+} acts as a network former elements as shown by increases in viscosity following oxidation of Fe (Fig. 37). Interestingly, the change of the melt viscosity at given temperature as a function of the Fe oxidation state further depends on the metal cation in the melt, illustrating the interplay between metal cations and Fe^{3+} for charge compensation. Regarding the local environment of Fe in silicate glasses, data on simplified and multicomponent systems, often, are contradictory. Ferric iron is mainly regarded as 4-fold coordinated, although higher coordination numbers have been also reported [260, 269, 287–298]. Regarding Fe^{2+} , its structural environment is even more debated. The presence of Fe^{2+} in CN 6 has been reported [241, 288, 299], along with trigonal bipyramidal and tetrahedral coordination [300–304]. Usually, it is safe to assert that Fe^{2+} has an average coordination close to 5, which could be possibly translated as a coexistence of 4-, 5- and 6- or just the simultaneous presence of 4- and 6- fold Fe^{2+} coordinated [260].

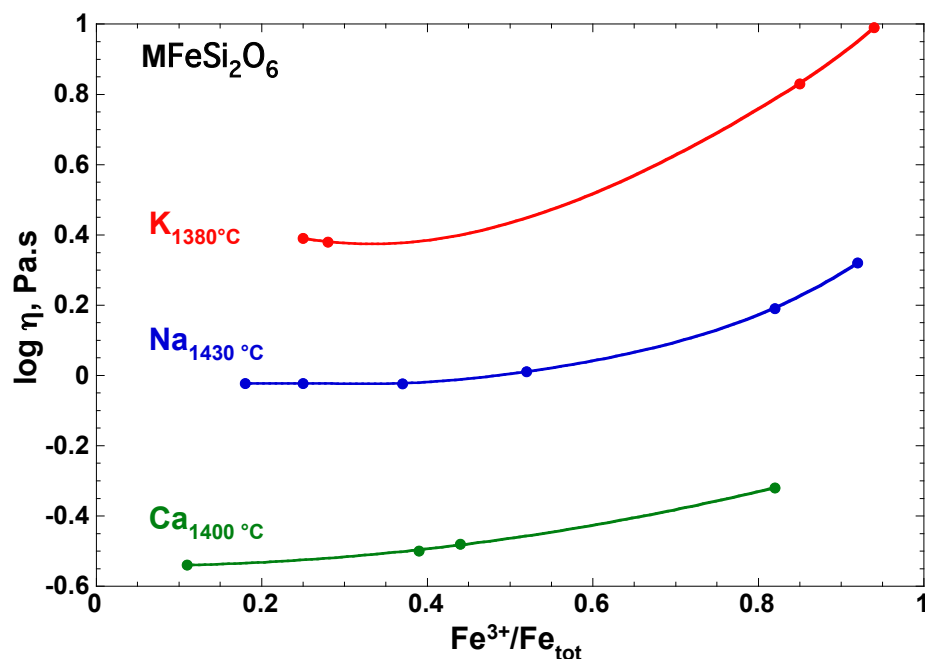


Figure 37: Viscosity (log Pa s) of alkali ferrosilicate melts at constant temperature as a function of the Fe^{3+}/Fe_{TOT} ratio. Data from [241, 243, 246].

1.6 Volatile elements in silicate glasses

Volatile elements are present in glasses in various proportions, depending on the origin of the glass. Industrial melts formed by batch melting (Sec. 1.1) contain gaseous inclusions of CO_2 , CO , N_2 and H_2O . In addition, those gaseous inclusions are usually removed by chemical fining, through introduction of sodium chloride, antimony oxide and sodium sulfate as fining agents in the melts. As a result, industrial glasses typically contain the H, C, S and Cl volatile elements at the ppm to hundreds of ppm concentration levels. For instance, we saw in sec. 1.2 that different levels of enrichments in those elements partly define the different categories of industrial silica glasses, because of their effect on the glass properties (*e.g.*, Fig. 13).

Geologic glasses also contain the H, C, S and Cl volatile elements, and also possibly F. Water and carbone dioxide concentrations in geologic melts can reach values as high as 6-7 wt% and a few thousands ppm, respectively. The knowledge of their concentrations in volcanic products is critical in assessing the eruptive history of volcanic edifices, because of the strong effect of water on the viscosity and density of

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silicate melts [305–315]. The F, Cl and S volatile elements are present at lower levels of concentrations, but still can be present at concentrations of hundreds to thousands of ppm in geologic glasses. In this section, we will give a brief overview of the exiting knowledge on those common volatile elements in industrial and geologic silicate glasses.

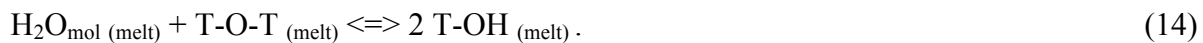
1.6.1 Water

Large variations of water concentrations between glasses entrapped at depth in geologic systems, as glass inclusions in crystals, and glasses formed at the Earth surface testify of the extensive degassing of magmas upon their journey to the Earth surface, this degassing being associated with a dramatic evolution of their viscosity [305–310]. The effect of dissolved water on magma viscosity implies a correlation between the fraction of degassed water during volcanic eruptions and their explosive nature, as observed, for instance, for phonolitic magmas [316, 317]. For example, the famous eruption of 79 CE of the Mt Vesuvius (Italy) that destroyed the cities of Pompeii and Herculaneum and killed their inhabitants [318, 319] is associated with the eruption of a magma that lost ~5-6 wt% of water during its journey to surface [320, 321], this dramatically increasing its viscosity [317], yielding to its explosive rupture in the volcanic conduit [322–324]. In contrast, the current eruption at the summit of the Mt Erebus (Antarctica) of a magma presenting a similar chemical composition but that contain only ~0.2 wt% water results in the establishment of a permanent lava lake in the volcanic crater [e.g., 317 and references cited therein]. Such dramatic contrast illustrates the importance of water in Earth science problems.

The solution mechanism of water in silicate glasses involves the presence of H₂O molecules (H₂O_{mol}), which can react with the oxygens O of the silicate network to form OH groups:



¹H and ²⁹Si NMR as well as Raman and IR spectroscopic studies observed and documented the occurrence of the reaction depicted by eq. (13) in silicate glasses and melts [114, 156, 325–341]. In this simple model, H₂O_{mol} reacts predominantly with bridging oxygens (BO), breaking T-O-T bonds (with T = Si, Al), so that equation (13) can be rewritten as:



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Equation (14) implies formation of non-bridging oxygens (NBO) when the reaction shifts to the right. This reaction usually does not proceed to completion. Indeed, both H_2O_{mol} and OH groups coexist in silicate melts, and hence glasses, in proportions defined by temperature [342–345], water concentration [e.g., 330, 346] and melt composition [156, 338, 344, 346–348].

From eq. (14), water addition in silicate melts and glasses results in their depolymerisation. The ratio of non-bridging oxygens over tetrahedral units, NBO/T, can be used to quantify such effect (see sec. 1.3.1., eq. 3): the higher it is, the higher the fraction of NBO in the melt, the lower the melt polymerization. For instance, solution of ~10 wt% of water in Na-bearing aluminosilicate melts can result in NBO/T increases of ~60% [339]. This, *in fine*, results in a great decrease of the glass T_g . Figure 38 illustrates this effect: reporting the reduced glass transition temperature T_g^* of hydrous glass as a function of their water content, a general trend is observed with large changes of the T_g that decreases by a factor of 2 with addition of ~ 10 wt% water.

Deubener et al. [349] showed that the T_g of hydrous glasses can be calculated from the knowledge of the glasses concentrations in H_2O_{mol} and OH^- and anhydrous T_g^{GN} values (i.e. the T_g of glasses synthesized at ambient conditions, which typically contain ~0.02 wt% water according to Deubener et al. [349]). Indeed, the reduced T_g^* is equal to:

$$T_g^* = \frac{C_G \times \frac{T_g^G}{T_g^{GN}} + A \times C_{OH^-} \times \frac{T_g^{OH^-}}{T_g^{GN}} + B \times C_{H_2O_{mol}} \times \frac{T_g^{H_2O_{mol}}}{T_g^{GN}}}{C_G + A \times C_{OH^-} + B \times C_{H_2O}}, \quad (14g)$$

with $C_G = (1 - C_w)$ where C_w is the total glass water content, C_{OH^-} the concentration of hydroxyl groups OH^- and $C_{H_2O_{mol}}$ the concentration of H_2O_{mol} species in the glass. T_g^G , T_g^{GN} , $T_g^{OH^-}$ and $T_g^{H_2O_{mol}}$ respectively represent the T_g of the strictly anhydrous glass, the T_g of the glass synthesized under air conditions (usually containing ~0.02 wt% water), and the virtual T_g of the OH^- groups and H_2O_{mol} species.

From the refinement of the model performed by Deubener et al. [349] that is visible in Fig. 38, $\frac{T_g^G}{T_g^{GN}} =$

1.01 and $\frac{T_g^{OH^-}}{T_g^{GN}} = \frac{T_g^{H_2O_{mol}}}{T_g^{GN}} = 0.22$. The A and B parameters depend on the melt NBO/T as $A = 35 - 7.58 \times NBO/T$ and $B = 7 - 3.03 \times NBO/T$. The hydrous glass T_g can thus be obtained from the knowledge of T_g^{GN} . This model allows simple estimations of the T_g of hydrous glasses, as shown in the study of Behrens

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and Yamashita [344] of the speciation of water in Na silicate glasses and melts. It further corroborates the observation that the glass composition has an effect on the influence of water on the hydrous glass T_g , because the parameters A and B depend on the glass NBO/T. Besides, the dependence of water speciation on glass composition yielded Deubener et al. [349] to introduce a calculation of the glass OH⁻ and H₂O_{mol} concentrations that depends on the glass anhydrous NBO/T.

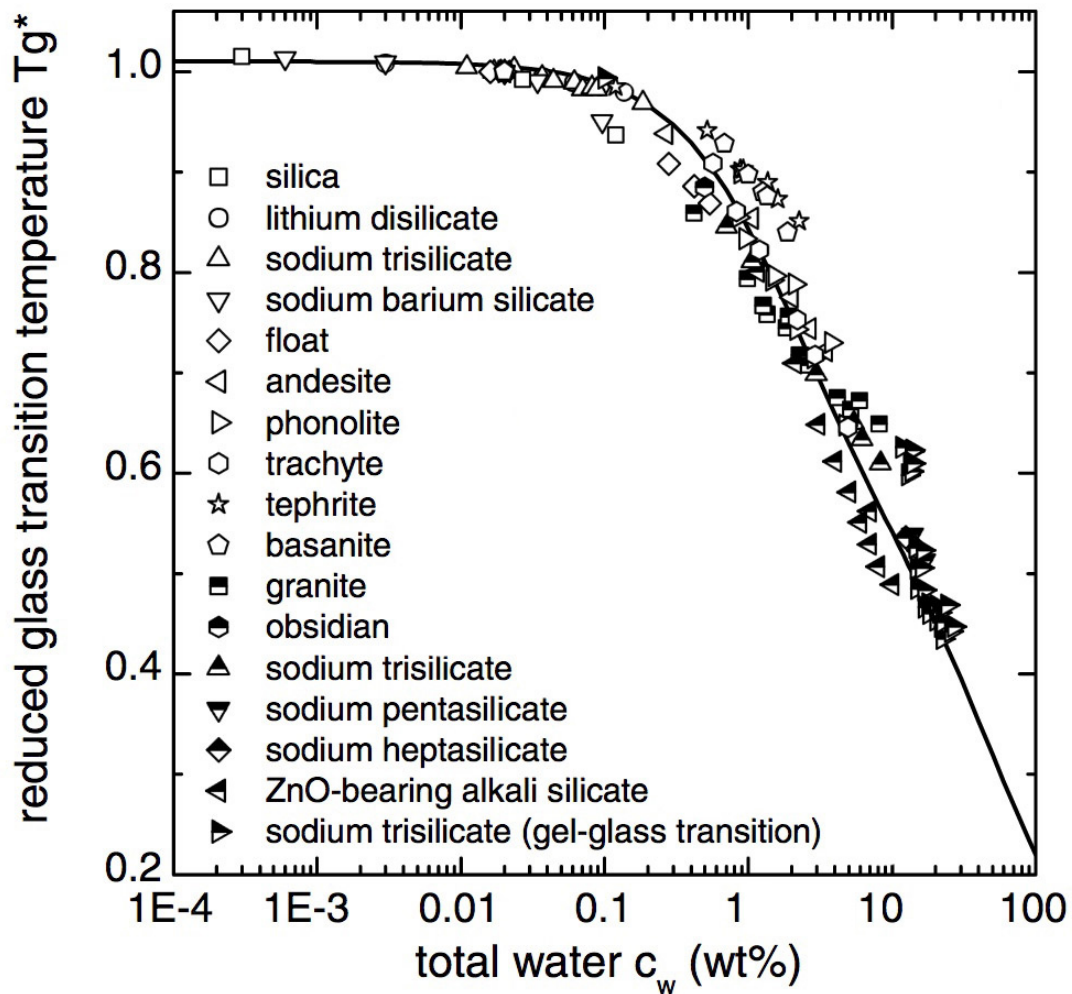


Figure 38: Reduced glass transition temperature T_g^* as a function of the total content of water in the glasses in wt%. $T_g^* = T_g^{\text{hydrous}}/T_g^{\text{GN}}$, with T_g^{GN} the T_g of the glass with 0.02 wt% water, a content typical of glass synthesized at ambient conditions. See Fig. 1 for examples of compositions of the basanite, tephrite, trachyte, phonolite, andesite, granite and obsidian geologic glasses. The curve represents the model of Deubener et al. [349]. Reproduced with minor alterations from Deubener J, Müller R, Behrens H, Heide G (2003) Water and the glass transition temperature of silicate melts, *Journal of Non-Crystalline Solids* 330:268–273, Copyright (2003), with permission from Elsevier.

While the modeling approach described previously is practical and coherent, the effect of the glass composition on the speciation of water actually is more complex than a simple dependence on the glass NBO/T. For instance, the IFS of alkali elements strongly affect the water speciation and bonding in Al-free silicate glasses (Fig. 39): the higher the alkali IFS, the higher the ratio $[H_2O_{mol}]/[OH^-]$. This translates a direct effect of the IFS of the alkali element on the equilibrium constant of the disproportionation reaction of H_2O_{mol} species in the glass network (eq. 13). ^{29}Si NMR and Raman spectroscopy results on alkali silicate glasses further reveal a change in the bonding of the OH^- groups within the glass structure [156]: alkali cations with high IFS favor the formation of M-OH bonds (M the alkali cation) in the glass network (Fig. 39). For instance, in hydrous $Li_2Si_4O_9$ glasses, up to ~50% of the OH^- groups can be bonded to Li^+ , the remaining fraction forming Si-O-H bonds [[156]

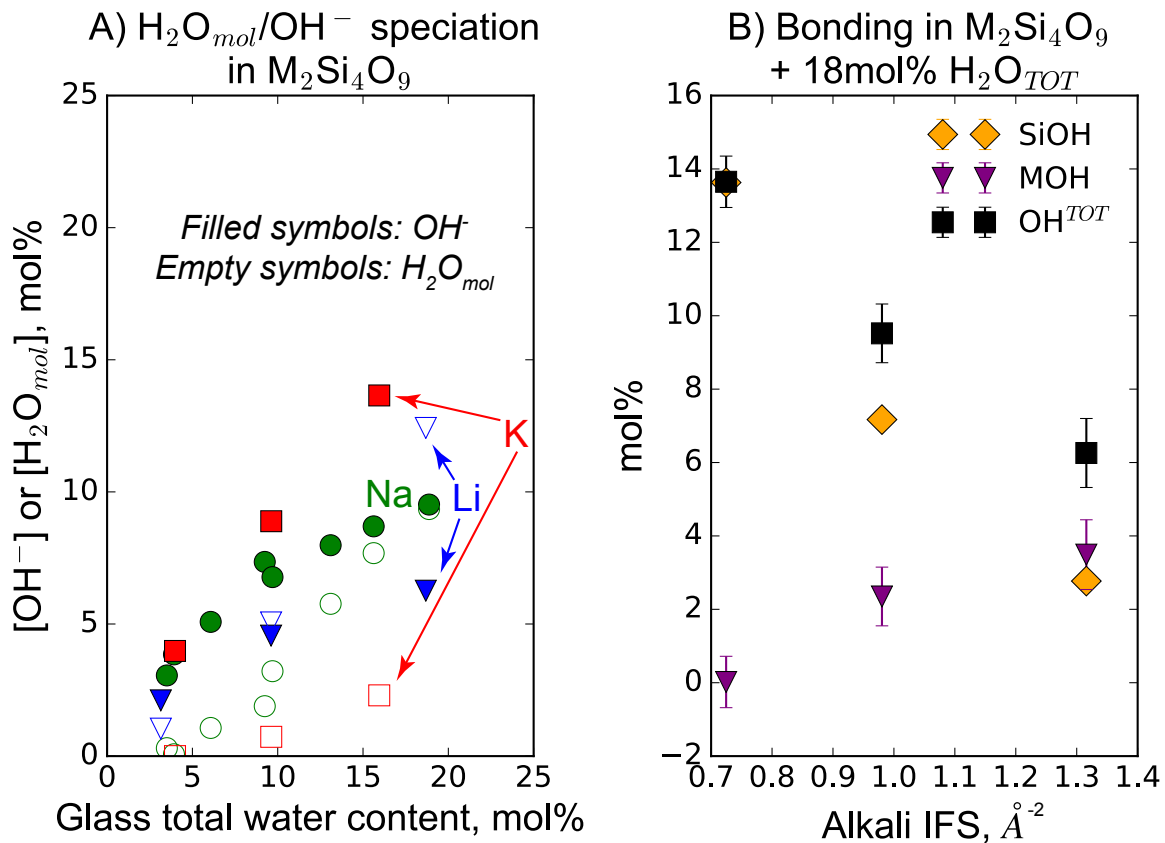


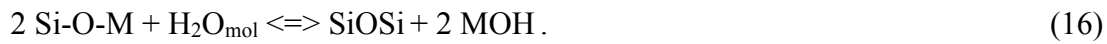
Figure 39: A) Concentration in mol% of OH^- (filled symbols) or H_2O_{mol} (empty symbols) species in $M_2Si_4O_9$ glasses; B) Concentrations of OH^- groups bonded to Si (SiOH) or the alkalis M (MOH) in $M_2Si_4O_9$ glasses represented as a function of the ionic field strength of the alkali element. The total content of OH^- groups in the $M_2Si_4O_9$ glasses is also reported and is equal to the sum of the SiOH and MOH contents. Data from [[156, 344]

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Spectroscopic data thus indicate the occurrence of other mechanisms for the solution of water in silicate glasses. The formation of M-OH groups, documented by ^1H and ^{29}Si NMR spectroscopy as well as Raman spectroscopy and molecular dynamic simulations studies in silicate and aluminosilicate glasses [156, 338, 350–354], should occur through the reaction of $\text{H}_2\text{O}_{\text{mol}}$ species with Si-O-M bonds (M a network modifier element), following the mechanisms [350]:



and



The reactions depicted by eqs. (14), (15) and (16) result in different outcomes for the melt polymerisation. Eq. (14) implies that water solution induces glass depolymerisation, but eq. (15) implies that water solution has no effect on the melt polymerisation (constant NBO/T). On the contrary, eq. (16) even implies a polymerisation of the melt as water solution subtracts the M elements from their network modifier role. Xue and Kanzaki [350] indicate that eq. (16) occurs only in very depolymerized melts. Eq. (15) should thus be the main process that account for the reaction of water with Si-O-M bonds. Interestingly, it predicts that 1 mol of $\text{H}_2\text{O}_{\text{mol}}$ should give equal amounts of SiOH and MOH entities. Such ratio corresponds to that observed by Le Losq et al. [156] in an $\text{Li}_2\text{Si}_4\text{O}_9$ glass with ~18 mol% total H_2O (Fig. 39), such that eq. (15) may be the main equilibrium accounting for the speciation of water in this glass.

In Al-bearing glasses, the mechanisms depicted by eqs. (14), (15) and (16) should remain valid. Reactions between $\text{H}_2\text{O}_{\text{mol}}$ species and charge-compensator cations as well as Si-O-Al and Al-O-Al bonds are also possible. In Na tectosilicate glasses, a strong reaction of water with charge-compensator cations was proposed by Kohn et al. [332] on the basis of ^{29}Si , ^{27}Al , ^{23}Na and ^1H NMR spectroscopy results from a hydrous $\text{NaAlSi}_3\text{O}_8$ glass. Indeed, those authors observed that the ^{23}Na NMR signals from $\text{NaAlSi}_3\text{O}_8$ glasses change significantly with water addition while their ^{29}Si and ^{27}Al NMR signals seemed barely affected. As a result, they proposed that water was mostly reacting with Na^+ charge-compensator cations. Following this study, further data were acquired on $\text{NaAlSi}_3\text{O}_8$, NaAlSiO_4 and $\text{CaAl}_2\text{Si}_{6.6}\text{O}_{17.3}$ glasses, leading Kohn and co. [355, 356] to propose that water reacts with charge-compensator elements rather than with T-O-T bonds in tectosilicate glasses. They explained the strong decrease in T_g and viscosity of

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such glasses and melts [e.g., 357] as resulting from a weakening of T-O-T bonds due to their protonation. Such hypothesis is questioned by recent ^1H - ^{29}Si and ^1H - ^{27}Al heteronuclear correlation NMR experiments [340, 351, 358, 359], which indicate formation of SiOH and AlOH bonds in tectosilicate glasses, with a repartition of proton between them that is a function of the Al/(Al+Si) ratio as well as of the degree of Al avoidance.

While the relationships between glass chemical composition and water solution mechanisms are complex, as highlighted by the previous discussion, some other properties intriguingly depend on the glass water content in a simple fashion. For instance, the partial molar volume of water, $V_m(\text{H}_2\text{O})$, seems to tend toward a unique value in silicate glasses [311–313]. Using density measurements on various water-bearing geologic and simple silicate glasses (mostly Al-bearing), Richet et al. [[313] showed that $V_m(\text{H}_2\text{O}) = 12.0 \pm 0.5 \text{ cm}^3 \text{ mol}^{-1}$. Using such value, it is possible to calculate the density of hydrous silicate glasses, $\rho(G_{\text{Hyd}})$, from the knowledge of their anhydrous density, $\rho(G_{\text{Anh}})$, using the following equations:

$$V_m(G_{\text{Anh}}) = M_w(G_{\text{Anh}})/\rho(G_{\text{Anh}}) , \quad (17)$$

$$V_m(G_{\text{Hyd}}) = x_G V_m(G_{\text{Anh}}) + x_W V_m(\text{H}_2\text{O}) , \quad (18)$$

$$M_w(G_{\text{Hyd}}) = x_G M_w(G_{\text{Anh}}) + x_W M_w(\text{H}_2\text{O}) , \quad (19)$$

$$\rho(G_{\text{Hyd}}) = M_w(G_{\text{Hyd}})/ V_m(G_{\text{Hyd}}) , \quad (20)$$

where $V_m(G_{\text{Anh}})$ and $V_m(G_{\text{Hyd}})$ represent the molar volumes of the anhydrous and hydrous glasses, $M_w(G_{\text{Anh}})$ and $M_w(G_{\text{Hyd}})$ their molar masses, $M_w(\text{H}_2\text{O}) = 18.02$, and x_G and x_W the relative fractions of anhydrous glass and water ($x_G + x_W = 1$).

Despite slight differences between different studies in the provided estimations of $V_m(\text{H}_2\text{O})$, *ex* and *in situ* experiments on silicate melts tend to corroborate the assumption that $V_m(\text{H}_2\text{O})$ is independent of glass composition. [312, 314, 315, 360]. However, from water solubility measurements, Mysen [339] indicated that, if $V_m(\text{H}_2\text{O})$ was constant at a value of $\sim 12.5 \text{ cm}^3 \text{ mol}^{-1}$, the H_2O activity coefficient must increase with the glass Al/(Al+Si) to explain water solubility values observed in Na silicate and aluminosilicate glasses. On the other hand, keeping the H_2O activity coefficient constant would imply changing $V_m(\text{H}_2\text{O})$ of a few $\text{cm}^3 \text{ mol}^{-1}$. Small variations in $V_m(\text{H}_2\text{O})$ in glasses actually may be coherent with spectroscopic observations. Indeed, ^1H NMR spectroscopy data reveal strong differences in the O-

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O distances around protons in alkali silicate glasses (Fig. 40), such that $V_m(H_2O)$ in such glasses must depend on their composition [[341] in particular, the alkali IFS seems to exert a significant influence on $V_m(H_2O)$.

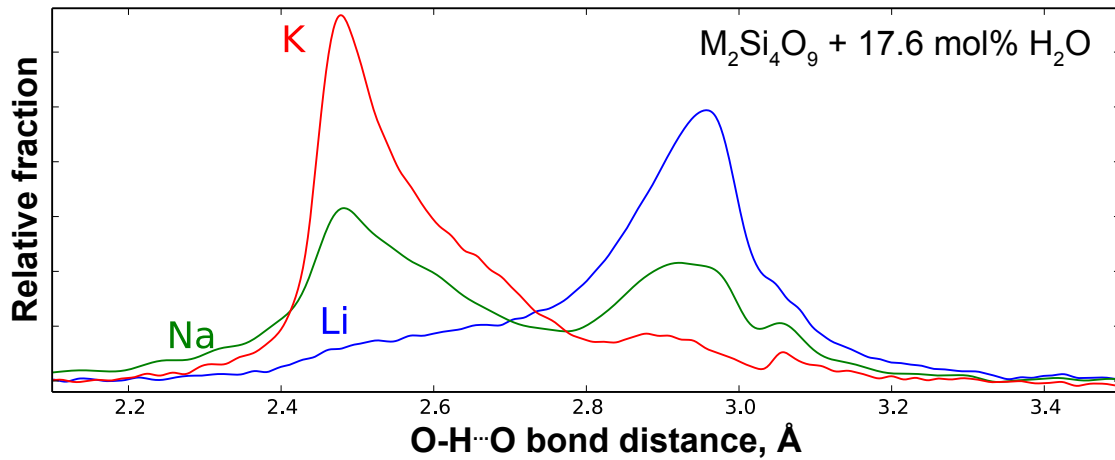


Figure 40: Relative fraction of protons in alkali tetrasilicate glasses as a function of the O-H...O bond distance. Glasses contain 17.6 mol% nominal water; data are from [341].

Therefore, while $V_m(H_2O)$ appears to be constant in most aluminosilicate glasses, as suggested by the comparable O-H stretching Raman signals from various Al-bearing geologic glasses for instance [361, 362], compositions rich in metal cations may depart from such trend, and, hence, present different $V_m(H_2O)$ as a function of the IFS of the metal cations.

The partial molar volume of water in glasses is not the only property that does not depend on water speciation or glass composition. Indeed, analysing different Al-bearing silicate glasses, Bouhifd et al. [363] showed that the partial molar heat capacity of water in glasses, $C_p(H_2O)$, do not depend on glass composition and can be calculated as:

$$C_p(H_2O) = -122.319 + 341.631 \times 10^{-3}T + 63.4426 \times 10^{-5}/T^2. \quad (21)$$

Calculating $C_p(H_2O)$ with eq. (21) and using partial C_p values for the other glass oxide components [e.g., 112] allows one to calculate the heat capacity of hydrous silicate glasses. Further recent calorimetric data on water-bearing tephrite and foidite glasses [364], Al-bearing silicate glasses (~40-50 mol% SiO₂) rich

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in Ca and Mg, corroborate the independence of $C_p(H_2O)$ on melt composition. However, those data suggest that the C_p values of the OH^- and H_2O_{mol} species are different, and vary with the degree of polymerisation. Such remarks highlight, again, the complex effect of composition on the speciation and bonding of water in silicate glasses.

1.6.3 Carbon and Sulphur

Carbon is much less soluble than water in silicate glasses: usual concentrations range between a few ppm to a few thousands of ppm [e.g., see the model of 365 and references cited therein]. Increasing glass synthesis pressure largely promotes carbon solubility, whereas the effect of temperature is weaker and not well known [366]. Under redox conditions typical of geologic environments or industrial furnaces, CO_2 and CO_3^{2-} are the most common form of carbon molecules found in silicate glasses, as evidenced by molecular dynamic simulation [367], Raman spectroscopy [368, 369] and experiments [370, 371], for instance. Formation of CO molecules is observed only in very reducing conditions [372].

Like water, the interactions of carbon with the structure of silicate glasses depend on the carbon speciation as well as on glass composition. Raman spectroscopy data indicate that CO_2 enters in the glass structure mostly as CO_3^{2-} molecules, which interact mostly with NBO and network modifier cations [329]. Raman bands due to CO_2 and CO_3^{2-} in silicate glasses actually have been known for long [372–376]: a weak band at $\sim 650\text{ cm}^{-1}$ corresponds to a bending mode of CO_3^{2-} (out of the molecule plane), a strong band at $\sim 1080\text{--}1090\text{ cm}^{-1}$ is associated with the symmetric stretching mode (ν_1) of CO_3^{2-} and a doublet in the region of $1400\text{--}1500\text{ cm}^{-1}$ due to the asymmetric stretching vibrations (ν_3) of CO_3^{2-} . Molecular CO_2 , when present, gives rise to two bands in the $1270\text{--}1400\text{ cm}^{-1}$ range (see [[[377]; for comparison, the molecular CO_2 doublet in gas hydrates is at $1274\text{--}1382\text{ cm}^{-1}$ [[378]]. Combining Raman spectroscopy with IR and NMR spectroscopy, Brooker et al. [[372] showed that increasing the glass $Al/(Al+Si)$ ratio along the $SiO_2\text{--}NaAlO_2$ binary join promotes solution of carbon as CO_3^{2-} , at the expense of molecular CO_2 . IR spectroscopy results further indicate that the availability of NBO promotes solution of CO_3^{2-} in the glass [[379]. From the Molecular Dynamic simulation data of Guillot and Sator [[367], this results from a preferential association of CO_3^{2-} carbonate groups with NBO in the melts, such that the availability of NBO controls the dissolution of carbonate ions. In addition, the MD simulation results show that network modifier cations are not randomly distributed in close vicinity of the carbonate groups, explaining why the solubility of carbon also varies as a function of the metal cation present in the glass [370, 379].

In addition to melt composition, the synthesis temperature also affect the ratio of $\text{CO}_3^{2-}/\text{CO}_2$ in the glasses. In agreement with molecular simulation data [367], results from synchrotron infrared experiments in diamond anvil cells [380] reveal that, in Al-bearing silicate melts, CO_3^{2-} carbonate groups convert in CO_2 molecules with increasing the melt temperature. Therefore, CO_2 -bearing glasses quenched at different rates may present differences in their $\text{CO}_3^{2-}/\text{CO}_2$ ratios.

The important structural changes resulting from the solution of CO_2 in glasses result in variations of the glass properties. For instance, Bourgue and Richet [381] showed that both the relative density and T_g of a potassium silicate glass decrease upon CO_2 solution (Fig. 41).

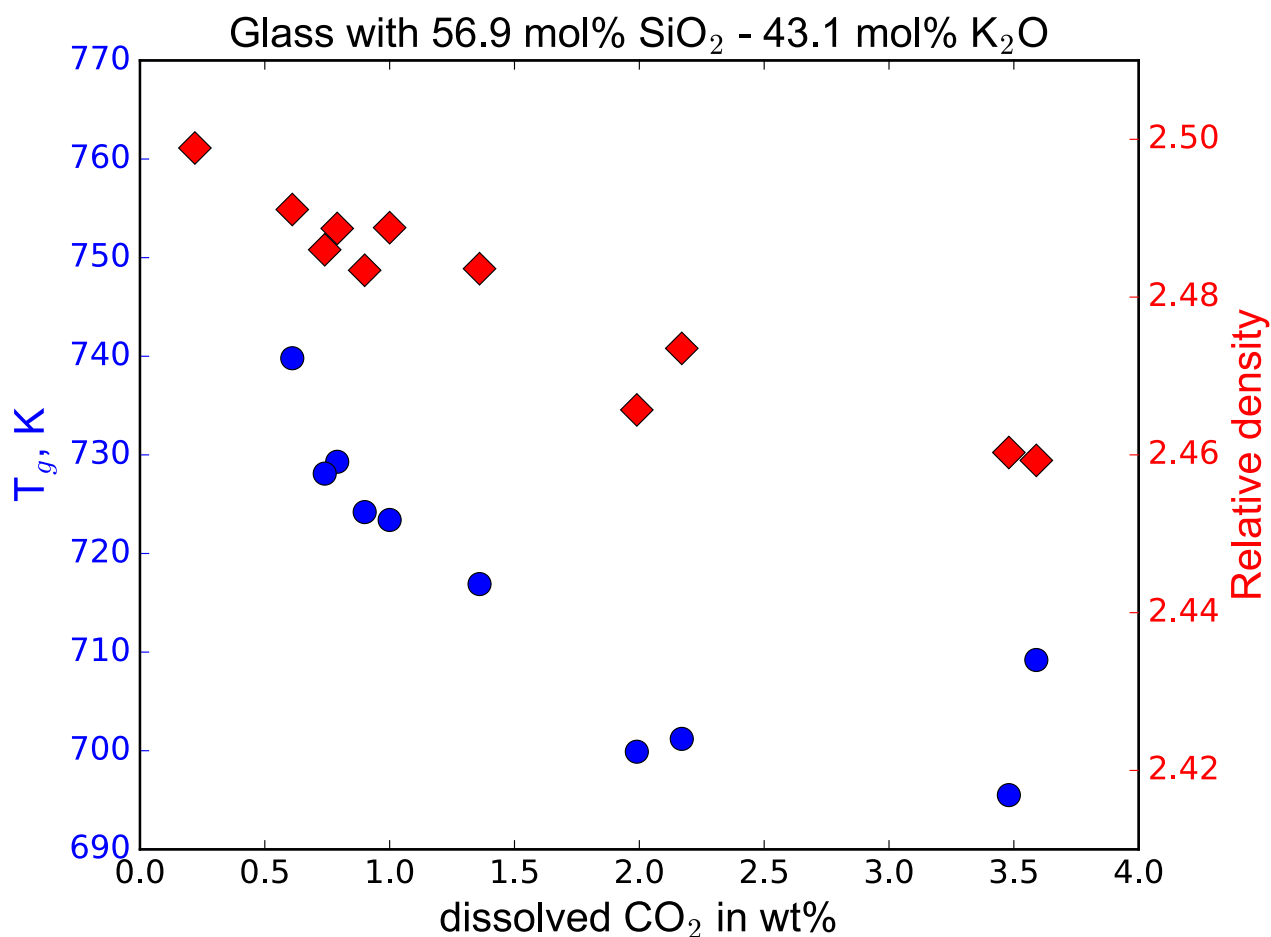


Figure 41: T_g and relative density of a potassium silicate glass, as a function of the fraction of CO_2 dissolved as CO_3^{2-} . Data from [381].

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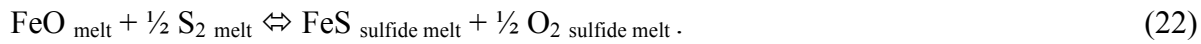
From the density measurement presented in Fig. 41, Bourgue and Richet [381] derived a constant partial molar volume of CO₂, $V_m(\text{CO}_2)$, for silicate glasses of $25.6 \pm 0.8 \text{ cm}^3 \text{ mol}^{-1}$. They further indicate that CO₂ seems to have no effect on the glass thermal expansion coefficient, at least up to 2.2 wt% dissolved CO₂. The $V_m(\text{CO}_2)$ indicated by Bourgue and Richet [381] agrees with other estimates for $V_m(\text{CO}_2)$ in melts at high temperature, which range between 20 and 30 cm³ mol⁻¹ [367, 382, 383].

In parallel to Carbon, Sulfur is the other element that is commonly found in concentrations of hundreds to thousands of ppm in both geologic and industrial glasses. In the latter, S is introduced during the addition of fining agents in the glass-making chain process. Those fining agents are introduced to aid refining the glass as well as the dissolution of silica. But S is not only present in glasses aimed at commercial application. Indeed, it is an important element of interest in the glass matrix used to confine nuclear wastes. As a result, the understanding of the reaction between S and the molecular structure of such glass confining matrix is an important goal for the long-term storage of nuclear waste [384]. In Earth sciences, sulphur is of wide interest because of its role in geological and biogeochemical processes [385].

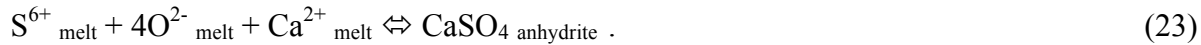
A difficulty in understanding the solution mechanisms of sulphur in silicate glasses is its multivalent character. Indeed, sulfur present four different oxidation states: S²⁻, S⁰, S⁴⁺ and S⁶⁺. However, mostly S²⁻ and S⁶⁺ have been detected in silicate glasses, as a function of their synthesis redox. For instance, S K-edge XANES data suggest that S is dissolved as S²⁻ and S⁶⁺ in natural and industrial glasses [386–390]. S⁴⁺ was reported in early studies [[e.g., 387] but actually this results from an artifact due to a change of S oxidation state under the irradiation of a synchrotron X-ray beam [390]. Presence of S⁶⁺ as SO₄²⁻ in silicate glasses is corroborated by Raman data: S-O stretching in SO₄²⁻ yields a characteristic signal at ~ 990-1000 cm⁻¹ [391–394]. Raman signals at ~ 2574 cm⁻¹ and ~ 400 cm⁻¹ have been detected in S²⁻ bearing glasses and have been assigned to H-S and Fe-S vibrations, respectively [394]. Therefore, S²⁻ bonds with H⁺ in Fe-free hydrous glasses and with Fe²⁺ in Fe-bearing glasses [395]. Sulphur thus reacts differently in the melt as a function of its oxidation state and of the melt composition.

The S solubility in glasses will vary depending on S oxidation state as well as on temperature, pressure at glass formation, and glass chemical composition [396]. The solubility of sulfide, S²⁻, is strongly affected by the iron concentration of the melt because of the importance of sulfide saturation on the sulfur solution process following the equation [[397]

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Presence of dissolved water and the melt composition affect the equilibrium constant of eq. (22), and, hence, the solubility of S^{2-} [398]. The solubility of sulfate, S^{6+} , is further driven by the concentration of Ca in the melt as its solution mechanism can be described by the equation [396]:



Various empirical models have been proposed to calculate the sulfur solubility at sulfide and sulfate saturation, either based on empirical approaches [396–399] or thermodynamic modeling [396, 400, 401], the latest one [396, 398] allowing one to calculate the sulfur solubility in glass-forming silicate melts with an accuracy of typically $\pm 10\%$ (relative).

1.6.2 Halogens

F and Cl are halogen elements that are commonly found in silicate glasses in low concentrations (≤ 1 wt%). In geosciences, F and Cl are of interest because they both can affect, for instance, the viscosity and liquidus of magmas. They further are used as geochemical tracers, allowing to reconstruct the eruptive dynamics of volcanic edifices [402]. In industrial glasses, F and Cl can be added to silica to modify its photonic properties [403, 404]. For instance, F reduces the silica glass refractive index (Fig. 42), and, hence, can be introduced for loss reduction in optical fibers. The Cl_2 gas is used for dehydration during the fabrication process of silica glass, leading to Cl impurities that affect the glass refractive index.

In silica glass, termination of Si-O-Si bonds by F and Cl seem to be the preferred interpretation for the solution mechanism of F and C[404]. The solution of F in Al-free and Al-bearing glasses is a more complex process. Raman and NMR spectroscopy data indicate the occurrence of Si-F, Al-F and M-F complexes [405–410], in proportions that depend on the glass Al/(Al+Si) and metal cation IFS. From ^{19}F NMR data on Mg, Na and La Al-free and Al-bearing silicate glasses, increasing the metal cation IFS promotes the formation of M-F bonds [407]. Such trend explains the findings of Bassner et al. [409], who noticed that F was preferentially associated with Ca in Ca-Na aluminosilicate glasses, despite the much lower concentration of Ca in the investigated glass. The latter authors further indicate that glass water content affect the F-Al and F-M fractions: it seems that F^- and OH^- groups occupy similar Al

environments, and that F cannot compete with OH groups for populating those environments. Small fractions of Si-F and high fractions of Al-F bonding are further evident from Raman spectroscopy and NMR results [405–408, 410]. The proportions of the different M-F, Al-F and, in a smaller extent, Si-F complexes depend on the Al/(Al+Si) ratio of the glasses [406, 410], resulting in a control of the solubility of F by the glass Al/(Al+Si) ratio [410]. In addition, the formation of Al-F and possibly Si-F bonds in the network of alumino-silicate glasses result in their depolymerisation, as observed by Raman spectroscopy [406, 410].

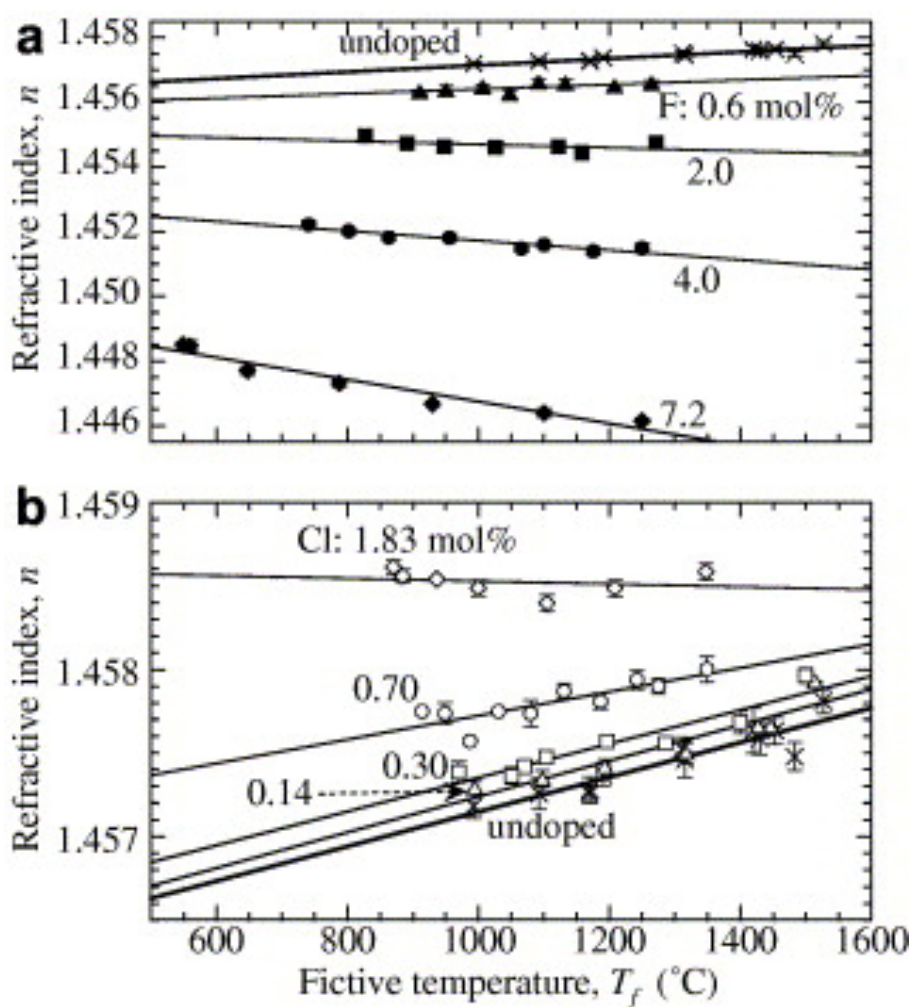


Figure 42: Refractive indices of halogen-doped silica glasses at 633 nm as a function of fictive temperature: (a) F- and (b) Cl-doped silica glasses. Solid lines are approximate linear fittings. Reprinted from Kakiuchida H, Sekiya EH, Shimodaira N, Saito K, Ikushima AJ (2007) Refractive index and density changes in silica glass by halogen doping. *Journal of Non-Crystalline Solids* 353:568–572, doi:10.1016/j.jnoncrysol.2006.10.025, Copyright (2007), with permission from Elsevier.

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Dissolution of Cl leads to a different pattern: according to Dalou et al. [410], it slightly depolymerizes the structure in Al-free melts, while it tends to polymerize that of Al-bearing melts. Such variations in the effect of Cl on melt polymerization may explain Cl effect on melt viscosity. Indeed, Baasner et al. [411] report an increase of melt viscosity for peralkaline $\text{Na}_2\text{O}-\text{CaO}-\text{Al}_2\text{O}_3-\text{SiO}_2$ melts; a decrease of melt viscosity is observed in peraluminous melts, and may be assigned to Al-Cl bonding favored by the absence of Ca and Na network modifiers available for bonding with Cl. Indeed, formation of M-Cl complexes seems to be the main mechanisms of Cl solution in the glasses [410, 412–416]. Such solution mechanism is corroborated by the fact that the metal cation IFS influence the solubility of Cl: at given $\text{Al}/(\text{Al}+\text{Si})$, cations with higher IFS promote solution of Cl [410]. Cl thus may have the tendency to bond with metal cations that present high IFS, as shown by the Cl XANES data of Evans et al. [415]. Indeed, the latter suggest that, in $\text{CaO}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2$ glasses with variables amounts of Na_2O , K_2O and MnO , Cl preferentially associates with Ca and Mg.

1.7 Conclusion

- The backbone of the molecular disordered structure of silicate glasses is the tri-dimensionnal network of interconnected Q^4 tetrahedral units containing Si at the center and O at the summits. At medium range order, tetrahedral units can form tri-dimensionnal molecular arrangements (*e.g.*, rings and cages) that influence the glass elastic properties.
- Metal cations play a role of network modifiers, breaking Si-O-Si inter-tetrahedral bonds, decreasing the network polymerization as they produce the transformation of bridging oxygens to non-bridging oxygens. Q^0 , Q^1 , Q^2 , and Q^3 units are formed upon addition of network modifiers, in proportion that depends on the fraction of added modifiers and of its Ionic Field Strength. Indeed, as described by the Modified Random Network model, modifiers tend to cluster in the network, potentially forming percolation channels that percolates through the disrupted tetrahedral network. The fractions of Q^n units as well as the medium range order organization of the network controls the glass properties, such as its entropy, density and T_g .
- Aluminum enters mostly in four-fold coordination in Q^4 units. Metal cations in vicinity do not modify the network but charge-compensate the electric charge deficit of AlO_4^- units. AlO_5 and AlO_6 units are present in low fractions in the glasses, depending on the $\text{Al}/(\text{Al}+\text{M})$ ratio as well as on the charge compensator IFS; similar conclusions are drawn for the presence of non-

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stoichiometric non-bridging oxygens. Presence of non-stoichiometric NBO and AlO_5 units influence the glass properties related to its plastic flow, such as its T_g and Vickers hardness.

- Multivalent elements in silicate glasses present oxidation states that mostly depend on temperature, oxygen fugacity and melt composition applied during the glass formation process. Coordination state of multivalent elements can be a complex function of melt composition, as shown by the various findings for Fe.
- Volatile elements are commonly present in small concentrations in geologic and industrial silicate glasses. Despite such low concentrations, they play determining roles on the glass structure and on its rheological properties. It is found that most volatile elements actually react with metal cations M as well as Al and Si cations, in proportions that depends on the M cation IFS and the glass $\text{Al}/(\text{Al}+\text{Si})$ and $\text{Al}/(\text{Al}+\text{M})$ ratios. In addition, glass composition also affects the speciation of the volatile elements in different species, as does the oxygen fugacity and temperature during the glass synthesis.

1.X References

1. Henrivaux MJ (1897) *Le verre et le cristal*, P. Vicq-Dunod et Cie. Flammarion, Paris
2. Pliny TE (1855) Chap. 65: THE ORIGIN OF GLASS. *Nat. Hist.* XXXVI:
3. Rasmussen SC (2012) Origins of Glass: Myth and Known History. In: *Glass Chang. World*. Springer Berlin Heidelberg, Berlin, Heidelberg, pp 11–19
4. Rasmussen SC (2012) How Glass Changed the World. doi: 10.1007/978-3-642-28183-9
5. Gandy W (1898) *The Romance of Glass-making*. S. W. Partridge & co., London
6. Neri A, Merret C, Kunckel J (1752) *L'Art de la Verrerie*, Druant et Pissot. Paris
7. Oppenheim AL, Brill RH, Barag D, Von Saldern A (1970) *Glass and Glassmaking in Ancient Mesopotamia*. An edition of the cuneiform texts which contain instructions for glassmakers, with a catalogue of surviving objects. The Corning Museum of Glass Press, Corning, N.Y
8. Kurkjian CR, Prindle WR (1998) Perspectives on the history of glass composition. *J Am Ceram Soc* 81:795–813.
9. Brill RH, Rising BA (1999) *Chemical analyses of early glasses*. Corning Museum of Glass, Corning, N.Y
10. Brill RH, Rising BA (1999) *Chemical analyses of early glasses*. Corning Museum of Glass, Corning, N.Y

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

11. Fontaine SD, Foy D (2007) L'épave Ouest-Embiez 1, Var : le commerce maritime du verre brut et manufacturé en Méditerranée occidentale dans l'Antiquité. *Rev Archéologique Narbonnaise* 40:235–265. doi: 10.3406/ran.2007.1182
12. Rasmussen SC (2012) Reinventing an Old Material: Venice and the New Glass. In: *Glass Chang. World*. Springer Berlin Heidelberg, pp 37–50
13. Vernioles J (2013) Les sites vitrifiés de France: inventaire et mode de réalisation. Ph.D., University Paris 1
14. Neuville DR, Cormier L, Caurant D, Montagne L (2017) From Glass to Crystal. Nucleation, growth and phase separation: from research to applications. EDP Sciences, France
15. Mozzi RL, Warren BE (1969) The structure of vitreous silica. *J Appl Crystallogr* 2:164–172.
16. Neuefeind J, Liss K-D (1996) Bond angle distribution in amorphous germania and silica. *Berichte Bunsenges Für Phys Chem* 100:1341–1349. doi: 10.1002/bbpc.19961000812
17. Clark TM, Grandinetti PJ, Florian P, Stebbins JF (2004) Correlated structural distributions in silica glass. *Phys Rev B*. doi: 10.1103/PhysRevB.70.064202
18. Tucker MG, Keen DA, Dove MT, Trachenko K (2005) Refinement of the Si–O–Si bond angle distribution in vitreous silica. *J Phys Condens Matter* 17:S67. doi: 10.1088/0953-8984/17/5/008
19. Malfait WJ, Halter WE, Verel R (2008) ²⁹Si NMR spectroscopy of silica glass: T1 relaxation and constraints on the Si–O–Si bond angle distribution. *Chem Geol* 256:269–277. doi: 10.1016/j.chemgeo.2008.06.048
20. Le Losq C, Neuville DR (2017) Molecular structure, configurational entropy and viscosity of silicate melts: Link through the Adam and Gibbs theory of viscous flow. *J Non-Cryst Solids* 463:175–188. doi: 10.1016/j.jnoncrsol.2017.02.010
21. Le Losq C (2012) Rôle des éléments alcalins et de l'eau sur les propriétés et la structure des aluminosilicates fondus et vitreux : implications volcanologiques. University Paris VII - Denis Diderot
22. Neuville DR, Richet P (1991) Viscosity and mixing in molten (Ca, Mg) pyroxenes and garnets. *Geochim Cosmochim Acta* 55:1011–1019. doi: 10.1016/0016-7037(91)90159-3
23. Neuville DR (2006) Viscosity, structure and mixing in (Ca, Na) silicate melts. *Chem Geol* 229:28–41. doi: 10.1016/j.chemgeo.2006.01.008
24. Doweidar H (1996) The density of alkali silicate glasses in relation to the microstructure. *J Non-Cryst Solids* 194:155–162.
25. Le Losq C, Neuville DR, Florian P, Henderson GS, Massiot D (2014) The role of Al³⁺ on rheology and structural changes of sodium silicate and aluminosilicate glasses and melts. *Geochim Cosmochim Acta* 126:495–517. doi: 10.1016/j.gca.2013.11.010
26. Zachariasen WH (1932) The atomic arrangement in glass. *J Am Chem Soc* 54:3841–3851.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

27. Warren BE (1934) X-Ray determination of the structure of glass.
28. Huang PY, Kurasch S, Srivastava A, Skakalova V, Kotakoski J, Krasheninnikov AV, Hovden R, Mao Q, Meyer JC, Smet J, Muller DA, Kaiser U (2012) Direct Imaging of a Two-Dimensional Silica Glass on Graphene. *Nano Lett* 12:1081–1086. doi: 10.1021/nl204423x
29. Henderson GS, Fleet ME, Bancroft GM (1984) An x-ray scattering study of vitreous KFeSi_3O_8 and $\text{NaFeSi}_3\text{O}_8$ and reinvestigation of vitreous SiO_2 using quasi-crystalline modelling. *J Non-Cryst Solids* 68:333–349. doi: 10.1016/0022-3093(84)90015-2
30. Seifert F, Mysen BO, Virgo D (1982) Three-dimensional network structure of quenched melts (glass) in the systems SiO_2 - NaAlO_2 , SiO_2 - CaAl_2O_4 and SiO_2 - MgAl_2O_4 . *Am Mineral* 67:696–717.
31. Bell RJ, Dean P (1972) The structure of vitreous silica: Validity of the random network theory. *Philos Mag* 25:1381–1398. doi: 10.1080/14786437208223861
32. Shintani H, Tanaka H (2008) Universal link between the boson peak and transverse phonons in glass. *Nat Mater* 7:870–877.
33. Bucheneau U, Prager M, Nücker N, Dianoux AJ, Ahmad N, Phillips WA (1986) Low-frequency modes in vitreous silica. *Phys Rev B* 34:5665–5673.
34. Chumakov AI, Monaco A, Crichton WA, Bosak A, Rüffer R, Meyer A, Kargl F, Comez L, Fioretto D, Giefers H, Roitsch S, Wortmann G, Manghnani MH, Hushur A, Williams Q, Balogh J, Parlinski K, Jochym P, Piekarczyk P (2011) Equivalence of the Boson Peak in glasses to the transverse acoustic van Hove singularity in crystals. *Phys Rev Lett* 106:225501–1:5.
35. Courtens E, Foret M, Hehlen B, Vacher R (2001) The vibrational modes of glasses. *Solid State Commun* 117:187–200.
36. Bell RJ, Bird NF, Dean P (1968) The vibrational spectra of vitreous silica, germania and beryllium fluoride. *J Phys C Proc Phys Soc* 1:299–303.
37. Bell RJ, Dean P, Hibbins-Butler DC (1970) Bell-JPC-70-3.pdf. *J Phys Chem Solid State Phys* 3:2111–2118.
38. Bell RJ, Dean P, Hibbins-Butler DC (1971) Normal mode assignments in vitreous silica, germania and beryllium fluoride. *J Phys C Solid State Phys* 4:1214–1220.
39. Sharma SK, Mammone JF, Nicol MF (1981) Raman investigation of ring configurations in vitreous silica. *Nature* 292:140–141.
40. McMillan PF, Piriou B (1982) The structures and vibrational spectra of crystals and glasses in the silica-alumina system. *J Non-Cryst Solids* 53:279–298.
41. Galeener FL (1982) Planar rings in glasses. *Solid State Commun* 44:1037–1040.
42. Galeener FL (1982) Planar rings in vitreous silica. *J Non-Cryst Solids* 49:53–62.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

43. McMillan PF (1984) Structural studies of silicate glasses and melts - Applications and limitations of Raman spectroscopy. *Am Mineral* 69:622–644.
44. McMillan PF, Wolf GH, Poe BT (1992) Vibrational spectroscopy of silicate liquids and glasses. *Chem Geol* 96:351–366.
45. McMillan PF, Poe BT, Gillet P, Reynard B (1994) A study of SiO₂ glass and supercooled liquid to 1950 K via high-temperature Raman spectroscopy. *Geochim Cosmochim Acta* 58:3653–3664.
46. Pasquarello A, Car R (1998) Identification of Raman defect lines as signatures of ring structures in vitreous silica. *Phys Rev Lett* 80, No. 23:5145–5147.
47. Umari P, Gonze X, Pasquarello A (2003) Concentration of small ring structures in vitreous silica from a first-principles analysis of the Raman spectrum. *Phys Rev Lett* 90:1–4.
48. Rahmani A, Benoit M, Benoit C (2003) Signature of small rings in the Raman spectra of normal and compressed amorphous silica: A combined classical and ab initio study. *Phys. Rev. B* 68:
49. Galeener FL, Barrio RA, Martinez E, Elliott RJ (1984) Vibrational decoupling of rings in amorphous solids. *Phys Rev Lett* 53:2429–2432.
50. Sharma SK, Philpotts JA, Matson DW (1985) Ring distributions in alkali- and alkaline-earth aluminosilicate framework glasses - A Raman spectroscopic study. *J Non-Cryst Solids* 71:403–410.
51. Umari P, Pasquarello A (2002) Modeling of the Raman spectrum of vitreous silica: concentration of small ring structures. *Phys B* 316–317:572–574.
52. Kalampounias AG (2006) A low-frequency Raman study of glassy, supercooled and molten silica and the preservation of the Boson peak in the equilibrium liquid state. *J Non-Cryst Solids* 352:4619–4624.
53. Galeener FL (1979) Band limits and the vibrational spectra of tetrahedral glasses. *Phys Rev B* 19:4292–4397.
54. Sen PN, Thorpe MF (1977) Phonons in AX₂ glasses: From molecular to band-like modes. *Phys Rev B* 15:4030–4038.
55. Sarnthein J, Pasquarello A, Car R (1997) Origin of the High-Frequency Doublet in the Vibrational Spectrum of Vitreous SiO₂. *Science* 275:1925–1927. doi: 10.1126/science.275.5308.1925
56. Spiekermann G, Steele-MacInnis M, Schmidt C, Jahn S (2012) Vibrational mode frequencies of silica species in SiO₂-H₂O liquids and glasses from ab initio molecular dynamics. *J Chem Phys* 136:154501. doi: 10.1063/1.3703667
57. Le Losq C, Neuville DR (2013) Effect of the Na/K mixing on the structure and the rheology of tectosilicate silica-rich melts. *Chem Geol* 346:57–71. doi: <http://dx.doi.org/10.1016/j.chemgeo.2012.09.009>

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

58. Mysen BO, Finger LW, Virgo D, Seifert FA (1982) Curve-fitting of Raman spectra of silicate glasses. *Am Mineral* 67:686–695.
59. Pasquarello A, Sarnthein J, Car R (1998) Dynamic structure factor of vitreous silica from first principles: Comparison to neutron-inelastic-scattering experiments. *Phys Rev B* 14133–14140.
60. Rino JP, Ebbsjö I, Kalia RK, Nakano A, Vashishta P (1993) Structure of rings in vitreous $\{\mathrm{SiO}\}_2$. *Phys Rev B* 47:3053–3062. doi: 10.1103/PhysRevB.47.3053
61. Benoit M, Ispas S, Jund P, Jullien R (2000) Model of silica glass from combined classical and *ab initio* molecular-dynamics simulations. *Eur Phys J B - Condens Matter Complex Syst* 13:631–636. doi: 10.1007/s100510050079
62. Yuan X, Cormack AN (2003) Si–O–Si bond angle and torsion angle distribution in vitreous silica and sodium silicate glasses. *J Non-Cryst Solids* 319:31–43. doi: 10.1016/S0022-3093(02)01960-9
63. Poulsen HF, Neuefeind J, Neumann H-B, Schneider JR, Zeidler MD (1995) Amorphous silica studied by high energy X-ray diffraction. *J Non-Cryst Solids* 188:63–74. doi: 10.1016/0022-3093(95)00095-X
64. Wright AC (1994) Neutron scattering from vitreous silica. V. The structure of vitreous silica: What have we learned from 60 years of diffraction studies? *J Non-Cryst Solids* 179:84–115. doi: 10.1016/0022-3093(94)90687-4
65. Mauri F, Pasquarello A, Pfrommer BG, Yoon Y-G, Louie SG (2000) Si-O-Si bond-angle distribution in vitreous silica from first-principles $\{^29\}\mathrm{Si}$ NMR analysis. *Phys Rev B* 62:R4786–R4789. doi: 10.1103/PhysRevB.62.R4786
66. Huang L, Yuan F, Guerette M, Zhao Q, Sundararaman S (2017) Tailoring structure and properties of silica glass aided by computer simulation. *J Mater Res* 32:174–182. doi: 10.1557/jmr.2016.397
67. Van Ginhoven RM, Jónsson H, Corrales LR (2005) Silica glass structure generation for *ab initio* calculations using small samples of amorphous silica. *Phys Rev B*. doi: 10.1103/PhysRevB.71.024208
68. Pasquarello A (2001) First-principles simulation of vitreous systems. *Curr Opin Solid State Mater Sci* 5:503–508.
69. Richet P, Bottinga Y (1984) Glass transitions and thermodynamic properties of amorphous SiO_2 , $\mathrm{NaAlSi}_n\mathrm{O}_{2n+2}$ and $\mathrm{KAlSi}_3\mathrm{O}_8$. *Geochim Cosmochim Acta* 48:453–470. doi: 10.1016/0016-7037(84)90274-6
70. Leko, V. K., Gusakova, N. J., Meshcheryakova, E. V., Prokhorova, T. I. (1977) The effect of impurity alkali oxides, hydroxyl groups, $\mathrm{Al}_2\mathrm{O}_3$, and $\mathrm{Ga}_2\mathrm{O}_3$ on the viscosity of vitreous silica. *Sov J Glass Phys Chem* 3:204–210.
71. Angell CA (1985) Spectroscopy simulation and scattering, and the medium range order problem in glass. *J Non-Cryst Solids* 73:1–17.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

72. Angell CA (1991) Relaxation in liquids, polymers and plastic crystals—strong/fragile patterns and problems. *J Non-Cryst Solids* 131:13–31.
73. Hofmeister AM, Whittington AG (2012) Effects of hydration, annealing, and melting on heat transport properties of fused quartz and fused silica from laser-flash analysis. *J Non-Cryst Solids* 358:1072–1082. doi: 10.1016/j.jnoncrsol.2012.02.012
74. Brückner R (1970) Properties and structure of vitreous silica. I. *J Non-Cryst Solids* 5:123–175. doi: 10.1016/0022-3093(70)90190-0
75. Heili M, Poumellec B, Burov E, Gonnet C, Le Losq C, Neuville DR, Lancry M (2016) The dependence of Raman defect bands in silica glasses on densification revisited. *J Mater Sci* 51:1659–1666. doi: 10.1007/s10853-015-9489-8
76. Hemley RJ, Mao HK, Bell PM, Mysen BO (1986) Raman Spectroscopy of SiO_2 Glass at High Pressure. *Phys Rev Lett* 57:747–750. doi: 10.1103/PhysRevLett.57.747
77. Weigel C, Foret M, Hehlen B, Kint M, Clément S, Polian A, Vacher R, Rufflé B (2016) Polarized Raman spectroscopy of SiO_2 under rare-gas compression. *Phys Rev B* 93:224303. doi: 10.1103/PhysRevB.93.224303
78. Burgin J, Guillon C, Langot P, Vallée F, Hehlen B, Foret M (2008) Vibrational modes and local order in permanently densified silica glasses: Femtosecond and Raman spectroscopy study. *Phys Rev B*. doi: 10.1103/PhysRevB.78.184203
79. Hehlen B (2010) Inter-tetrahedra bond angle of permanently densified silicas extracted from their Raman spectra. *J Phys Condens Matter* 22:025401. doi: 10.1088/0953-8984/22/2/025401
80. Dietzel A (1942) Die Kationfeldstärken und ihre Beziehungen zu entlasungs-vorgängen, zur verbindunsbildung und zu den Smelzpunkten von Silikaten. *Z Für Elektrochem* 48:9.
81. Ispas S, Charpentier T, Mauri F, Neuville DR (2010) Structural properties of lithium and sodium tetrasilicate glasses: Molecular dynamics simulations versus NMR experimental and first-principles data. *Solid State Sci* 12:183–192. doi: 10.1016/j.solidstatesciences.2009.06.033
82. Virgo D, Mysen BO, Kushiro I (1980) Anionic Constitution of 1-Atmosphere Silicate Melts: Implications for the Structure of Igneous Melts. *Science* 208:1371–1373. doi: 10.1126/science.208.4450.1371
83. Engelhardt G, Nofz M, Forkel K, Wihsmann FG, Mägi M, Samoson A, Lippmaa E (1985) Structural studies of calcium aluminosilicate glasses by high resolution solid state ^{29}Si and ^{27}Al magic angle spinning nuclear magnetic resonance. *Physic Chem Glas* 26:157–165.
84. Maekawa H, Maekawa T, Kawamura K, Yokokawa T (1991) The structural groups of alkali silicate glasses determined from ^{29}Si MAS-NMR. *J Non-Cryst Solids* 127:53–64. doi: 10.1016/0022-3093(91)90400-Z

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
85. Schramm CM, de Jong BHWS, Parziale VE (1984) ^{29}Si Magic Angle Spinning NMR study on local silicon environments in amorphous and crystalline lithium silicates. *J Am Chem Soc* 106:4396–4402.
 86. Dupree R, Holland D, Williams DS (1986) The structure of binary alkali silicate glasses. *J Non-Cryst Solids* 81:185–200. doi: 10.1016/0022-3093(86)90269-3
 87. Emerson, J. F., Stallworth, P. E., Bray, P. J. (1989) High-Filed ^{29}Si NMR studies of alkali silicate glasses. *J Non-Cryst Solids* 113:253–259.
 88. Sen S, Youngman RE (2003) NMR study of Q-speciation and connectivity in K_2O – SiO_2 glasses with high silica content. *J Non-Cryst Solids* 331:100–107. doi: 10.1016/j.jnoncrysol.2003.08.071
 89. Zhang P, Dunlap C, Florian P, Grandinetti PJ, Farnan I, Stebbins JF (1996) Silicon site distributions in an alkali silicate glass derived by two-dimensional ^{29}Si nuclear magnetic resonance. *J Non-Cryst Solids* 204:294–300. doi: 10.1016/S0022-3093(96)00601-1
 90. Davis MC, Kaseman DC, Parvani SM, Sanders KJ, Grandinetti PJ, Massiot D, Florian P (2010) Q^n Species Distribution in K_2O · 2SiO_2 Glass by ^{29}Si Magic Angle Flipping NMR. *J Phys Chem A* 114:5503–5508. doi: 10.1021/jp100530m
 91. Brandriss ME, Stebbins JF (1988) Effects of temperature on the structures of silicate liquids: ^{29}Si NMR results. *Geochim Cosmochim Acta* 52:2659–2669. doi: 10.1016/0016-7037(88)90034-8
 92. Mysen BO, Frantz JD (1993) Structure of silicate melts at high temperature: In-situ measurements in the system BaO – SiO_2 . *Am Mineral* 78:699–709.
 93. Mysen BO, Frantz JD (1993) Structure and properties of alkali silicate melts at magmatic temperatures. *Eur J Mineral* 5:393–407.
 94. Mysen BO (1999) Structure and properties of magmatic liquids: From haplobasalt to haploandesite. *Geochim Cosmochim Acta* 63:95–112.
 95. Mysen BO, Virgo D, Seifert FA (1982) The structure of silicate melts: Implications for chemical and physical properties of natural magma. *Rev Geophys* 20:353–383. doi: 10.1029/RG020i003p00353
 96. Mysen, B. O., Virgo, D., Seifert, F. A. (1985) Relationships between properties and structure of aluminosilicate melts. *Am Mineral* 70:88–105.
 97. Dupree R, Holland D, McMillan PW, Pettifer RF (1984) The structure of soda-silica glasses: A mas NMR study. *J Non-Cryst Solids* 68:399–410. doi: 10.1016/0022-3093(84)90020-6
 98. Shakhmatkin BA, Vedishcheva NM, Shultz MM, Wright AC (1994) The thermodynamic properties of oxide glasses and glass-forming liquids and their chemical structure. *J Non-Cryst Solids* 177:249–256. doi: 10.1016/0022-3093(94)90538-X
 99. Vedishcheva NM, Shakhmatkin BA, Shultz MM, Wright AC (1996) The thermodynamic modelling of glass properties: a practical proposition? *J Non-Cryst Solids* 196:239–243. doi: 10.1016/0022-3093(95)00593-5

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

100. Schneider J, Mastelaro VR, Zanutto ED, Shakhmatkin BA, Vedishcheva NM, Wright AC, Panepucci H (2003) Qn distribution in stoichiometric silicate glasses: thermodynamic calculations and ^{29}Si high resolution NMR measurements. *J Non-Cryst Solids* 325:164–178. doi: 10.1016/S0022-3093(03)00332-6
101. Gaddam A, Montagne L, Ferreira JMF (2016) Statistics of silicate units in binary glasses. *J Chem Phys* 145:124505. doi: 10.1063/1.4963341
102. Olivier L, Yuan X, Cormack AN, Jäger C (2001) Combined ^{29}Si double quantum NMR and MD simulation studies of network connectivities of binary $\text{Na}_2\text{O} \cdot \text{SiO}_2$ glasses: new prospects and problems. *J Non-Cryst Solids* 293:53–66.
103. Machacek J, Gedeon O, Liska M (2006) Group connectivity in binary silicate glasses. *J Non-Cryst Solids* 352:2173–2179. doi: 10.1016/j.jnoncrysol.2006.01.036
104. Doweidar H (1998) Density-structure correlations in $\text{Na}_2\text{O}-\text{Al}_2\text{O}_3-\text{SiO}_2$ glasses. *J Non-Cryst Solids* 240:55–65.
105. Doweidar H (1999) Density-structure correlations in silicate glasses. *J Non-Cryst Solids* 249:194–200. doi: 10.1016/S0022-3093(99)00310-5
106. Doweidar H, Feller S, Affatigato M, Tischendorf B, Ma C, Hammarsten E (1999) Density and molar volume of extremely modified alkali silicate glasses. *Phys Chem Glas* 40:339–344.
107. Shannon RD (1976) Revised effective ionic radii and systematic studies of interatomic distances in halides and chalcogenides. *Acta Crystallogr A* 32:751–767. doi: 10.1107/S0567739476001551
108. Stebbins JF (1998) Cation sites in mixed-alkali oxide glasses: correlations of NMR chemical shift data with site size and bond distance. *Solid State Ion* 112:137–141. doi: 10.1016/S0167-2738(98)00224-0
109. Cormier L, Ghaleb D, Neuville DR, Delaye J-M, Calas G (2003) Chemical dependence of network topology of calcium aluminosilicate glasses: a computer simulation study. *J Non-Cryst Solids* 332:255–270. doi: 10.1016/j.jnoncrysol.2003.09.012
110. Neuville DR, Cormier L, Flank A-M, Prado RJ, Lagarde P (2004) Na K-edge XANES spectra of minerals and glasses. *Eur J Mineral* 16:809–816.
111. George AM, Stebbins JF (1996) Dynamics of Na in sodium aluminosilicate glasses and liquids. *Phys Chem Miner* 23:526–534.
112. Richet P (1987) Heat capacity of silicate glasses. *Chem Geol* 62:111–124. doi: 10.1016/0009-2541(87)90062-3
113. Uhlig H, Hoffmann MJ, Lamparter H-P, Aldinger F, Bellissent R, Steeb S (1996) Short-Range and Medium-Range Order in Lithium Silicate Glasses, Part I: Diffraction Experiments and Results. *J Am Ceram Soc* 79:2833–2838. doi: 10.1111/j.1151-2916.1996.tb08716.x

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

114. Zotov N, Keppler H (1998) The structure of sodium tetrasilicate glass from neutron diffraction, reverse Monte Carlo simulations and Raman spectroscopy. *Phys Chem Miner* 25:259–267.
115. Ispas S, Benoit M, Jund P, Jullien R (2002) Structural properties of glassy and liquid sodium tetrasilicate: comparison between ab initio and classical molecular dynamics simulations. *J Non-Cryst Solids* 307–310:946–955. doi: 10.1016/S0022-3093(02)01549-1
116. Du J, Corrales LR (2006) Compositional dependence of the first sharp diffraction peaks in alkali silicate glasses: A molecular dynamics study. *J Non-Cryst Solids* 352:3255–3269. doi: 10.1016/j.jnoncrysol.2006.05.025
117. Richet P (1984) Viscosity and configurational entropy of silicate melts. *Geochim Cosmochim Acta* 48:471–483. doi: 10.1016/0016-7037(84)90275-8
118. Toplis MJ (2001) Quantitative links between microscopic properties and viscosity of liquids in the system $\text{SiO}_2\text{--Na}_2\text{O}$. *Chem Geol* 174:321–331. doi: 10.1016/S0009-2541(00)00323-5
119. Mysen BO (1995) Experimental, in situ, high-temperature studies of properties and structure of silicate melts relevant to magmatic processes. *Eur J Mineral* 7:745–766.
120. Lee SK, Stebbins JF (2003) Nature of Cation Mixing and Ordering in Na-Ca Silicate Glasses and Melts. *J Phys Chem B* 107:3141–3148. doi: 10.1021/jp027489y
121. Lee SK, Mysen BO, Cody GD (2003) Chemical order in mixed-cation silicate glasses and melts. *Phys Rev B*. doi: 10.1103/PhysRevB.68.214206
122. George AM, Stebbins JF (1998) Structure and dynamics of magnesium in silicate melts: A high-temperature 25Mg NMR study. *Am Mineral* 83:1022–1029.
123. Li D, Peng M, Murata T (1999) Coordination and local structure of magnesium in silicate minerals and glasses: Mg K-Edge XANES study. *Can Mineral* 37:199–206.
124. Neuville DR, Cormier L, Flank A-M, Briois V, Massiot D (2004) Al speciation and Ca environment in calcium aluminosilicate glasses and crystals by Al and Ca K-edge X-ray absorption spectroscopy. *Chem Geol* 213:153–163.
125. Cormier L, Neuville DR (2004) Ca and Na environments in $\text{Na}_2\text{O--CaO--Al}_2\text{O}_3\text{--SiO}_2$ glasses: influence of cation mixing and cation-network interactions. *Chem Geol* 213:103–113.
126. Greaves GN, Fontaine A, Lagarde P, Raoux D, Gurman SJ (1981) Local structure of silicate glasses. *Nature* 293:611–616.
127. Greaves GN (1985) Exafs and the structure of glass. *J Non-Cryst Solids* 71:203–217.
128. Greaves GN (1989) EXAFS, glass structure and diffusion. *Philos Mag Part B* 60:793–800.
129. Greaves GN, Ngai KL (1995) Reconciling ionic-transport properties with atomic structure in oxide glasses. *Phys Rev B* 52:6358–6380. doi: 10.1103/PhysRevB.52.6358
130. Lee SK, Stebbins JF (2003) The distribution of sodium ions in aluminosilicate glasses: a high-field Na-23 MAS and 3Q MAS NMR study. *Geochim Cosmochim Acta* 67:1699–1709.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

131. Lee SK (2005) Microscopic origins of macroscopic properties of silicate melts and glasses at ambient and high pressure: Implications for melt generation and dynamics. *Geochim Cosmochim Acta* 69:3695–3710. doi: 10.1016/j.gca.2005.03.011
132. Lee S, Stebbins J (2009) Effects of the degree of polymerization on the structure of sodium silicate and aluminosilicate glasses and melts: An ^{17}O NMR study. *Geochim Cosmochim Acta* 73:1109–1119.
133. Poole, J.P. (1948) Viscosité à basse température des verres alcalino-silicatés. *Verres Refract* 2:222–228.
134. Poole JP (1949) Low-temperature viscosity of alkali silicate glasses. *J Am Ceram Soc* 32:230–233.
135. Mazurin O, Borisovskii E (1957) Neutralization reduction of electrical conductivity in silica glasses. *Sov Phys-Tech Phys* 2:243–254.
136. Isard JO (1969) The mixed alkali effect in glass. *J Non-Cryst Solids* 1:235–261.
137. Day DE (1976) Mixed alkali glasses - Their properties and uses. *J Non-Cryst Solids* 21:343–372.
138. Roling B, Ingram MD (2000) Mixed alkaline–earth effects in ion conducting glasses. *J Non-Cryst Solids* 265:113–119. doi: 10.1016/S0022-3093(99)00899-6
139. Calahoo C, Zwanziger JW (2017) The mixed modifier effect in ionic conductivity and mechanical properties for $x\text{MgO}-(50-x)\text{CaO}-50\text{SiO}_2$ glasses. *J Non-Cryst Solids* 460:6–18. doi: 10.1016/j.jnoncrysol.2017.01.017
140. Hummel W, Arndt J (1985) Variation of viscosity with temperature and composition in the plagioclase system. *Contrib Mineral Petrol* 90:83–92.
141. Lapp JC, Shelby JE (1987) The mixed alkali effect in lithium-sodium aluminosilicate glasses. *J Non-Cryst Solids* 95–96:889–896.
142. Kjeldsen J, Smedskjaer MM, Mauro JC, Youngman RE, Huang L, Yue Y (2013) Mixed alkaline earth effect in sodium aluminosilicate glasses. *J Non-Cryst Solids* 369:61–68. doi: 10.1016/j.jnoncrysol.2013.03.015
143. Smedskjaer MM, Rzoska SJ, Bockowski M, Mauro JC (2014) Mixed alkaline earth effect in the compressibility of aluminosilicate glasses. *J Chem Phys* 140:054511. doi: 10.1063/1.4863998
144. Frischat GH, Poggemann JF, Heide G (2004) Nanostructure and atomic structure of glass seen by atomic force microscopy. *J Non-Cryst Solids* 345–346:197–202.
145. Vessal B, Greaves GN, Marten PT, Chadwick AV, Mole R, Houde-Walter S (1992) Cation microsegregation and ionic mobility in mixed alkali glasses. *Nature* 356:504–506. doi: 10.1038/356504a0
146. Du J, Cormack AN (2004) The medium range structure of sodium silicate glasses: a molecular dynamics simulation. *J Non-Cryst Solids* 349:66–79. doi: 10.1016/j.jnoncrysol.2004.08.264

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
147. Meyer A, Horbach J, Kob W, Kargl F, Schober H (2004) Channel formation and intermediate range order in sodium silicate melts and glasses. *Phys Rev Lett* 93:1–4.
 148. Kargl F, Meyer A (2004) Inelastic neutron scattering on sodium aluminosilicate melts: sodium diffusion and intermediate range order. *Chem Geol* 213:165–172. doi: 10.1016/j.chemgeo.2004.08.040
 149. Kargl F, Meyer A, Koza MM, Schober H (2006) Formation of channels for fast-ion diffusion in alkali silicate melts: A quasielastic neutron scattering study. *Phys Rev B*. doi: 10.1103/PhysRevB.74.014304
 150. Kargl F, Meyer A (2008) Na-relaxation and intermediate range structure in sodium–potassium silicate melts. *Chem Geol* 256:278–285. doi: 10.1016/j.chemgeo.2008.06.046
 151. Adam G, Gibbs JH (1965) On the temperature dependence of cooperative relaxation properties in glass-forming liquids. *J Chem Phys* 139–146.
 152. Florian P, Vermillion KE, Grandinetti PJ, Farnan I, Stebbins JF (1996) Cation distribution in mixed alkali disilicate glasses. *J Am Chem Soc* 118:3493–3497.
 153. Angeli F, Villain O, Schuller S, Ispas S, Charpentier T (2011) Insight into sodium silicate glass structural organization by multinuclear NMR combined with first-principles calculations. *Geochim Cosmochim Acta* 75:2453–2469. doi: 10.1016/j.gca.2011.02.003
 154. Balasubramanian S, Rao KJ (1995) A molecular dynamics study of the mixed alkali effect in silicate glasses. *J Non-Cryst Solids* 181:157–174. doi: 10.1016/0022-3093(94)00489-7
 155. Matson DW, Sharma SK, Philpotts JA (1983) The structure of high-silica alkali-silicate glasses. A Raman spectroscopic investigation. *J Non-Cryst Solids* 58:323–352. doi: 10.1016/0022-3093(83)90032-7
 156. Le Losq C, Mysen BO, Cody GD (2015) Water and magmas: insights about the water solution mechanisms in alkali silicate melts from infrared, Raman, and ^{29}Si solid-state NMR spectroscopies. *Prog Earth Planet Sci*. doi: 10.1186/s40645-015-0052-7
 157. Brawer SA, White WB (1975) Raman spectroscopic investigation of the structure of silicate glasses. I. The binary alkali silicates. *J Chem Phys* 63:2421–2432. doi: 10.1063/1.431671
 158. Brawer SA, White WB (1977) Raman spectroscopic investigation of the structure of silicate glasses (II). Soda-alkaline earth-alumina ternary and quaternary glasses. *J Non-Cryst Solids* 23:261–278. doi: 10.1016/0022-3093(77)90009-6
 159. Furukawa T, Fox KE, White WB (1981) Raman spectroscopic investigation of the structure of silicate glasses. III. Raman intensities and structural units in sodium silicate glasses. *J Chem Phys* 75:3226–3237. doi: 10.1063/1.442472
 160. Voigt U, Lammert H, Eckert H, Heuer A (2005) Cation clustering in lithium silicate glasses: Quantitative description by solid-state NMR and molecular dynamics simulations. *Phys Rev B* 72:064207.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

161. Saal AE, Hart SR, Shimizu N, Hauri EH, Layne GD (1998) Pb Isotopic Variability in Melt Inclusions from Oceanic Island Basalts, Polynesia. *Science* 282:1481–1484. doi: 10.1126/science.282.5393.1481
162. Berry AJ, Danyushevsky LV, O'Neill HSC, Newville M, Sutton SR (2008) Oxidation state of iron in komatiitic melt inclusions indicates hot Archaean mantle. *Nature* 455:960–963. doi: 10.1038/nature07377
163. Spandler C, O'Neill HSC, Kamenetsky VS (2007) Survival times of anomalous melt inclusions from element diffusion in olivine and chromite. *Nature* 447:303–306. doi: 10.1038/nature05759
164. Roedder E, Weiblen PW (1970) Silicate Liquid Immiscibility in Lunar Magmas, Evidenced by Melt Inclusions in Lunar Rocks. *Science* 167:641–644. doi: 10.1126/science.167.3918.641
165. Hauri EH, Weinreich T, Saal AE, Rutherford MC, Orman JAV (2011) High Pre-Eruptive Water Contents Preserved in Lunar Melt Inclusions. *Science* 333:213–215. doi: 10.1126/science.1204626
166. Elkins-Tanton LT (2012) Magma Oceans in the Inner Solar System. *Annu Rev Earth Planet Sci* 40:113–139. doi: 10.1146/annurev-earth-042711-105503
167. Day DE, Rindone GE (1962) Properties of soda aluminosilicate glasses: III, coordination of aluminum ions. *J Am Ceram Soc* 45:579–581.
168. Riebling EF (1966) Structure of sodium aluminosilicate melts containing at least 50 mol% $\text{SiO}^{\text{circ}}_2$. *J Chem Phys* 44:2857–2865.
169. Taylor TD, Rindone GE (1970) Properties of Soda Aluminosilicate Glasses: V, Low-Temperature Viscosities. *J Am Ceram Soc* 53:692–695.
170. McKeown DA, Galeener FL, Brown Jr GE (1984) Raman studies of Al coordination in silica-rich sodium aluminosilicate glasses and some related minerals. *J Non-Cryst Solids* 68:361–378.
171. McMillan PF, Kirkpatrick RJ (1992) Al coordination in magnesium aluminosilicate glasses. *Am Mineral* 77:898–900.
172. Stebbins JF, Kroeker S, Lee SK, Kiczinski TJ (2000) Quantification of five- and six-coordinated aluminum ions in aluminosilicate and fluoride-containing glasses by high-field, high-resolution ^{27}Al NMR. *J Non-Cryst Solids* 275:1–6.
173. Toplis MJ, Kohn SC, Smith ME, Poplett JF (2000) Fivefold-coordinated aluminum in tectosilicate glasses observed by triple quantum MAS NMR. *Am Mineral* 85:1556–1560.
174. Neuville DR, Cormier L, Massiot D (2006) Al coordination and speciation in calcium aluminosilicate glasses: Effects of composition determined by ^{27}Al MQ-MAS NMR and Raman spectroscopy. *Chem Geol* 229:173–185.
175. Florian P, Sadiki N, Massiot D, Coutures JP (2007) ^{27}Al NMR Study of the Structure of Lanthanum- and Yttrium-Based Aluminosilicate Glasses and Melts. *J Phys Chem B* 111:9747–9757. doi: 10.1021/jp072061q

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

176. Stebbins JF, Xu Z (1997) NMR evidence for excess non-bridging oxygen in an aluminosilicate glass. *Nature* 390:60–62.
177. Mysen BO, Toplis MJ (2007) Structural behavior of Al^{3+} in peralkaline, metaluminous, and peraluminous silicate melts and glasses at ambient pressure. *Am Mineral* 92:933–946.
178. Thompson LM, Stebbins JF (2011) Non-bridging oxygen and high-coordinated aluminum in metaluminous and peraluminous calcium and potassium aluminosilicate glasses: High-resolution ^{17}O and ^{27}Al MAS NMR results. *Am Mineral* 96:841–853.
179. Allwardt JR (2005) Aluminum coordination and the densification of high-pressure aluminosilicate glasses. *Am Mineral* 90:1218–1222. doi: 10.2138/am.2005.1836
180. Kelsey KE, Stebbins JF, Singer DM, Brown Jr. GE, Mosenfelder JL, Asimow PD (2009) Cation field strength effects on high pressure aluminosilicate glass structure: Multinuclear NMR and La XAFS results. *Geochim Cosmochim Acta* 73:3914–3933. doi: 10.1016/j.gca.2009.03.040
181. Thompson LM, Stebbins JF (2012) Non-stoichiometric non-bridging oxygens and five-coordinated aluminum in alkaline earth aluminosilicate glasses: Effect of modifier cation size. *J Non-Cryst Solids* 358:1783–1789. doi: 10.1016/j.jnoncrysol.2012.05.022
182. Morin EI, Wu J, Stebbins JF (2014) Modifier cation (Ba, Ca, La, Y) field strength effects on aluminum and boron coordination in aluminoborosilicate glasses: the roles of fictive temperature and boron content. *Appl Phys A* 116:479–490. doi: 10.1007/s00339-014-8369-4
183. Takahashi S, Neuville DR, Takebe H (2015) Thermal properties, density and structure of percalcic and peraluminous $\text{CaO-Al}_2\text{O}_3\text{-SiO}_2$ glasses. *J Non-Cryst Solids* 411:5–12. doi: 10.1016/j.jnoncrysol.2014.12.019
184. Novikov AN, Neuville DR, Hennet L, Gueguen Y, Thiaudière D, Charpentier T, Florian P Al and Sr environment in tectosilicate glasses and melts: Viscosity, Raman and NMR investigation. *Chem Geol*. doi: 10.1016/j.chemgeo.2016.11.023
185. Iftekhhar S, Pahari B, Okhotnikov K, Jaworski A, Svensson B, Grins J, Edén M (2012) Properties and Structures of $\text{RE}_2\text{O}_3\text{-Al}_2\text{O}_3\text{-SiO}_2$ (RE = Y, Lu) Glasses Probed by Molecular Dynamics Simulations and Solid-State NMR: The Roles of Aluminum and Rare-Earth Ions for Dictating the Microhardness. *J Phys Chem C* 116:18394–18406. doi: 10.1021/jp302672b
186. Neuville DR, Cormier L, Massiot D (2004) Al environment in tectosilicate and peraluminous glasses: A ^{27}Al MQ-MAS NMR, Raman, and XANES investigation. *Geochim Cosmochim Acta* 68:5071–5079. doi: 10.1016/j.gca.2004.05.048
187. Allwardt JR, Poe BT, Stebbins JF (2005) The effect of fictive temperature on Al coordination in high-pressure (10 GPa) sodium aluminosilicate glasses. *Am Mineral* 90:1453–1457.
188. Kanehashi K, Stebbins JF (2007) In situ high temperature ^{27}Al NMR study of structure and dynamics in a calcium aluminosilicate glass and melt. *J Non-Cryst Solids* 353:4001–4010. doi: 10.1016/j.jnoncrysol.2007.06.030

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

189. Neuville DR, Cormier L, de Ligny D, Roux J, Flank A-M, Lagarde P (2008) Environments around Al, Si, and Ca in aluminate and aluminosilicate melts by X-ray absorption spectroscopy at high temperature. *Am Mineral* 93:228–234. doi: 10.2138/am.2008.2646
190. Malfait WJ, Verel R, Ardia P, Sanchez-Valle C (2012) Aluminum coordination in rhyolite and andesite glasses and melts: Effect of temperature, pressure, composition and water content. *Geochim Cosmochim Acta* 77:11–26. doi: 10.1016/j.gca.2011.11.011
191. Bista S, Stebbins JF, Hankins WB, Sisson TW (2015) Aluminosilicate melts and glasses at 1 to 3 GPa: Temperature and pressure effects on recovered structural and density changes. *Am Mineral* 100:2298–2307. doi: 10.2138/am-2015-5258
192. Stebbins JF, Dubinsky EV, Kanehashi K, Kelsey KE (2008) Temperature effects on non-bridging oxygen and aluminum coordination number in calcium aluminosilicate glasses and melts. *Geochim Cosmochim Acta* 72:910–925. doi: 10.1016/j.gca.2007.11.018
193. Thompson LM, Stebbins JF (2013) Interaction between composition and temperature effects on non-bridging oxygen and high-coordinated aluminum in calcium aluminosilicate glasses. *Am Mineral* 98:1980–1987. doi: 10.2138/am.2013.4511
194. Wu J, Stebbins JF (2010) Quench rate and temperature effects on boron coordination in aluminoborosilicate melts. *J Non-Cryst Solids* 356:2097–2108.
195. Kiczinski TJ, Du L-S, Stebbins J (2005) The effect of fictive temperature on the structure of E-glass: A high resolution, multinuclear NMR study. *J Non-Cryst Solids* 351:3571–3578. doi: 10.1016/j.jnoncrysol.2005.09.026
196. Lee SK, Cody GD, Fei Y, Mysen BO (2006) The effect of Na/Si on the structure of sodium silicate and aluminosilicate glasses quenched from melts at high pressure: A multi-nuclear (Al-27, Na-23, O-17) 1D and 2D solid-state NMR study. *Chem Geol* 229:162–172. doi: 10.1016/j.chemgeo.2006.01.018
197. Toplis MJ, Dingwell DB, Lenci T (1997) Peraluminous viscosity maxima in Na₂O-Al₂O₃-SiO₂ liquids: The role of triclusters in tectosilicate melts. *Geochim Cosmochim Acta* 61:2605–2612. doi: 10.1016/S0016-7037(97)00126-9
198. Gruener G, Odier O, De Sousa Meneses D, Florian P, Richet P (2001) Bulk and local dynamics in glass-forming liquids: A viscosity, electrical conductivity, and NMR study of aluminosilicate melts. *Phys Rev B* 64:024206–1:5.
199. Neuville D, Cormier L, Montouillout V, Massiot D (2007) Local Al site distribution in aluminosilicate glasses by ²⁷Al MQMAS NMR. *J Non-Cryst Solids* 353:180–184.
200. Loewenstein W (1954) The distribution of aluminium in the tetrahedra of silicates and aluminates. *Am Mineral* 39:92–97.
201. Lee SK, Stebbins JF (1999) The degree of aluminum avoidance in aluminosilicate glasses. *Am Mineral* 84:937–945.

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
202. Lee SK, Stebbins JF (2002) Extent of intermixing among framework units in silicate glasses and melts. *Geochim Cosmochim Acta* 66:303–309.
 203. Mysen BO, Lucier A, Cody GD (2003) The structural behavior of Al^{3+} in peralkaline melts and glasses in the system $\text{Na}_2\text{O}-\text{Al}_2\text{O}_3-\text{SiO}_2$. *Am Mineral* 88:1668–1678.
 204. Lee SK, Stebbins J (2000) Al-O-Al and Si-O-Si sites in framework aluminosilicate glasses with Si/Al=1: quantification of framework disorder. *J Non-Cryst Solids* 270:260–264.
 205. Xiang Y, Du J, Smedskjaer MM, Mauro JC (2013) Structure and properties of sodium aluminosilicate glasses from molecular dynamics simulations. *J Chem Phys* 139:044507. doi: 10.1063/1.4816378
 206. Bauchy M (2014) Structural, vibrational, and elastic properties of a calcium aluminosilicate glass from molecular dynamics simulations: The role of the potential. *J Chem Phys* 141:024507. doi: 10.1063/1.4886421
 207. Lee SK, Kim H-I, Kim EJ, Mun KY, Ryu S (2016) Extent of Disorder in Magnesium Aluminosilicate Glasses: Insights from ^{27}Al and ^{17}O NMR. *J Phys Chem C* 120:737–749. doi: 10.1021/acs.jpcc.5b10799
 208. Neuville DR, Mysen BO (1996) Role of aluminium in the silicate network: In situ, high-temperature study of glasses and melts on the join $\text{SiO}_2\text{-NaAlO}_2$. *Geochim Cosmochim Acta* 60:1727–1737.
 209. Toplis MJ, Dingwell DB, Hess K-U, Lenci T (1997) Viscosity, fragility, and configurational entropy of melts along the join $\text{SiO}_2\text{-NaAlSiO}_4$. *Am Mineral* 82:979–990.
 210. Vuilleumier R, Sator N, Guillot B (2009) Computer modeling of natural silicate melts: What can we learn from ab initio simulations. *Geochim Cosmochim Acta* 73:6313–6339. doi: 10.1016/j.gca.2009.07.013
 211. Navrotsky A, Geisinger KL, McMillan P, Gibbs GV (1985) The tetrahedral framework in glasses and melts — inferences from molecular orbital calculations and implications for structure, thermodynamics, and physical properties. *Phys Chem Miner* 11:284–298. doi: 10.1007/BF00307406
 212. Li W, Garofalini SH (2004) Molecular dynamics simulation of lithium diffusion in $\text{Li}_2\text{O}-\text{Al}_2\text{O}_3-\text{SiO}_2$ glasses. *Solid State Ion* 166:365–373. doi: 10.1016/j.ssi.2003.11.015
 213. Zirl DM, Garofalini SH (1990) Structure of Sodium Aluminosilicate Glasses. *J Am Ceram Soc* 73:2848–2856. doi: 10.1111/j.1151-2916.1990.tb06685.x
 214. Jackson WE, Brown GE, Ponader CW (1987) X-ray absorption study of the potassium coordination environment in glasses from the $\text{NaAlSi}_3\text{O}_8\text{-KAlSi}_3\text{O}_8$ binary. *J Non-Cryst Solids* 93:311–322. doi: 10.1016/S0022-3093(87)80177-1
 215. Petkov V, Gerber T, Himmel B (1998) Atomic ordering in $\text{Ca}_{x/2}\text{Al}_x\text{Si}_{1-x}\text{O}_2$ glasses ($x=0, 0.34, 0.5, 0.68$) by energy-dispersive X-ray diffraction. *Phys Rev B* 58:11982.

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
216. Petkov V, Billinge SJL, Shastri SD, Himmel B (2000) Polyhedral units and network connectivity in calcium aluminosilicate glasses from high-energy X-ray diffraction. *Phys Rev Lett* 85:3436.
 217. Cormier L, Neuville DR, Calas G (2000) Structure and properties of low-silica calcium aluminosilicate glasses. *J Non-Cryst Solids* 274:110–114. doi: 10.1016/S0022-3093(00)00209-X
 218. Guignard M, Cormier L (2008) Environments of Mg and Al in MgO–Al₂O₃–SiO₂ glasses: A study coupling neutron and X-ray diffraction and Reverse Monte Carlo modeling. *Chem Geol* 256:111–118. doi: 10.1016/j.chemgeo.2008.06.008
 219. Iuga D, Morais C, Gan Z, Neuville DR, Cormier L, Massiot D (2005) NMR Heteronuclear Correlation between Quadrupolar Nuclei in Solids. *J Am Chem Soc* 127:11540–11541.
 220. Stebbins JF, Lee SK, Oglesby JV (1999) Al–O–Al oxygen sites in crystalline aluminates and aluminosilicate glasses; high-resolution oxygen-17 NMR results. *Am Mineral* 84:983–986. doi: 10.2138/am-1999-5-635
 221. Oglesby JV, Zhao P, Stebbins JF (2002) Oxygen sites in hydrous aluminosilicate glasses: The role of Al–O–Al and H₂O. *Geochim Cosmochim Acta* 66:291–301.
 222. Harvey J-P, Asimow PD (2015) Current limitations of molecular dynamic simulations as probes of thermo-physical behavior of silicate melts. *Am Mineral* 100:1866–1882. doi: 10.2138/am-2015-5159
 223. Lacy ED (1963) Aluminium in glasses and melts. *Physic Chem Glas* 4:234–238.
 224. Deer WA, Howie RA, Zussman J (1992) An introduction to the rock-forming minerals. Longman Group Limited, London
 225. Kushiro I (1978) Viscosity and structural changes of albite (NaAlSi₃O₈) melt at high pressures. *Earth Planet Sci Lett* 41:87–90.
 226. Lee SK, Cody GD, Fei Y, Mysen BO (2004) Nature of polymerization and properties of silicate melts and glasses at high pressure. *Geochim Cosmochim Acta* 68:4189–4200. doi: 10.1016/j.gca.2004.04.002
 227. Taylor M, Brown GE (1979) Structure of mineral glasses—I. The feldspar glasses NaAlSi₃O₈, KAlSi₃O₈, CaAl₂Si₂O₈. *Geochim Cosmochim Acta* 43:61–75.
 228. Urbain G, Bottinga Y, Richet P (1982) Viscosity of liquid silica, silicates and aluminosilicates. *Geochim Cosmochim Acta* 46:1061–1072. doi: 10.1016/0016-7037(82)90059-X
 229. Kubicki JD, Sykes D (1993) Molecular orbital calculations of vibrations in three-membered aluminosilicate rings. *Phys Chem Miner* 19:381–391. doi: 10.1007/BF00202975
 230. Hater W, Müller-Warmuth W, Meier M, Frischat GH (1989) High-resolution solid-state NMR studies of mixed-alkali silicate glasses. *J Non-Cryst Solids* 113:210–212.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

231. Ali F, Chadwick AV, Greaves GN, Jermy MC, Ngai KL, Smith ME (1995) Examination of the mixed-alkali effect in (Li,Na) disilicate glasses by nuclear magnetic resonance and conductivity measurements. *Solid State Nucl Magn Reson* 5:133–143.
232. Bray PJ, Emerson JF, Lee D, Feller SA, Bain DL, Feil DA (1991) NMR and NQR studies of glass structure. *J Non-Cryst Solids* 129:240–248.
233. Emerson JF, Bray PJ (1994) Nuclear magnetic resonance and transmission electron microscopy studies of mixed-alkali silicate glasses. *J Non-Cryst Solids* 169:87–95.
234. Kelsey KE, Allwardt JR, Stebbins JF (2008) Ca–Mg mixing in aluminosilicate glasses: An investigation using ^{17}O MAS and $^3\text{QMAS}$ and ^{27}Al MAS NMR. *J Non-Cryst Solids* 354:4644–4653. doi: 10.1016/j.jnoncrysol.2008.05.049
235. Weigel C, Le Losq C, Vialla R, Dupas C, Clément S, Neuville DR, Rufflé B (2016) Elastic moduli of XAlSiO_4 aluminosilicate glasses: effects of charge-balancing cations. *J Non-Cryst Solids* 447:267–272. doi: 10.1016/j.jnoncrysol.2016.06.023
236. Greaves GN, Greer AL, Lakes RS, Rouxel T (2011) Poisson's ratio and modern materials. *Nat Mater* 10:823–837. doi: 10.1038/nmat3134
237. Hunault MOJY, Galois L, Lelong G, Newville M, Calas G (2016) Effect of cation field strength on Co^{2+} speciation in alkali-borate glasses. *J Non-Cryst Solids* 451:101–110. doi: 10.1016/j.jnoncrysol.2016.06.025
238. Wang C, Peng M, Jiang N, Jiang X, Zhao C, Qiu J (2007) Tuning the Eu luminescence in glass materials synthesized in air by adjusting glass compositions. *Mater Lett* 61:3608–3611. doi: 10.1016/j.matlet.2006.11.133
239. Farias AM, Sandrini M, Viana JRM, Baesso ML, Bento AC, Rohling JH, Guyot Y, De Ligny D, Nunes LAO, Gandra FG, Sampaio JA, Lima SM, Andrade LHC, Medina AN (2015) Emission tunability and local environment in europium-doped OH–free calcium aluminosilicate glasses for artificial lighting applications. *Mater Chem Phys* 156:214–219. doi: 10.1016/j.matchemphys.2015.03.002
240. Cicconi MR, Veber A, de Ligny D, Rocherullé J, Lebullenger R, Tessier F (2017) Chemical tunability of europium emission in phosphate glasses. *J Lumin* 183:53–61. doi: 10.1016/j.jlumin.2016.11.019
241. Dingwell DB, Virgo D (1987) The effect of oxidation state on the viscosity of melts in the system $\text{Na}_2\text{O}-\text{FeO}-\text{Fe}_2\text{O}_3-\text{SiO}_2$. *Geochim Cosmochim Acta* 51:195–205.
242. Dingwell DB (1989) Shear viscosities of ferrosilicate liquids. *Am Mineral* 74:1038–1044.
243. Dingwell DB (1991) Redox viscometry of some Fe-bearing silicate melts. *Am Mineral* 76:1560–1562.
244. Fudali RF (1965) Oxygen fugacities of basaltic and andesitic magmas. *Geochim Cosmochim Acta* 29:1063–1075.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

245. Mysen BO, Scarfe CM, Cronin DJ (1985) Viscosity and structure of iron- and aluminum-bearing calcium silicate melts at 1 atm. *Am Mineral* 70:487–498.
246. Dingwell, D. B., Virgo, D. (1988) Viscosities of melts in the Na₂O-FeO-Fe₂O₃-SiO₂ system and factors controlling relative viscosities of fully polymerized silicate melts. *Geochim Cosmochim Acta* 52:395–403.
247. Schreiber HD (1977) Redox states of Ti, Zr, Hf, Cr, and EU in basaltic magmas - an experimental study. *Lunar Planet Sci Conf Proc* 8:1785–1807.
248. Schreiber HD, Thanyasiri T, Lach JJ (1978) Redox equilibria of Ti, Cr, and Eu in silicate melts--reduction potentials and mutual interactions. *Phys Chem Glas* 19:126–139.
249. Schreiber HD, Lauer HV, Thanyasiri T (1980) The redox state of cerium in basaltic magmas: an experimental study of iron-cerium interactions in silicate melts. *Geochim Cosmochim Acta* 44:1599–1612. doi: 10.1016/0016-7037(80)90120-9
250. Schreiber HD (1986) Redox processes in glass-forming melts. *J Non-Cryst Solids* 84:129–141.
251. Schreiber HD (1987) An electrochemical series of redox couples in silicate melts: A review and applications to geochemistry. *J Geophys Res Solid Earth* 92:9225–9232. doi: 10.1029/JB092iB09p09225
252. Schreiber HD, Kochanowski BK, Schreiber CW, Morgan AB, Coolbaugh MT, Dunlap TG (1994) Compositional dependence of redox equilibria in sodium silicate glasses. *J Non-Cryst Solids* 177:340–346. doi: 10.1016/0022-3093(94)90548-7
253. Schreiber HD, Wilk Jr NR, Schreiber CW (1999) A comprehensive electromotive force series of redox couples in soda–lime–silicate glass. *J Non-Cryst Solids* 253:68–75. doi: 10.1016/S0022-3093(99)00344-0
254. Johnston WD (1964) Oxidation-Reduction Equilibria in Iron-Containing Glass. *J Am Ceram Soc* 47:198–201. doi: 10.1111/j.1151-2916.1964.tb14392.x
255. Mysen BO (1987) Redox equilibria and coordination of Fe²⁺ and Fe³⁺ in silicate glasses from ⁵⁷Fe Mossbauer spectroscopy. *J Non-Cryst Solids* 95:247–254.
256. Mysen BO, Virgo D (1989) Redox equilibria, structure, and properties of Fe-bearing aluminosilicate melts: Relationships among temperature, composition, and oxygen fugacity in the system Na₂O-Al₂O₃-SiO₂-Fe-O. *Am Mineral* 74:58–76.
257. Kress VC, Carmichael IS (1991) The compressibility of silicate liquids containing Fe₂O₃ and the effect of composition, temperature, oxygen fugacity and pressure on their redox states. *Contrib Mineral Petrol* 108:82–92.
258. Lauer HV, Morris RV (1977) Redox Equilibria of Multivalent Ions in Silicate Glasses. *J Am Ceram Soc* 60:443–451. doi: 10.1111/j.1151-2916.1977.tb15530.x
259. Mysen BO, Virgo D, Seifert F (1984) Redox equilibria of iron in alkaline earth silicate melts: relationships between melt structure, oxygen fugacity, temperature and properties of iron-bearing silicate liquids. *Am Mineral* 69:834–847.

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
260. Cicconi MR, Giuli G, Ertel-Ingrisch W, Paris E, Dingwell DB (2015) The effect of the [Na/(Na+K)] ratio on Fe speciation in phonolitic glasses. *Am Mineral* 100:1610–1619. doi: 10.2138/am-2015-5155
 261. Drickamer HG, Bastron VC, Fisher DC, Grenoble DC (1970) The high-pressure chemistry of iron. *J Solid State Chem* 2:94–104. doi: 10.1016/0022-4596(70)90038-1
 262. Wang PJ, Drickamer HG (1973) Reduction of Cu(II) at high pressure. *J Chem Phys* 59:713–717. doi: 10.1063/1.1680080
 263. Gibbons RV, Ahrens TJ, Rossman GR (1974) A spectrographic interpretation of the shock-produced color change in rhodonite (MnSiO₃): The shock-induced reduction of Mn(III) to Mn(II). *Am Mineral* 59:177–182.
 264. O'Neill HSC (2006) An experimental determination of the effect of pressure on the Fe³⁺/Fe ratio of an anhydrous silicate melt to 3.0 GPa. *Am Mineral* 91:404–412. doi: 10.2138/am.2005.1929
 265. Zhang HL, Hirschmann MM, Cottrell E, Withers AC (2017) Effect of pressure on Fe³⁺/ΣFe ratio in a mafic magma and consequences for magma ocean redox gradients. *Geochim Cosmochim Acta* 204:83–103. doi: 10.1016/j.gca.2017.01.023
 266. Burnham AD, Berry AJ, Halse HR, Schofield PF, Cibir G, Mosselmans JFW (2015) The oxidation state of europium in silicate melts as a function of oxygen fugacity, composition and temperature. *Chem Geol* 411:248–259. doi: 10.1016/j.chemgeo.2015.07.002
 267. Paul A, Douglas RW (1965) Ferrous-ferric equilibrium in binary alkali silicate glasses. *Phys Chem Glas* 6:207.
 268. Dickenson MP, Hess PC (1982) Redox equilibria and the structural role of iron in aluminosilicate melts. *Contrib Mineral Petrol* 78:352–357. doi: 10.1007/BF00398931
 269. Mysen, B. O., Virgo, D., Neumann, E.-R., Seifert, F. A. (1985) Redox equilibria and the structural states of ferric and ferrous iron in melts in the system CaO-MgO-Al₂O₃-SiO₂-FeO: relationships between redox equilibria, melt structure and liquidus phase equilibria. *Am Mineral* 70:317–331.
 270. Berry AJ, O'Neill HSC, Scott DR, Shelley JMG (2006) The effect of composition on Cr²⁺/Cr³⁺ in silicate melts. *Am Mineral* 91:1901–1908.
 271. O'Neill HSC, Eggins SM (2002) The effect of melt composition on trace element partitioning: an experimental investigation of the activity coefficients of FeO, NiO, CoO, MoO₂ and MoO₃ in silicate melts. *Chem Geol* 186:151–181. doi: 10.1016/S0009-2541(01)00414-4
 272. Giuli G, Paris E, Mungall J, Romano C, Dingwell D (2004) V oxidation state and coordination number in silicate glasses by XAS. *Am Mineral* 89:1640–1646.
 273. Farges F, Linnen RL, Brown GE (2006) REDOX AND SPECIATION OF TIN IN HYDROUS SILICATE GLASSES: A COMPARISON WITH Nb, Ta, Mo AND W. *Can Mineral* 44:795–810. doi: 10.2113/gscanmin.44.3.795

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
274. Morris RV, Haskin LA (1974) EPR measurement of the effect of glass composition on the oxidation states of europium. *Geochim Cosmochim Acta* 38:1435–1445. doi: 10.1016/0016-7037(74)90097-0
275. Cicconi MR, Giuli G, Paris E, Ertel-Ingrisch W, Ulmer P, Dingwell DB (2012) Europium oxidation state and local structure in silicate glasses. *Am Mineral* 97:918–929.
276. Duffy JA, Ingram MD (1976) An interpretation of glass chemistry in terms of the optical basicity concept. *J Non-Cryst Solids* 21:373–410. doi: 10.1016/0022-3093(76)90027-2
277. Duffy JA (1993) A review of optical basicity and its applications to oxidic systems. *Geochim Cosmochim Acta* 57:3961–3970. doi: 10.1016/0016-7037(93)90346-X
278. Leboutteiller A, Courtine P (1998) Improvement of a Bulk Optical Basicity Table for Oxidic Systems. *J Solid State Chem* 137:94–103. doi: 10.1006/jssc.1997.7722
279. Ottonello G, Moretti R, Marini L, Vetushi Zuccolini M (2001) Oxidation state of iron in silicate glasses and melts: a thermochemical model. *Chem Geol* 174:157–179. doi: 10.1016/S0009-2541(00)00314-4
280. Moretti R (2005) Polymerisation, basicity, oxidation state and their role in ionic modelling of silicate melts. *Ann Geophys.* doi: 10.4401/ag-3221
281. Sack RO, Carmichael ISE, Rivers M, Ghiorso MS (1980) Ferric-Ferrous equilibria in natural silicate liquids at 1 bar. *Contrib Mineral Petrol* 75:369–376.
282. Kilinc A, Carmichael ISE, Rivers ML, Sack RO (1983) The Ferric-Ferrous ratio of natural silicate liquids equilibrated in air. *Contrib Mineral Petrol* 83:136–140.
283. Burnham AD, Berry AJ (2014) The effect of oxygen fugacity, melt composition, temperature and pressure on the oxidation state of cerium in silicate melts. *Chem Geol* 366:52–60. doi: 10.1016/j.chemgeo.2013.12.015
284. Tangeman JA, Lange R, Forman L (2001) Ferric-ferrous equilibria in K₂O-FeO-Fe₂O₃-SiO₂ melts. *Geochim Cosmochim Acta* 65:1809–1819. doi: 10.1016/S0016-7037(00)00561-5
285. Kress VC, Carmichael ISE (1988) Stoichiometry of the iron oxidation reaction in silicate melts. *Am Mineral* 73:1267–1274.
286. Thornber CR, Roeder PL, Foster JR (1980) The effect of composition on the ferric-ferrous ratio in basaltic liquids at atmospheric pressure. *Geochim Cosmochim Acta* 44:525–532. doi: 10.1016/0016-7037(80)90048-4
287. Dyar MD (1985) A review of Moessbauer data on inorganic glasses; the effects of composition on iron valency and coordination. *Am Mineral* 70:304–316.
288. Virgo D, Mysen BO (1985) The structural state of iron in oxidized vs. reduced glasses at 1 atm: A⁵⁷Fe Mössbauer study. *Phys Chem Miner* 12:65–76. doi: 10.1007/BF01046829

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
289. Hannoyer B, Lenglet M, Dürr J, Cortes R (1992) Spectroscopic evidence of octahedral iron (III) in soda-lime silicate glasses. *J Non-Cryst Solids* 151:209–216. doi: 10.1016/0022-3093(92)90031-E
 290. Galois L, Calas G, Arrio MA (2001) High-resolution XANES spectra of iron in minerals and glasses: structural information from the pre-edge region. *Chem Geol* 174:307–319.
 291. Partzsch GM, Lattard D, McCammon C (2004) Mossbauer spectroscopic determination of $\text{Fe}^{3+}/\text{Fe}^{2+}$ in synthetic basaltic glass: a test of empirical fO_2 equations under superliquidus and subliquidus conditions. *Contrib Mineral Petrol* 145:565–580.
 292. Farges F, Lefrère Y, Rossano S, Berthereau A, Calas G, Brown Jr. GE (2004) The effect of redox state on the local structural environment of iron in silicate glasses: a combined XAFS spectroscopy, molecular dynamics, and bond valence study. *J Non-Cryst Solids* 344:176–188. doi: 10.1016/j.jnoncrysol.2004.07.050
 293. Jackson WE, Farges F, Yeager M, Mabrouk PA, Rossano S, Waychunas GA, Solomon EI, Brown GE (2005) Multi-spectroscopic study of Fe(II) in silicate glasses: Implications for the coordination environment of Fe(II) in silicate melts. *Geochim Cosmochim Acta* 69:4315–4332. doi: 10.1016/j.gca.2005.01.008
 294. Métrich N, Susini J, Foy E, Farges F, Massare D, Sylla L, Lequien S, Bonnin-Mosbah M (2006) Redox state of iron in peralkaline rhyolitic glass/melt: X-ray absorption micro-spectroscopy experiments at high temperature. *Chem Geol* 231:350–363. doi: 10.1016/j.chemgeo.2006.02.001
 295. Wilke M, Farges F, Partzsch GM, Schmidt C, Behrens H (2007) Speciation of Fe in silicate glasses and melts by in-situ XANES spectroscopy. *Am Mineral* 92:44–56. doi: 10.2138/am.2007.1976
 296. Rossano S, Behrens H, Wilke M (2008) Advanced analyses of ^{57}Fe Mössbauer data of aluminosilicate glasses. *Phys Chem Miner* 35:77–93. doi: 10.1007/s00269-007-0200-8
 297. Giuli G, Paris E, Hess K-U, Dingwell DB, Cicconi MR, Eeckhout SG, Fehr KT, Valenti P (2011) XAS determination of the Fe local environment and oxidation state in phonolite glasses. *Am Mineral* 96:631–636. doi: 10.2138/am.2011.3464
 298. Giuli G, Alonso-Mori R, Cicconi MR, Paris E, Glatzel P, Eeckhout SG, Scaillet B (2012) Effect of alkalis on the Fe oxidation state and local environment in peralkaline rhyolitic glasses. *Am Mineral* 97:468–475. doi: 10.2138/am.2012.3888
 299. Calas G, Petiau J (1983) Coordination of iron in oxide glasses through high-resolution K-edge spectra: information from the pre-edge. *Solid State Commun* 48:625–629.
 300. Waychunas GA, Brown GE, Ponader CW, Jackson WE (1988) Evidence from X-ray absorption for network-forming Fe^{2+} in molten alkali silicates. *Nature* 332:251–253. doi: 10.1038/332251a0
 301. Keppler HN (1992) Crystal field spectra and geochemistry of transition metal ions in silicate melts and glasses. *Am Mineral* 77:62–75.

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

302. Jackson WE, Leon JM de, Brown GE, Waychunas GA, Conradson SD, Combes J-M (1993) High-Temperature XAS Study of Fe₂SiO₄ Liquid: Reduced Coordination of Ferrous Iron. *Science* 262:229–233. doi: 10.1126/science.262.5131.229
303. Rossano S, Balan E, Morin G, Bauer J-P, Calas G, Brouder C (1999) ⁵⁷Fe Mössbauer spectroscopy of tektites. *Phys Chem Miner* 26:530–538. doi: 10.1007/s002690050216
304. Rossano S, Ramos A, Delaye J-M, Creux S, Filipponi A, Brouder C, Calas G (2000) EXAFS and Molecular Dynamics combined study of CaO – FeO – 2SiO₂ glass. New insight into site significance in silicate glasses. *EPL Europhys Lett* 49:597. doi: 10.1209/epl/i2000-00192-1
305. Leonteva A (1940) Measurements of the viscosity of obsidian and of hydrated glasses. *Izv Akad Nauk SSSR Ser Geol* 2:44–54.
306. Saucier H (1952) Quelques expériences sur la viscosité à haute température de verres ayant la composition d'un granite. Influence de la vapeur d'eau sous pression. *Bull Société Fr Minéralogie Cristallogr* 75:1–45.
307. Friedman I, Long W, Smith RL (1963) Viscosity and water content of rhyolite glass. *J Geophys Res* 68:6523–6535.
308. Shaw HR (1963) Obsidian-H₂O viscosities at 1000 and 2000 bars in the temperature range 700^{°C} to 900^{°C}. *J Geophys Res* 68:6337–6343.
309. Dingwell DB, Romano C, Hess K-U (1996) The effect of water on the viscosity of a haplogranitic melt under P-T-X conditions relevant to silicic volcanism. *Contrib Mineral Petrol* 124:19–28. doi: 10.1007/s004100050170
310. Richet P, Lejeune A-M, Holtz F, Roux J (1996) Water and the viscosity of andesite melts. *Chem Geol* 128:185–197. doi: 10.1016/0009-2541(95)00172-7
311. Richet P, Polian A (1998) Water as a Dense Icelike Component in Silicate Glasses. *Science* 281:396–398. doi: 10.1126/science.281.5375.396
312. Ochs FA, Lange RA (1999) The Density of Hydrous Magmatic Liquids. *Science* 283:1314–1317. doi: 10.1126/science.283.5406.1314
313. Richet P, Whittington A, Holtz F, Behrens H, Ohlhorst S, Wilke M (2000) Water and the density of silicate glasses. *Contrib Mineral Petrol* 138:337–347.
314. Malfait WJ, Seifert R, Petitgirard S, Mezouar M, Sanchez-Valle C (2014) The density of andesitic melts and the compressibility of dissolved water in silicate melts at crustal and upper mantle conditions. *Earth Planet Sci Lett* 393:31–38. doi: 10.1016/j.epsl.2014.02.042
315. Malfait WJ, Seifert R, Petitgirard S, Perrillat J-P, Mezouar M, Ota T, Nakamura E, Lerch P, Sanchez-Valle C (2014) Supervolcano eruptions driven by melt buoyancy in large silicic magma chambers. *Nat Geosci* 7:122–125. doi: 10.1038/ngeo2042
316. Andújar J, Scaillet B (2012) Relationships between pre-eruptive conditions and eruptive styles of phonolite–trachyte magmas. *Lithos* 152:122–131. doi: 10.1016/j.lithos.2012.05.009

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.

317. Le Losq C, Neuville DR, Moretti R, Kyle PR, Oppenheimer C (2015) Rheology of phonolitic magmas – the case of the Erebus lava lake. *Earth Planet Sci Lett* 411:53–61. doi: 10.1016/j.epsl.2014.11.042
318. Luongo G, Perrotta A, Scarpato C (2003) Impact of the AD 79 explosive eruption on Pompeii, I. Relations amongst the depositional mechanisms of the pyroclastic products, the framework of the buildings and the associated destructive events. *J Volcanol Geotherm Res* 126:201–223. doi: 10.1016/S0377-0273(03)00146-X
319. Luongo G, Perrotta A, Scarpato C, De Carolis E, Patricelli G, Ciarallo A (2003) Impact of the AD 79 explosive eruption on Pompeii, II. Causes of death of the inhabitants inferred by stratigraphic analysis and areal distribution of the human casualties. *J Volcanol Geotherm Res* 126:169–200. doi: 10.1016/S0377-0273(03)00147-1
320. Balcone-Boissard H, Boudon G, Villemant B (2011) Textural and geochemical constraints on eruptive style of the 79 AD eruption at Vesuvius. *Bull Volcanol* 73:279–294.
321. Shea T, Hellebrand E, Gurioli L, Tuffen H (2014) Conduit- to Localized-scale Degassing during Plinian Eruptions: Insights from Major Element and Volatile (Cl and H₂O) Analyses within Vesuvius ad 79 Pumice. *J Petrol* 55:315–344. doi: 10.1093/petrology/egt069
322. Dingwell DB (1996) Volcanic Dilemma - Flow or Blow? *Science* 273:1054–1055. doi: 10.1126/science.273.5278.1054
323. Papale P (1999) Strain-induced magma fragmentation in explosive eruptions. *Nature* 397:425–428. doi: 10.1038/17109
324. Zhang Y (1999) A criterion for the fragmentation of bubbly magma based on brittle failure theory. *Nature* 402:648–650. doi: 10.1038/45210
325. Scholze H (1960) Zur Frage der unterscheidung zwischen H₂O-Molekeln und OH-gruppen in gläsern und mineralen. *Naturwissenschaften* 47:226–227.
326. Moulson AJ (1958) Entry of Water into Silica Glass. *Nature* 182:200–201.
327. Moulson AJ, Roberts JP (1961) Water in silica glass. *Trans Faraday Soc* 57:1208–1216.
328. Bartholomew RF, Butler BL, Hoover HL, Wu CK (1980) Infrared Spectra of a Water-Containing Glass. *J Am Ceram Soc* 63:481–485. doi: 10.1111/j.1151-2916.1980.tb10748.x
329. Mysen BO, Virgo D, Harrison WJ, Scarfe CM (1980) Solubility mechanisms of H₂O in silicate melts at high pressures and temperatures: a Raman spectroscopy study. *Am Mineral* 65:900–914.
330. Stolper A (1982) Water in silicate glasses: An infrared spectroscopic study. *Contrib Mineral Petrol* 81:1–17.
331. Mysen BO, Virgo D (1986) Volatiles in silicate melts at high pressure and temperature 1. Interaction between OH groups and Si⁴⁺, Al³⁺, Ca²⁺, Na⁺ and H⁺. *Chem Geol* 57:303–331.
332. McMillan P, Remmele RLJ (1986) Hydroxylsites in SiO₂ glass: A note on infrared and Raman spectra. *Am Mineral* 71:772–778.

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
333. Mysen BO, Virgo D (1986) Volatiles in silicate melts at high pressure and temperature 2. Water in melts along the join $\text{NaAlO}_2\text{-SiO}_2$ and a comparison of solubility mechanisms of water and fluorine. *Chem Geol* 57:333–358.
334. Kohn SC, Dupree R, Smith ME (1989) Proton environments and hydrogen-bonding in hydrous silicate glasses from proton NMR. *Nature* 337:539–541.
335. Kohn SC, Dupree R, Smith ME (1989) A multinuclear magnetic resonance study of the structure of hydrous albite glass. *Geochim Cosmochim Acta* 53:2925–2935.
336. Davis KM, Tomozawa M (1996) An infrared spectroscopic study of water-related species in silica glasses. *J Non-Cryst Solids* 201:177–198.
337. Schmidt BC, Behrens H, Riemer T, Kappes R, Dupree R (2001) Quantitative determination of water speciation in aluminosilicate glasses: a comparative NMR and IR spectroscopic study. *Chem Geol* 174:195–208.
338. Cody GD, Mysen BO, Lee SK (2005) Structure vs. composition: A solid-state ^{29}Si NMR study of quenched glasses along the $\text{Na}_2\text{O-SiO}_2\text{-H}_2\text{O}$ join. *Geochim Cosmochim Acta* 69:2373–2384.
339. Mysen BO (2007) The solution behavior of H_2O in peralkaline aluminosilicate melts at high pressure with implications for properties of hydrous melts. *Geochim Cosmochim Acta* 71:1820–1834. doi: 10.1016/j.gca.2007.01.007
340. Malfait WJ, Xue X (2010) The nature of hydroxyl groups in aluminosilicate glasses: Quantifying Si-OH and Al-OH abundances along the $\text{SiO}_2\text{-}^{27}\text{Al-}^{29}\text{Si}$ NMR spectroscopy. *Geochim Cosmochim Acta* 74:719–737.
341. Le Losq C, Cody GD, Mysen BO (2015) Alkali influence on the water speciation and the environment of protons in silicate glasses revealed by ^1H MAS NMR spectroscopy. *Am Mineral* 100:466–473. doi: 10.2138/am-2015-5004
342. Nowak M, Behrens H (1995) The speciation of water in haplogranitic glasses and melts determined by in situ near-infrared spectroscopy. *Geochim Cosmochim Acta* 59:3445–3450.
343. Sowerby JR, Keppler H (1999) Water speciation in rhyolitic melt determined by in-situ infrared spectroscopy. *Am Mineral* 84:1843–1849.
344. Behrens H, Yamashita S (2008) Water speciation in hydrous sodium tetrasilicate and hexasilicate melts: Constraint from high temperature NIR spectroscopy. *Chem Geol* 256:306–315.
345. Chertkova N, Yamashita S (2015) In situ spectroscopic study of water speciation in the depolymerized $\text{Na}_2\text{Si}_2\text{O}_5$ melt. *Chem Geol* 409:149–156. doi: 10.1016/j.chemgeo.2015.05.012
346. Silver LA, Ihinger PD, Stolper E (1990) The influence of bulk composition on the speciation of water in silicate glasses. *Contrib Mineral Petrol* 104:142–162. doi: 10.1007/BF00306439
347. Ohlhorst S, Behrens H, Holtz F (2001) Compositional dependence of molar absorptivities of near-infrared OH- and H_2O bands in rhyolitic to basaltic glasses. *Chem Geol* 174:5–20.

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
348. Le Losq C, Moretti R, Neuville DR (2013) On the speciation and amphoteric behavior of water in aluminosilicate melts and glasses: High-Temperature Raman spectroscopy and reaction equilibria. *Eur J Mineral* 25:777–790. doi: <http://dx.doi.org/10.1127/0935-1221/2013/0025-2322>
 349. Deubener J, Müller R, Behrens H, Heide G (2003) Water and the glass transition temperature of silicate melts. *J Non-Cryst Solids* 330:268–273.
 350. Xue X, Kanzaki M (2004) Dissolution mechanisms of water in depolymerized silicate melts: Constraints from ^{29}Si NMR spectroscopy and ab initio calculations. *Geochim Cosmochim Acta* 68:5027–5057.
 351. Xue X, Kanzaki M (2006) Depolymerization effect of water in aluminosilicate glasses: Direct evidence from ^{27}Al heteronuclear correlation NMR. *Am Mineral* 91:1922–1926.
 352. Xue X, Kanzaki M (2008) Structure of hydrous aluminosilicate glasses along the diopside–anorthite join: A comprehensive one- and two-dimensional ^1H and ^{27}Al NMR study. *Geochim Cosmochim Acta* 72:2331–2348. doi: 10.1016/j.gca.2008.01.022
 353. Mysen BO, Cody G (2005) Solution mechanisms of H_2O in depolymerized peralkaline melts. *Geochim Cosmochim Acta* 69:5557–5566.
 354. Mookherjee M, Stixrude L, Karki B (2008) Hydrous silicate melt at high pressure. *Nature* 452:983–986. doi: 10.1038/nature06918
 355. Kohn SC, Dupree R, Mortuza MG (1992) The interaction between water and aluminosilicate magmas. *Chem Geol* 96:399–409.
 356. Kohn SC, Smith ME, Dirken PJ, Van Eck ERH, Kentgens APM, Dupree R (1998) Sodium environment in dry and hydrous albite glasses: Improved ^{23}Na solid state NMR data and their implications for water dissolution mechanisms. *Geochim Cosmochim Acta* 62:79–87.
 357. Whittington AG, Bouhifd MA, Richet P (2009) The viscosity of hydrous $\text{NaAlSi}_3\text{O}_8$ and granitic melts: Configurational entropy models. *Am Mineral* 94:1–16.
 358. Xue X (2009) Water speciation in hydrous silicate and aluminosilicate glasses: Direct evidence from ^{29}Si – ^{27}Al – ^1H double-resonance NMR. *Am Mineral* 94:395–398.
 359. Malfait WJ, Xue X (2014) Hydroxyl speciation in felsic magmas. *Geochim Cosmochim Acta* 140:606–620. doi: 10.1016/j.gca.2014.05.017
 360. Bouhifd MA, Whittington AG, Richet P (2015) Densities and volumes of hydrous silicate melts: New measurements and predictions. *Chem Geol* 418:40–50. doi: 10.1016/j.chemgeo.2015.01.012
 361. Mercier M, Muro AD, Métrich N, Giordano D, Belhadj O, Mandeville CW (2010) Spectroscopic analysis (FTIR, Raman) of water in mafic and intermediate glasses and glass inclusions. *Geochim Cosmochim Acta* 74:5641–5656. doi: 10.1016/j.gca.2010.06.020

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
362. Le Losq C, Neuville DR, Moretti R, Roux J (2012) Determination of water content in silicate glasses using Raman spectrometry: Implications for the study of explosive volcanism. *Am Mineral* 97:779–790. doi: 10.2138/am.2012.3831
 363. Bouhifd MA, Whittington A, Roux J, Richet P (2006) Effect of water on the heat capacity of polymerized aluminosilicate glasses and melts. *Geochim Cosmochim Acta* 70:711–722. doi: 10.1016/j.gca.2005.09.012
 364. Bouhifd MA, Whittington AG, Withers AC, Richet P (2013) Heat capacities of hydrous silicate glasses and liquids. *Chem Geol* 346:125–134. doi: 10.1016/j.chemgeo.2012.10.026
 365. Papale P, Moretti R, Barbato D (2006) The compositional dependence of the saturation surface of H₂O+CO₂ fluids in silicate melts. *Chem Geol* 229:78–95.
 366. Ni H, Keppler H (2013) Carbon in Silicate Melts. *Rev Mineral Geochem* 75:251–287. doi: 10.2138/rmg.2013.75.9
 367. Guillot B, Sator N (2011) Carbene dioxide in silicate melts: a molecular dynamics simulation study. *Geochim Cosmochim Acta* 75:1829–1857.
 368. Pedeché S, Simon P, Matzen G, Moulin B, Blanchard K, Querel G (2003) Probing gas bubbles inside industrial glasses by Raman scattering. *J Raman Spectrosc* 34:248–252. doi: 10.1002/jrs.984
 369. Morizet Y, Brooker RA, Iacono-Marziano G, Kjarsgaard BA (2013) Quantification of dissolved CO₂ in silicate glasses using micro-Raman spectroscopy. *Am Mineral* 98:1788–1802. doi: 10.2138/am.2013.4516
 370. Mysen BO, Virgo D (1980) The solubility behavior of CO₂ in melts on the join NaAlSi₃O₈-CaAl₂Si₂O₈-CO₂ at high pressures and temperatures: a Raman spectroscopic study. *Am Mineral* 65:1166–1175.
 371. Dasgupta R, Mallik A, Tsuno K, Withers AC, Hirth G, Hirschmann MM (2013) Carbon-dioxide-rich silicate melt in the Earth's upper mantle. *Nature* 493:211–215. doi: 10.1038/nature11731
 372. Brooker RA, Kohn SC, Holloway JR, McMillan PF, Carroll MR (1999) Solubility, speciation and dissolution mechanisms for CO₂ in melts on the NaAlO₂-SiO₂ join. *Geochim Cosmochim Acta* 63:3549–3565. doi: 10.1016/S0016-7037(99)00196-9
 373. Verweij H, Van Den Boom H, Breemer RE (1977) Raman scattering of carbonate ions dissolved in potassium silicate glasses. *J Am Ceram Soc* 60:529–534.
 374. Sharma SK (1979) Structure and solubility of carbon dioxide in silicate glasses of diopside and sodium melilite compositions at high pressures from Raman spectroscopic data. *Carnegie Inst Wash Year b* 78:532–537.
 375. Mysen BO, Virgo D (1980) Solubility mechanisms of carbon dioxide in silicate melts: a Raman spectroscopic study. *Am Mineral* 65:885–899.

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
376. Rai CS, Sharma SK, Muenow DW, Matson DW, Byers CD (1983) Temperature dependence of CO₂ solubility in high pressure quenched glasses of diopside composition. *Geochim Cosmochim Acta* 47:953–958.
377. Burke EAJ, Lustenhouwer WJ (1987) The application of a multichannel laser Raman microprobe (Microdil-28®) to the analysis of fluid inclusions. *Chem Geol* 61:11–17.
378. Nakano S, Moritoki M, Ohgaki K (1998) High-Pressure Phase Equilibrium and Raman Microprobe Spectroscopic Studies on the CO₂ Hydrate System. *J Chem Eng Data* 43:807–810. doi: 10.1021/je9800555
379. Brooker RA, Kohn SC, Holloway JR, McMillan PF (2001) Structural controls on the solubility of CO₂ in silicate melts: Part II: IR characteristics of carbonate groups in silicate glasses. *Chem Geol* 174:241–254. doi: 10.1016/S0009-2541(00)00318-1
380. Korschak A, Keppler H (2014) The speciation of carbon dioxide in silicate melts. *Contrib Mineral Petrol* 167:998. doi: 10.1007/s00410-014-0998-2
381. Bourgue E, Richet P (2001) The effects of dissolved CO₂ on the density and viscosity of silicate melts: a preliminary study. *Earth Planet Sci Lett* 193:57–68. doi: 10.1016/S0012-821X(01)00491-5
382. Ghosh S, Ohtani E, Litasov K, Suzuki A, Sakamaki T (2007) Stability of carbonated magmas at the base of the Earth's upper mantle. *Geophys Res Lett* 34:L22312. doi: 10.1029/2007GL031349
383. Liu Q, Lange RA (2003) New density measurements on carbonate liquids and the partial molar volume of the CaCO₃ component. *Contrib Mineral Petrol* 146:370–381. doi: 10.1007/s00410-003-0505-7
384. Kot WK, Gan H, Pegg IL Sulfur incorporation in waste glass melts of various compositions. In: *Environ. Issues Waste Manag. Technol.* pp 441–449
385. Métrich N, Mandeville CW (2010) Sulfur in Magmas. *Elements* 6:81–86. doi: 10.2113/gselements.6.2.81
386. Paris E, Giuli G, Carroll MR (2001) paris-CM-01-39-331.pdf. *Can Mineral* 39:331–339.
387. Bonnin-Mosbah M, Métrich N, Susini J, Salome M, Massare D, Menez B (2002) Micro X-ray absorption near edge structure at the sulfur and iron K-edges in natural silicate glasses. *Spectrochim Acta Part B At Spectrosc* 57:711–725.
388. McKeown D., Muller I., Gan H, Pegg I., Stolte W. (2004) Determination of sulfur environments in borosilicate waste glasses using X-ray absorption near-edge spectroscopy. *J Non-Cryst Solids* 333:74–84. doi: 10.1016/j.jnoncrystol.2003.09.035
389. Fleet ME (2005) XANES SPECTROSCOPY OF SULFUR IN EARTH MATERIALS. *Can Mineral* 43:1811–1838.
390. Wilke M, Jugo PJ, Klimm K, Susini J, Botcharnikov R, Kohn SC, Janousch M (2008) The origin of S⁴⁺ detected in silicate glasses by XANES. *Am Mineral* 93:235–240. doi: 10.2138/am.2008.2765

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
391. McKeown DA, Muller IS, Gan H, Pegg IL, Kendziora CA (2001) Raman studies of sulfur in borosilicate waste glasses: sulfate environments. *J Non-Cryst Solids* 288:191–199.
 392. Manara D, Grandjean A, Pinet O, Dussossoy JL, Neuville DR (2007) Sulfur behavior in silicate glasses and melts: Implications for sulfate incorporation in nuclear waste glasses as a function of alkali cation and V₂O₅ content. *J Non-Cryst Solids* 353:12–23. doi: 10.1016/j.jnoncrsol.2006.09.041
 393. Lenoir M, Grandjean A, Poissonnet S, Neuville DR (2009) Quantitation of sulfate solubility in borosilicate glasses using Raman spectroscopy. *J Non-Cryst Solids* 355:1468–1473. doi: 10.1016/j.jnoncrsol.2009.05.015
 394. Klimm K, Botcharnikov RE (2015) The determination of sulfate and sulfide species in hydrous silicate glasses using Raman spectroscopy. *Am Mineral* 95:1574–1579. doi: 10.2138/am.2010.3590
 395. Klimm K, Kohn SC, O'Dell LA, Botcharnikov RE, Smith ME (2012) The dissolution mechanism of sulphur in hydrous silicate melts. I: Assessment of analytical techniques in determining the sulphur speciation in iron-free to iron-poor glasses. *Chem Geol* 322–323:237–249. doi: 10.1016/j.chemgeo.2012.04.027
 396. Baker DR, Moretti R (2011) Modeling the Solubility of Sulfur in Magmas: A 50-Year Old Geochemical Challenge. *Rev Mineral Geochem* 73:167–213. doi: 10.2138/rmg.2011.73.7
 397. O'Neill HSC, Mavrogenes JA (2002) The Sulfide Capacity and the Sulfur Content at Sulfide Saturation of Silicate Melts at 1400°C and 1 bar. *J Petrol* 43:1049–1087. doi: 10.1093/petrology/43.6.1049
 398. Liu Y, Samaha N-T, Baker DR (2007) Sulfur concentration at sulfide saturation (SCSS) in magmatic silicate melts. *Geochim Cosmochim Acta* 71:1783–1799. doi: 10.1016/j.gca.2007.01.004
 399. Wykes JL, O'Neill HSC, Mavrogenes JA (2015) The Effect of FeO on the Sulfur Content at Sulfide Saturation (SCSS) and the Selenium Content at Selenide Saturation of Silicate Melts. *J Petrol* 56:1407–1424. doi: 10.1093/petrology/egv041
 400. Moretti R, Papale P, Ottonello G (2003) A model for the saturation of C-O-H-S fluids in silicate melts. *Geol Soc Lond Spec Publ* 213:81–101. doi: 10.1144/GSL.SP.2003.213.01.06
 401. Moretti R, Ottonello G (2005) Solubility and speciation of sulfur in silicate melts: The Conjugated Toop-Samis-Flood-Grjotheim (CTSFG) model. *Geochim Cosmochim Acta* 69:801–823. doi: 10.1016/j.gca.2004.09.006
 402. Aiuppa A, Baker DR, Webster JD (2009) Halogens in volcanic systems. *Chem Geol* 263:1–18. doi: 10.1016/j.chemgeo.2008.10.005
 403. Butov OV, Golant KM, Tomashuk AL, van Stralen MJN, Breuls AHE (2002) Refractive index dispersion of doped silica for fiber optics. *Opt Commun* 213:301–308. doi: 10.1016/S0030-4018(02)02087-4

- Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.
404. Kakiuchida H, Sekiya EH, Shimodaira N, Saito K, Ikushima AJ (2007) Refractive index and density changes in silica glass by halogen doping. *J Non-Cryst Solids* 353:568–572. doi: 10.1016/j.jnoncrysol.2006.10.025
 405. Zeng Q, Stebbins JF (2000) Fluoride sites in aluminosilicate glasses: High-resolution ^{19}F NMR results. *Am Mineral* 85:863–867.
 406. Mysen BO, Cody GD, Smith A (2004) Solubility mechanisms of fluorine in peralkaline and meta-aluminous silicate glasses and in melts to magmatic temperatures¹. *Geochim Cosmochim Acta* 68:2745–2769. doi: 10.1016/j.gca.2003.12.015
 407. Kiczinski TJ, Du L-S, Stebbins JF (2004) F-19 NMR study of the ordering of high field strength cations at fluoride sites in silicate and aluminosilicate glasses. *J Non-Cryst Solids* 337:142–149. doi: 10.1016/j.jnoncrysol.2004.03.123
 408. Kiczinski TJ, Stebbins JF (2006) The Effect of Fictive Temperature on the Structural Environment of Fluorine in Silicate and Aluminosilicate Glasses. *J Am Ceram Soc* 89:57–64. doi: 10.1111/j.1551-2916.2005.00677.x
 409. Baasner A, Schmidt BC, Dupree R, Webb SL (2014) Fluorine speciation as a function of composition in peralkaline and peraluminous $\text{Na}_2\text{O}-\text{CaO}-\text{Al}_2\text{O}_3-\text{SiO}_2$ glasses: A multinuclear NMR study. *Geochim Cosmochim Acta* 132:151–169. doi: 10.1016/j.gca.2014.01.041
 410. Dalou C, Le Losq C, Mysen BO, Cody GD (2015) Solubility and solution mechanisms of chlorine and fluorine in aluminosilicate melts at high pressure and high temperature. *Am Mineral* 100:2272–2283. doi: 10.2138/am-2015-5201
 411. Baasner A, Schmidt BC, Webb SL (2013) Compositional dependence of the rheology of halogen (F, Cl) bearing aluminosilicate melts. *Chem Geol* 346:172–183. doi: 10.1016/j.chemgeo.2012.09.020
 412. Stebbins JF, Du L-S (2002) Letters. Chloride ion sites in silicate and aluminosilicate glasses: A preliminary study by ^{35}Cl solid-state NMR. *Am Mineral* 87:359–363.
 413. Sandland TO, Du L-S, Stebbins JF, Webster JD (2004) Structure of Cl-containing silicate and aluminosilicate glasses: A ^{35}Cl MAS-NMR study. *Geochim Cosmochim Acta* 68:5059–5069. doi: 10.1016/j.gca.2004.07.017
 414. Zimova M (2006) The effect of chlorine on the viscosity of $\text{Na}_2\text{O}-\text{Fe}_2\text{O}_3-\text{Al}_2\text{O}_3-\text{SiO}_2$ melts. *Am Mineral* 91:344–352. doi: 10.2138/am.2006.1799
 415. Evans KA, Mavrogenes JA, O'Neill HS, Keller NS, Jang L-Y (2008) A preliminary investigation of chlorine XANES in silicate glasses. *Geochem Geophys Geosystems* 9:Q10003. doi: 10.1029/2008GC002157
 416. Baasner A, Hung I, Kemp TF, Dupree R, Schmidt BC, Webb SL (2014) Constraints on the incorporation mechanism of chlorine in peralkaline and peraluminous $\text{Na}_2\text{O}-\text{CaO}-\text{Al}_2\text{O}_3-\text{SiO}_2$ glasses. *Am Mineral* 99:1713–1723. doi: 10.2138/am.2014.4717

Le Losq, Cicconi, Greaves, Neuville. Silicate Glasses. In: Springer Handbook of Glass.