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Bayesian joint muographic and gravimetric inversion applied to volcanoes

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SUMMARY
Gravimetry is a technique widely used to image the structure of the Earth. However, inversions are ill-posed and the imaging power of the technique rapidly decreases with depth. To overcome this limitation, muography, a new imaging technique relying on high energy atmospheric muons, has recently been developed. Because muography only provides integrated densities above the detector from a limited number of observation points, inversions are also ill-posed. Previous studies have shown that joint muographic and gravimetric inversions better reconstruct the 3-D density structure of volcanic edifices than independent density inversions. These studies address the ill-posedness of the joint problem by regularizing the solution with respect to a priori density model. However, the obtained solutions depend on some hyperparameters, which are either determined relative to a single test case or rely on ad-hoc parameters. This can lead to inaccurate retrieved models, sometimes associated with artefacts linked to the muon data acquisition. In this study, we use a synthetic example based on the Puy de Dôme volcano to determine a robust method to obtain the resulting model closest to the synthetic model and devoid of acquisition artefacts. We choose a Bayesian approach to include an a priori density model and a smoothing by a Gaussian spatial correlation function relying on two hyperparameters: an a priori density standard deviation and an isotropic spatial correlation length. This approach has the advantage to provide a posteriori standard deviations on the resulting densities. Using our synthetic volcano, we investigate the most reliable criterion to determine the hyperparameters. Our results suggest that k-fold Cross-Validation Sum of Squares and the Leave One Out methods are more robust criteria than the classically used L-curves. The determined hyperparameters allow to overcome the artefacts linked to the data acquisition geometry, even when only a limited number of muon telescopes is available. We also illustrate the behaviour of the inversion in case of offsets in the a priori density or in the data and show that they lead to recognizable structures that help identify them.

Key words: Gravity anomalies and Earth structure; Joint inversion; Tomography.

1 INTRODUCTION
The 3-D density structure of volcanic edifices is classically inferred from the inversion of gravimetric data (e.g. Camacho et al. 1997; Cella et al. 2007; Linde et al. 2014). Gravimetry provides measurements of the gravity field at multiple locations throughout the study area, corresponding to the integrated effect of the whole Earth and sensitive to the local density variations. The inversion of gravimetric data is a non-unique process, requiring strong a priori geological information to constrain the models or a combination with other geophysical data such as seismic traveltimes (e.g. Onizawa et al. 2002; Coutant et al. 2012).

Muography is a method that emerged from the field of particle physics. Using atmospheric muons, which result from the interaction of cosmic rays with the atmosphere, muography provides 2-D images of integrated densities. Muons are charged leptons, alike
to electrons, but ~200 times heavier. They interact with matter through various stochastic processes, depending on their energies and on the medium composition (e.g. Groom et al. 2001; Nagamine 2003). The higher the muon energy, the farther the muon travels, but also the lower the muon flux. High energy muons can penetrate up to kilometres of rocks at TeV energies. Furthermore, due to their strong relativistic boost, they travel along straight (ballistic) paths. Muon telescopes, or muon detectors, are used to detect them and reconstruct their trajectories after the muons have crossed their target (Nagamine et al. 1995). The backward extrapolation of the muon trajectory up to the target is referred to as the line of sight. The rate of muons crossing the target along this direction depends, to first order, on the subsurface density integrated along the line of sight. Rock composition can modify this rate by up to 10–15 per cent (Lechmann et al. 1995) published the first muographic measurements that were used to probe the inner structure of a volcano, from experiments on Mount Tsukuba and Mount Asama in Japan. In the last two decades, muography has been developed for the density imaging (e.g. Tanaka et al. 2001; Lesparre et al. 2012; Kusagaya & Tanaka 2013) and monitoring (e.g. Tanaka et al. 2014; Jourde et al. 2015b) of volcanoes. Muon tomography, consisting of combining muographies from several viewpoints, allows the reconstruction the 3-D density distribution. To conduct a muon tomography, Nagahara & Miyamoto (2018) used the method of filtered back projection, which has the advantage of not relying on any a priori information. They show that the method requires data from at least a dozen of viewpoints to allow a proper 3-D reconstruction, which is for now impractical. Indeed, the main limitations of muon tomography are the number of available telescopes, the acquisition duration and the ability to only image densities above the horizontal plane passing through the telescope. Because of these limitations, it is helpful to combine muography with other types of data.

Both gravimetry and muography being independent and complementary methods sensitive to the subsurface densities, they have the potential to help each other through a joint inversion scheme to better constrain the 3-D density models and perform a density tomography. Both types of observations can be linearly related to densities, but the inverse problem is ill-posed, requiring additional constraints, such as a regularity of the solution relative to a priori density model. Whether the problem searched is formulated in a Bayesian framework or as a misfit minimization problem, there is a need to determine some a priori parameters tuning the regularization, further referred to as hyperparameters. So far, the determination of the hyperparameters in joint inversions of muography and gravimetric data has not been fully addressed. Nishiyama et al. (2014b) use a Bayesian approach and determine the a priori density standard deviation and correlation length by maximizing the similarity of the reconstructed densities with a given checkerboard model. Because the determined a priori values depend on the density contrast and the length of the checkerboard cells, there is no guarantee that the determined parameters apply to in-situ reconstructions. Nishiyama et al. (2017) choose the a priori density mean and standard deviation according to the muographic observations and arbitrarily fix the correlation length. Jourde et al. (2015a) and Rosas-Carbajal et al. (2017) use a non-Bayesian framework, regularizing the inversion with a prior model, weighted to counteract the physical decrease of sensitivity away from the data. Rosas-Carbajal et al. (2017) invert for a density model as well as a constant possible offset between the density inferred by muographic data and the density inferred by gravimetric data. The authors determine the weight of the smoothing by using a classical L-curve scheme in which the best compromise between the data fit and the weight of the prior model is obtained. Noteworthy, the inversion results show artefacts related to the muographic acquisition geometry: resulting densities tend to smear out along the telescopes’ lines of sight. The determination of the regularization parameters should be further studied. Indeed it is recognized as a key issue in many geophysical problems with similar formulations such as the determination of slip distribution of faults (Fukuda & Johnson 2008), or the joint determination of density and seismic slowness (Coutant et al. 2012). Comparing solutions determined using L-curves, Akaike Bayesian information criteria, cross-validation and fully Bayesian criteria, Fukuda & Johnson (2008) have shown that there is no rational way of selecting a single value from the L-curve, and that the cross-validation and the fully Bayesian criteria provide more reliable results.

In this paper, we present a robust workflow to linearly invert muographic and gravimetric data using a Bayesian framework (Taranota & Valette 1982), suitable even when a limited number of muon telescopes is available. A realistic synthetic model adopting the topography of the Puy de Dôme volcano in the Chaîne des Puys in France and the acquisition geometry of gravimetric and muographic campaigns conducted on this volcano is used (Fig. 1). We use a homogeneous a priori density model and impose smoothing using a Gaussian spatial correlation function. We systematically explore the two inversion hyperparameters that tune the result of the inversion, that is, the density a priori standard deviation and the spatial correlation length, and we discuss the use of L-curves, Leave One Out (LOO) criterion and k-fold Cross-Validation Sum of Squares (CVSS; e.g. Walha 1990; Augier 2011; James et al. 2013) to determine the optimal set of hyperparameters in a robust way and with limited artefacts related to the data acquisition. A realistic synthetic model adopting the topography of the Puy de Dôme volcano in the Chaîne des Puys in France and the acquisition geometry of gravimetric (Portal et al. 2016) and muographic (Ambrosino et al. 2015) campaigns conducted on this volcano is used (Fig. 1). We perform the following inversions: (i) independent inversion of gravimetric data, (ii) independent inversion of muographic data from a single viewpoint, (iii) joint inversion of gravimetric data and muographic data from a single viewpoint, (iv) independent inversion of muographic data from three viewpoints and (v) joint inversion of gravimetric data and muographic data from three viewpoints. We show the potential of the Bayesian framework for the inversion of muographic data. By comparing joint inversion results to independent gravimetric inversion results, we also assess the contribution of the muographic data to the recovery of the true density distribution. We illustrate the behaviour of the inversions for both one and three muographic viewpoints. Finally, we test the influence of an offset between the averaged density inferred from muography and the averaged density inferred from gravimetry.

2 Method

We describe the medium using a 3-D cubic mesh of equally spaced nodes of densities \( p \). The density at any point with coordinates \((x, y, z)\) in the medium is obtained by trilinear interpolation of the densities of the eight surrounding nodes so that the density varies continuously in the medium.
2.1 Forward modelling

2.1.1 Gravimetry

The gravimetric anomaly $g$ produced by a volume $V$ of density $\rho(x, y, z)$ at a given data location $(x_0, y_0, z_0)$ is given by

$$g(x_0, y_0, z_0) = G \int \int \int_V \frac{\rho(x, y, z)(z - z_0)}{\sqrt{(x - x_0)^2 + (y - y_0)^2 + (z - z_0)^2}} \, dx \, dy \, dz,$$

where $G = 6.67 \times 10^{-11} \text{m}^3 \text{kg}^{-1} \text{s}^{-2}$ is the universal gravitational constant. The forward gravimetric computation is made using the method and the associated code of Coutant et al. (2012) described in Barnoud et al. (2016). The integration over the volume is made numerically over $x$ and $y$ and analytically over $z$. The topography is taken into account with the same resolution as the input digital terrain model in the vicinity of the gravimetric stations and is averaged with a larger step farther away.

As $\rho(x, y, z)$ can be expressed as a linear combination of densities at nodes $\rho$, the modelled gravimetric data $g$ are also linearly related to the densities at nodes $\rho$:

$$G \rho = g$$

via the sensitivity matrix $G$ of elements $G_{ij}$ that contain the contribution of each node $j$ to each gravimetric data $i$ (Fig. 2a).

2.1.2 Muography

Muography provides estimates of averaged densities along lines of sight. These density estimates are retrieved from the flux of muons crossing the medium and observed in conic bins of given widths of azimuth and elevation, assuming that the estimates are uniform averages of the densities over the solid angles. To take into account the sensitivity to all the density nodes in a given cone, we integrate numerically in azimuth and elevation over the cone using a beam of lines of sight along which the density is integrated. The averaged density $\varrho$ as seen from a muon detector located in $(x_0, y_0, z_0)$ and in a cone with the azimuths and elevations centred on $(\alpha_0, \beta_0)$ is expressed as

$$\varrho(x_0, y_0, z_0, \alpha_0, \beta_0) = \frac{\sum_{\alpha} \sum_{\beta} \int \rho(\alpha, \beta, r) \, dr}{\sum_{\alpha} \sum_{\beta} \int dr},$$

where $\alpha$ and $\beta$ are the azimuths and elevations of the beam lines of sight and $r$ is the length travelled in the medium. The integration along each line of sight can be performed either numerically or analytically.

With the density in the mesh represented as a linear combination of densities at nodes $\rho$, the averaged densities $\varrho$ inferred from muography can also be written as a linear combination of the densities at the nodes, similar to the gravimetric case

$$M \rho = \varrho$$

via the sensitivity matrix $M$ whose elements $M_{ij}$ contain the contribution of each node of density $j$ to the averaged density in each cone $i$. The sensitivity of a muographic measurement to a node of density decreases with the distance of the node to the muon detector (Fig. 2b). $M$ is a sparse matrix, in contrast to the gravimetric sensitivity matrix $G$ which is full.

Figure 1. Shaded topography map of the Puy de Dôme volcano with the location of the gravimetric data (black dots) and of the muographic viewpoints CDC, TDF and BDR used in this study. The grey dotted line shows the location of the section used in Fig. 2.
The normal distribution of the data errors $\epsilon$ malism, following Tarantola (2005), which has the advantage of where $\epsilon_d$ Setting the vector of observed data where $\sigma^2_\rho$ includes the standard deviation of density nodes. $G_i$ is the $i$th line of the sensitivity matrix $G$, corresponding to the $i$th datum. $M_i$ is the $i$th line of the sensitivity matrix $M$, corresponding to the $i$th datum. The sensitivities $S_{\text{grav}}$ and $S_{\text{muog}}$ of all data are computed by summing the lines of the sensitivity matrices, following eqs (16) and (17), respectively.

2.2 Inversion method

Both the gravimetric and the muographic problems being linear, with respect to the same physical property of rocks, in this case density $\rho$, the joint problem is expressed as

$$[G \ M] \rho = g.$$  \hspace{1cm} (5)

Setting the vector of observed data $d = [g_{\text{obs}} \ e_{\text{obs}}]^T$ and the sensitivity matrix with the contribution of each node of density to each data $A = [G \ M]^T$, the inverse problem to solve is

$$d = A \rho + \epsilon$$ \hspace{1cm} (6)

where $\epsilon$ accounts for measurement and modelling errors.

To solve the inverse problem in eq. (6), we use a Bayesian formalism, following Tarantola (2005), which has the advantage of taking into account the errors on the data and to easily include an a priori model to regularize the inversion. We assume a zero-mean normal distribution of the data errors $\epsilon$ with a covariance matrix $C_D$, $\mathcal{N}(0, C_D)$, and we write the likelihood as

$$p(d|\rho) \propto \exp \left( -\frac{1}{2} (d - A \rho)^T C_D^{-1} (d - A \rho) \right),$$  \hspace{1cm} (7)

where $C_D$, the data covariance matrix, contains the data variances $\sigma^2_{\epsilon_j}$ on its diagonal and optionally covariances between data on its off-diagonal terms. This might be the case for muographic data depending on the processing applied to obtain the averaged densities. In this paper, we work with synthetic data sets and only consider independent data, leading to a diagonal data covariance matrix. We also assume an a priori normal probability density distribution for the densities at nodes, $\mathcal{N}(\rho_{\text{prior}}, C_P)$, with an average density $\rho_{\text{prior}}$ and a covariance matrix $C_P$

$$p(\rho) \propto \exp \left( -\frac{1}{2} (\rho - \rho_{\text{prior}})^T C_P^{-1} (\rho - \rho_{\text{prior}}) \right).$$  \hspace{1cm} (8)

The a priori density covariance matrix $C_P$ is a full matrix that includes the standard deviation $\sigma_\rho$ on the a priori densities $\rho_{\text{prior}}$ and a Gaussian spatial correlation function

$$c_r(d) = \exp \left( -\frac{d^2}{\lambda^2} \right)$$ \hspace{1cm} (9)

where $d$ is the distance between two nodes of the model and $\lambda$ is the spatial correlation length that controls the smoothness of the density model.

Bayes theorem states that the a posteriori probability distribution of the model densities is

$$p(\rho|d) = \frac{p(d|\rho)p(\rho)}{p(d)}$$ \hspace{1cm} (10)

where $p(d)$ is the marginal likelihood that does not depend on $\rho$. Hence, the a posteriori probability density function of densities is also normal and such that

$$p(\rho|d) \propto \exp \left( -\frac{1}{2} \phi(\rho) \right)$$ \hspace{1cm} (11)

with the objective function $\phi$ given by

$$\phi(\rho) = (d - A \rho)^T C_D^{-1} (d - A \rho) + (\rho - \rho_{\text{prior}})^T C_P^{-1} (\rho - \rho_{\text{prior}})$$

$$= ||d - A \rho||^2_D + ||\rho - \rho_{\text{prior}}||^2_P.$$  \hspace{1cm} (12)

The first term represents the fit of the data. In the following, we will note $\chi^2 = \frac{1}{2} ||d - A \rho||^2_D$ the fit normalized by the number of data $n_d$, and we will refer to $\chi^2$ as the data misfit. The second term accounts for the proximity of the model to the prior $\rho_{\text{prior}}$, and the correlation between nearby parameters. The second term will be further referred to as the model regularization.

Maximizing the posterior probability density function is equivalent to minimizing the objective function $\phi(\rho)$ of eq. (12) and leads to a unique solution that can be written (Tarantola 2005)

$$\bar{\rho} = \rho_{\text{prior}} + C_P A^T (A C_P A^T + C_D)^{-1} (d - A \rho_{\text{prior}}).$$  \hspace{1cm} (13)

The estimated density $\bar{\rho}$ is the centre of the a posteriori density distribution $p(\rho|d)$, with the associated a posteriori density covariance matrix (Tarantola 2005)

$$\bar{C}_P = C_P - C_P A^T (A C_P A^T + C_D)^{-1} A C_P.$$

Figure 2. Sensitivities of the gravimetric and muographic data of the Puy de Dôme. The west–east cross-section is located on Fig. 1. (a) Normalized sensitivity of a single gravimetric measurement. Black triangle: gravity station location. (b) Normalized sensitivity of a single muographic measurement. Black square: location of a muographic measurement looking in a $3^\circ \times 3^\circ$ cone. (c) Normalized cumulative sensitivity of all gravimetric data to density nodes. (d) Normalized cumulative sensitivity of all muographic data to density nodes. $G_i$ is the $i$th line of the sensitivity matrix $G$, corresponding to the $i$th datum. $M_i$ is the $i$th line of the sensitivity matrix $M$, corresponding to the $i$th datum. The sensitivities $S_{\text{grav}}$ and $S_{\text{muog}}$ of all data are computed by summing the lines of the sensitivity matrices, following eqs (16) and (17), respectively.
The diagonal terms of $\tilde{C}_p$ are the variances $\tilde{\sigma}_p^2$ of the estimated densities at nodes $\tilde{\rho}$.

The solution of the inversion, eqs (13) and (14), is tuned by the two hyperparameters included in the a priori covariance matrix $C_p$: the standard deviation $\sigma_p$ of the a priori density distribution and the spatial correlation length $\lambda$. To estimate the optimal set of hyperparameters, $(\sigma_p, \lambda)$, we explore them in a systematic way and compare different criteria. On the one hand, we use classical L-curves, where the optimal set of $(\sigma_p, \lambda)$ usually corresponds to the corner of the curve representing the model regularization versus the data misfit, indicating that the best balance between the data fit and the model regularity is achieved. Typically, the larger $\lambda$ or the smaller $\sigma_p$, the smoother the solution, but the larger the data misfit. Following Harris & Segall (1987), we also use a modified L-curve, where we assume that the optimal $(\sigma_p, \lambda)$ minimizes the a posteriori density standard deviation as well as the data misfit. On the other hand, we use k-fold CVSS, including the LOO case (e.g. Wahba 1990; Augier 2011; James et al. 2013). The CVSS method is based on the assumption that appropriate values of the hyperparameters lead to a resulting model that is able to accurately predict new data. The k-fold CVSS consists in separating the data in $k$ subsets and performing $k$ inversions excluding one of the subsets each time. The excluded subset is modelled using the inversion result. To assess how well we recover the initial synthetic model, the Mean Square Error (MSE) between modelled and observed data of this subset is evaluated. The value to minimize is then the sum or the average of the $k$ successive MSEs. Here, we use a weighted MSE that includes a normalization by the data errors, to account for the different physical quantities involved in the joint inversion, that is, densities for the muographic data and accelerations for the gravimetric data. The quantity to minimize is therefore (James et al. 2013)

$$C_k(\sigma_p, \lambda) = \frac{1}{k} \sum_{i=1}^{k} \sum_{l=1}^{k} ||d - A_{i+l}||^2_D,$$

where $n_k$ is the total number of gravity and muography observed data in each data subset $D_i$. We compute $C_k(\sigma_p, \lambda)$ for $k$ equal to a quarter of the data ($k = 4$) and to all but one data ($k = n$), the latter case corresponding to the so-called LOO solution. The data are randomly distributed between the subsets.

### 3 Application to a Realistic Synthetic Case

#### 3.1 Setup

Here, we apply the inverse scheme to a realistic synthetic case in which we use the topography and acquisition geometry of recent gravimetry and muography campaigns at the Puy de Dôme volcano. The gravimetric data (Fig. 1) consist of 650 points (Portal et al. 2016). The synthetic gravimetric data correspond to a free-air anomaly, without any regional field component. Three muography viewpoints are used. Two of the muographic viewpoints (CDC: Col de Ceyssat and TDF: telecommunications room) are actual measuring stations from preliminary campaigns on the Puy de Dôme volcano (Ambrosino et al. 2015) while the third one (BDR: Bois de Rochetoux) is a fictitious viewpoint added for the purposes of this study. The muographic synthetic data consist in 2067 averaged densities in cones of 1° x 1°.

The volume is discretized using a mesh of $n_v = 209$ 525 density nodes with a 25 m spacing in the three dimensions, only 107

![Figure 3](https://academic.oup.com/gji/article-abstract/218/3/2179/5530762)

**Figure 3.** Synthetic density model $\rho_{true}$ used to compute the synthetic gravimetric and muographic data. The densities are constructed using a Gaussian random field with a standard deviation $\sigma_{true} = 100$ kg m$^{-3}$ and a Gaussian spatial covariance with a correlation length $\lambda_{true} = 200$ m. 164 nodes with non-zero sensitivities to be taken into account for the inversion. This implies that the covariance matrix $C_p$ used for the model regularization would account for $\sim 85$ Go of memory.

The order of operations has to be handled appropriately to reduce memory needs. Based on eq. (13), the largest matrix we compute and store is the matrix product $C_p A^T$ (Barnoud et al. 2016), only accounting for $\sim 2$ Go of memory.

The sensitivity of a gravimetric datum to the density nodes is given by the $i$th line $G_{i}$ of the absolute values of the sensitivity matrices $G$ (Fig. 2a). Similarly, the sensitivity of a muographic datum to the density nodes is given by the $i$th line $M_i$ of the absolute values of the sensitivity matrices $M$ (Fig. 2b). In Figs 2(c) and (d), we show the cumulative sensitivity of each node to all data. For the $n_{grav}$ gravimetric data, it is expressed as

$$S_{grav} = \sum_{i=1}^{n_{grav}} G_{i},$$

and for the $n_{muog}$ muographic data, it is expressed as

$$S_{muog} = \sum_{i=1}^{n_{muog}} M_i.$$
length $\lambda_{\text{true}} = 200 \text{ m}$, as indicated in the covariance matrix $C_{\rho_{\text{true}}}$. In the Appendix, we recall that a synthetic density model with the required statistical properties can be constructed using

$$\rho_{\text{true}} = \rho_{\text{true}} + LX,$$

(18)

where $X \sim \mathcal{N}(0, I)$ is a random vector drawn from a zero-mean normal distribution with an identity covariance matrix and $L$ is the Cholesky decomposition of the covariance matrix $C_{\rho_{\text{true}}} = LL^T$ with a Gaussian correlation function of standard deviation $\sigma_{\rho_{\text{true}}}$ and of correlation length $\lambda_{\text{true}}$. For computational reasons, the random realization is generated on a subsampled mesh with a node spacing of 100 m and a spline interpolation is used to obtain the model over the 25 m mesh.

We generate synthetic data to which we add Gaussian noise with realistic standard deviations (Fig. 4). For the gravimetric data, we use a standard deviation of 0.1 mGal, which has a comparable order of magnitude as the errors from the gravimetric measurements (Portal et al. 2016). For the muographic data, we use a standard deviation of 100 kg m$^{-3}$, corresponding to the order of magnitude of errors estimated for acquisition campaigns on the Puy de Dôme with an exposure of about 50 d m$^2$ (Cărloganu et al. 2016; Cărloganu & the TOMUVOL collaboration 2018; Niess et al. 2018b) and determined for similar campaigns on other volcanoes as well (e.g. Lesparre et al. 2012; Nishiyama et al. 2014a, 2016; Oláh et al. 2018). We assume that the data errors are accurately estimated when dealing with real data and use these values as data standard deviations in the data covariance matrix $C_D$.

3.3 Inversion of the synthetic data

3.3.1 Determination of the inversion hyperparameters

We perform the independent and joint inversions for 128 sets of hyperparameters $(\sigma_\rho, \lambda)$, with an a priori density standard deviation $\sigma_\rho$ ranging from 5 to 400 kg m$^{-3}$ and a spatial correlation length $\lambda$ ranging from 50 to 800 m. In order to identify the most suitable criterion, the L-curve and CVSS criteria are compared for the independent and joint inversions of the gravimetric and the muographic data from the three viewpoints. Comparison of the density model from the joint inversion with the true density model (Fig. 5) shows that the closest model in terms of root mean square error (RMSE) and mean absolute error (MAE) is obtained for hyperparameters $(\sigma_\rho, \lambda) = (100 \text{ kg m}^{-3}, 200 \text{ m})$, corresponding to the parameters used to construct the synthetic density model (Fig. 3). We use this set of hyperparameters as reference in the remainder of this paper, indicated with a black dot in Figs 5–7. Note that the RMSE and MAE cannot be used as criteria to determine the hyperparameters for real data inversion as the true densities are not known.

L-curves of the model regularization versus the data misfit lead to $(\sigma_\rho, \lambda) = (400 \text{ kg m}^{-3}, 50 \text{ m})$ for three configurations of inversions, that is, the largest tested $\sigma_\rho$ and the lowest tested $\lambda$ (Fig. 6, top). The associated density models appear underconstrained: they are not smooth enough leading to density anomalies spread out along the muographic lines of sight and small areas of high density intensity anomalies located right beneath the gravimetric data. As the L-curves of the model regularization versus the data misfit do not show typical L-shapes and are not conclusive to select the hyperparameters, we also plot L-curves of the a posteriori density standard deviation $\tilde{\sigma}_\rho$ (averaging the variances at all nodes) versus the data misfit (Fig. 6, bottom) and the lowest tested $\lambda$. Similarly to the L-curves used by Harris & Segall (1987) for instance. This representation leads to $(\sigma_\rho, \lambda) = (50 \text{ kg m}^{-3}, 150 \text{ m})$ for the gravimetric inversion, $(150 \text{ kg m}^{-3}, 50 \text{ m})$ for the muographic inversion and $(100 \text{ kg m}^{-3}, 100 \text{ m})$ for the joint inversion. In this case, the corresponding density model is much smoother for the gravimetric inversion, but not for the muographic inversion which is still badly constrained, with anomalies spread out along the muographic lines of sight. For the joint inversion, the resulting model is better constrained than with the previous L-curve, but not as smooth as the true density model (Fig. 3).

Figure 4. Synthetic data computed from the synthetic density model shown in Fig. 3. (a) Gravimetric data on the shaded topographic map. A Gaussian noise with a standard deviation of 0.1 mGal has been added to the gravimetric data. Black squares: location of the muon detectors. The dotted square shows the limits of the inversion volume. The solid black lines show the locations of cross-sections AA’ and BB’ shown in this paper. (b) Muographic data from CDC, TDF and BDR viewpoints. A Gaussian noise with a standard deviation of 0.1 mGal has been added to the gravimetric data. Black squares: location of the muon detectors. The dotted square shows the limits of the inversion volume. The solid black lines show the locations of cross-sections AA’ and BB’ shown in this paper.
Fig. 7 shows maps of the LOO criterion \( C_n(\sigma_\rho, \lambda) \) and of the fourfold CVSS criterion \( C_4(\sigma_\rho, \lambda) \) computed by eq. (15). The optimal density models are obtained for the hyperparameters \( \sigma_\rho \) and \( \lambda \) that minimize the functions \( C_n(\sigma_\rho, \lambda) \) and \( C_4(\sigma_\rho, \lambda) \) of eq. (15). Both the LOO and the fourfold CVSS criteria give hyperparameters \( \sigma_\rho, \lambda \) (white dots) that slightly overestimate the reference value of (100 kg m\(^{-3}\), 200 m) (black dots) by at most 100 kg m\(^{-3}\) and 50 m. The fourfold CVSS leads to values closer to the reference values than the LOO. In the three configurations of inversion, the LOO and CVSS maps have areas of minimum \( C_n \) and \( C_4 \) that are elongated along the \( \sigma_\rho \) axis, indicating that the spatial correlation length \( \lambda \) is better constrained than the \( \sigma_\rho \) a posteriori density standard deviation \( \sigma_\rho \) (Fig. 7). However, we observe that the density models associated with low LOO or CVSS criteria (areas in violet on hyperparameters maps of Fig. 7) are very similar to each other. Indeed, even when the LOO and fourfold CVSS criteria do not give the exact same hyperparameter values, the corresponding models are indistinguishable (Fig. 7) and acceptably recover the amplitudes and shapes of the synthetic density model (Fig. 3). This observation is supported by the RMSE and MAE maps that have similar shapes as the LOO and CVSS maps (Fig. 5), suggesting that the corresponding resulting models are expected to be similar to each other (Fig. 5). Besides, we no longer observe artefacts linked to the data acquisition geometry in the resulting density models, compared to the models obtained using the L-curves (Fig. 6). The LOO and CVSS criteria are therefore robust criteria and helpful tools to determine optimal values for hyperparameters.

3.3.2 Inversion results

In order to allow the comparison between the joint and independent inversions, we retain the results obtained using the reference hyperparameter values \( (\sigma_\rho, \lambda) = (100 \text{ kg m}^{-3}, 200 \text{ m}) \), keeping in mind that density models obtained with hyperparameters inferred from LOO or CVSS criteria are very close. We show the inversion results for the five following configurations of inversion (Figs 8 and 9): the independent inversion of the gravimetric data (Figs 8b and 9b); the independent inversion of the muographic data from the CDC viewpoint (Figs 8c and 9c); the joint inversion of the gravimetric data and the muographic data from the CDC viewpoint (Figs 8d and 9d); the independent inversion of the muographic data from CDC, TDF and BDR viewpoints (Figs 8e and 9e); and the joint inversion of the gravimetric data and the muographic data from the three viewpoints (Figs 8f and 9f), see Fig. 1 for the locations. Horizontal and vertical cross-sections extracted from the 3-D resulting models are shown in terms of central density \( \overline{\rho} \) and standard deviation \( \sigma_\rho \) of the \( a \ posteriori \) density distribution (Figs 8 and 9, top and middle, respectively), as well as a random realization of densities drawn within this \( a \ posteriori \) distribution (Figs 8 and 9, bottom). The density models randomly drawn from the \( a \ posteriori \) distribution allow for a better visualization of the resolution of the model because well resolved areas show some spatial correlation while unresolved areas display randomly distributed densities with no spatial correlation. To quantify the differences between the results of the five inversion configurations, we compute several estimators: the RMSE and the MAE between the inverted models and the synthetic density model, the mean of the \( a \ posteriori \) density standard deviation at all nodes, and the data misfits (Table 1).

The independent inversion of the gravimetric data alone allows retrieving lateral density variations with the lowest \( a \ posteriori \) standard deviation in the upper 200 m below the topographic surface (Fig. 9b). The standard deviation on the \( a \ posteriori \) density (Fig. 9b, middle) as well as the model constructed by randomly sampling the \( a \ posteriori \) density distribution according to their standard deviation (Fig. 9b, bottom) show that the gravimetric inversion suitably reproduces the anomalies at shallow depth but has difficulties retrieving the shapes of the structures at depth.

The independent inversion of the muographic data from the CDC viewpoint alone spreads out the density variations along the lines of sight along which the muography is blind (Fig. 8c). The \( a \ posteriori \) density standard deviation shows that the densities far away from the detector are not constrained (Fig. 8c, middle). The retrieved density model is quantitatively less good than the one obtained with the independent gravimetric inversion, the RMSE and MAE with respect to the synthetic density model increasing by 49.5 and 52.8 per cent respectively (Table 1).

The result of the joint inversion of the gravimetric and the muographic data from CDC viewpoint alone (Figs 8d and 9d) shows the complementarity of the two types of data. The resolution at depth is improved and the limit of low \( a \ posteriori \) standard deviation at depth is extended from ~200 m to ~300 m below the summit. In this case, muography brings resolution at depth in the central part of the dome and helps to better localize the anomalies in the core of the dome (compare Fig. 9d middle with Figs 9b middle and 9c middle), while gravimetry constrains the parts of the model away from the muon detector and the lateral variations (Fig. 8d). Quantitatively
Figure 6. L-curve criteria to select the inversion hyperparameters $\sigma_\rho$ and $\lambda$, along with the associated optimal models (horizontal cross-sections extracted at 1100 m of altitude). The dot colour indicates the prior value $\sigma_\rho$, while the dot size is proportional to the correlation length $\lambda$. The white dots indicate the optimal values obtained for the L-curves. The corresponding $\sigma_\rho$ and $\lambda$ values are indicated in the legend. The black dots indicate the reference model for $(\sigma_\rho, \lambda) = (100 \text{ kg m}^{-3}, 200 \text{ m})$ (Figs 3 and 5).

(a) Gravimetric inversion (b) Muographic inversion (c) Joint inversion

Independent inversion of the gravimetric data. (b) Independent inversion of the muographic data from CDC, BDR and TDF viewpoints. (c) Joint inversion of the gravimetric and the muographic data from CDC, BDR and TDF viewpoints.

speaking, the model RMSE and MAE and the a posteriori standard deviation on density are improved with respect to the independent inversions, the addition of the muographic data improving RMSE and MAE by 2.8 and 3.1 per cent compared to the independent gravimetric inversion (Table 1).

The independent inversion of the muographic data from the three viewpoints CDC, TDF and BDR adequately retrieves the amplitudes and locations of the density anomalies in the upper part of the dome above $\sim 1000$ m of altitude, that is, in the upper $\sim 400$ m below the summit (Fig. 9e). This corresponds to the part of the dome illuminated by the muographic data (Fig. 2d). This result is shown as well by the a posteriori standard deviation on density (Fig. 9e, middle). Quantitatively, the RMSE and the MAE are increased by 37.2 per cent and 37.7 per cent respectively with respect to the gravimetric independent inversion (Table 1).

The joint inversion of the complete gravimetric and muographic data sets (Fig. 9f) provides the best reconstruction with RMSE and MAE improved by 7.7 per cent and 9.2 per cent respectively with respect to the gravimetric independent inversion (Table 1). Similarly, the mean a posteriori standard deviation on density is lower than
3.4 Influence of density shifts between gravimetric data, muographic data and *a priori* model

The inversion method takes into account the statistical errors on the data and on the *a priori* density, but it does not take into account possible systematic errors, that is, biases. Up to now, we have therefore assumed that the *a priori* density corresponds to the true averaged density and that gravimetric and muographic data are not biased. However, the *a priori* density $\rho_{\text{prior}}$, the average density inferred from gravimetric data $\overline{\rho}_{\text{grav}}$ and the averaged density seen by muographic data $\overline{\rho}_{\text{muog}}$ might differ. Indeed, gravimetry and muography are not sensitive to the densities in the same way and data coverage illuminate the volume differently (Figs 1b and c). Real gravimetric data are sensitive to the bulk rock densities of the whole Earth and are likely to be affected by the regional field that cannot be perfectly corrected, leading to a possible inaccurate determination of the averaged density, either underestimated or overestimated depending...
Figure 8. Horizontal cross-sections at 1100 m depth extracted from the 3-D density models resulting from inversions using \((\sigma_{\rho}, \lambda) = (100 \text{ kg m}^{-3}, 200 \text{ m})\). (a) True density model. (b–f) A posteriori density distribution retrieved by the inversions. Top: mean density \(\bar{\rho}\). Middle: standard deviation \(\sigma_{\bar{\rho}}\) of the a posteriori density distribution. Bottom: an example of random realization of a density model from a normal distribution centred on \(\bar{\rho}\) and with a standard deviation of \(\sigma_{\bar{\rho}}\). (b) Independent inversion of the gravimetric data. (c) Independent inversion of the muographic data from the CDC viewpoint only. (d) Joint inversion of the gravimetric data and the muographic data from CDC. (e) Independent inversion of the muographic data from CDC, BDR and TDF viewpoints. (f) Joint inversion of the gravimetric and the muographic data from CDC, TDF and BDR viewpoints. See Fig. 4 for the locations of the muographic viewpoints.

Figure 9. West–east and south–north vertical cross-sections extracted from the 3-D density models resulting from inversions using \((\sigma_{\rho}, \lambda) = (100 \text{ kg m}^{-3}, 200 \text{ m})\). Sections AA’ and BB’ are located in Fig. 3. (a) True density model. (b–f) A posteriori density distribution retrieved by the inversions. Top: mean density \(\bar{\rho}\). Middle: standard deviation \(\sigma_{\bar{\rho}}\) of the a posteriori density distribution. Bottom: an example of random realization of a density model from a normal distribution centred on \(\bar{\rho}\) and with a standard deviation of \(\sigma_{\bar{\rho}}\). (b) Independent inversion of the gravimetric data. (c) Independent inversion of the muographic data from CDC viewpoint only. (d) Joint inversion of the gravimetric data and the muographic data from CDC. (e) Independent inversion of the muographic data from CDC, BDR and TDF viewpoints. (f) Joint inversion of the gravimetric and the muographic data from CDC, TDF and BDR. See Fig. 4 for locations of the sections and of the muographic viewpoints.
Table 1. Comparison of the independent and joint inversion results (Figs 8 and 9) in terms of root-mean-square error (RMSE) and mean absolute error (MAE) with respect to the true synthetic model (Fig. 3), mean $\sigma_\rho$ of the standard deviation on the a posteriori density model and data misfits $\chi^2$, $\chi^2_{\text{grav}}$ and $\chi^2_{\text{muog}}$. The muographic and joint inversions use either a single muographic viewpoint CDC or the three muographic viewpoints CDC, TDF and BDR. The percentages indicate the gain with respect to the independent gravimetric inversion. A negative percentage indicates an improvement in the density model recovery, compared to the gravimetric inversion.

<table>
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<td>MAE (kg m$^{-3}$)</td>
<td>50.9</td>
<td>77.8 (52.8 per cent)</td>
<td>49.3 (−3.1 per cent)</td>
<td>70.1 (37.7 per cent)</td>
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<tr>
<td>$\sigma_\rho$ (kg m$^{-3}$)</td>
<td>62.8</td>
<td>95.6 (86.5 per cent)</td>
<td>60.9 (−3.0 per cent)</td>
<td>86.5 (37.7 per cent)</td>
<td>57.4 (−8.6 per cent)</td>
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4 DISCUSSION

4.1 Complementarity of gravimetry and muography

The joint inversion of the gravimetric and muographic data significantly improves the resulting density models compared with the independent inversions, especially when using three muographic viewpoints (Figs 8 and 9). The final density model (Figs 8 and 9) suitably reproduces the true density and gives the lowest averaged a posteriori density standard deviation (57.4 kg m$^{-3}$). Note that the improvement of the joint inversion with respect to the independent inversions and the ability to recover the true density model strongly depends on the errors of the data. We used realistic data errors given the state-of-the-art of muography. However, future improvements in muography data acquisition and analysis are likely to decrease the expected errors, hence increase the contribution of the muographic data in the resolution of the density model. Jourde et al. (2015a) observed a gain in resolution below the volume sounded by muography when adding gravimetric data, but we do not observe any significant gain, possibly due to the different regularization used and the limited spatial extension of the gravimetric data with respect to the inverted volume. However, our results show that the gravimetric data overcome the lack of spatial coverage even when only one viewpoint is available in muography: gravimetry constrains the lateral variations while muography brings resolution at depth (Fig. 9d).

When imaging a volcano, we are interested in both the shallow density distribution in the real Earth and on the data coverage. Muography is sensitive to second order to the medium composition, such as the water content (e.g. Lechmann et al. 2018). Real muographic data might be contaminated by non-ballistic muons (Nishiyama et al. 2014a, 2016), arising from the interactions of muons with matter. This contamination strongly depends on the geometry and the amplitudes of the density anomalies are equally well reproduced in the upper part of the model which is well resolved and the densities are overestimated in the bottom part of the model to compensate for the low upper density and to force the overall density model to be in accordance with the gravimetric data and the a priori density. When the gravimetric data overestimate the density by 200 kg m$^{-3}$ (Fig. 11c), the recovered densities are underestimated in the upper part of the dome which is adequately resolved by the muographic data, while the densities are overestimated in the bottom part of the model to compensate for the low upper density and to force the overall density model to be in accordance with the gravimetric data and the a priori density. When the gravimetric data overestimate the density by 200 kg m$^{-3}$ (Fig. 11d), the shallower density variations are not affected a lot as the muographic data also contribute to the resolution. But high densities are recovered in the bottom part of the model to explain the gravimetric data. The last case is a compromise between the two previous cases: muography underestimates the density by 100 kg m$^{-3}$, gravimetry overestimates the density by 100 kg m$^{-3}$ and the true and a priori densities lie in between (Fig. 11e). In this case, the top densities are underestimated while the bottom densities are overestimated for the same reasons as in the previous case. In terms of geometry of the recovered anomalies, the shapes of the structures are strongly altered (Figs 11c–e).
Figure 10. Fourfold CVSS criterion maps for the inversions with density shifts between the density inferred from gravimetric data $\rho_{\text{grav}}$, the averaged density measured by muography $\rho_{\text{muog}}$ and the a priori density $\rho_{\text{prior}}$ with respect to the true averaged density $\rho_{\text{true}}$. The white dots indicate the optimal values obtained from the fourfold CVSS criterion. The black dots indicate the reference model for $(\sigma_\rho, \lambda) = (100 \text{ kg m}^{-3}, 200 \text{ m})$ (Figs 3 and 5). Left: joint inversion of the gravimetric data and the muographic data from the CDC viewpoint only. Right: joint inversion of the gravimetric data and the muographic data from the three viewpoints CDC, TDF and BDR.
Figure 11. Effect of density shifts between the density inferred from gravimetric data \( \rho_{\text{grav}} \), the averaged density measured by muography \( \rho_{\text{muog}} \) and the a priori density \( \rho_{\text{prior}} \) with respect to the true averaged density \( \overline{\rho}_{\text{true}} \). The density models were obtained using the inversion hyperparameters determined with the fourfold CVSS criterion (white dots on Fig. 10). All colour bars are centred on \( \overline{\rho}_{\text{true}} \). The range of the colour bars of panels (a) and (b) is the same as in the other figures of the paper (\( \pm 200 \text{ kg m}^{-3} \)), but the colour bars in panels (c)–(e) (\( \pm 300 \text{ kg m}^{-3} \)) are different. Left: joint inversion of the gravimetric data and the muographic data from the CDC viewpoint only. Right: joint inversion of the gravimetric data and the muographic data from the three viewpoints CDC, TDF and BDR. See Fig. 4 for the locations of the sections and of the muon detectors.
structure and the structure at larger depth, so that we recommend using these two types of data jointly as previously established by Jourde et al. (2015a): muography improves the determination of the density structure at shallow elevations above the muons telescope, while gravimetry has the ability to resolve the density structure deeper than the muons telescope with a suitable network coverage.

4.2 A priori density model and determination of associated hyperparameters

The a priori density model to regularize the inversion was introduced using a Bayesian formalism. The a priori model is described as normally distributed densities with a Gaussian spatial correlation with a given correlation length as in eq. (9), resulting in smooth recovered density models. In previously published inversion of muographic data (e.g. Rosas-Carbajal et al. 2017), we could observe smearing out of the density anomalies along the muographic lines of sight. With our model regularization and three muographic viewpoints, these artefacts disappear (Figs 8e–f). This makes our approach particularly suitable for case studies with as few as three muon telescopes to image a kilometre scale volcano.

Another advantage of our regularization is that it avoids using weighting to counteract for the decreasing sensitivity. As shown by the sensitivities on Fig. 2, a given density node of the model is sensitive in different ways to the gravimetric and to the muographic data so that counteracting the sensitivity decay by weighting requires an arbitrary compromise between the two methods (e.g. Rosas-Carbajal et al. 2017). Imposing a spatial correlation to the model is a self-consistent approach in the sense that the a priori density model information does not depend on the data.

We consider an isotropic density smoothing using a unique constant isotropic correlation length. Cosburn et al. (2019) use both vertical and horizontal correlation lengths to account for the horizontal layering of the studied structures. In the case of a volcanic dome, such as the Puy de Dôme volcano, we expect spherical shapes rather than layered structures, therefore using an isotropic correlation length is appropriate.

In this study, we designed a synthetic model (Fig. 3) consistent with the assumption of a smooth Gaussian density distribution. However, the real Earth is likely to encompass anomalies at several scales and to present some discontinuities. Our approach is suitable for applications with little geological a priori knowledge or as a first step towards building a more complex model. In a second step, one could use the resulting smooth model as an initial model for nonlinear inversions looking for geometrical features like the precise location of interfaces between structures of presumed densities (e.g. Camacho et al. 2007; Celli et al. 2007; Lelièvre & Farquharson 2013; Linde et al. 2014).

The same regularization approach was used by (Nishiyama et al. 2014b) for real data, with one muographic viewpoint. To the difference of the approach of Nishiyama et al. (2014b), where the hyperparameters were determined based on a given arbitrary synthetic model and then used for the real data inversion, our approach is generic as determination of the hyperparameters only relies on the data. For k-fold CVSS, usage shows that typical values for k lie between 5 and 10 (e.g. James et al. 2013). We use a random distribution of the data to mix the gravimetric and the muographic data. In our case, a k-fold CVSS criterion with k = 4 and the LOO methods are shown to give satisfying results. For such Bayesian joint inversion of gravimetric and muographic data, we therefore advocate using the LOO criterion or the CVSS criterion with k ≥ 4.

4.3 Dealing with biases

In real data, the average density seen by muography is usually lower than the average density estimated with gravimetric data (e.g. Nishiyama et al. 2016; Rosas-Carbajal et al. 2017). Rosas-Carbajal et al. (2017) invert for a density model as well as a possible constant offset between the densities inferred by muographic data and gravimetric data. Lelièvre et al. (2019) investigate several methods to invert for a constant offset and suggest that the best approach is the automatic determination by least-squares minimization of a constant offset added to the observed muographic data.

The Bayesian approach presented in this paper has the advantage of taking into account the statistical errors on the data (variances and covariances), but it does not take into account any systematic errors. In Section 3.4, we show that biases in the a priori density, in the muographic data, or in the gravimetric data lead to density models that are likely to be misinterpreted.

When the mean a priori density is biased, the bias is difficult to detect in the resulting density model, but it affects only the parts of the model where the a posteriori density standard deviation is high thanks to the data resolving power (Figs 9d and f middle, and 11a and b). Where densities standard deviations are high, densities are close to the a priori. Therefore both the resulting densities and the associated a posteriori standard deviations should be considered for any geological interpretation of the resulting models.

Biases between the gravimetric and the muographic data sets are likely to alter not only the amplitudes of the recovered density anomalies, but also their shapes (Figs 11c–e), leading to misinterpretations. In particular, we obtain overestimated densities in the bottom part of the model to compensate either for muography underestimating densities either gravimetry overestimating densities, in accordance with the synthetic tests of Lelièvre et al. (2019). The occurrence of extreme densities in parts of the model badly resolved by the data could therefore be used as an indication of some shift between densities inferred from gravimetric and muographic data. This shift can be taken into account either automatically as shown in Lelièvre et al. (2019) or by hand performing several inversions with extreme values for the data to visualize the resulting range of possible density models. Gravimetric data can be shifted by adding the effect of a constant density offset in the model, computed using the forward formulation. Muographic data can easily be shifted (real data plus or minus the estimated bias). In real muographic data, the bias due to scattered low energy particles is likely not constant. It depends on the observation direction, on the topography shape and on the target surface structure, see for instance Gómez et al. (2017) for the effect of the muon incidence elevation angle, Niess et al. (2018b) for the effect of the topography or Cărloganu & the TOMUVOL collaboration (2018) for the effect of the local environment. In addition, uncertainties on the target composition result in non-uniform systematics as well. Detailed simulations can be used to estimate these effects. Reverse Monte-Carlo techniques are particularly efficient for this purpose (e.g. Niess et al. 2018a). The estimation of these biases would allow to get the possible range for the data. Hence, the range of resulting density models could be evaluated by inverting the data, plus and minus the biases estimates.

5 CONCLUSION

We show that the Bayesian formalism is well suited for the joint inversion of gravimetric and muographic data, even with a limited number of muon acquisition viewpoints.
The inversion results outline the potential of joint muography and gravimetry inversions to retrieve the 3-D density distribution of geological structures such as volcanic edifices. In particular, the muographic data improve the resolution of structures at depth above the altitude of the muon detectors. They are able to constrain the vertical extension of structures, in contrast to gravimetric data which better constrain the lateral variations. When as few as one to three muographic viewpoints are available, gravimetric data bring complementary information at shallow depth and constrain density variations in the direction of the muography lines of sight, along which muography is blind, as well as away from the muon detectors. Below the muon detectors, gravimetry has the ability to resolve densities depending on the network coverage, with a resolving power rapidly decreasing with depth. We also determined that offsets between densities inferred by gravimetric and muographic data lead to recognizable artefacts in the retrieved densities, which should be taken into account in the inversions of real data.

The Bayesian formalism has the advantage to produce a model of densities with the associated a posteriori standard deviations that should be used to avoid misinterpretation of the resulting structures. We show that, for the joint inversion of gravimetric and muographic data, cross-validation criteria (k-fold CVSS or LOO) are more robust criteria than the classically used L-Curves to select the inversion a priori hyperparameters for such joint inversion of gravimetric and muographic data. We also show that the regularization overcomes the artefacts due to the acquisition geometry: when three muographic viewpoints are used, we have no more smearing out along the muographic lines of sight as it can be observed in previously published studies. Therefore our method presents a significant improvement to the robustness of the inversions.

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REFERENCES


**APPENDIX: CONSTRUCTION OF THE GAUSSIAN RANDOM DENSITY MODEL**

A Gaussian random density model following the distribution defined by $N(\mu_{\text{true}}, \Sigma_{\text{true}})$ can be constructed from

$$\rho_{\text{true}} = \rho_{\text{true}} + LX,$$  \hspace{1cm} (A1)

where $X \sim N(0, I)$ is a random vector drawn from a zero-mean normal distribution with an identity covariance matrix and $L$ is the Cholesky decomposition of the covariance matrix $\Sigma_{\text{true}} = LL^T$ with a Gaussian correlation function of standard deviation $\sigma_{\text{true}}$ and of correlation length $\lambda_{\text{true}}$. Indeed, given the mean and variance of $X$

$$\begin{cases}
E[X] = 0 \\
E[XX^T] - E[X]E[X^T] = I,
\end{cases}$$  \hspace{1cm} (A2)

the mean and variance of $LX$ are (e.g. Gentle 2009):

$$\begin{cases}
E[LX] = LE[X] = 0 \\
\end{cases}$$  \hspace{1cm} (A3)

Hence, $LX \sim N(0, \Sigma_{\text{true}})$ and $\rho_{\text{true}} = \rho_{\text{true}} + LX \sim N(\mu_{\text{true}}, \Sigma_{\text{true}})$. 

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