Blind Marine Seismic Deconvolution Using Statistical MCMC Methods

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Abstract—In order to improve the resolution of seismic images, a blind deconvolution of seismic traces is necessary, since the source wavelet is not known and cannot be considered as a stationary signal. The reflectivity sequence is modeled as a Gaussian mixture, depending on three parameters (high and low reflector variances and reflector density), on the wavelet impulse response, and on the observation noise variance. These parameters are unknown and must be estimated from the recorded trace, which is the reflectivity convolved with the wavelet, plus noise. Two methods are compared in this paper for the parameter estimation. Since we are considering an incomplete data problem, we first consider maximum likelihood estimation by means of a stochastic expectation maximization (SEM) method. Alternatively, proper prior distributions can be specified for all unknown quantities. Then, a Bayesian strategy is applied, based on a Monte Carlo Markov Chain (MCMC) method. Having estimated the parameters, one can proceed to the deconvolution. A maximum posterior mode (MPM) criterion is optimized by means of an MCMC method. The deconvolution capability of these procedures is checked first on synthetic signals and then on the seismic data of the IFREMER ESSR4 campaign, where the wavelet duration blurs the reflectivity, and on the SMAVH high-resolution marine seismic data.

Index Terms—Blind deconvolution, EM algorithm, Markov Chain Monte Carlo (MCMC), maximum likelihood, maximum posterior mode (MPM) method, seismic signals, stochastic expectation maximization (SEM) algorithm.

I. INTRODUCTION

THE AIM of high-resolution marine seismic exploration is to obtain an image of the sea-bottom lithography from reflected acoustic waves emitted by various seismic sources. It uses high-frequency waves with frequencies ranging from 500 Hz to 2 KHz, which penetrates in the ground as far as 100 m with a resolution of 1 m. The seismic source, which produces the seismic wavelet, is tracked by a ship Fig. 1, which also drags a multitrace streamer consisting in an array of hydrophones. The hydrophones record the reflected acoustic signals. The traveling distance between two shots is equal to the half of the distance between two sensors, which means that each sensor will receive signals coming from the same common reflection point of the sea bed during a shot sequence. This allows to average the signals to increase the signal-to-noise ratio (SNR). The received signals can be modeled as the convolution of a source wavelet and a reflectivity sequence, which describes the most important interfaces met by the emitted wave [1]. The source wavelet

cannot be considered as a stationary signal for three main reasons: the usual sources as air guns or sparkers cannot be sufficiently controlled in operational conditions at sea to obtain calibrated wavelets; the received signal is blurred by the ghost phenomenon due to multiple reflections on the free sea surface and also by averaging hydrophone signals in the common point method; the layered geological structure of the sea bottom acts as a lowpass filter, depending on the localization of the common point. Furthermore, the problem can be modeled as the filtering of the reflectivity sequence by a lowpass filter, whose impulse response is the wavelet source. So, since the wavelet is unknown, blind deconvolution must be applied to recover the reflectivity sequence from the recorded trace.

The first deconvolution methods [2], [3] assumed that the reflectivity sequence was a white noise and that the wavelet was described by a deterministic ARMA model. In this case, usual parametric methods based on second-order statistics can be used, but they systematically lead to a minimum phase solution. If the wavelet is nonminimum phase, which occurs in most experiments, one may then use higher order statistics (third- or fourth-order cumulants in general) to retrieve the signal phase [4], [5]. However, this kind of techniques requires a large amount of data to ensure efficient estimation. Such conditions are not satisfied in marine reflection seismology, where the observation vectors contain less than 1000 samples. Moreover, these higher order approaches do not really take into account the statistical information contained in the reflectivity signal. In particular, one simple way to characterize the seismic signal is to consider that the sea floor is composed of several layers that are more or less homogenous and separated by interfaces where the reflections occur. This simple hypothesis permits us to statistically model the reflectivity sequence as a Bernoulli–Gaussian process [6], [7] or, more generally, as a mixture of Gaussian distributions [8]. For each sample of the reflectivity sequence, a Bernoulli variable characterizes the presence or the absence of a reflector. Then, conditionally to that label, the amplitudes follow a Gaussian distribution. Once this type of *a priori* information is given, it is easy to incorporate it into the convolution model.

The estimation can then be achieved by maximizing the posterior likelihood functional. In practice, maximizing the likelihood functional of incomplete data models (here, the model is completed by incorporating the Bernoulli reflectivity sequence) can be done by using the expectation–maximization (EM) algorithm [9]. However, for our problem, this algorithm requires high-dimensional integrations that are very difficult to calculate by means of classical numerical methods. MCMC methods

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Fig. 1. Acquisition system.

Fig. 2. Synthetic trace.

[15] enable us to circumvent this difficulty by solving the integration and optimization problems by simulating random variables. This approach leads to stochastic versions of the EM algorithm [stochastic expectation maximization (SEM), SAEM] [12]. Moreover, we investigate the problem in a totally Bayesian framework in which prior information is introduced upon the parameters [10]. Then, the estimation of the Bernoulli–Gaussian model and of the wavelet is solved by the simulation of random variables via MCMC algorithms such as the Gibbs sampler [10].

Once the model has been estimated, the next step is the deconvolution itself. This problem is said to be “ill-posed” because, in the presence of noise, different reflectivity sequences can lead to similar seismic data. Therefore, it is necessary to use as much prior information as possible upon the reflectivity to limit the set of acceptable solutions. Thus, it is natural to account for the Bernoulli–Gaussian assumption introduced above. Then, we need an adequate procedure to achieve the detection of reflectors and the estimation of their amplitude. Actually, the problem can be solved using either the maximum *a posteriori* (MAP) criterion, which can be optimized using the simulated annealing technique, or by the suboptimum maximum posterior mode (MPM) method, which involves optimization by means of a MCMC technique.

II. PROBLEM MODELING

A seismic trace can be modeled as the convolution of the unknown wavelet h with the reflectivity sequence r , yielding the observation

$$y(k) = \sum_{i=0}^L h(i)r(k-i) + w(k). \quad (1)$$

$\underline{y}$ and $\underline{r}$ will, respectively, denote the trace vector of dimension N and the reflectivity sequence of length N . The source wavelet $\underline{h}$ is represented by a moving average (MA) model with L coefficients. The noise is assumed to be white Gaussian with variance σ_w^2 . The reflectivity is represented by two random variables $z(k) = (q, r)$, which define a generalized Bernoulli–Gaussian process [1]. If $q(k)$ is equal to one, it indicates the position of a high reflector position and a small reflector position for zero. $r(k)$ is the amplitude of the reflected signal and its distribution is a Gaussian mixture.

$$p(q(k) = 1) = 1 - p(q(k) = 0) = \lambda \quad (2)$$

$$p(r(k)) = \lambda \mathcal{N}(0, \sigma_1^2) + (1 - \lambda) \mathcal{N}(0, \sigma_0^2). \quad (3)$$

λ is the density of the reflectivity sequence. σ_1^2 is the variance of high reflectors, while σ_0^2 represents the variance of secondary reflections due to the inhomogeneity of each layer. It is assumed that $\sigma_0^2 \ll \sigma_1^2$.

Fig. 3. True and estimated wavelets by using maximum likelihood (right images) and Bayesian methods (left images) for various SNR (from top to bottom: 17, 13, and 6 db).

The aim of the problem is to estimate $\theta = (\underline{h}, \lambda, \sigma_0^2, \sigma_1^2, \sigma_w^2)$ from the data $\underline{y}$ and then to proceed to the deconvolution with the estimated source wavelet $\underline{h}$.

III. PARAMETER ESTIMATION

A. Maximum Likelihood Approach

A classical method consists in using the maximum likelihood (ML) criterion as follows:

$$\hat{\theta}_{MV} = \arg \max_{\theta} \ln p(\underline{y}; \theta). \quad (4)$$

As mentioned above, this is in fact an incomplete data problem [8], [9] and it must be solved by introducing the missing variables $\underline{z}$. The joint probability density is expressed by

$$p(\underline{y}, \underline{z}; \theta) = p(\underline{y}|\underline{z}; \theta)p(\underline{z}; \theta). \quad (5)$$

As $\underline{z} = (q, r)$, it can also be written as

$$p(\underline{y}, \underline{z}; \theta) = p(\underline{y}|\underline{z}; \theta)p(r|\underline{q}; \theta)p(\underline{q}; \theta). \quad (6)$$

Each component of this equation can be easily expressed. As q is a Bernoulli variable, $p(\underline{q}) = \prod_{k=1}^N p(q(k))$ with

$$p(q(k); \theta) = \lambda^{q(k)}(1 - \lambda)^{1-q(k)}. \quad (7)$$

The variables $r(k)$ are independent conditionally to the variables $q(k)$

$$p(r|\underline{q}; \theta) = \prod_{k=1}^N p(r(k)|q(k); \theta) \quad (8)$$

$$p(y|\underline{z}; \theta) = \frac{1}{(2\pi\sigma_w^2)^{\frac{N}{2}}} \exp \left[-\frac{\sum_{k=1}^N (y(k) - h^*r(k))^2}{2\sigma_w^2} \right]. \quad (9)$$

The complete Log likelihood is expressed by

$$\begin{aligned} L(\underline{y}, \underline{z}; \theta) = & -\frac{((\underline{y} - \underline{H}r)^T(\underline{y} - \underline{H}r))}{2\sigma_w^2} - \frac{s_1(z)}{2\sigma_1^2} \\ & - \frac{s_0(z)}{2\sigma_0^2} - \left(\frac{N}{2}\right) \ln \sigma_w^2 + n_1(q) \ln \left(\frac{\lambda}{\sigma_1}\right) \\ & + n_0(q) \ln \left(\frac{(1-\lambda)}{\sigma_0}\right) \end{aligned} \quad (10)$$

with $\underline{Q} = \text{diag}(q)$, $s_1(z) = r^T \underline{Q} r$, $s_0(z) = r^T (\underline{I} - \underline{Q}) r$, $n_1(q) = q^T q$. We note that $\underline{H}r = \underline{R}h$, where $\underline{H}$ and $\underline{R}$ are the convolution matrices associated with r and h .

When the complete data $\underline{z}$ are known, the optimization problem solution is straightforward. A Gibbs sampler [10], [11] is used to obtain the complete data by simulation, which consists of a random sampling from the probability distribution $p(z(k)|\underline{y}; \underline{z}_{-k})$, where $\underline{z}_{-k}$ means all the vector values except

TABLE I

PERFORMANCES OF THE BAYESIAN AND ML DECONVOLUTION MSEw MEAN SQUARE ERROR BETWEEN TRUE AND ESTIMATED WAVELETS. MSEr MEAN SQUARE ERROR BETWEEN TRUE AND ESTIMATED IMPULSES OF THE REFLECTIVITY SEQUENCE FA NUMBER OF FALSE ALARMS (DETECTED IMPULSES OF HIGH-AMPLITUDE, WHICH DO NOT EXIST). D = NUMBER OF DETECTED IMPULSES OF HIGH AMPLITUDE (ON 25 GENERATED IMPULSES), LE1 = NUMBER OF IMPULSES WHOSE LOCALIZATION HAS AN ERROR OF ONE SAMPLE LEFT OR RIGHT, LE2 = NUMBER OF IMPULSES WHOSE LOCALIZATION HAS AN ERROR OF TWO SAMPLES LEFT OR RIGHT, LE3 = NUMBER OF IMPULSES WHOSE LOCALIZATION HAS AN ERROR OF THREE SAMPLES LEFT OR RIGHT

SNR(dB)		MSEw	MSEr	FA	D	LE1	LE2	LE3
17	Bayesian	0.0009	2.9	2	23	1	0	0
	ML	0.0018	2.6	18	22	2	1	0
13	Bayesian	0.0029	2.4	4	22	1	0	0
	ML	0.0026	3.5	22	21	3	1	1
6	Bayesian	0.027	2.7	7	19	1	0	0
	ML	0.06	3.4	11	18	1	2	0

$z(k)$. In this case, it can be shown that the posterior density function becomes [8]

$$p(q(k) = 1 | \underline{y}, \underline{z}_{-k}) = \left[1 + \frac{1 - \lambda}{\lambda} \sqrt{\frac{V_0 \sigma_1^2}{V_1 \sigma_0^2}} \exp\left(\frac{m_0^2}{2V_0} - \frac{m_1^2}{2V_1}\right) \right]^{-1} \quad (11)$$

with

$$\begin{aligned} V_1 &= \left(\frac{1}{V_w} + \frac{1}{\sigma_1^2} \right)^{-1}, \quad V_0 = \left(\frac{1}{V_w} + \frac{1}{\sigma_0^2} \right)^{-1} \\ m_1 &= \frac{V_1}{V_w} m_w, \quad m_0 = \frac{V_0}{V_w} m_w, \quad V_w = \frac{\sigma_w^2}{E} \\ m_w &= \frac{1}{E} \sum_{i=0}^L h(i) \left[y(k+i) - \sum_{\substack{j=0 \\ j \neq i}}^L h(j) r(i+k-j) \right] \\ E &= \sum_{i=0}^L h^2(i). \end{aligned} \quad (12)$$

For the simulation of the vector $\underline{z}$, the Gibbs sampler has the following iterative steps:

For $i \geq 1$ and $k = 1, \dots, N$

- Choice of k ;
- Computation of $p(q(k) = 1 | \underline{y}; \underline{z}_{-k})$;
- Random sampling of u in $[0, 1]$ and new value of $q(k)$:

$$q(k) = \begin{cases} 1 & \text{if } p(q(k) = 1 | \underline{y}; \underline{z}_{-k}) > u; \\ 0 & \text{if } p(q(k) = 1 | \underline{y}; \underline{z}_{-k}) < u; \end{cases}$$
- Estimation: simulation of $r(k)$ with $\mathcal{N}(m_q, \sigma_q^2)$;
- New value: $z^{(i)}(k) = (q(k), r(k))$;

The SEM [12] is then used to solve the estimation problem:

Initialization:

- choice of $\theta^{(0)}$ and $\underline{z}^{(0)}$;

For $i = 1, \dots, M$:

- E step:
 - Simulation of $\underline{z}^{(i)}$ by the Gibbs sampler according to $p(\underline{z} | \theta^{(i-1)})$;
- M step:
 - Estimation of the parameters

$$\begin{aligned} \hat{h}^{(i)} &= M_1(\underline{r})^{-1} M_2(\underline{r}) \text{ with} \\ M_1(\underline{r}) &= \underline{R}^T \underline{R}, \quad M_2(\underline{r}) = \underline{R}^T \underline{y} \\ \hat{\sigma}_w^{2(i)} &= \frac{1}{N} (\underline{y}^T \underline{y} - M_2(\underline{r})^T M_1(\underline{r})^{-1} M_2(\underline{r})) \\ \hat{\sigma}_1^{2(i)} &= \frac{s_1(\underline{z})}{n_1(\underline{q})} \\ \hat{\sigma}_0^{2(i)} &= \frac{s_0(\underline{z})}{n_0(\underline{q})} \\ \hat{\lambda}^{(i)} &= \frac{n_1(\underline{q})}{N}. \end{aligned}$$

B. Bayesian Approach

In Bayesian estimation, the hyper parameters θ are random and have a prior density $p(\theta)$. When no prior distribution is available, it is possible to specify noninformative prior distributions that are easy to handle, such as conjugate priors [10]. Then, the objective is to estimate the joint posterior distribution $p(\underline{y}, \theta | \underline{z})$, which can be expressed by using the Bays rule by

$$p(\underline{y}, \theta | \underline{z}) = \frac{p(\underline{z} | \underline{y}, \theta) p(\underline{y} | \theta) p(\theta)}{p(\underline{z})} \quad (13)$$

with

$$p(\underline{z}) = \int \dots \int p(\underline{z} | \underline{y}, \theta) p(\underline{y} | \theta) p(\theta) d\underline{y} d\theta. \quad (14)$$

Fig. 4. True reflectivity and reflectivities.

This requires integrations that are generally impossible to perform. To overcome this difficulty, Monte Carlo methods are used to build a Markov chain whose equilibrium distribution coincides with the desired joint posterior distribution of the unknown parameters. This is performed by the Gibbs sampler,

which generates estimates by sampling from conditional distributions.

A detection-estimation procedure is used to simulate the missing variable, because the direct sampling from $p(\underline{r}, \underline{q} | \underline{z}, \theta^{k-1})$ is impossible.

Fig. 5. Convergence speed of the SEM method (SNR = 15 dB).

Fig. 6. Convergence speed for the Bayesian method (SNR = 15 dB).

- Detection: simulate $q(n)$ according to the distribution $p(q(n) = 1|\underline{z}, \underline{q}_{-n}, \underline{r}_{-n})$, where any vector $\underline{v}_{-n}$ is defined by

$$\underline{v}_{-n} = [v(1), \dots, v(n-1), v(n+1), \dots, v(N)]^T.$$

- Estimation: simulate $r(n)$ according to $p(r(n)|\underline{z}, \underline{q}, \underline{r}_{-n})$.

Each parameter is associated to a prior and classical conjugate priors, such as Gaussian distributions $\mathcal{N}$; Beta distributions or Inverse Gamma distributions $\mathcal{Be}$ are used [10]

$$\underline{h} : \mathcal{N}(\underline{0}_{(L+1) \times 1}, \Sigma_0 \underline{I}_{(L+1)}) \quad (15)$$

$$\lambda : \mathcal{Be}(\zeta, \tau) \quad (16)$$

$$\sigma_1^2 : \mathcal{IG}\left(\frac{v_1}{2}, \frac{\gamma_1}{2}\right) \quad (17)$$

$$\sigma_w^2 : \mathcal{IG}\left(\frac{v_w}{2}, \frac{\gamma_w}{2}\right). \quad (18)$$

The posterior distribution of $\underline{h}$ is expressed by

$$\underline{h}|\underline{r}, \underline{z}, \sigma_w^2 : \mathcal{N}(m, \Sigma) \quad (19)$$

where

$$\Sigma^{-1} = \frac{1}{\sigma_w^2} M_1(\underline{r}) + \Sigma_0^{-1} \underline{I}_{(L+1)} \quad (20)$$

Fig. 7. Registered seismic trace n° 1.

Fig. 8. Estimated wavelets of ten successive traces from 1 to 10 of shot 602 by maximum likelihood with the maximum value position at $d = 68$ ms. Fig. 9. Estimated wavelets of ten successive traces from 1 to 10 of shot 602 by Bayesian method with the maximum value position at $d = 68$ ms.

and

$$m = \frac{\sum}{\sigma_w^2} M_2(\underline{r}). \quad (21)$$

The posterior distribution of the noise variance is an Inverse Gamma distribution, as follows:

$$\sigma_w^2 | \underline{z}, \underline{r}, \underline{h}^{(k)} : \mathcal{IG} \left(\frac{\nu_w + N}{2}, \frac{\gamma_w + s_w^2}{2} \right) \quad (22)$$

with

$$s_w^2 = \sum_{n=1}^N \left(y(n) - \underline{h}^T \underline{x}(n) \right)^2. \quad (23)$$

The posterior distribution of the primary reflectivity variance follows:

$$\sigma_1^2 | \underline{z} : \mathcal{IG} \left(\frac{\nu_1 + N}{2}, \frac{\gamma_1 + s_1^2}{2} \right) \quad (24)$$

with

$$s_1^2 = \sum_{n=1}^N r^2(n) q(n). \quad (25)$$

The posterior distribution of the secondary reflectivity variance follows:

$$\sigma_0^2 | \underline{z} : \mathcal{IG} \left(\frac{\nu_0 + N}{2}, \frac{\gamma_0 + s_0^2}{2} \right) \quad (26)$$

with

$$s_0^2 = \sum_{n=1}^N r^2(n) (1 - q(n)). \quad (27)$$

The density of the reflectivity is given by

$$\lambda | \underline{q} : \mathcal{Be}(\zeta + n_1, \tau + N - n_1) \quad (28)$$

with

$$n_1 = \sum_{n=1}^N q(n). \quad (29)$$

From these distributions, an algorithm [13] similar to the Gibbs sampler leads to the parameter estimation

Random initialization of $\theta^{(0)}$ and $\underline{z}^{(0)}$;

For $i = 1, \dots, I$,

- simulate $\underline{z}^{(i)}$ from $p(\underline{z} | \theta^{(i-1)})$
- simulate $\underline{h}^{(i)}$ from $p(\underline{h} | \underline{y}, \underline{r}^{(i)}, \sigma_w^{2(i-1)})$

Fig. 10. Convergence speed for the SEM method.

Fig. 11. Convergence speed for the Bayesian method.

- simulate $\sigma_w^{2(i)}$ from $p(\sigma_w^2 | y, \underline{r}^{(i)}, \underline{h}^{(i)})$
- simulate $\sigma_1^{2(i)}$ from $p(\sigma_1^2 | \underline{z}^{(i)})$
- simulate $\sigma_0^{2(i)}$ from $p(\sigma_0^2 | \underline{z}^{(i)})$
- simulate $\lambda^{(i)}$ from $p(\lambda | \underline{q}^{(i)})$.

IV. DECONVOLUTION

As the parameters θ and, in particular, $\underline{h}$ are now known, the deconvolution problem is reduced to an estimation of the hidden variables $\underline{z}$ from the observation. The *a posteriori* log likelihood is expressed by $L(y, \underline{z}; \hat{\theta})$.

The separation principle [1] allows us to solve this maximization problem in two steps:

$$\text{Detection : } \hat{\underline{q}} = \arg \max_{\underline{q}} p(\underline{q} | y) \quad (30)$$

$$\text{Estimation : } \hat{\underline{r}} = \arg \max_{\underline{r}} p(\underline{r} | y, \underline{q}). \quad (31)$$

The estimation step leads to a mean square method because it is a quadratic form when $\underline{q}$ is fixed

$$\hat{\underline{r}}(\underline{q}) = \underline{S}^{-1} \underline{H}^T y, \quad \text{with} \quad \underline{S} = \underline{H}^T \underline{H} + \underline{M} \quad \text{and} \\ \underline{M} = \gamma \left[\underline{Q} + \frac{1}{\kappa} (\underline{I} - \underline{Q}) \right], \quad \kappa = \frac{\sigma_0^2}{\sigma_1^2}, \quad \gamma = \frac{\sigma_w^2}{\sigma_1^2}.$$

Fig. 12. Estimated reflectivity sequence corresponding to trace n°1.

Fig. 13. Seismic image of ESSR4 data.

Fig. 14. Deconvolved image of ESSR4 data.

The detection problem is not so easy because the vector $\underline{q}$ has 2^N discrete configurations, and testing all of them is not possible. The solution consists in the applying a simpler criterion called the maximum posterior mode (MPM) [14], which maximizes the marginal distribution $p(z(k) | y)$. As it cannot be explicitly optimized, it is simulated by means of a Monte Carlo method

Fig. 15. Seismic image with SMAVH data.

using the Gibbs sampler. At each iteration, a sequence $\underline{z}^{(i)}$ is generated according to the distribution $p(z(k)|y; \underline{z}_{-k})$.

The MPM algorithm follows the steps:

1) Initialization:

- Choice of $\underline{q}^{(0)}$ and $\underline{r}^{(0)}$;

2) Iteration $i < I$ with a learning period $i \leq I_0$ and a steady-stage period $I_0 < i \leq I$;

- Computation of $d(k) = P(q(k) = 1|y; \underline{z}_{-k})$
- Detection: random sampling of U in $[0,1]$

$$q(k) = \begin{cases} 1, & \text{if } d(k) > u \\ 0, & \text{elsewhere} \end{cases}$$

- Estimation: random uniform sampling of $r(k)$ with

$$r(k) = \begin{cases} \mathcal{N}(m_1, V_1), & \text{if } q(k) = 1 \\ \mathcal{N}(m_0, V_0), & \text{if } q(k) = 0 \end{cases}$$

3) Decision for $1 \leq k \leq N$;

- Detection

$$q(k) = \begin{cases} 1, & \text{if } \frac{1}{I-I_0} \sum_{i=I_0+1}^I q^{(i)}(k) > q_{\min} \\ 0, & \text{elsewhere} \end{cases}$$

- Estimation

$$r(k) = \begin{cases} \frac{\sum_{i=I_0+1}^I q^{(i)}(k) r^{(i)}(k)}{\sum_{i=I_0+1}^I q^{(i)}(k)}, & \text{if } q(k) = 1 \\ \frac{\sum_{i=I_0+1}^I (1-q^{(i)}(k)) r^{(i)}(k)}{\sum_{i=I_0+1}^I (1-q^{(i)}(k))}, & \text{if } q(k) = 0. \end{cases}$$

V. APPLICATION

A. Synthetic Traces

We build a synthetic trace (Fig. 2) by convolving a white Gaussian mixture impulse signal ($\sigma_1^2 = 1$, $\sigma_0^2 = \kappa \sigma_1^2$, $\kappa = 0.001$) with a synthetic Ricker wavelet (Fig. 3), plus additive white noise, whose variance σ_w^2 is adjusted to obtain a variable SNR. The SNR is defined as

$$\text{SNR} = 10 \log \left(\frac{\lambda \sigma_1^2 E_h}{\sigma_w^2} \right)$$

where $E_h = \sum_{n=0}^L h^2(n)$ stands for the wavelet energy.

In this simulation, 25 impulses of high amplitude corresponding to high reflectors were generated. We compare the deconvolution performances of the two methods with respect

TABLE II
ESTIMATED PARAMETERS FOR SMAVH DATA

λ	σ_w^2	σ_1^2	σ_0^2
0.333	0.011	1.12	$9.72 \cdot 10^{-3}$

Fig. 16. Deconvolved image with SMAVH data.

to SNR. Fig. 3 shows the true and estimated wavelets. It can be seen in Table I that the error MSE_w is a bit higher with the ML method than with the Bayesian method at each SNR.

After applying the algorithms, it may occur that two neighboring estimated impulses are found, either successive or separated by one sample. Then, a postprocessing procedure must be used to fuse these impulses by replacing them by their gravity center. Fig. 4 shows the deconvolved reflectivity sequences and the true reflectivity sequence for both algorithms at SNR 17, 13, and 6 db. Table I compares the performances of the ML and Bayesian approaches for the reflectivity sequence in terms of good detection, false alarms, localization errors, and estimation error of the amplitude of detected impulses.

It can be seen on this example that the MCMC method gives a better deconvolution than the ML method. At high SNR, the localization of impulses by MCMC is more accurate with less misses and false detection than by ML, where more false detections occur. If SNR decreases, the impulse miss number increases, because low-amplitude impulses are not detected. High-amplitude impulses continue to be well localized. It can be remarked that both methods show good robustness with regard to the SNR decrease.

The SEM convergence speed of the θ parameters is shown in Fig. 5 for a SNR of 15 db. The convergence speed of the Bayesian approach is shown in Fig. 6. In this last case, the prior distribution parameters were $\delta_h = 0.01$, $\gamma_w = v_w = 10^{-10}$, $\gamma_1 = v_1 = 10^{-10}$, $\gamma_0 = v_0 = 10^{-10}$, and $\varsigma = \tau = 1$. We remark that it is faster for the Bayesian method, while the parameter variance is higher. The wavelet estimation is more accurate with the Bayesian method when the SNR decreases.

B. Real Traces

1) *ESSR4 Data*: In the case of the ESSR4 data acquired during the ZAIANGO campaign by the IFREMER, the seismic

Fig. 17. Seismic trace along line 1 with SMAVH data.

Fig. 18. Estimated reflectivity sequence along line 1 with SMAVH data.

data set was obtained using an innovative manner for synchronizing a cluster of GI guns in order to achieve deeper penetration. The sources are no more synchronized on their firing order but on their time of maximum energy, therefore decreasing the signal bandwidth but increasing significantly the low-frequency content and, thus, the penetration. The source was composed of 11 synchronized airguns shooting every 20 s with a source wavelet whose spectrum is in the band of 0–128 Hz. The multi-trace streamer was composed of 360 clusters of 16 hydrophones, whose role was to improve the SNR by summing the received signals. The length of the steamer was 4.5 km and each hydrophone cluster was separated by 12.5 m. The sampling rate was 4 ms and each trace was composed of 2500 samples (Fig. 7).

The length of the MA model is a parameter that must be fixed. In our case, sufficient knowledge about the source wavelet generation suggested the choice of a MA model with 35 coefficients, giving a source duration of 140 ms.

When estimating the source wavelet, a problem of phase shift occurs because the convolution of shifted versions of the estimated wavelet with the reflectivity shifted the other way lead to the same seismic trace. As the wavelet is modeled by an MA process, there are some differences in the estimation when the position of the wavelet maximum value is changed. We assume that this position belongs to the interval. Then, the following algorithm is applied for neighboring traces, where the hypothesis of the wavelet stationary is made ($J \leq 10$).

For trace $j = 1$:

- Wavelet estimation of $\hat{h}_{\min}^1$ with the hypothesis that the maximum is at position $d_{\min}$;
- For $d = d_{\min} + 1, \dots, d_{\max}$
 - Wavelet shift of one sample of $\hat{h}_{d-1}^1$ toward the right;
 - Wavelet estimation of $\hat{h}_d^1$;

For trace $j = 2, \dots, J$

- Wavelet estimation with an initialization by $\hat{h}_d^{j-1}$ for $d = d_{\min}, \dots, d_{\max}$.

The decision is given by the minimum of variance, which gives

$$\bar{h}_d = J^{-1} \sum_{j=1}^J \hat{h}_d^j.$$

$$\hat{d}_o = \underset{d}{\operatorname{argmin}} \frac{1}{J} \sum_{j=1}^J \left\| \hat{h}_d^j - \bar{h}_d \right\|^2.$$

In this experiment, all maximum positions from 20 to 80 ms are explored with a step of 8 ms. The minimum variance is obtained for $d = 68$ ms with a superposition of ten estimated wavelets from neighboring traces at the same maximum value. The ML wavelet estimation for the seismic trace of Fig. 7 is shown in Fig. 8. The Bayesian approach is presented in Fig. 9. The learning time has been experimentally set to 700 iterations. We remark that the variance of the estimated wavelet is lower with the Bayesian approach than with the ML approach, which suggests that introducing a prior knowledge regularizes estimation. The empirical mean of the other parameters is computed from the next 400 samples. The convergence speed of these parameters for the ML and Bayesian methods is, respectively, in Figs. 10 and 11.

After convergence, we remark a higher variance of the parameters with the Bayesian method. One reason for this is that the Bayesian approach uses simulations from probability distributions more extensively than the ML approach. Nevertheless, the estimation means are similar. Fig. 12. shows the reflectivity sequence obtained by applying the MPM deconvolution method.

Putting the traces of all the shots of the same hydrophone cluster side by side gives an image of the sea bottom stratigraphy (Figs. 13 and 14). Transitions between layers are thinner and less blurred in the deconvolved image.

2) *High-Resolution SMAVH Data*: These data come from a high-resolution marine seismic experiment undertaken by IFREMER in shallow water. The source is an air gun whose bandwidth is 200 Hz. The streamer is composed of 24 hydrophones separated by a length of 25 m. Fig. 15 shows the seismic image obtained from 401 successive shots. The MCMC algorithm is applied on the raw data and the estimated parameters are given in Table II. Fig. 16 shows the deconvolved image. Figs. 17 and 18 show, respectively, a raw seismic trace and the corresponding estimated reflectivity sequence along the vertical line 1. Comparing Figs. 15 and 16 shows that blind deconvolution improves the information quality given by the seismic image: New layers, which were blurred in the seismic image, appear clearly in the deconvolved image; layer interfaces are thinner and better localized in the deconvolved image. A contribution of blind deconvolution here is to let appear layers immediately under the sea-floor line, which were masked in the original image because of the high energy of the sea-floor reflection. This can also be seen by comparing Figs. 17 and 18.

VI. CONCLUSION

We compared an SEM implementation of the ML method with an MCMC implementation of a Bayesian method for parameter estimation in blind deconvolution of seismic traces. These methods were applied to synthetic and real seismic traces. In the last case, practical problems, such as the wavelet shift, have been experimentally solved. We showed that the Bayesian approach leads to slightly improved results, which implies that use of prior information leads to a more regularized wavelet estimation. We also showed on real data experiments that these blind deconvolution methods enable to improve seismic image resolution: Better localization of layer interfaces is achieved and new layers can be detected after processing in the deconvolved image.

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