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Evaluation and analysis of deep percolation losses of drip irrigated citrus crops under non-saline and saline conditions in a semi-arid area

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Abstract
In arid and semi-arid regions, irrigation management is important to avoid water loss by soil evaporation and deep percolation (DP). In this context, estimating the irrigation water demand has been investigated by many studies in the Haouz plain. However, DP losses beneath irrigated areas in the plain have not been quantified. To fill the gap, this study evaluated DP over two drip-irrigated citrus orchards (Agafay and Saada) using both water balance and direct fluxmeter measurement methods, and explored the simple FAO-56 approach to optimise irrigation in order to both avoid crop water stress and reduce DP losses in case of non-saline and saline soils. The experimental measurements determined different terms of the water balance by using an Eddy-Covariance system, fluxmeter, soil moisture sensors and a meteorological station. Using the water balance equation and fluxmeter measurements, results showed that about 37% and 45% of supplied water was lost by DP in Saada and Agafay sites, respectively. The main cause of DP losses was the mismatch between irrigation and the real crop water requirement. For Agafay site, it was found that increased over-irrigation had the effect of reducing soil salinity by leaching salts.

The applied FAO-56 model suggested an optimal irrigation scheduling by taking into account both rainfall and soil salinity. The recommended irrigations could save about 39% of supplied water in non-saline soil at Saada and from 30% to 47% in saline soil at Agafay.

Key words: Saline soil; water balance; fluxmeter; FAO-56 approach; irrigation scheduling.

1. Introduction
In the southern Mediterranean region, as in many arid and semi-arid regions of the world, water scarcity is one of the main factors limiting crop development, growth and yield (Kharrou et al., 2013). In this region, irrigation is a major component of water demand. It is estimated that about 83% of available resources is dedicated to agriculture with an efficiency lower than 50% (Chehbouni et al., 2008). In Morocco, this waste of water has several origins including leakage during water routing, but also a lack of irrigation efficiency in the field (Khabba et al., 2013; Belaqziz et al., 2014). Indeed, scheduling in timing and amount of irrigation is mostly determined according to the water availability so that the actual plant water needs are generally not taken into account. In addition, the traditional flooding systems are predominantly leading to significant water loss by soil evaporation and deep percolation (DP) (Kharrou et al., 2011). Currently, the Moroccan government has set up an ambitious
program for irrigation conversion from flood to drip (PMV, 2013). However, the obtained results for DP are surprising as an inadequate use of the drip technique may lead to substantial water losses (Khabba et al., 2013). It has been shown that the DP losses for the drip irrigation sites are in the range 29-41 % of water input while they are relatively lower for flood irrigation, ranging from 26 to 31 % (Khabba et al., 2013). Moreover, as the soil salinity may limits plant growth (Ayars et al., 2012), the farmers usually apply excessive irrigation to leach soil salinity out of root layer or root zone and hence avoid salinity stress (Visconti et al., 2012). In general, over-irrigation amounts are arbitrary and largely estimated. In our study basin, this situation leads to an overexploitation of groundwater, with a level decreasing from 1 to 3 m year\(^{-1}\) (Le Page et al., 2012; Boukhari et al., 2015).

In order to preserve water resources, the rationalisation of irrigation water use is necessary. An accurate estimation of the water consumed by evapotranspiration (ET) and lost by DP would provide a basis for improving irrigation efficiency (Kharrou et al., 2013; Belaqziz et al., 2014; Xianwen et al., 2016).

Regarding ET, crop coefficients and associated measurements have been reported in the literature and used to test, develop and calibrate a range of ET models (Allen et al., 2011; Er-Raki et al., 2013). DP is commonly assessed by the soil water balance equation when both ET and water supply irrigation and rainfall are available (Sammis et al., 1983). The method has been used under different irrigation techniques and for various crops (Vázquez et al., 2006; Wang et al., 2012). DP can also be measured by direct methods such as lysimeters (Allen et al., 1991; Kim et al., 2011; Duncan et al., 2016), or fluxmeters (Deurer et al., 2008; Gee et al., 2009). However, these methods are expensive (Upreti et al., 2015) and may disrupt flow, causing errors in the measured drainage (Gee et al., 2009).

Other indirect methods have also been used such as the hydraulic method (Qinbo et al., 2011; Allman et al., 2015), temperature measurements in the unsaturated zone (Constantz et al., 2003; Landon et al., 2016), and geochemical tracers (Stonestrom et al., 2003; Stephens et al., 2006).

In Morocco, citrus is one of the main components of agricultural systems (Boubker, 2004). Currently, it covers a total area of about 120,000 ha (MAPM, 2013), but it is in rapidly expanding. However, in Morocco, the key physiological stages of this crop (from flowering to maturation) coincide with the dry period (March-October). Thus, intensive irrigation is necessary for citrus development (El Hari et al., 2010). As far as we know such a study on the estimation of DP for citrus in Morocco conditions has never been performed before.
In this context, the objective of this study is twofold: 1) the evaluation and analysis of DP for citrus orchards irrigated by drip system and grown under semi-arid conditions, and 2) the exploration of FAO-56 simple approach to optimise irrigation in order to both control crop water stress and to reduce DP in the case of saline and non-saline soils. The analysis of DP losses can potentially provide useful information for optimising citrus irrigation schedules under non-saline and saline soil conditions. To our best knowledge it is the first time that the FAO-56 method has been tested for citrus under saline soil conditions.

2. MATERIALS AND METHODS

2.1. Sites description

The study was conducted on two citrus orchards: Agafay and Saada (Figure 1). The Agafay site covers an area of 38 ha approximately 44 km southwest of Marrakech city (31°29' 50.19"N, 008°25' 02"W). The field experiment was carried out from 2006 to 2013 in a mandarin orchard planted in July 2000. The trees were planted at a spacing of 4 m in rows and 6 m between rows, which is about 35% ground cover. The height of trees was about 3 m and the depth of the root zone around 0.6 m. This depth was determined by making five pits near to the tree's root zone. The crop was maintained in over-irrigated conditions by drip irrigation: the irrigation frequency was almost every day without taking into account rainfall events. Moreover, the amount of water applied by the farmers during each irrigation event varied between 2 and 9 mm day\(^{-1}\) depending on climatic conditions. The Agafay site is divided into three sectors and irrigated within 24 hours at a rate varying from 28 to 60 m\(^3\) ha\(^{-1}\). Note that irrigation is applied during rainfall events in order to leach the soil salinity from the root zone. The soil type is homogeneous, with high sand and low clay contents (18% clay, 32% silt and 50% sand). According to the pedo-transfer function of Wosten et al. (1997) the soil moisture at field capacity (\(\theta_{fc}\)) and wilting point (\(\theta_{wp}\)) are 0.26 and 0.12 mm\(^3\) mm\(^{-3}\), which corresponds to 156 and 72 mm at the root zone at the experimental plot, respectively.

The Saada site is located approximately 15 km west of Marrakech city (31°37'36"N, 08°09'35"W). It covers an area of 128 ha and is planted with 13-year old mandarin trees; the field experiment was conducted during 2004. The orange trees were placed at a spacing of 5 m in rows and 3 m between rows, with 70% of the ground cover fraction. The average height of the trees was about 3.15 m and the depth of the root zone was about 0.6 m. The Saada citrus was maintained in over-irrigated conditions by daily irrigation varied from 2 to 5 mm, depending on climatic conditions. The orchard is divided into several sectors, but all irrigated on the same day with a rate varying from 4 to 56 m\(^3\) ha\(^{-1}\). The soils have high sand and low
clay contents (12% clay, 38% silt, and 50% sand). More details about the two sites can be found in the studies by Er-Raki et al. (2009) and Er-Raki et al. (2012).

For both sites, the farmers use the reference evapotranspiration ($ET_0$) and the values of crop coefficients ($K_c$) for citrus provided in Table 12 in FAO-56 (Allen et al. 1998) for determining the amount of irrigation as: irrigation = $K_c \times ET_0$. This can under or overestimate the irrigation water needs (Er-Raki et al. 2009). The drip irrigation system used in both sites is installed such that about 40% of the soil field is wetted corresponding to the irrigated part, while 60% of the soil field remains as dry fraction which corresponds to the non-irrigated part. These areas are differentiated by their soil salinity degrees. In Agafay the soil has an electrical conductivity greater than 4 dS m$^{-1}$ which qualifies as a highly saline soil based on Mathieu and Pieltain, (2003). By contrast, at the Saada orchard, the soil is considered non-saline (2 dS m$^{-1}$) (Sefiani et al., 2017).

The climate of these areas is typically Mediterranean. It is characterised by low and irregular rainfall with an annual average of about 240 mm (Khabba et al., 2013). For the period 1972-2012, the average temperature is high in summer (37 °C) and low in winter (5 °C) (Kharrou et al., 2013). The rainy season is commonly from November to May with a maximum rainfall in November and February. The dry season is about five months, from May to September.

2.2. Experimental Data

Meteorological parameters (rainfall, temperature, humidity, wind speed and direction and solar radiation) were measured by an automatic weather station installed at 2 m above over-irrigated clipped grass. Half-hourly measurements of these parameters are obtained by monitoring wind speed and direction using an anemometer A100R (R. M. Young Company, USA), air temperature and humidity using HMP45AC (Vaisala Oyj, Helsinki, Finland), incoming solar radiation using radiometer (Kipp & Zonen CNR1, Netherlands) and rainfall using a rain gauge. Daily average values of meteorological data were calculated from the half hourly values in order to compute the daily $ET_0$ (mm day$^{-1}$), according to the FAO-56 Penman–Monteith (Allen et al. 1998). The temporal evolution of daily $ET_0$ (Figure 2) is typical for the semi-arid climate; low values in winter (1-2 mm), high values in summer (6-8 mm) and an annual average of about 1600 mm. For Agafay orchard, the study concerns two years contrasting in their rainfall; 2007 and 2010 are considered as dry (148.3 mm) and rainy (342.6 mm) years respectively.

For the two sites Agafay and Saada, actual crop evapotranspiration ($ET_{c,act}$) was measured by an eddy-covariance system. It consists of a 3D sonic anemometer (CSAT3, Campbell...
Scientific Ltd.) and a fast response hygrometer (Campbell Scientific Inc., USA). Raw data were sampled at a rate of 20 Hz and were recorded using data loggers (CR5000, Campbell Scientific Ltd). The size of both fields was large enough to meet the required fetch conditions for the eddy covariance system. The reliability of eddy covariance measurements in our studied sites was assessed by analysing the energy balance closure. Indeed, the obtained daily errors of this balance closure were generally less than 10% of available energy (Er-Raki et al., 2012), which can be considered acceptable for the eddy covariance measurements over the tree-orchards (Ezzahar et al., 2007; Er-Raki et al., 2009; Er-Raki et al., 2012).

In addition, soil moisture content was measured by 12 Time Domain Reflectometry sensors (TDR) (CS616, Campbell Scientific Ltd.). The TDR are installed in two different locations: 6 in the non-irrigated part between the rows and 6 in irrigated part under tree canopies, at depths of 5, 10, 20, 30, 40, and 60 cm, in order to measure the soil water content in different conditions. Measurements were taken at 1 Hz, and 30 min averages stored data loggers (CR23X, Campbell Scientific Ltd.). Finally, a fluxmeter for direct measurement of DP was installed in the Agafay site at a depth of 80 cm, beneath the root zone. Note that the fluxmeter was not installed in the Saada site.

2.3. Methodology

In our study, the deep percolation (DP) losses are defined as the water amounts flowing downwards below root zone. They are evaluated for the two studied sites by using the following soil water balance equation:

\[ DP = P + I - \text{ET}_{c,\text{act}} - R + CR \pm \Delta W \]  

(1)

where DP, P, I, ET, \text{ET}_{c,\text{act}}, R and CR are deep percolation, precipitation, irrigation, actual crop evapotranspiration, runoff and capillary rise from water table respectively. These terms are all in mm.

Because the studied sites are both flat and the precipitations are not heavy, R is neglected in the water balance equation. Furthermore, as the water table is deep, depth varied from 15 to 80 m (Boukhari et al., 2015), CR was considered to be zero.

\[ \Delta W \] is the variation of soil water content in the root zone, defined as:

\[ \Delta W = W(t) - W(t - 1) \]  

(2)

where W (t) is the water storage at time t in the root zone derived from the layer-wise soil moisture values (\(\theta\)), measured by TDR. The value of \(\Delta W\) (eq. 2) is positive when water is
added to the root zone; otherwise it is null or negative. At a given moment, water storage $W(t)$ was computed from the values of $\theta_i$ ($i = 1, \ldots, 6$) and the thickness of each layer ($\delta Z$, in mm) as:

$$W(t) = \sum_{i=1}^{6} (\delta Z_i \cdot \theta_i)$$

(3)

The relative deep percolation ($D_P^R$) as a percentage of the rate of irrigation or total water supply (irrigation and rainfall) is expressed respectively as follows:

$$D_P^{R(I)} = \frac{DP}{I} \cdot 100$$

(4)

$$D_P^{R(I+P)} = \frac{DP}{I+P} \cdot 100$$

(5)

The calculated DP losses and the actual allocated amounts of drip irrigation are analysed at different time scales, especially for different growing stages: induction, flowering, fruit set, fruit drop, fruit growth, slowdown growth, maturation and harvest (El Hari, 1992). Note that this analysis was repeated for two years, 2007 and 2010, in Agafay and during 2004 in Saada site where the measurements are available. According to Bouazzama and Bahiri, (2008) and Domingo et al., (2007), one can note that high amount of irrigation is critical in fruit growth phase (June -October) to increase juice content and fruit size. The water stress during the fruit drop (May-June) decreases the number of fruit per tree, and this is why water availability in the soil in this period is important to reduce fruitlet fall (Bouazzama and Bahiri, 2008). Water stress can also influence fruit quality such as acidity in maturation stage, and reduce fruit numbers in flowering and fruit set phases (Lado et al., 2014; Käthner et al., 2017). Bellvert et al. (2016) has investigated the evaluation of water stress throughout different growing seasons for several fruit tree species by using remotely-sensed indicators.

Additionally, in order to assess the adequacy of the water supply at different growing stages, two indicators are used: depleted fraction (DF) and relative evapotranspiration (RET) indices (Eq. 4). DF represents the part of the water supply ($I+P$) that is consumed by the standard evapotranspiration ($ET_c$), whereas RET allows assessing the occurrence of water deficits (Bos et al., 2005; Kharrou et al., 2013) defined as ratio of actual crop evapotranspiration ($ET_{c,act}$) and ($ET_c$):

$$DF = \frac{ET_c}{I+P} \quad \text{and} \quad \text{RET} = \frac{ET_{c,act}}{ET_c}$$

(6)

The value of $ET_c$ is estimated using the FAO-56 simple approach (Allen et al., 1998) as the product of crop coefficient ($Kc$) and $ET_0$. The value of $Kc$ for citrus is taken as 0.65 which is
the average value of the ratio between measured (ET_{c,act}) by eddy covariance system and ET_{0}.

The obtained value of Kc is corroborated by the work of Er-Raki et al. (2009) when they calibrated Kc values for citrus in the same region.

For orange trees in arid and semi-arid areas, the critical value of DF is equal to 0.6 (Kharrou et al., 2013). Values of DF between 0.6 and 1.1 are assumed to have no significant effect on growth and yield.

The RET index allows the occurrence of water deficits to be assessed (Roerink et al., 1997; Bos et al., 2005), with an acceptable range from 0.75 to 1 (Roerink et al., 1997). The critical value of RET, which is taken to be 0.75, corresponds to the economic threshold suggesting it is reached when water stress has caused a 25% decrease in crop ET. This threshold can be considered acceptable for irrigated agriculture if it does not lead to meaningful quality and quantity losses for farmers, whereas RET values that are lower than 0.75 are considered to involve a water stress that affects the agricultural development of the crop.

In this study, the combined analysis of RET and DF indicators is used to identify how irrigation water management allows better crop development through the reduction of water stress (Kharrou et al., 2013).

Thus, the diagram (RET, DF) allows the identification of four zones (Kharrou et al., 2013):

Zone A: “farmer satisfaction”, there is no water stress but irrigation is excessive,

Zone B: “water manager’s task”, there is no water stress and the irrigation is adequate.

Zone C: “risk”, there are water stress and excessive irrigation.

Zone D: “survival”, water stress is induced by a wrong irrigation scheduling.

In order to minimise DP, the FAO-56 approach is used to predict water requirement on a daily basis, by calculating the right amount of irrigation (I) without recording any water stress (Ks) and no losses by DP:

\[ I(i) = DP(i) - P(i) + ETc(i) + Dr(i-1) \quad \text{with} \quad DP=0 \]  \hspace{1cm} (7)

\[ Dr(i) = Dr(i-1) - P(i) - I(i) + ETc(i) + Dp(i) \]  \hspace{1cm} (8)

\[ Dr(i-1) = 1000 \times (\theta_{fc} - \theta_{i-1}) \times Zr \]  \hspace{1cm} (9)

where i is the number of the day, θ_{fc} is the soil moisture content at field capacity (m³ m⁻³), θ_{i-1} is the average soil moisture content in effective root zone (m³ m⁻³), and Zr is the rooting depth (m). Dr is the root zone depletion (mm), which measures the difference between the total available water (TAW) and the actual available water. At field capacity, Dr is equal to zero.
and no water stress occurs (Eq. 10). When soil water is extracted by evapotranspiration, Dr increases until it exceeds the readily available water (RAW), and water stress will be induced (Eq. 11).

Water stress coefficient (Ks) is expressed as (Allen et al., 1998):

\[ \text{Dr}(i) \leq \text{RAW} \quad \text{Ks} = 1 \tag{10} \]
\[ \text{Dr}(i) > \text{RAW} \quad \text{Ks} = \frac{TAW - \text{Dr}(i)}{TAW - \text{RAW}} \tag{11} \]

\[ \text{TAW} = 1000 \ast (\theta_{tc} - \theta_{wp}) \ast Zr \tag{12} \]
\[ \text{RAW} = p \ast \text{TAW} \tag{13} \]

RAW and TAW are the readily available water and the total available water in the root zone (mm), p is the depletion fraction which is equal to 0.5 for citrus orchards according to FAO-56 (Table 22, Allen et al., 1998). The TAW depends on the type of soil and the rooting depth (Kelly et al., 2010; Troy et al., 2013). Recently, Rosa et al., (2016) adjusted both TAW and RAW under saline conditions, for evapotranspiration partitioning of maize and sweet sorghum by applying the SIMDualKc approach.

In the case of saline soil, such as the Agafay site, the leaching of salinity is essential for crop growth, development and yield (Visconti et al., 2012). Salt leaching requires adequate irrigation management, which is based on adding sufficient amounts of water beyond the crop water requirement for evapotranspiration and photosynthesis (Russo et al., 2009). Salt is continually added to soils when the irrigation water salinity is higher than tolerable water salinity value (Naidu et al., 1996; Yoseph and Jim, 2004) which varies for citrus from 0.75 to 2.25 dS m\(^{-1}\) (Richards, 1954).

Salts in the soil can reduce evapotranspiration by making soil water less "available" for plant root extraction (Allen et al., 1998). Salinity stress occurs when the salt concentration, evaluated by the electrical conductivity of the saturated-soil-paste extract (ECe), is higher than a given concentration threshold ECe\(_{\text{threshold}}\) equal to 1.7 dS m\(^{-1}\). In this case we estimate salinity effect on evapotranspiration reduction by varying the soil electric conductivity from non-saline condition (1.7 dS m\(^{-1}\)) to the high salinity (10 dS m\(^{-1}\)), using the following formulae (Allen et al., 1998):

When salinity stress occurs without water stress: \(\text{Dr} < \text{RAW} \)

\[ Ks_{\text{salinity}} = 1 - \frac{b}{(Ky+100)} \ast (\text{ECe} - \text{ECe}_{\text{threshold}}) \tag{14} \]
When soil water stress occurs in addition to salinity stress: Dr > RAW, the total stress ($K_{\text{tot}}$) can be given by:

$$K_{\text{tot}} = \left(1 - \frac{b}{(K_y+100)} \right) \times (EC_e - EC_{\text{threshold}}) \times \left( \frac{TAW-Dr}{TAW-RAW} \right) = K_{\text{salinity}} \times K_s$$  \hspace{1cm} (15)

where $b$ is the percent yield reduction per unit increase in EC$_e$ (dS m$^{-1}$), and $K_y$ is the crop yield response factor being equal to 1.2 (Table 24, Allen et al., 1998).

Citrus yields decrease by about 16% for each 1.0 dS m$^{-1}$ over the EC$_{\text{threshold}}$ (Visconti et al., 2012; FAO, 2003). The crop yield response factor was estimated according to Allen et al., (1998) as:

$$1 - \frac{Y_r}{Y_m} = K_y \times (1 - \frac{ET_{c,act}}{ET_c})$$  \hspace{1cm} (16)

$$\frac{Y_r}{Y_m} = 1 - (EC_e - EC_{\text{threshold}}) \times \frac{b}{100}$$  \hspace{1cm} (17)

where $Y_r$ and $Y_m$ are the real and maximum yields (kg ha$^{-1}$), respectively.

The amount of water applied to wash out excess salts from the root zone is known as the leaching fraction (LF) (Plaut et al., 2013; FAO, 2003) and is commonly expressed using the following relationship (Ayers and Westcot, 1985):

$$LF = \frac{EC_{iw}}{5 \times EC_{\text{threshold}} - EC_{iw}}$$  \hspace{1cm} (18)

where $EC_{iw}$ is electrical conductivity of irrigation water (dS m$^{-1}$).

### 3. Results and Discussion

#### 3.1. Evaluation of deep percolation (DP)

##### 3.1.1. DP estimations

The value of DP is estimated by using water balance equation and fluxmeter. For the three years of this study, the annual ET$_{c,act}$ of citrus varied between 786.3 and 818 mm (Table 1). These values are similar to those reported for citrus orchards in the Haouz region (El Hari et al., 2010; Er-Raki et al., 2009, 2012). The difference between the annual ET$_{c,act}$ and rainfall was 490 mm year$^{-1}$ in Saada, 606 mm year$^{-1}$ (2007) and 476 mm year$^{-1}$ (2010) in Agafay (Table 1), with an average value of 524 mm year$^{-1}$. This large gap makes irrigation very critical for citrus growth and yield. However, using the water balance equation shows that the generated DP is very high, varying from 420.3, 454.8 to 708.6 mm year$^{-1}$ with an average of...
about 501 mm year\(^{-1}\). These correspond to a DP\(_{R(I+P)}\) of 38.3, 49.3 and 46.3\%, with an average value of 43\%.

For more analysis of DP, we split the whole season into dry and wet periods. During the dry period, from June to October, about 59.8\% (2004) of irrigation was supplied in Saada, and 62.8\% (2007) and 67.3\% (2010) in Agafay. This generates respectively DP rates of about 270.5, 240 and 417.5 mm dry-period\(^{-1}\) as well as DP\(_{R(I)}\) of 24.5\%, 34.7\% and 48.1\% (Table 2).

During the wet period (November-May), the DP is relatively lower, recording an average of about 284 mm, varying from 279.8, 216.7 to 355.5 mm wet-period\(^{-1}\), respectively, which represents DP\(_{R(I)}\) values of about 46.1, 41.9 and 53.4\% (Table 2). According to those results, the annual DP values are about 500 mm, corresponding to a DP\(_{R(I+P)}\) of 43\% and equivalent to almost double the habitual annual rainfall in the study region (250 mm).

Several studies have estimated DP using water balance method in arid and semi-arid climates by recording annual values of DP\(_{R(I+P)}\) of about 11\% (García and Castel, 2007). Compared to this value, our two citrus orchards presented high DP values. This result does not affect the validity of the drip irrigation technique, which is known by its high efficiency ranged between 80 and 90\% (Boman, 2002; Kelly et al., 2010). The observed losses are rather due to the combined effects of an inadequate use of this technique and over-irrigation.

Furthermore, direct measurement of DP was performed in 2010 by a fluxmeter installed just under the root zone in the Agafay orchard. Figure 3 shows the cumulative deep percolation measured by the fluxmeter (DP\(_f\)) and the one calculated by water balance (DP). The evolution of the cumulative values of DP\(_f\) is systematically lower than that of DP. However, the annual values of DP and DP\(_f\) are similar; 708.6 and 638.9 mm for DP and DP\(_f\) respectively. This difference could be explained by two reasons:

- The scale used by both methods, since the water balance method calculates DP at parcel scale, while the fluxmeter measures DP at local scale;
- The rainfall interception and runoff, which are not taken into account by the water balance equation (Eq.1). As reported by García and Castel, (2007), rainfall interception by trees is an important component of the water balance as it could present 8\% of rainfall.

After a preliminary evaluation of deep percolation by using direct measurement (fluxmeter) and soil water balance, the question left to address is to determine the most important factors such as soil water content, physiological and management responsible on the DP losses.

### 3.1.2. Soil water content and DP losses
The evaluation of DP losses is assessed by studying the variation of soil moisture content ($\theta_i$) in the root zone. For bare soil between the rows in the un-irrigated part of the field, Figure 4 presents an example of the temporal evolution of the soil moisture in 60-80 cm depth (lower limit of root zone). In this part, which covers about 60% of the bare soil, the average soil moisture in the root zone varies between the wilting point ($\theta_{wp} = 0.12$ mm$^3$ mm$^{-3}$) and the field capacity ($\theta_c = 0.26$ mm$^3$ mm$^{-3}$). DP takes place only on rainy days when the soil moisture exceeds $\theta_c$.

For the irrigated part of the field along tree rows in Agafay and Saada sites, the soil moisture in 60-80 cm depth is frequently above $\theta_c$ (Figure 5). Therefore, DP losses were occurring almost throughout both dry and wet periods due to the high frequency of irrigation events.

### 3.1.3. DP as a function of growth stages

Knowing that some part of water supply can be used by the plant and other part can be lost by direct soil evaporation and/or by DP, it is crucial to estimate the amount of water lost by DP which is the main objective of this present study. The supplied water is split into crop water use ($ET_{c,act}$) and DP in each crop stage. Figure 6 illustrates an example of irrigation amount according to the phenological phases. The results show that the higher irrigation values are recorded during fruit growth stage, occurring in summer. In this stage, irrigation were 552 mm and 582.4 mm in 2007 and 2010 in Agafay and 239.7 mm in 2004 in Saada site, representing about 54.4%, 50.52% and 33.4% of annual irrigation. In descending order, the irrigation rate of the other phases at Agafay (2007, 2010) and Saada (2004) are: 16.2%, 17.2%, 19.7% for fruit drop phase; 9.6%, 10.7%, and 10.4% for maturation, 9.6%, 9.9%, 9.7% for flowering and fruit set; 2.4%, 2.7%, 3.3% for induction; and 2%, 1.5%, 2.9% for harvest.

The higher values of DP losses which equal 200 mm, with $DP_{R(0)}$ of about 35%, are recorded during fruit growth stage (Figure 6). Though this is an important growth period and water demand is high ($ET_{c,act} \approx 390$ mm), the irrigation in this stage is largely overdosed (about 560 mm). During the other phases, the DP quantities (mm) are relatively lower, but still important by comparison to the irrigation amounts (Figure 7). This explains that a major part of supplied water is lost by DP when the phenological activities and evapotranspiration are low.

### 3.1.4. Evaluation of water supply adequacy

The evaluation of water supply adequacy by using the diagram RET-DF shows that all phenological phases are in zone A and B of Figure 8, which is consistent with the above findings (Figure 7):
Zone A includes the phases characterised by excess irrigation, recording a DF value less than 0.6. For Agafay, these phases are fruit growth, slowdown growth, maturation and harvest for 2007 and 2010 seasons. In Saada case, the phases were flowering, maturation and harvest. This results shows that the orange trees were largely over-irrigated. Consequently, these phases need removed to zone B.

Zone B includes the phases characterised by an adequate use of irrigation water, recording DF values between 0.6 and 1.1. These phases are induction, flowering, fruit set, and fruit drop for Agafay, and induction, fruit drop, fruit growth, slowdown growth and fruit set, for Saada.

According to these results, the Saada site is characterised by more adequate use of irrigation during the majority of its growth stages than the Agafay site. This is expected since the Agafay site has a high level of soil salinity that needs more water for salt leaching and then more DP losses. This aspect is discussed further in the following section.

### 3.2. Impact of rainfall and soil salinity on DP

#### 3.2.1. Effects of rainfall on DP losses

To analyse the effective role of rainfall in citrus water supply and its impact on DP losses, we compare the data of Agafay recorded in the dry year 2007 (148 mm) and the rainy year 2010 (343 mm). Despite the high rainfall in 2010, the supplied irrigation (1188 mm) was higher than 2007 (1061 mm). Consequently the DP losses are greater (708.6 mm) in 2010 than in 2007 (454.8 mm) (Table 1). Even in the wet year of 2010 the drip irrigation alone exceeds the ET$_{c,act}$ by 31% (Figure 9).

At a weekly scale, the farmers supplied irrigations in five weeks in 2007 and nine in 2010 (examples giving in Table 3), although the rainfall amounts were adequate to fulfil crop water need (ET$_{c,act}$). Such behaviour attests to the inappropriate application of irrigation: the wrong quantity is delivered at the wrong moment. Consequently, the DP reached high values, with an average of 34.6 mm week$^{-1}$ (Figure 9). The way irrigation is applied by the farmer was not appropriate for controlling DP losses.

#### 3.2.2. Impacts on DP of measures to control soil salinity

In Agafay site, the high soil salinity (about 4 dS m$^{-1}$) could explain the applied over-irrigation. In this case, the irrigation water is used in excess for leaching soil salts in order to avoid root dieback and leaf loss of citrus (Sheng et al., 2002; Ayars et al., 2012). For Agafay orange, the irrigation water exceeds ET$_{c,act}$ by about 306.8 and 370.5 mm in dry and wet years,
respectively. These DP amounts correspond to about 29 and 31% of irrigation water, respectively. However, Robert and Richard (1999) reported that only 17% irrigation excess is needed for salt leaching. Also Barnard et al., (2010) and Plaut et al. (2013) found that in sandy loam soil, similar to the Agafay site, 20% of irrigation excess is sufficient to leach salts from the root zone. This result shows that, in Agafay orchard, the applied water quantities are still higher than is needed to address the salinity issue.

3.3. Irrigation water management to control DP losses and crop water stress

The optimisation of irrigation water scheduling, in time and quantity, was performed by using FAO-56 simple approach for the two orange sites in order to avoid both water stress and deep percolation (i.e. Ks=1 and DP=0 at all times). Based on the equation (7), the obtained results show that annual values of recommended irrigations without taking into account the salinity issue are 685 and 530 mm for Agafay (2007 and 2010) and 558 mm for Saada. By adding the amount of rainfall (149, 343 and 296 mm, respectively), the total supplied coincides with the adequate citrus water requirement (845 mm year\(^{-1}\) = Kc * ET\(_0\) = 0.65 * 1300 mm year\(^{-1}\)) (FAO, 2003; Er-Raki et al., 2009; El Hari et al., 2010). The recommended irrigation can then save approximately 39 and 45% of the irrigation in Saada and Agafay, respectively.

As the Agafay site is under high soil salinity, it is of interest to quantify the effect of this parameter on the stress and on the amount of irrigation needed. The total stress (Ks\(_{tot}\)) increases (Figure 10) when soil electrical conductivity exceeds the tolerable salinity threshold of citrus (1.7 dS m\(^{-1}\)). In this case, the impact of salinity is remedied by adding an additional irrigation to the recommended amount. This supplemental irrigation is used for salinity leaching and not used by the crop, and then lost by DP. The question addressed is how much amount of water should be used as a leaching fraction (LF).

Based on equation (18), when the soil electrical conductivity is 2.5 dS m\(^{-1}\), the leaching fraction is 17% of water supply. This fraction equals an additional irrigation of about 116 and 90 mm year\(^{-1}\) for 2007 and 2010, respectively. By considering these amounts, the annual recommended irrigation needed to avoid water and salinity stress ranges between 766 mm (2007) and 624 mm (2010).

Taking into account the water and salinity stress and DP losses, the recommended irrigations in Saada and Agafay allowed us to plot all the growing phases in zone B with DF=1. This is likely to ensure an effective irrigation strategy for optimising citrus irrigation schedules, while avoiding water stress and DP losses. The recommended irrigation can then save approximately 39 and 37% of the irrigation in Saada and Agafay, respectively. Saving
irrigation water in such proportions is equivalent to save 3520 m³ ha⁻¹ year⁻¹ in Saada and 2950 m³ ha⁻¹ year⁻¹ (2007) to 5640 m³ ha⁻¹ year⁻¹ (2010) in Agafay. The average quantity of irrigation water that could be saved is about 4295 m³ ha⁻¹ year⁻¹. If we consider 6000 ha of citrus in the Haouz plain cultivated and irrigated in the same conditions of Agafay site, the overall saved amount of water would be 25.8 × 10⁶ m³ year⁻¹.

4. Conclusion

The paper investigates deep percolation (DP) in citrus orchards with drip irrigation under semi-arid climate and develops a method based on the FAO-56 model to define irrigation schemes and to optimise irrigation in non-saline and saline soil conditions. The results obtained show that, under the irrigation conditions of the study, the DP calculated by the water balance equation is very high; varying from 420 to 709 mm year⁻¹ with a relative DP_{R(I)} value of about 38.3 and 49.2% for non-saline and saline soils, respectively. Direct measurements of DP confirmed these estimations by recording a DP of about 638.9 mm year⁻¹. This is in accordance with high values of root zone moisture, which almost exceed the soil moisture at field capacity. The evolution of DP across the phenological phases has given additional information on the higher value of DP losses of about 200 mm with DP_{R(I)} of 35% recorded during the fruit growth stage.

The FAO-56 simple approach was used to assess the appropriate irrigation amount in case of saline and non-saline soils. The results shows that, by taking into account the rainfall, this model recommends an amount of irrigation much lower than that actually applied by the farmers. Following the model simulations, it seems possible to save about 39% of water supply at the Saada site, with non-saline soil, and about 30% to 47% at the Agafay site, with saline soil. This study has demonstrated that a reasonable drip irrigation scheduling is necessary for water saving.

As a main perspective of the study, the evaluation of DP losses would be important information to estimate groundwater recharge beneath the irrigated fields. However, regarding the heterogeneous lithology of the Haouz plain and its deep groundwater, the DP contribution to the aquifer recharge remains an important scientific issue.

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References


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Figure captions

Figure 1: Location of the study sites Agafay and Saada.

Figure 2: Daily reference evapotranspiration (dotted line: ET<sub>0</sub>) calculated following the FAO–Penman–Monteith equation and precipitation events (vertical bars: P) in the study orchards: Agafay during 2007 (a), 2010 (b) and Saada during 2004 (c).

Figure 3: Cumulative deep percolation measured by fluxmeter (solid line: DP<sub>f</sub>) and deep percolation calculated by water balance (dashed line: DP) in Agafay for 2010.

Figure 4: Daily rainfall (vertical bars: P), deep percolation (dashed line: DP) and soil moisture (continuous line: θ<sub>f</sub>) for bare soil in depth 60-80 cm (example of Agafay site, 2007). The upper and the lower horizontal continuous lines corresponds to the soil moisture at field capacity (θ<sub>fc</sub>) and at wilting point (θ<sub>wp</sub>), respectively. The areas of soil moisture variation above θ<sub>fc</sub>, between θ<sub>fc</sub> and θ<sub>wp</sub> and below θ<sub>wp</sub> correspond to the deep percolation, available soil water and unavailable soil water zones, respectively.

Figure 5: Daily water supply (vertical bars: I+P), deep percolation (dashed line: DP) and soil moisture (continuous line: θ<sub>f</sub>) for irrigated part in depth 60-80 cm (example of Agafay site, 2010). The upper and the lower horizontal continuous lines corresponds to the soil moisture at field capacity (θ<sub>fc</sub>) and at wilting point (θ<sub>wp</sub>), respectively. The areas of soil moisture variation above θ<sub>fc</sub>, between θ<sub>fc</sub> and θ<sub>wp</sub> and below θ<sub>wp</sub> correspond to the deep percolation, available soil water and unavailable soil water zones, respectively.

Figure 6: Variation of deep percolation (dashed line: DP), actual crop evapotranspiration (—ET<sub>c_act</sub>—) and irrigation (vertical bars: I) according to the phenological stages during 2007 in Agafay.

Figure 7: Relative deep percolation evolution according to the phenological stages during (grey vertical bars: 2007) and (white vertical bars: 2010) in Agafay and (black vertical bars: 2004) in Saada.

Figure 8: Combined analysis of relative evapotranspiration (RET) and depleted fraction (DF) for all phenological phases: (♦: induction), (■: flowering), (▲: Fruit set), (△: Fruit drop), (○: Fruit growth), (●: slowdown growth), (□: maturation) and (◇: harvest) for Agafay (2007: (a)), (2010: (b)) and Saada in (2004: (c)).

Figure 9: Weekly evolution of precipitation (black vertical bars: P), irrigation (grey vertical bars: I), actual crop evapotranspiration (♦: ET<sub>c_act</sub>) and deep percolation (dashed line: DP) during (2007: (a)) and (2010: (b)) in Agafay.

Figure 10: Effect of the variation of soil electric conductivity (Ece) on the total stress coefficient (dashed line: K<sub>s,tot</sub>). The required irrigation for each value of Ece is also shown (♦). The horizontal continuous line corresponds to the amount of recommended irrigation without salinity issue.
Figure 1
Figure 2
Figure 3
Figure 4
Figure 5
I, ET$_{act}$, DP (mm stage$^{-1}$)

Phenological stages

- Induction
- Flowering
- Fruit set
- Fruit drop
- Slowdown growth
- Maturation
- Harvest

Figure 6
Figure 7

Phenological stages

Induction  Flowering  Fruit set  Fruit drop  Fruit growth  Slowdown growth  Maturation  Harvest

$DP_{RLI}$ (%)
Figure 8
Figure 9
Figure 10

Total stress coefficient ($K_{s_{tot}}$)
Electric conductivity $E_{ce}$ (dS m$^{-1}$)
Irrigation required (mm)
Table 1: Annual values of actual crop evapotranspiration ($ET_{c,act}$), rainfall ($P$), Irrigation ($I$), deep percolation ($DP$) and relative deep percolation ($DP_{R}$) in non-saline (Saada) and saline soil (Agafay) sites.

<table>
<thead>
<tr>
<th>orchards</th>
<th>years</th>
<th>$ET_{c,act}$ (mm)</th>
<th>$P$ (mm)</th>
<th>$I$ (mm)</th>
<th>DP (mm)</th>
<th>$DP_{R(I+P)}$ (%)*</th>
</tr>
</thead>
<tbody>
<tr>
<td>Saada</td>
<td>2004</td>
<td>786.3</td>
<td>295.8</td>
<td>910.4</td>
<td>420.3</td>
<td>38.3</td>
</tr>
<tr>
<td>Agafay</td>
<td>2007</td>
<td>754.8</td>
<td>148.3</td>
<td>1061.3</td>
<td>454.8</td>
<td>49.3</td>
</tr>
<tr>
<td></td>
<td>2010</td>
<td>818.1</td>
<td>342.6</td>
<td>1188.6</td>
<td>708.6</td>
<td>46.3</td>
</tr>
<tr>
<td>Average</td>
<td></td>
<td>786.4</td>
<td>245.4</td>
<td>1124.9</td>
<td>581.7</td>
<td>47.8</td>
</tr>
<tr>
<td>Overall</td>
<td>Average</td>
<td>786.4</td>
<td>270.6</td>
<td>1017.7</td>
<td>501</td>
<td>43</td>
</tr>
</tbody>
</table>

* $DP_{R(I+P)}$ is calculated in daily basis

Table 2: Seasonal values of irrigation ($I$), deep percolation ($DP$) and relative deep percolation ($DP_{R}$) in non-saline (Saada) and saline soil (Agafay) orchards.

<table>
<thead>
<tr>
<th>orchards</th>
<th>Variables</th>
<th>saada</th>
<th>Agafay</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Years</td>
<td>2004</td>
<td>2007</td>
</tr>
<tr>
<td>season</td>
<td>I (mm)</td>
<td>DP(mm)</td>
<td>$DP_{R(I)}$ (%)</td>
</tr>
<tr>
<td>wet</td>
<td>365.7</td>
<td>279.8</td>
<td>46.1</td>
</tr>
<tr>
<td>dry</td>
<td>544.7</td>
<td>270.5</td>
<td>24.6</td>
</tr>
<tr>
<td>dry</td>
<td>388.2</td>
<td>355.5</td>
<td>53.4</td>
</tr>
</tbody>
</table>

Table 3: Weekly rates of irrigation ($I$), actual crop evapotranspiration ($ET_{c,act}$), deep percolation ($DP$) and relative deep percolation ($DP_{R}$) in presence of precipitations in saline soil (Agafay).

<table>
<thead>
<tr>
<th>weeks</th>
<th>P (mm)</th>
<th>I (mm)</th>
<th>$ET_{c,act}$ (mm)</th>
<th>DP (mm)</th>
<th>$DP_{R(I+P)}$ (%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>17-23/12/2010</td>
<td>25</td>
<td>13.6</td>
<td>7.6</td>
<td>34.6</td>
<td>89</td>
</tr>
<tr>
<td>21-27/12/2007</td>
<td>6.8</td>
<td>8.6</td>
<td>6.7</td>
<td>12.8</td>
<td>83</td>
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</tbody>
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