Incision and aggradation in proglacial rivers: Post-Little Ice Age long-profile adjustments of Southern Iceland outwash plains

Erwan Roussel, Philip Marren, Etienne Cossart, Jean-Pierre Toumazet, Marie Chenet, Delphine Grancher, Vincent Jomelli

To cite this version:
Erwan Roussel, Philip Marren, Etienne Cossart, Jean-Pierre Toumazet, Marie Chenet, et al.. Incision and aggradation in proglacial rivers: Post-Little Ice Age long-profile adjustments of Southern Iceland outwash plains. Land Degradation and Development, Wiley, 2018, 10.1002/ldr.3127. hal-01897894

HAL Id: hal-01897894
https://hal.archives-ouvertes.fr/hal-01897894
Submitted on 5 Nov 2019

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L’archive ouverte pluridisciplinaire HAL, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d’enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.
INCISION AND AGGRADATION IN PROGLACIAL RIVERS:
POST-LITTLE ICE AGE LONG-PROFILE ADJUSTMENTS OF
SOUTHERN ICELAND OUTWASH PLAINS

Journal: Land Degradation & Development

Manuscript ID LDD-17-1101.R1

Wiley - Manuscript type: Special Issue Article

Date Submitted by the Author: 05-Jun-2018

Complete List of Authors: Roussel, Erwan; Universite Clermont Auvergne, GEOLAB UMR6042 CNRS
Marren, Philip; University of Chester, Department of Geography and International Development
Cossart, Etienne; Université de Lyon (Jean Moulin, Lyon 3), UMR 5600 CNRS (Environnement Ville Société)
Toumazet, Jean-Pierre; Universite Clermont Auvergne, GEOLAB UMR6042 CNRS
Chenet, Marie; Université Paris 1 Panthéon-Sorbonne, CNRS-LGP, 92195 Meudon, France, LGP
Grancher, Delphine; Université Paris 1 Panthéon-Sorbonne, CNRS-LGP, 92195 Meudon, France, LGP
Jomelli, Vincent; Université Paris 1 Panthéon-Sorbonne, CNRS-LGP, 92195 Meudon, France, LGP

Keywords: proglacial river, glacial retreat, proglacial landforms, fluvial geomorphology, paraglacial processes
Abstract:

The retreat of glaciers in response to climate warming leads to substantial changes in meltwater and sediment yield. Glacial shrinkage also induces the emergence and growth of proglacial margin landforms which strongly affect water and sedimentary transfers from the glacier to the outwash plains. On a decadal-timescale, field observations show that outwash plains of retreating glaciers typically exhibit proximal incision which decreases in magnitude downstream and stops at an inflection point where aggradation begins. Nevertheless, there is a lack of knowledge about the rates and magnitude of this fluvial adjustment and the effects of the proglacial margin configuration on the temperance or the aggravation of this fluvial adjustment to glacier retreat. This paper investigates the proglacial rivers of 14 retreating glaciers in southeast Iceland over a post-Little Ice Age timescale, combining fluvial deposits mapping, lichenometric dating and long-profile measurements of proglacial fluvial terraces. Our results demonstrate that: (1) proximal incision, associated with distal aggradation and downstream migration of the inflection point is the dominant pattern of proglacial river response to post-LIA glacier retreat in Iceland; (2) estimated mean rates of downstream migration of the inflection point range between 5 and 46 m.a\(^{-1}\); (3) the downstream migration rate of the inflection point is positively correlated with the proportion of proglacial lakes within the glacier foreland. These findings suggest that proglacial margins dominated by proglacial lakes intensify the rates of proximal incision and inflection point migration.

Keywords: proglacial river, glacial retreat, proglacial landforms, fluvial geomorphology, paraglacial processes

1. INTRODUCTION

Since the end of the Little Ice Age, rising temperatures have elevated the equilibrium lines altitude of many glacial systems and intensified ice melting (Barry, 2006; Vaughan et al., 2013). This implies an adjustment of glacier geometry, which is usually reflected by the retreat of the glacier front. Glacial retreat leads to the emergence and growth of proglacial margin landforms at the interface between the glacier and the fluvial...
system (Heckmann et al., 2016; Carrivick & Heckmann, 2017). The water discharge of proglacial rivers is largely
supplied by meltwater and they are directly connected to the glacier, such that sediment released by glacial
melt is immediately available to be carried as fluvial sediment load. The supply of both water and sediment is
highly sensitive to changes in glacier mass balance. The primary hydrological impact of glaciers on downstream
river systems is to modulate the timing and seasonality of hydrological fluxes (Röthlisberger & Lang, 1987). A
shift to deglacial conditions leads to marked changes in the seasonality of river flow (e.g., a shift of the
hydrograph to an earlier peak flow in spring months) and the amount of annual glacier runoff (Bliss et al., 2014;
Milner et al., 2017). Current trends in annual glacial runoff differ among world regions and glacier hypsometry
but there is a consensus on the typical sequence of meltwater changes during catchment deglaciation (Bliss et
al. 2014): at the beginning of deglaciation, when the glacier is still sufficiently voluminous, the glacial meltwater
tends to increase for a period lasting several decades, before decreasing when the ice stock runs out (Braun et
al., 2000; Jansson et al., 2003; Marren, 2005; Huss et al., 2008). The successive growth and decay of the glacial
meltwater supply during deglaciation periods is both supported by palaeohydraulic studies (Maizels, 1986) and
simulations of meltwater discharge (Bliss et al., 2014). In Iceland, modelling approaches predict increased
meltwater runoff until 2050. According to these simulations, glacier runoff declines afterwards (Flowers et al.,
2005; Jóhannesson et al., 2006; Bliss et al., 2014). A shift to deglacial conditions also affects the sediment
supply to proglacial rivers in three ways: (1) The erosive capacity of a glacier is a function of basal sliding such
that there is a dependence of erosion rate on the ice thickness which rely on the ice mass balance driven by
climate conditions. Thus, in the long term, the loss of glacier thickness and volume limits their own ability to
produce glacial material (Hallet et al., 1996; de Winter et al., 2012); (2) Deglaciation produces a lagged
paraglacial sediment pulse caused by the reworking of stored englacial, subglacial and marginal sediments
(Church & Ryder, 1972; Church & Slaymaker, 1989; Ballantyne, 2002; Mercier, 2008). Paraglacial sediment
supply is highest during the initiation of glacial retreat and declines as sediment sources are exhausted or
stabilized; (3) However, glacier retreat also induces the emergence and growth of proglacial margin landforms,
like moraine ridges or proglacial lakes, which trap and store sediments and decrease the sediment connectivity
in the glacier foreland (Cossart, 2008; Cossart & Fort, 2008; Heckmann & Schwanghart, 2013; Carrivick &
Tweed, 2013; Cossart & Fressard, 2017; Cordier et al., 2017, Lane et al., 2017). As a result, the downstream
sediment flux of retreating glaciers may be significantly reduced, particularly where lasting proglacial lakes
develop (Geilhausen et al., 2013; Bogen et al., 2015, Staines et al., 2015). Nevertheless, sediment connectivity
may temporarily or durably be restored by rare and extreme floods triggered by rainstorms (Marren &
Toomath, 2013; Owczarek et al., 2014), glacial outbursts (Westoby et al., 2014; Worni et al., 2014) or iceberg-
jam failures (Roussel et al., 2016).

Changes in meltwater and sediment fluxes from retreating glaciers lead to substantial geomorphic adjustments
of proglacial rivers. Understanding the behaviour of proglacial rivers during periods of rapid change of glacier
mass balance is especially important at present, given that climate change projections indicate that most
glaciated regions will undergo a rapid decrease in glacier extent over the coming century (Marzeion et al.,
2012; Radić et al., 2014), coupled with major hydrogeomorphological changes (Bliss et al., 2014; Milner et al.,
2017). Reduced sediment flux from a retreating glacier usually induces incision of the upper reach of the
proglacial river (proximal outwash plain) and the formation of a single deep channel (Fahnestock, 1969; Maizels, 1979, 1983; Germanoski & Schumm, 1993; Gurnell et al., 1999; Marren, 2005; Wilkie & Clague, 2009, Owczearek et al., 2014). In Iceland, this fluvial response has been observed on the proglacial margins of the southern glacial outlets of the Vatnajökull icecap, especially on the proglacial rivers of Skaftafellsjökull and Svínafellsjökull, (Thompson & Jones, 1986; Thompson, 1988; Marren, 2002; Roussel et al., 2008; Marren & Toomath, 2013, 2014). During the incision phase, fluvial bars tend to merge into larger sedimentary units, which are not in equilibrium with the water and sedimentary flux (Germanoski & Schumm, 1993). The incision of the proximal river results in the establishment of glacio-fluvial terraces whose morphosedimentary features reflect the former hydrological conditions before the incision (Marren, 2005). Flume experiments show that the long profile of an incising braiding river is usually characterized by an inflection point, downstream of which aggradation is dominant (Germanoski & Schumm, 1993). Indeed, the reworked material of the incising upper reach is transported and deposited downstream induces an increase of the braiding intensity of the lower reach of the river. When the sediment delivery deficit continues, flume experiments have shown that the inflection point, which separates the incision from the aggradation reach, tends to migrate downstream (Germanoski & Harvey, 1993). At the same time, glacial retreat promotes the formation of proglacial lakes that trap the coarser fraction of the sediment flux and may strengthen the proximal river incision and the aggradation of the distal river reach (Embleton & King, 1968; Chew & Ashmore, 2001; Benn et al., 2003; Gardarsson & Eliasson, 2006; Schomaker, 2010; Carrivick & Tweed, 2013; Bogen et al., 2015). Nevertheless, this pattern of fluvial response to glacier retreat can vary over space and time due to topographical constrains induced by marginal landforms: (1) the spatial and temporal evolution of the overall sediment connectivity within the proglacial margin can lead to damming and localised river aggradation (Cossart, 2008; Cossart & Fort, 2008); (2) Channel pattern changes can be controlled by moraine ridges between which the river is forced to find a course (Marren & Toomath, 2013, 2014).

Field observations report that most proglacial rivers associated with retreating glaciers in Iceland exhibit proximal incision which decreases in magnitude downstream and stops at an inflection point where aggradation begins (Marren, 2005; Roussel et al., 2008; Marren & Toomath, 2013). However, no studies have measured longitudinal profile adjustments of this type on a large sample of proglacial rivers and over a sufficiently long time period to determine whether this behaviour is a representative response to glacial retreat. Furthermore, there is a lack of information about the rates and magnitude of this fluvial adjustment and the effects of the glacial retreat rate and the proglacial margin configuration on the temperance or the aggravation of the fluvial response to glacier retreat.

Therefore, this paper investigates the proglacial landscapes and the outwash plains of 14 retreating glaciers in southeast Iceland over a post-Little Ice Age timescale. It combines quantitative assessment of the glacier retreat, the landform assemblage of the proglacial margin and the post-LIA outwash plains changes through lichenometric dating and longitudinal profile measurements of active outwash and fluvial terraces. The aims of this paper are threefold: (1) to test the representativeness of the outwash plains response to glacier retreat described above over the time-scale of centuries; (2) to assess the rates of this fluvial response through the
reconstruction of the downstream migration of the inflection point; (3) to explore the influence of the
landforms assemblage of the proglacial margin on the rate of the outwash plains response.

2. Materials and Methods

Site conditions and history

Three southern Icelandic marginal areas were selected: the Sólheimajökull area in the south of the
Mýrdalsjökull icecap, the Óræfajökull and southeastern Vatnajökull areas, both located on the southern margin
of the Vatnajökull icecap (Figure 1). These three study areas present a sample of 14 glacier tongues which feed
12 proglacial rivers (Table 1). Two pairs of formerly coalescent glaciers feed proglacial rivers which merge
within the proglacial margin (Hrútárjökull/Fjallsjökull and Skálafellsjökull/Heinabergsjökull). None of the
studied glaciers exhibit surge dynamics (Thórarinsson, 1969; Björnsson et al., 2003; Ingólfsson et al., 2016).

Three glaciers (Sólheimajökull, Kotárjökull and Heinabergsjökull) are known to be jökulhlaup outlets (Thompson
& Jones, 1986; Sigurðsson, 1998; Björnsson, 1992; Roberts et al., 2003; Russell et al., 2003, 2010). For
convenience, the investigated glaciers and outwash plains are called afterward by their ID number (Table 1)
instead of their Icelandic name. Additional information on site conditions and history is available in the
appendix.

Methods

In order to address the aims of this study, we followed three stage methodological workflow (Figure 2)
including field surveying and GIS mapping of proglacial landforms, lichenometric dating of outwash terraces,
and topographic surveys of the active proglacial rivers and outwash terraces.

Proglacial landform mapping

The proglacial margin plays an essential role in sediment connectivity and may impose sediment flux
discontinuities into the glaciofluvial continuum. Therefore, characterizing the landform assemblage of the
proglacial margins is essential to fully assess the fluvial response of outwash plains to post-LIA glacial retreat.
The mapping work is based on the recognition of 11 typical classes of proglacial landforms in the field (Figure 2-
A): Glacier front, fresh proximal till deposits (including dirty cones, hummocky debris, dead-ice with
supraglacial debris), moraine ridge, active and abandoned intra-morainic plain, intra-morainic lake, proglacial
lake, alluvial fan, rocky outcrop, active outwash and fluvioglacial terraces. Artificial dams and embankments
built in the proglacial margins were also mapped as they may act on the sediment connectivity and marginal
landforms evolution (e.g. inhibiting basal erosion of morainic ridges). LIA moraines mark the maximum glacier
snout position during the LIA and their locations are well known for most of the glaciers we investigated. They
have already been examined and mapped by several authors (Thórarinsson, 1956; Evans et al., 1999, 2017;
Bradwell, 2004; Chenet et al., 2010, Hannesdóttir et al., 2015). In the field, the locations of LIA moraines and
outwash terraces were collected using a handheld GPS (Trimble Geoexplorer XH). Post-processing of the GPS
signal enabled a planimetric precision of 1.5 m on average. The fieldwork phase ensured the robustness of the
subsequent mapping work based on the photo-interpretation of recent photographs with GIS software.

Coupling this field database to 2009 and 2012 georeferenced orthophotos (2 m resolution on average)
For Peer Review

provided by DigitalGlobe® within ArcGIS software (10.4), the landform assemblages in the proglacial margin, the outwash terraces, and the active outwash were mapped. The proglacial margin areas provide a quantitative assessment of the deglaciated surface since the LIA. An average linear retreat of the glacier fronts was computed by dividing the proglacial margin area by the glacier front length for each glacier. A manual measurement of the linear retreat of glacier fronts was also performed, along the centerline of the glacier tongues, between the maximum LIA moraines and the glacier front in 2009.

Lichenometric dating

Lichenometric dating of outwash terraces (Figure 2-B) was performed using the generalized extreme value approach (GEV) described in Jomelli et al. (2007) for the statistical treatment of the measurements of the longest axis of the thalli of Rhizocarpon Geographicum. This procedure has already been successfully applied in Iceland to date the LIA moraines of the south Vatnajökull (Chenet et al., 2010). This method provides a more statistically robust estimate of the age of the deposits because it takes into account the bias induced by the strategy of data collection (i.e. the selection of the largest thalli whose distribution is similar to the model GEV). In addition, this method provides confidence intervals to assess the quality of the estimated age.

In the field, measurement of lichen long axes was carried out according to the procedure recommended by Jomelli et al. (2007): (i) a random selection of 50 boulders distributed over the whole surface of the deposit; (ii) on each boulder, the long axis of the largest thallus of Rhizocarpon Geographicum was measured with an accuracy of 0.5 mm. Coalescing lichens and thalli with irregular shape are systematically rejected. Consistent with all lichenometric methods, two datasets of thalli long axis were collected in the field: on already dated surfaces and on the deposits requiring dating. Eight well-known dated surfaces in an area of ca. 45 km² in south foreland of Vatnajökull were used to ensure that climatic conditions do not vary excessively (Chenet et al., 2010). Dated surfaces included jökulhlaup deposits, lava flows, rockfalls, dams, and moraines; all made of basalt to ensure lithological homogeneity. Surfaces were dated using historical descriptions, aerial photographs (available since 1945), and personal communications from the Skaftafell Park staff. Additional consideration of the statistical treatments applied to the lichenometric database are available in several papers (Cooley et al., 2006; Naveau et al., 2007; Rabatel et al., 2008).

Topographic survey

Longitudinal topographic profiles of active outwash and outwash terraces were surveyed in the field. Immediately downstream of LIA moraines of each investigated glacier, valley transverse profiles were also undertaken in order to capture the nesting of outwash terraces (Figure 2-C). Following the main direction of flow, longitudinal profiles of the active outwash and terraces were surveyed from the LIA moraines to the downstream edge of the fluvial deposit (confluence with a tributary or the shoreline). Transverse topographic profiles located just downstream of LIA moraines enable the computation of the relative altitude of each outwash terraces in respect to the active outwash. They are essential to align longitudinal profiles between them and to capture the downstream migration of the inflection points separating the upstream incising reach from the downstream aggrading reach of the proglacial rivers. These profiles were achieved using a laser telemeter (Impulse 200LR from Laser Technology) with a centimeter accuracy and a tripod target. The average
distance between two measurements was 20 meters. Two measurements by stations (foresight and backsight) were carried out in order to systematically check that the difference between measurements were below the telemeter accuracy (0.03 m) and to limit error propagation. As the objective was to assess the topography of the outwash plains, measurements were carried out between topographic stations located on the heads of fluvial bars. Therefore, fine topography at the channel scale was not captured. Each topographic station was located using a handheld GPS (Trimble Geoexplorer XH).

Reconstruction of the downstream migration of the inflection point (IP)

Inflection points (sensu Germanoski & Harvey, 1993) separate upstream incising reaches from downstream aggrading reaches of a river. The reconstruction of changes in the location of the inflection point (IP) characterizes the evolution of the extent of proximal outwash affected by incision. IP changes were reconstructed based on the identification of intersection points of the longitudinal profiles of two successive (dated) terrace generations. These intersections are the former IP location separating the incision and the aggradation reach of the outwash plain at the date of the youngest of the intersecting terrace profiles (Figure 2-D). For each reconstructed IP, its location (distance from LIA moraines and elevation) was validated in the field and recorded with the handheld GPS (Trimble Geoexplorer XH). The initial IP location (distance = 0 m) at the LIA glacial maximum coincides with the contact between the LIA moraines and the oldest and highest outwash terrace (T1), if lichenometric dating of the latter is younger than LIA moraine ridges. In other words, at the LIA glacial maximum, aggradation of the outwash is likely predominant and there is no upstream incision because there is no outwash upstream to the LIA moraines. The reconstruction of the former location of the inflection point when T1 was incised and the IP location observed in 2007 enabled the computation of average rates of longitudinal migration (increasing distance of the IP from the LIA moraines) and altitudinal lowering (decrease in the IP elevation) of the IP through time. A synthetic metric of average rate of downstream migration of the inflection point was also computed as the Euclidean norm of the vector (square root of the sum of the square of coordinate values) defined by the two values of longitudinal migration and altitudinal lowering of the inflection points. Note that this average rate of IP downstream migration does not necessarily reflect a gradual progression of the inflection point as IP migration may occur suddenly during extreme events. Finally, a stepwise multiple linear regression was performed in order to test the effect of proglacial landform assemblage on the downstream migration rates of the inflection point.

3. RESULTS

Proglacial landform assemblages

The proportion of each landform unit within the post-LIA proglacial margin has been quantified (Table 2 and Figure 3). Glacigenic deposits, combining moraine ridges and fresh proximal till deposits, are the most abundant proglacial landforms (respectively 44.8 % and 5.2% on average of the investigated proglacial margins). The marginal landscape is obviously dominated by glacial deposits associated with seasonal fluctuations of glacier fronts and the continuous re-advances of the 1970s and 1980s. Proglacial margins of the steepest, Öraefajökull glaciers (glaciers 2 to 10), show a higher proportion of moraine deposits. Steep proximal
topography is associated with a large amounts of moraine deposits that may be related to large production rates of glaciogenic sediments and/or limited fluvial reworking. Abandoned and active intra-morainic plains form respectively around 16.8 % and 8.9 % on average of the proglacial landforms. The abundance of these two types of landforms testify the involvement of fluvial processes in the morphogenesis of the proglacial margin. Rocky outcrops (11.4% on average) are only present in the proglacial margins of piedmont glaciers (Glaciers 1, 11, 12 and 13) whose tendency to overdeepening would generally exhume more resistant areas of the glacial bed, and in the foreland of glacier 6 where steep proximal relief tends to rapidly clean the glacial bed and remove till layers via gravity and runoff processes. Proglacial and intra-morainic lakes (respectively 10.4% and 1.5% on average) are totally absent from the proglacial margins of the steepest glaciers of the Öræfajökull sector (glaciers 6 and 7), most likely due to the extreme steepness of the ablation zone and the glacier forefield. Alternatively, proglacial lakes are common on the margin of piedmont glaciers of the southeastern sector of Vatnajökull where their large frontal lobes probably favored overdeepening of the marginal areas during the LIA glacier advance. Alluvial fans (0.6 % on average) are only present in steep marginal settings (glaciers 7 and 9). This association seems logical since development of alluvial fans is strongly dependent on the presence of steep marginal slopes. Embankments (0.3 % on average) are only present on the margins of two piedmont glaciers (glaciers 3 and 13). Artificial embankments are associated either with tourist trails running through the proglacial margin (in the case of glacier 13), or with consolidation structures of the main Icelandic road (road number 1) built through the LIA moraines (glacier 3).

**Outwash terrace mapping and dating**

Outwash terrace mapping and lichenometric dating (Figure 3) emphasizes four types of fluvial processes affecting the proglacial river responses to post-LIA glacier retreat: (1) post-LIA transient aggradation; (2) incision and contraction of the proximal outwash; (3) shifts in the location of the fluvial activity due to stream capture occurring within the proglacial margin and (4) outwash incision prior to LIA glacial maximum. Based on the analysis of glacier foreland mapping and lichenometric dating of outwash terraces, an overview of successive aggrading and incising stages of proximal outwash response to post-LIA glacial retreat has been produced (Figure 4). The outwash terraces of the proglacial rivers 6 and 7 could not be dated due to the absence of lichens: the oldest terrace (T1) of the outwash plain 6 is the historic jökulhlaup deposit of 1727 and is exclusively composed of fine materials devoid of lichens. The terrace (T1) of the outwash plain 7 is lightly vegetated and its abandonment by proglacial flows might be recent and artificial, relating to the upstream embankment structure (Figure 3). The youth of this fluvial deposit may explain the absence of lichens on the proximal surfaces.

**Post-LIA transient aggradation of proximal outwash**

First generation outwash terraces (T1 on Figure 3) are adjacent to the moraines ridges marking the maximal glacier extent during the LIA. They are interpreted as fill terraces built up by the aggrading proglacial river during the LIA glacier advances. Dating reveals that proximal aggradation, or at least the maintenance of the outwash plain in a steady state, has persisted after the glaciers reached their maximum LIA extent and their fronts likely initiated their retreat or stagnation phase. This post-LIA aggradation stage (or steady state phase)
affected the proximal outwash of the proglacial rivers 1, 2, 5, 8, 9, 11, 13 and 14. The duration of this
aggradation or steady state stage (computed as the difference between the T1 lichenometric dating and the
age of LIA moraines) is highly variable and ranges between 5 (±14) and 147 (±15) years for the outwash 11 and
5 respectively (Figure 4 and Table 3). The outwash 7 is the only one which seems to persist in the aggradation
stage (Figure 4). Nevertheless, as its proximal reach is affected by an artificial embankment, its fluvial response
is not fully controlled by natural drivers. Based on the other investigated outwash surfaces, aggradation may be
considered as a transient stage of post-LIA outwash response.

Contraction and incision of the proximal outwash

The post-LIA contraction of the proximal outwash width is detectable in the successive abandonment of
outwash terraces (T2 to T5 on Figure 3) by proglacial rivers. These fluvial deposits are interpreted as nested cut
terraces and reveal the incising dynamic affecting proximal outwash. This glaciofluvial evolution follows the
transient aggradation stage and is still currently the dominant behaviour in most of the rivers. The outwash
plain 7 is the only one which appears to still be aggrading or remaining in a steady state, likely due to the
artificial constraints described above (Figure 4). On the other outwash plains, the proximal fluvial activity was
gradually concentrated downstream of a single breach in the LIA moraines. Currently, the active outwash plain
is framed by multiple paired or unpaired terraces whose preservation is dependent on the lateral mobility
and/or the rate of proximal river incision (Charlton, 2008).

Stream capture and shift in the outwash plain location

The mapping of glacier forelands (Figure 3) and the sequence of outwash terrace abandonment (Figure 4) also
shows a typical marginal fluvial process: stream capture within the proglacial margin leading to a shift of the
outwash plain location. This process affected the outwash plain 3 in 1939 (±3), 5 in 1891 (±3), 11 in 1914 (±8)
and 1959 (±5), and 13 in 1918 (±7). Stream capture seems to preferentially occur within proglacial margins
where two ice lobes were coalescent at the LIA glacial maximum (Figure 3). This is the case for the glaciers 3, 11
and 13 (Figure 3) whose glacial fronts were respectively coalescing with those of 4, 12 and the western arm of
13 at the LIA glacial maximum. Stream capture and location shift of the proximal outwash plain can occur
slowly, and follows a sequence of proximal outwash narrowing: In the foreland of glaciers 11 and 12(Figure 3),
three outwash terraces (T1, T2 and T3) were formed due to the contraction of the proximal braiding belt of the
outwash plain shared by the two coalescing glaciers. In 1914, about 50 years after the LIA glacial maximum, a
first stream capture occurred within the glacier 11 proglacial margin causing the abandonment of the outwash
plain T3 and the development of the outwash plain T4. After 45 years of activity (during which T4 may have
eroded the outwash terrace T3), a new stream capture occurred in the proglacial margin of glacier 11. T4 is
abandoned in 1959 in favour of glacier 11 intra-morainic plains and the outwash plain which is still currently
active. Ultimately, proximal incision and stream capture processes do not exclude each other. Moreover,
according to our results, stream captures occur exclusively in the incising stages of the proximal outwash.

Proximal outwash incision prior to the LIA maximum

The lichenometric dating of the first terrace generation (T1) of the proglacial rivers 3 and 4 predates the LIA
glacial maximum (M1). According to these lichenometric ages, the abandonment of outwash T1 downstream of
the glacier 3 may have occurred 16 years (±25 yr.) before the LIA maximum. But, as the absolute uncertainty exceed the difference between the ages of M1 and T1, proximal outwash incision and LIA maximum could also have occurred synchronously between 1864 and 1873. Dating results are much more surprising for the river 4, since the abandonment of the oldest terrace occurred almost two centuries (179 ±21 years) before the LIA maximum (Figure 3 and Figure 4). Thus, proximal incision, or at least lateral contraction, of the outwash of the rivers 3 and 4 began before the glacier reached its maximum LIA extent. Early incision also affected the outwash 6 whose T1 terrace (Figure 3) was built by the 1727 jökulhlaup. The incision of this deposit may be independent of the fluctuations of the front of the glacier 6 and is instead related to a progressive return of the proglacial river to pre-jökulhlaup conditions (Thompson & Jones, 1986).

Outwash long profile adjustment

Outwash longitudinal profiles reveal a dominant pattern of fluvial adjustment characterized by a proximal incision often accompanied by distal aggradation (Figure 5). The inflection point (IP) connecting the incising upper reach and the aggrading lower reach tends to migrate downstream through time.

Proximal incision, slope decrease and distal aggradation

Post-LIA proximal incision can be quantitatively assessed for the rivers 1, 2, 8, 9 and 14, which were not affected by stream captures or early incision prior to the LIA glacial maximum. These proglacial rivers present several nested outwash terraces, the oldest of which (T1) is always backed by the LIA moraine ridges (Figure 5). More recent outwash terraces can be disconnected from these moraines. This is the case for the rivers 8, 9 and 14. This observation reflects either a downstream migration of the proximal outwash contraction or the destruction of the upstream part of recent terraces due lateral erosion and avulsion of the active river channel. Following the transient aggradation stage (or steady state stage), the initiation of proximal incision implies the abandonment of the oldest outwash terrace (T1) by river flows. As proximal degradation continues, proglacial rivers abandon younger terraces. This is particularly the case for the rivers 2 and 14 where six and four generations of terraces are distinguishable, respectively. The rate of proximal outwash incision, computed over the period from the abandonment of the oldest terrace until 2007, averages 0.073 ± 0.03 m.a⁻¹ and ranges from 0.045 ± 0.006 m.a⁻¹ to 0.131 ± 0.012 m.a⁻¹ for the rivers 2 and 14 respectively (Table 3).

The slope of the proximal rivers has decreased almost systematically since the LIA maximum. This is qualitatively detectable on Figure 5 by a visual comparison of the long profiles of the first generation of post-LIA outwash terraces (T1) with those active in 2007 (A). The mean slope of the proximal rivers (the first 500 meters from the maximum LIA moraines) were calculated using a simple linear regression: results indicate that proximal slope decreased by 0.011 m.m⁻¹ (SD =0.009 m.m⁻¹) on average between the first-generation of post-LIA outwash terraces and the active rivers in 2007. However, there is a high variability in the post-LIA lowering of the proximal slope. For example, the proximal slope values of the active and the outwash terraces of the proglacial river 7 are identical, in agreement with the absence of proximal incision (Table 3). Conversely, the proximal slope of the river 9 decreased by 0.0311 m.m⁻¹, dropping from 0.0339 m.m⁻¹ in 1891 (79 years after the LIA maximum) to 0.0028 m.m⁻¹ in 2007 (195 years after the LIA maximum PAG). For the river 9, the
comparison of the old (terrace T1) and current (2007) proximal slopes is only based on the first 200 meters from the LIA maximum moraines due to the narrowness of the conserved fragment of T1 terrace. Consequently, the slope measurements on this outwash plain reflect more closely the proximal concavities and are difficult to compare with the other rivers. Decrease of the proglacial rivers slopes is mainly due to the proximal incision of the outwash plains. However, proximal incision is often associated with a distal aggradation of the outwash plain, which contributes to the overall decrease of the proglacial river slope. Indeed, nine of the twelve studied outwash plains present this type of post-LIA adjustment of their longitudinal profile: these include the outwash plains of glaciers 2, 3, 4, 5, 8, 9, 11, 13 and 14 (Figure 5).

**Downstream migration of the inflection point (IP)**

Post-LIA changes in the location of the inflection point (IP) were reconstructed for the proglacial rivers 2, 3, 4, 5, 8, 9, 11, 13 and 14 (Figure 5). The longitudinal profiles of terraces and active outwash of the proglacial rivers 1, 6 and 7 do not exhibit inflection points. Results demonstrate the post-LIA downstream migration of the inflection points, indicating a longitudinal increase of the extent of proximal outwash area affected by incision (Figure 4). This evolution may indicate a perennial sediment deficit in the marginal area leading to a downstream progression of the proximal incision. According to these results, the sediment recharge of the proglacial rivers is primarily carried out by the reworking of bed material (proximal incision) rather than lateral erosion of the former proximal outwash terraces (proximal contraction).

Average rates of longitudinal migration, elevation lowering and downstream migration of the inflection points were also computed for the proglacial rivers 2, 3, 4, 5, 8, 9, 11, 13 and 14 (Table 3). Rates of IP longitudinal migration average 20.4 m.a\(^{-1}\) (SD=13.6 m.a\(^{-1}\)) and range from 5.2 m.a\(^{-1}\) for the river 5 to 46.4 for the river 14. Altitudinal lowering rates of the inflection points average 0.3 m.a\(^{-1}\) (SD=0.15 m.a\(^{-1}\)) and range from 0.09 m.a\(^{-1}\) for the river 9 to 0.5 m.a\(^{-1}\) for the river 14. The synthetic rates of IP downstream migration are similar to longitudinal migration rates (average = 20.4 m.a\(^{-1}\), SD=13.6 m.a\(^{-1}\) and values range from 5.2 m.a\(^{-1}\) for the Virkisá to 46.4 for the Hoffellsá).

A category of proglacial rivers in which the downstream migration of the inflection point is rather slow can be identified. These are rivers 4, 5, 8, and 9, whose post-LIA average rates of IP longitudinal migration and elevation lowering are respectively lower than 15 m.a\(^{-1}\) and 0.2 m.a\(^{-1}\). A second category of proglacial rivers (rivers 2, 11 and 14), exhibit the highest rates of IP longitudinal migration and elevation lowering, exceeding 28 m.a\(^{-1}\) and 0.40 m.a\(^{-1}\) respectively. Between these two categories, the rivers 13 and 3 display intermediate rates of IP longitudinal migration (16.85 and 26.5 m.a\(^{-1}\) respectively) and IP elevation lowering (0.25 and 0.32 m.a\(^{-1}\) respectively). These results suggest that the proximal sediment deficit of the rivers 4, 5, 8 and 9 is less severe than for the rivers 2, 11, and 14, whose inflection points, and the spatial extent of proximal incision, rapidly migrated downstream.

Linear regression analysis (Figure 6-A) indicates that average rates of longitudinal migration and altitudinal lowering are significantly and positively correlated (R\(^2\)=0.791, P=0.001). It suggests that the overall process of post-LIA downstream migration of proximal incision involves the same proportion of longitudinal migration and elevation lowering of the IP whatever the intensity of the downstream migration.
The stepwise multiple linear regression procedure test the dependency of the IP downstream migration rate on the landform assemblage of the proglacial margin and the post-LIA glacial retreat. The linear stepwise procedure only selected the proportion of proglacial lakes within the marginal landscape as a significant model parameter ($t=3.527$, $p=0.01$). The downstream migration rate of the inflection point is positively and significantly correlated with the percentage of proglacial lakes within the proglacial margin ($R^2=0.640$, $P=0.010$). According to this result (Figure 6-B), the development of large proglacial lakes in the glacier foreland increases the rate of that the IP migrates downstream.

Atypical outwash adjustments

The outwash plains of the glaciers 1, 6 and 7 exhibit different post-LIA adjustments of their longitudinal profile (Figure 5). The outwash plain of glacier 1 has experienced an incision of its proximal part since the LIA maximum. However, no evidence of distal aggradation was found in the field. In addition, the post-LIA decrease of the proximal slope remains extremely low (0.013 m.m$^{-1}$ for T1 terrace versus 0.012 m.m$^{-1}$ for the river in 2007). The deep incising efficiency of recurrent volcanic jökulhlaups which affect the Glacier 1 glacier foreland may produce this uniform degradation of the entire outwash plain and inhibits post-LIA distal aggradation. The glacier 6 is also affected by volcanic jökulhlaups. Moreover, its outwash terrace T1 (Figure 5) constitutes the deposit of the historical jökulhlaup of 1727. The incision of this deposit may be independent of the fluctuations of the glacier front and is rather related to a progressive return of the river 6 to the pre-jökulhlaup conditions (Thompson & Jones, 1986). Alternatively, the case of the outwash plain 7 is more surprising: no incision of the abandoned and vegetated outwash plain was observed on the field. As mentioned above, the vegetated outwash plain of glacier 7 seems extremely recent (absence of lichens). Its abandonment by proglacial flows seems artificial and related to the embankment structure located on the right bank of the upstream reach of the river 7 (Figure 3). As a consequence, the abandoned outwash plain of the river 7 and the decrease in the lateral extent of the fluvioglacial dynamic cannot be considered as a purely natural response of the proglacial river to glacial retreat. Alternatively, large alluvial fans located within the proglacial margin of this steep glacier may provide an abundant sediment supply that mitigates the marginal sediment deficit and inhibits proximal incision.

Local controls on the long profile of active outwash in 2007

Long profiles of the active outwash of the glacier 4, 8, 10 and 13 (in 2007) reveal the local controls exerted by embankment structures or rocky outcrops on the efficiency of sediment transport. Indeed, these long profiles display significant decrease of local river slope values in the upstream vicinity of bridgeworks (related to the Icelandic Ring Road) or rocky outcrops. These zones force sediment deposition, related to dam effects and lateral constraints exerted by embankment and outcrops, and can hinder the downstream migration of sedimentary waves and mitigate the standard pattern of post-LIA adjustment of the outwash long profile by moderating proximal incision. Alternatively, the large alluvial fan located on the left bank of the river 14 seems to exert a control over the river long profile. A local increase in the slope value can be observed in the vicinity of the fan toe, probably due to lateral sediment contribution coming from the alluvial fan (Knighton, 1998). Conversely, the alluvial fan located on the left bank of the distal part of the outwash plain 2 does not seem to
exert any control over the slope values of the active river in 2007. Nevertheless, it appears that presence of large active alluvial fans can also modify the post-LIA adjustment of the outwash long profile by exerting control over local slope values.

4. DISCUSSION

The dominant sequence of post-LIA outwash adjustment: a two-stage descriptive model

According to our results, the dominant response of Type I (sensu Maizels, 1993) Icelandic proglacial rivers to post-LIA glacier retreat follow two successive stages: (1) a first stage of transient aggradation, or steady state, of the proximal outwash, (2) a second stage characterising by proximal incision, distal aggradation and the downstream migration of the inflection point (IP) which separates the upstream incising reach and the downstream aggrading reach of proglacial rivers.

Stage I: Outwash transient aggradation (or in balance), a paraglacial signature?

During the first stage, even if the post-LIA glacial retreat is initiated, the outwash plain is still aggrading or, at least, kept in a steady state (Figure 7-A). Two explanations of this transient aggrading stage are possible: (1) The glacial front may have registered a slight retreat during stage I and experienced a steady period with minor seasonal fluctuations that maintain high glacigenic sediment supply to proglacial rivers in which aggradation continues (or remains in its LIA maximum state); (2) Alternatively, if the ice front registered a pronounced and continuous post-LIA retreat, the transient aggradation stage can only be related to efficient recharge of the marginal sediment stock by paraglacial processes and / or greater rates of subglacial sediment evacuation. In particular, in the early stages of deglaciation, thermokarstic processes, related to the dismantling of dead-ice moraines, may be able to compensate for the decrease of glacigenic sediment supply induced by glacial retreat for a time (Mercier et al., 2009, Irvine-Fynn et al., 2011). When the efficiency of subglacial sediment evacuation decrease, thermokarstic processes cease and the moraine ridges are stabilized, the paraglacial recharge of the proglacial fluvial system is no longer sufficient to maintain high sediment supply and stage I ends.

According to lichenometric dating of the first generation outwash terraces (T1), the duration of the transient aggrading stage is highly variable: it averages 61.8 years (SD= 40.48 years) and ranges between 5 (±14) and 147 (±15) years. However, the range of duration of post-LIA transient aggradation stage of the South Icelandic fluvio-glacial plains is in agreement with previous studies on the duration of paraglacial processes as effective sediment providers for the proglacial fluvial system. On the proglacial margin of the Small River Glacier in Canada, marginal deposits and mostly abandoned intra-morainic plains have been shown to cease to function as sediment sources after several decades (Orwin and Smart, 2004). Our results tends to support this hypothesis of a stabilization of morainic ridges and intra-morainic deposits within half a century or even a century. This period of time is too short to encompass all of what is commonly referred as the so-called paraglacial period (Ballantyne, 2002, Mercier, 2010, Cossart, 2005). More likely, the duration of stage I is only the time period during which the effectiveness of paraglacial processes is sufficient to, at least, keep the outwash plain in balance. As the marginal landscape differs widely from one glacier to another, the duration of
the transient aggrading period also varies accordingly. This variability may be related to the amount of the morainic sediment stock in the proglacial margin and its sensitivity to potential reworking processes. In particular, the presence of dead ice in the moraine deposits may play a major role in the initiation of debris flows eroding moraine ridges inducing the recharge of the proglacial fluvial system with paraglacial sediments (Church & Ryder, 1972, Church & Slaymaker 1989, Mercier, 2009).

Stage II: proximal incision, distal aggradation and downstream progression of incision wave

The incision stage of the proximal outwash plain is initiated when the sediment stock of the proglacial margin is deficient and / or disconnected and the efficiency of paraglacial remobilization of the morainic and intra-morainic deposits is no longer sufficient to maintain the outwash plain in balance (Figure 7-A). The erodibility of the marginal sediment stock and the continuation of glacial retreat results in the lowering of the local base level which promotes stream captures within the proglacial margin. The occurrence of stream captures implies the abandonment by proglacial flows of former intra-morainic plains, often to the benefit of a single moraine breach which constitutes the favoured outlet of water and sediment flows to the outwash plain. This post-LIA response of the marginal channel network was also observed in Spitsbergen on the Lovén glacier margins (Griselin 1982, Mercier, 2001, Mercier & Laffly, 2005). The response of south-Icelandic outwash plains to post-LIA glacial retreat is the consequence of this marginal sediment deficit and the concentration of water and sediment flux at a single moraine breach. The sediment recharge of the outwash plain is primarily carried out by the incision of the proximal fluvio-glacial plain, resulting in the abandonment of the previous aggrading outwash plain and the contraction of the proximal fluvial activity width. The production of glacio-fluvial sediments in the upstream incising reach caused further aggradation in the downstream reach. As the marginal sediment deficit continues, the degradation of the proximal outwash progresses downstream as an incision wave (Germanosky & Harvey, 1993). Rather than a uniform degradation of the outwash plain, the successive formation of different terrace generations reflect the crossing of geomorphological thresholds and the nonlinearity of the fluvio-glacial response to fluctuations of water and sediment fluxes. These crossing of geomorphological thresholds may be linked to abrupt variations in the efficiency of the marginal sediment connectivity caused especially by moraines dams and their failures (Cossart, 2008, Cossart & Fressard, 2017). Proglacial lakes within the marginal landscape also play a key role as they interrupt the delivery of water and sediment to the outwash plain (Schomacker, 2010; Carrivick & Tweed, 2013; Geilhausen, 2013; Bogen et al., 2015, Staines et al., 2015). According to our results, the abundance of proglacial lakes within the proglacial margin speed up the downstream progression of the incision wave affecting proximal outwash. They clearly reduce the marginal sediment connectivity in trapping the coarsest part of the glacigenic sediment production. Thus, in a similar way to artificial dams, proglacial lakes deliver at their outlets “hungry water” prone to channel incision (Kondolf, 1997) and bed armoring (Vericat et al., 2006). High-magnitude/low-frequency events emerging from proglacial lakes as outburst floods (Westoby et al., 2014; Worni et al., 2014), rainstorm floods (Marren & Toomath, 2013) and iceberg jam floods (Roussel et al., 2016) may be able to temporarily reconnect the sediment connectivity between the proglacial lake and the outwash plain.

Singularities in the dominant outwash adjustment model
The dominant sequence of post-LIA response of the South Icelandic type I outwash developed above has to be qualified on the basis of field observations. The reported singularities are not major divergences from the overall logic of the dominant outwash plain adjustment. They rather reflect the effects of individual marginal landscape assemblages on the response of outwash plains to post-LIA glacial retreat.

**Coalescent glacier fronts at the LIA maximum and stream capture**

The post-LIA response of the outwash plains 3, 5, 11 and 13 demonstrate that processes of marginal stream capture and large-scale shifts of the outwash location appear to be associated with a particular glacial configuration: Two coalescent lobes at the LIA maximum (Figure 7-B). As the contact between two glacier lobes is a preferential outlet for water and sediment fluxes (Thompson & Jones, 1986), the active proximal outwash plain is typically located downstream of the junction of the glacier fronts at the LIA glacial maximum. When the ice fronts of coalescent glaciers were disconnected following the post-LIA glacial retreat, the outwash plain shared by coalescent lobes was abandoned in favour of the intra-morainic plains of each of the two forefields which then developed two distinctive outwash plains. Downstream progression of incision waves and stream capture processes do not exclude each other. They may succeed each other in the history of outwash plain response to post-LIA glacial retreat.

**Outwash incision prior to the LIA maximum**

Evidence of proximal outwash plain incision prior to the LIA maximum (proglacial rivers 3 and 4) suggests two alternative interpretations: (1) a lateral contraction and/or a stream capture and a shift in the location of the fluvial activity may occur during periods of glacial advance or glacial stagnation. In periods of glacial advance, this kind of adjustment may be related to the damming effect of proglacial moraine ridges which spatially constrain and channelize water and sediments flows; (2) A lateral contraction and/or a stream capture and a shift in the location of the fluvial activity may be caused by a transient retreat of the ice front during the LIA. The later advance of the ice front, up to the maximum glacial extension of the LIA, does not produce a sufficient aggradation or widening of the outwash plain to destroy or bury the glaciofluvial deposit built prior to the temporary LIA glacial retreat. These scenarios may explain the conservation of outwash terraces older than the LIA maximum.

**Deviation to the dominant outwash adjustment model: Jökulhaups outwash (Type III)**

The glacial tongues of the glaciers 1 and 6 are known to be outlets of jökulhaups triggered by subglacial eruptions (Sigurðsson, 1998). Our results show an alternative post-LIA response of the longitudinal profile of the outwash plains 1 and 6. The river 1 experienced an incision of its upper reach since the LIA maximum. However, no evidence of distal aggradation was found in the field. In addition, the post-LIA decrease in the proximal outwash slope remains extremely low (0.013 m.m\(^{-1}\) for T1 and 0.012 m.m\(^{-1}\) for the active outwash). It suggests that the strong incisional power of the volcanic jökulhaups that affect the foreland of glacier 1 is responsible for this uniform degradation of the outwash plain and the absence of distal aggradation. The outwash plain 6 is also affected by volcanic jökulhaups (the outwash terrace T1 is the major deposit of the 1727 historical jökulhaup). Following Thompson & Jones (1986), we believe that the successive incisions of this inherited deposit are independent of the fluctuations of the glacier 6 and rather testify to the gradual return of
the outwash plain to pre-jökulhlaup conditions. In more general terms, and as Nicholas & Sambrook-Smith (1998) suggest, in terms of morphogenesis and reworking of the fluvioglacial sediment stock, the outwash plains impacted by jökulhlaups (Type III) are insensitive to the range of meltwater discharge associated with the "normal" glacial ablation. The size of the material deposited during these catastrophic floods is beyond the scope of seasonal proglacial flows. Consequently, the reshaping of fluvioglacial forms of type III outwash is totally independent of glacial fluctuations. Their fluvial morphogenesis seems to depend exclusively on the frequency and magnitude of volcanic jökulhlaups.

5. CONCLUSION

Mapping, lichenometric dating and long-profile surveys of Icelandic proglacial margins, glacio-fluvial terraces and active proglacial rivers permit to address the threefold aims of this study: (1) Proximal incision, associated with distal aggradation and downstream migration of the inflection point, is the dominant pattern of proglacial river response to post-LIA glacier retreat in Iceland. The post-LIA contraction and incision of Icelandic proximal outwash plains usually occurs after a first transient aggradation or steady state stage of proximal outwash (duration ranges between 5 ±14 and 147 ±15 years) which may be related to efficient recharge of the marginal sediment stock by paraglacial processes and / or greater rates of subglacial sediment evacuation. Proximal incision, distal aggradation and downstream migration of the inflection point can precede or be preceded by stream capture phenomena which particularly affects marginal landscapes where two glacier fronts were coalescent at the LIA maximum. Outwash plain incision prior to the LIA maximum have been observed and may be related to damming effect of proglacial moraine ridges or transient glacial retreat during the LIA. Two proglacial rivers affected by historical jökulhlaups exhibit an alternative adjustment rather controlled by extremes events than by glacier retreat. (2) The mean rates of downstream migration of the inflection point average 20.4 m.a⁻¹ (SD=13.6 m.a⁻¹) and range from 5.2 m.a⁻¹ for the river 5 to 46. 4 for the river 14. IP migration results suggest that the proximal sediment deficit of the rivers 4, 5, 8 and 9 is less severe than for the rivers 2, 11 and 14, whose inflection points, and the spatial extent of proximal incision, rapidly migrated downstream. (3) The downstream migration rate of the inflection point is positively correlated with the proportion of proglacial lakes within the glacier foreland. These findings suggest that proglacial margins dominated by proglacial lakes intensify the rates of the fluvial adjustment to glacial retreat.

Acknowledgements

Research was supported by the GEOLAB-CNRS 6042, and the GDR “Mutations Polaires” (Coord. M. Griselin). The authors thank the Skaftafell National Park which provided assistance and Pr. Johannes Steiger for his helpful suggestions. The contribution to the manuscript has to be shared between authors as follows: the investigation was outlined by the first author which carried out most of the field measurement activity. All authors contributed to analyse the results and to write the paper.

REFERENCES


Cossart E. 2005. Evolution géomorphologique du haut bassin durancien depuis la dernière glaciation (Briançonnais, Alpes française du Sud). Université Paris 7 - Denis Diderot, 382 pp

Cossart E. 2008. Landform connectivity and waves of negative feedbacks during the paraglacial period, a case study: the Tabuc subcatchment since the end of the Little Ice Age (massif des Écrins, France). *Géomorphologie : relief, processus, environnement* **14**: 249–260. DOI: 10.4000/geomorphologie.7430


Figures and Tables

Figure 1: Location of investigated glaciers and associated proglacial rivers
Table 1: Main characteristics of glaciers and associated proglacial rivers investigated in this study.

<table>
<thead>
<tr>
<th>ID</th>
<th>Name</th>
<th>LIA max (yr. AD)*</th>
<th>Conf. interval*</th>
<th>Glacier Area (Km²)</th>
<th>Glacier length (km)</th>
<th>Glacier slope (°)</th>
<th>ELA (m a.s.l.)***</th>
<th>River</th>
<th>Outwash Type***</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>Sólheimajökull</td>
<td>1851 (1839 - 1863)</td>
<td>47.6</td>
<td>15.4</td>
<td>5.1</td>
<td>1480</td>
<td>na</td>
<td>Jökulsá</td>
<td>Type III</td>
</tr>
<tr>
<td>2</td>
<td>Morsárjökull</td>
<td>1883 (1873 - 1903)</td>
<td>26.3</td>
<td>13.1</td>
<td>6.2</td>
<td>1421</td>
<td>1000-1130</td>
<td>Morsá</td>
<td>Type I</td>
</tr>
<tr>
<td>3</td>
<td>Skátafelljökull</td>
<td>1878 (1864 - 1892)</td>
<td>87.0</td>
<td>19.0</td>
<td>6.6</td>
<td>1889</td>
<td>1000-1160</td>
<td>Skátafellá</td>
<td>Type I</td>
</tr>
<tr>
<td>4</td>
<td>Svinafellsjökull</td>
<td>1765 (1749 - 1776)</td>
<td>31.7</td>
<td>12.7</td>
<td>12.4</td>
<td>2104</td>
<td>1000-1120</td>
<td>Svinafellá</td>
<td>Type I</td>
</tr>
<tr>
<td>5</td>
<td>Virkisjökull</td>
<td>1740 (1731 - 1755)</td>
<td>20.6</td>
<td>11.3</td>
<td>10.7</td>
<td>2076</td>
<td>na</td>
<td>Virkisá</td>
<td>Type I</td>
</tr>
<tr>
<td>6</td>
<td>Kotájökull</td>
<td>1819 (1809 - 1828)</td>
<td>6.1</td>
<td>6.6</td>
<td>12.8</td>
<td>1840</td>
<td>1000-1130</td>
<td>Kotá</td>
<td>Type III</td>
</tr>
<tr>
<td>7</td>
<td>Hólárjökull</td>
<td>1844 (1838 - 1857)</td>
<td>5.2</td>
<td>6.9</td>
<td>12.9</td>
<td>1838</td>
<td>na</td>
<td>Holá</td>
<td>Type I</td>
</tr>
<tr>
<td>8</td>
<td>Kvíárjökull</td>
<td>1810 (1798 - 1824)</td>
<td>18.6</td>
<td>11.7</td>
<td>10.3</td>
<td>2010</td>
<td>1010-1130</td>
<td>Kvíá</td>
<td>Type I</td>
</tr>
<tr>
<td>9</td>
<td>Hrútárjökull</td>
<td>1812 (1805 - 1820)</td>
<td>11.9</td>
<td>8.5</td>
<td>14.6</td>
<td>2047</td>
<td>880-910</td>
<td>Hóllsá</td>
<td>Type III</td>
</tr>
<tr>
<td>10</td>
<td>Fláajökull</td>
<td>1821 (1811 - 1834)</td>
<td>213.1</td>
<td>29.1</td>
<td>3.3</td>
<td>2040</td>
<td>870-960</td>
<td>Hóllsá</td>
<td>Type I</td>
</tr>
<tr>
<td>11</td>
<td>Skátafelljökull</td>
<td>1865 (1856 - 1876)</td>
<td>136.2</td>
<td>32.4</td>
<td>3.3</td>
<td>1520</td>
<td>910-1020</td>
<td>Kolgrímá</td>
<td>Type I</td>
</tr>
<tr>
<td>12</td>
<td>Heinabergsjökull</td>
<td>1851 (1835 - 1865)</td>
<td>93.8</td>
<td>23.5</td>
<td>4.8</td>
<td>1520</td>
<td>990-1100</td>
<td>Kolgrímá</td>
<td>Type I</td>
</tr>
<tr>
<td>13</td>
<td>Fláajökull</td>
<td>1844 (1838 - 1857)</td>
<td>5.2</td>
<td>6.9</td>
<td>12.9</td>
<td>1838</td>
<td>na</td>
<td>Holá</td>
<td>Type I</td>
</tr>
<tr>
<td>14</td>
<td>Hoffellsjökull</td>
<td>1888 (1874 - 1898)</td>
<td>48.8</td>
<td>13.4</td>
<td>9.1</td>
<td>2040</td>
<td>870-960</td>
<td>Hoffellsá</td>
<td>Type I</td>
</tr>
</tbody>
</table>

* after Chenet et al., 2010 and Roussel et al., 2008; ** after Chenet et al., 2010 and Hannesdóttir et al., 2014; *** after Maizels, 1993, 1997; ; na for not available
Table 3: Outwash responses to post-LIA glacial retreat

<table>
<thead>
<tr>
<th>ID</th>
<th>Glacier name</th>
<th>River name</th>
<th>Outwash Type*</th>
<th>Duration of aggradation stage (yr)</th>
<th>Post-LIA proximal incision (m.a.-1)</th>
<th>Post-LIA decrease of proximal slope (m.m.-1)</th>
<th>Average IP longitudinal migration (m.a.-1)</th>
<th>Average IP altitudinal lowering (m.a.-1)</th>
<th>Average IP downstream migration (m.a.-1)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>Sólheimajökull</td>
<td>Jökulsá</td>
<td>Type III</td>
<td>41 (±16)</td>
<td>0.047 (±0.002)</td>
<td>0.00057</td>
<td>na</td>
<td>na</td>
<td>na</td>
</tr>
<tr>
<td>2</td>
<td>Morsárjökull</td>
<td>Morsá</td>
<td>Type I</td>
<td>51 (±18)</td>
<td>0.045 (±0.006)</td>
<td>0.00584</td>
<td>28.7</td>
<td>0.44</td>
<td>28.68</td>
</tr>
<tr>
<td>3</td>
<td>Skaftafelljökull</td>
<td>Skaftállsá</td>
<td>Type I</td>
<td>na</td>
<td>na</td>
<td>0.0206</td>
<td>26.5</td>
<td>0.32</td>
<td>26.47</td>
</tr>
<tr>
<td>4</td>
<td>Svinafellsjökull</td>
<td>Svinállsá</td>
<td>Type I</td>
<td>na</td>
<td>na</td>
<td>0.0044</td>
<td>8.9</td>
<td>0.13</td>
<td>8.91</td>
</tr>
<tr>
<td>5</td>
<td>Virkisjökull</td>
<td>Virkísá</td>
<td>Type I</td>
<td>147 (±15)</td>
<td>na</td>
<td>0.0157</td>
<td>5.2</td>
<td>0.19</td>
<td>5.21</td>
</tr>
<tr>
<td>6</td>
<td>Kotárjökull</td>
<td>Kotá</td>
<td>Type III</td>
<td>na</td>
<td>na</td>
<td>0.0177</td>
<td>na</td>
<td>na</td>
<td>na</td>
</tr>
<tr>
<td>7</td>
<td>Hólárjökull</td>
<td>Holá</td>
<td>Type I</td>
<td>na</td>
<td>0</td>
<td>0</td>
<td>na</td>
<td>na</td>
<td>na</td>
</tr>
<tr>
<td>8</td>
<td>Kvíárjökull</td>
<td>Kviá</td>
<td>Type I</td>
<td>60 (±21)</td>
<td>0.076 (±0.004)</td>
<td>0.0122</td>
<td>6.6</td>
<td>0.20</td>
<td>6.57</td>
</tr>
<tr>
<td>9</td>
<td>Hrútárjökull</td>
<td>Fjallsá</td>
<td>Type I</td>
<td>79 (±25)</td>
<td>0.068 (±0.003)</td>
<td>na</td>
<td>14.0</td>
<td>0.09</td>
<td>13.97</td>
</tr>
<tr>
<td>10</td>
<td>Fjallsjökull</td>
<td>Kolgrímá</td>
<td>Type I</td>
<td>5 (±14)</td>
<td>na</td>
<td>na</td>
<td>30.3</td>
<td>0.47</td>
<td>30.34</td>
</tr>
<tr>
<td>11</td>
<td>Skálafellsjökull</td>
<td>Kolgrímá</td>
<td>Type I</td>
<td>na</td>
<td>na</td>
<td>0.0108</td>
<td>16.9</td>
<td>0.25</td>
<td>16.86</td>
</tr>
<tr>
<td>12</td>
<td>Heinabergsjökull</td>
<td>Kolgrímá</td>
<td>Type I</td>
<td>na</td>
<td>na</td>
<td>0.131 (±0.012)</td>
<td>46.4</td>
<td>0.5</td>
<td>46.38</td>
</tr>
</tbody>
</table>

* After Maizels, 1993, 1997, "na" for not applicable
Figure 1: Location of investigated glaciers and associated proglacial rivers

110x69mm (300 x 300 DPI)
Figure 2: Methodological workflow: A - Field prospection and GIS mapping of proglacial and glaciofluvial landforms; B - Lichenometric dating of outwash terraces; C - Topographic survey: Cross section and long profile of active outwash and terraces; D - Method for the determination of the Inflection Point migration.
Figure 3: Outwash terraces mapping and dating

275x414mm (300 x 300 DPI)
Figure 4: Overview of aggrading and incising stages in the post-LIA outwash adjustment

182x188mm (300 x 300 DPI)
Figure 5: Long profiles of dated outwash terraces (T1 to T4) and active outwash in 2007 (A)

202x233mm (300 x 300 DPI)
Figure 6: A - Linear regression between average rates of altitudinal lowering and longitudinal migration of the inflection points. B - Linear regression between average rates of IP downstream migration and the proportion of proglacial lakes in the marginal area.
Figure 7: A - The dominant sequence of post-LIA fluvial adjustment: a two-stages descriptive model. B - LIA coalescent glaciers specific case, with stream capture and shift in the outwash location.

192x209mm (300 x 300 DPI)