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Coralline algal structure is more sensitive to rate, rather than the magnitude, of ocean acidification

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Abstract

Marine pCO₂ enrichment via ocean acidification (OA), upwelling and release from carbon capture and storage (CCS) facilities is projected to have devastating impacts on marine biomineralisers and the services they provide. However, empirical studies using stable endpoint pCO₂ concentrations find species exhibit variable biological and geochemical responses rather than the expected negative patterns. In addition, the carbonate chemistry of many marine systems is now being observed to be more variable than previously thought. To underpin more robust projections of future OA impacts on marine biomineralisers and their role in ecosystem service provision, we investigate coralline algal responses to realistically variable scenarios of marine pCO₂ enrichment. Coralline algae are important in ecosystem function; providing habitats and nursery areas, hosting high biodiversity, stabilising reef structures and contributing to the carbon cycle. Red coralline marine algae were exposed for 80 days to one of three pH treatments: (i) current pH (control); (ii) low pH (7.7) representing OA change; and (iii) an abrupt drop to low pH (7.7) representing the higher rates of pH change observed at natural vent systems, in areas of upwelling and during CCS releases. We demonstrate that red coralline algae respond differently to the rate and the magnitude of pH change induced by pCO₂ enrichment. At low pH, coralline algae survived by increasing their calcification rates. However, when the change to low pH occurred at a fast rate we detected, using Raman spectroscopy, weaknesses in the calcite skeleton, with evidence of dissolution and molecular positional disorder. This suggests that, while coralline algae will continue to calcify, they may be structurally weakened, putting at risk the ecosystem services they provide. Notwithstanding evolutionary adaptation, the ability of coralline algae to cope with OA may thus be determined primarily by the rate, rather than magnitude, at which pCO₂ enrichment occurs.

Keywords: calcification, coralline algae, crustose coralline algae (CCA), maerl, ocean acidification, photosynthesis, raman, rate, respiration, rhodolith

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Introduction

Marine pCO₂ enrichment may occur via ocean acidification (OA) (Caldeira & Wickett, 2003), releases from natural CO₂ vents (Halling-Spencer et al., 2008), upwelling (Feely et al., 2008), and sudden releases from carbon capture and storage (CCS) facilities (Blackford et al., 2009). Natural CO₂ enrichment along with simulated OA and CCS leakages are used to investigate the responses of marine biomineralisers to projected OA. However, enrichment occurs at different rates with differing longevities in each release process.

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Ocean acidification is projected to occur slowly over centennial time scales causing oceanic pH to drop 0.3–0.5 units by the end of this century (Caldeira & Wickett, 2003). CO₂ releases from marine vent sites (Kroeker et al., 2012), biota (Anthony et al., 2012) and upwelling areas (Gruber et al., 2012) naturally alter marine pCO₂ but are highly variable even at diel time scales (Hall-Spencer et al., 2008, Kroeker et al., 2012). Sudden CO₂ leakage from CCS infrastructure (CO₂ pipelines and underground geological storage reservoirs) into the overlying sea water may occur at even faster time scales (hours), accompanied by a sudden fall in pH (Blackford et al., 2009) with changes in pH being most acute at small spatial scales (meters) (Agnew & Taylor, 1986). Thus, while pCO₂ enrichment will expose marine biomineralisers to
reduced pH, the rates of exposure will vary significantly. Such variability is expected to be characteristic of coastal pCO₂ enrichment over the coming century, unlike pelagic regions which are expected to remain more stable (Duarte et al., 2013). Therefore the rate, rather than the magnitude, of exposure may be critical in determining the ability of marine biomineralisers to cope with projected changes in carbonate chemistry and synchronous multiple stressors such as temperature. This may partially explain why organisms show variable responses to contemporary pCO₂ enrichment of similar magnitude (Burdeett et al., 2012, Melzner et al., 2009, Ries, 2011).

Red coralline algae (Fig. 1) have an established record as model biomineralisers for exploring the impact of high CO₂ on marine calcifying biota (Burdeett et al., 2012, Kroeker et al., 2012, Martin & Gattuso, 2009, Ragazzola et al., 2012, Ries, 2011). This is because (i) they use dissolved inorganic carbon to calcify and during photosynthesis (Martin & Gattuso, 2009), providing an excellent contrast to animal models; (ii) they are important in ecosystem function (e.g., Foster, 2001, Kamenos et al., 2004a, Kamenos et al., 2004b, Nelson, 2009); (iii) play a significant role in carbon cycling and reef stabilization (Nelson, 2009); and (iv) are utilised in ultra-high resolution palaeoenvironmental reconstructions (Burdeett et al., 2011, Kamenos, 2010, Kamenos et al., 2012, Williams et al., 2011).

Coralline algae have a high-Mg skeleton (7.7–28.8 mol% MgCO₃) (Chave, 1954, Kamenos et al., 2009), which is expected to make them particularly sensitive to marine pCO₂ enrichment as high-Mg calcite is the most soluble form of calcium carbonate (Martin & Gattuso, 2009). However, algal responses to pCO₂ enrichment do not show expected regular patterns. Some studies show negative patterns: for example, exposure to high pCO₂ conditions inhibits coralline algal settlement (Kuffner et al., 2007), leads to coralline algal dissolution (Hall-Spencer et al., 2008, Martin & Gattuso, 2009), surface lesions (Martin & Gattuso, 2009), epithelial cell damage (Burdeett et al., 2012) and modelled structural stress (Ragazzola et al., 2012). Conversely, other studies observe non-negative responses: for example, intracellular concentrations of the algal antioxidant dimethylsulphoniopropionate (DMSP) did not increase under gradual change to low pH (Burdeett et al., 2012), OA induced higher coralline algal calcification (Martin et al., 2013) and did not cause coralline algae to change the CaCO₃ polymorph deposited (a high Mg-calcite structure was maintained despite the higher stability of calcite or aragonite, Ries, 2011).

This study determines the effect of the rate and magnitude of pH change on the molecular structural integrity, calcification and respiration of living and dead coralline algae. It was hypothesised that, under faster rates of pH change, red coralline algae would exhibit the greatest response away from background levels; they are less likely to show beneficial phenotypic plastic responses to faster, compared to slower environmental change. Additionally, live red coralline algae were expected to be less impacted than dead red coralline algae due to the environmental buffering provided by the live epithelium.

Fig. 1 Lithothamnion glaciale thalli containing HCO₃⁻. Percentage of L. glaciale thalli (live and dead) where HCO₃⁻ was absent (white bars) or present (grey bars) after incubation for 80 days in control, low, stable pH, and low, abrupt pH treatments.
Materials and methods

Experimental setup

Live and dead Lithothamnion glaciale thalli were hand collected from Loch Sween, Scotland (56°01.99’ N, 05°36.13’ W), in November 2009 using SCUBA from a depth of 7 m (thalli size: 4–6 cm diam.). In situ, the algae experience an annual temperature range of 5–16 °C, light levels of 30–120 μmol photons m⁻² s⁻¹ photosynthetically active radiation (PAR) and pH 8.1 (Rix et al., 2012). Lithothamnion glaciale was identified as described in Irvine and Chamberlain (1994). Thalli were maintained in aerated sea water at ca. 10 °C during transportation to Plymouth Marine Laboratory (Plymouth, UK) and were transferred to a 3000 l flow-through seawater system within one day of collection. Approximately 400 g of L. glaciale were separated into 24 experimental mesocosms (6 l volume, 28 × 19 × 16 cm): 12 containing live algae and 12 containing dead algae. Thalli were maintained in the system for one week (water temperature 11.63 ± 0.32 °C, salinity 34.9 ± 0.36 [mean ± SD] prior the start of the ramping period to allow the thalli to recover from collection and adjust to laboratory conditions.

Coraline algae were incubated in three pH treatments for 80 days using the method developed by Findlay et al. (2008):

1. Control (pH 8.1, n = 12 [six with live algae and six with dead algae])
2. Low, stable pH change (pH 7.7, n = 6 [three with live algae and three with dead algae]), rate of change from ambient: 0.05 pH units d⁻¹ over 10 days prior to the experimental period) representing the A1FI IPCC year 2100 scenario (IPCC Core Writing Team Pachauri Rk & Reisinger a Eds, 2007).
3. Low, abrupt pH change (pH 8.1 to 7.7, n = 6 [three with live algae and three with dead algae]), rate of change from ambient: 0.25 pH units d⁻¹ over two days beginning on day 52. The low, abrupt pH treatment represented a sudden reduction in pH following an acute injection of CO₂ associated with CCS leaks, natural CO₂ vent systems or areas of upwelling.

Rate of pH change in the low, abrupt pH treatment was significantly greater than in the low, stable pH treatment over the 2 days of change beginning on day 52 (F₁,48 = 13.51, P = 0.001; multiple linear regression, assumptions met). Reduced pH levels were achieved by gradually increasing the bubbling of CO₂ in the mesocosms.

Carbonate system measurements

Seawater temperature, salinity (WTW LF187 combination temperature and salinity probe), pHₑ₈₆ (Metrohm, 826 pH mobile with a Metrohm glass electrode) and dissolved oxygen (1302 Oxygen Electrode; Strathkelvin Instruments, Glasgow, UK) were monitored daily. Nutrient concentrations (nitrate, nitrite, phosphate, silicate and ammonium) and total alkalinity (Aₚ) were monitored weekly to ensure seawater quality was maintained (Table 1 and Table S1). Mesocosms were maintained at ambient temperature (11.63 ± 0.32 °C, mean ± SD) and light (90 μmol photons m⁻² s⁻¹; ca. 10 h light: 14 h dark). Nutrients were analysed with an autoanalyzer (Bran+Luebbe, Norderstedt, Germany) using standard methods (Brewer & Riley, 1965, Grashoff, 1976, Kirkwood, 1989, Mantoura & Woodstedt, Germany) using standard methods (Brewer & Riley, 1965, Grasshoff, 1976, Kirkwood, 1989, Mantoura & Woodward, 1983, Zhang & Chi, 2002). Total alkalinity was measured by poisoning according to Dickson et al. (2007) then analysing via potentiometric titration using an Apollo SciTech Alkalinity Titrator Model AS-ALK2 (Apollo SciTech, Bogart, GA, USA) and Batch 100 certified reference materials from Andrew Dickson.

Carbonate system calculations

Measured pH and Aₚ data were used to calculate carbonate and bicarbonate ion concentrations and calcite and aragonite (Ωₑ₈₆ and Ωₑ₉₈ respectively) saturation states using CO2SYS (Pierrot et al., 2006) with dissociation constants from Mehrbach et al. (1973) refit by Dickson & Millero (1987) and KSO₄ using (Dickson, 1990). The calculations for these parameters also included temperature, salinity, silicate and phosphate data. pH and CO₂ levels remained relatively constant throughout the 80 days exposure period for the control and stable treatments (Table 1). Aₚ, temperature, salinity, silicate, and phosphate showed no differences between CO₂ treatments.

Sample preparation

Live and dead L. glaciale thalli from each treatment (live: n = 3, dead: n = 3) were sampled after 80 days exposure. Samples were selected at random from each mesocosm; within the control treatment three of the six mesocosms were chosen at random. Thalli were air dried, embedded in resin (Buehler EpoxyCure, Düsseldorf, Germany), transverse sectioned.

Table 1 Mean ± SD experimental system values over the 80 days experimental period for measured temperature (n = 80), salinity (n = 80), pH (n = 80) and alkalinity (n = 12) and calculated, dissolved inorganic carbon (DIC), bicarbonate (HCO₃⁻), carbonate (CO₃²⁻), pCO₂, calcite saturation state (Ωₑ₈₆) and aragonite saturation state (Ωₑ₉₈) for the three pH treatments.

<table>
<thead>
<tr>
<th></th>
<th>Control</th>
<th>Low, stable pH</th>
<th>Low, abrupt pH</th>
</tr>
</thead>
<tbody>
<tr>
<td>Temperature (°C)</td>
<td>11.74 ± 0.34</td>
<td>11.51 ± 0.28</td>
<td>11.64 ± 0.33</td>
</tr>
<tr>
<td>Salinity (‰)</td>
<td>34.9 ± 0.34</td>
<td>35.0 ± 0.32</td>
<td>34.9 ± 0.41</td>
</tr>
<tr>
<td>pH (T)</td>
<td>8.18 ± 0.10</td>
<td>7.70 ± 0.14</td>
<td>7.75 ± 0.40</td>
</tr>
<tr>
<td>Alkalinity (μmol kg⁻¹)</td>
<td>2975 ± 443</td>
<td>2964 ± 467</td>
<td>2991 ± 414</td>
</tr>
<tr>
<td>DIC (μmol kg⁻¹)</td>
<td>2717 ± 420</td>
<td>2850 ± 489</td>
<td>3023 ± 556</td>
</tr>
<tr>
<td>HCO₃⁻ (μmol kg⁻¹)</td>
<td>2498 ± 395</td>
<td>2696 ± 475</td>
<td>2810 ± 463</td>
</tr>
<tr>
<td>CO₃²⁻ (μmol kg⁻¹)</td>
<td>198 (±53)</td>
<td>110 (±26)</td>
<td>73 (±40)</td>
</tr>
<tr>
<td>pCO₂</td>
<td>498 ± 161</td>
<td>1081 ± 488</td>
<td>2778 ± 4047</td>
</tr>
<tr>
<td>Ωₑ₈₆</td>
<td>4.7 ± 1.26</td>
<td>2.62 ± 0.61</td>
<td>1.74 ± 0.94</td>
</tr>
<tr>
<td>Ωₑ₉₈</td>
<td>3.0 ± 0.81</td>
<td>1.67 ± 0.39</td>
<td>1.11 ± 0.60</td>
</tr>
</tbody>
</table>
Physical molecular structure

Changes in the physical structure of the skeleton at the molecular level were detected using Raman spectroscopy. Raman spectroscopy uses laser light to determine the molecular vibrational modes of substances, providing information on crystallographic structure, composition and stability. Raman spectroscopy was conducted using a Renishaw inVia Raman equipped with a Leica DM 2500M (Leica Microsystems GmbH, Wetzlar, Germany) microscope using a 785 nm laser and 1200 l mm\(^{-1}\) grating within the School of Geographical and Earth Sciences at the University of Glasgow (Glasgow, UK). Peak presence or absence at 1014 cm\(^{-1}\) Raman shift was recorded. Frequency and full width half maximum (FWHM, the peak width at half the peak height) were calculated for the \(V_1\) lattice mode (peaks nominally at a Raman shift of ca. 1089 cm\(^{-1}\)). Raman spectroscopy was conducted on the spring-time deposited portion of each growth band to ensure temperature-dependent Mg concentrations remained the same across treatments. The spring-time growth portion was identified from Alizarin red stained calibration individuals in parallel experiments (e.g., Kamenos et al., 2008). Peak parameters were compared using a general linear model (assumptions of normality and heterogeneity of variance were met).

Bicarbonate presence. The 1014 cm\(^{-1}\) peak in Raman spectra of high Mg-calcite indicates the presence of bicarbonate (HCO\(_3^−\)) within the analysed sample (Bischoff et al., 1985).

Carbonate ion positional disorder. Band width (FWHM) of the ca.1089 cm\(^{-1}\) Raman spectrum peak is positively related to Mg concentrations and most likely results from positional disorder of the carbonate ion (increasing rotation of CO\(_3^2−\) out of the basal plane) (Bischoff et al., 1985).

Composition. The frequency of the ca. 1089 cm\(^{-1}\) peak in spectra of biogenic calcite is primarily controlled by the Mg content (Bischoff et al., 1985). As the Mg content increases, the peak moves from 1085 cm\(^{-1}\) in calcite containing <3.9 mol% MgCO\(_3\) to 1094 cm\(^{-1}\) in Magnesite (25 mol% MgCO\(_3\)) (Bischoff et al., 1985, Urmos et al., 1991).

Net calcification, photosynthesis and respiration

Calcification and respiration of \(L.\) glaciale were determined in control and low, stable pH treatments after 80 days exposure to experimental conditions. Data are not available for the low, abrupt pH treatment. Live thalli \((n = 3\) per treatment) were incubated in 170 ml chambers for ca. 20 min during the day (PAR: ca. 90 photons \(\mu\)mol m\(^{-2}\) s\(^{-1}\)) and at night (>1 h after sunset). Dead thalli \((n = 3\) per treatment) were incubated in 170 ml chambers for ca. 10 h spanning both day and night. Water motion was maintained using magnetic stirrers. \(A_T\) and dissolved oxygen concentrations were determined at the start and end of each experimental incubation. Calcification was determined using the total alkalinity anomaly technique (Smith & Key, 1975). Oxygen consumption / production was determined following the technique of Martin et al. (2013) using an oxygen electrode (1302 Oxygen Electrode; Strathkelvin Instruments) attached to a calibrated oxygen meter (Oxygen Meter 781; Strathkelvin Instruments). Comparisons were conducted using one and two-way general linear models (assumptions of normality and heterogeneity of variance were met).

Results

Bicarbonate presence

Bicarbonate was not present in live \(L.\) glaciale coralline algal thalli cultured under control conditions, but was present in increasing proportions within live thalli cultured under stable low pH and abruptly changing pH conditions (Fig. 1). All dead thalli contained HCO\(_3^−\) (Fig. 1).

Carbonate ion positional disorder

There were no significant differences in \(V_1\) lattice mode (symmetric stretch at ca. 1089 cm\(^{-1}\) frequency) between treatments within live \((F_{2, 14} = 0.15, P = 0.865)\) or dead thalli \((F_{1, 14} = 3.01, P = 0.015)\) (Fig. 2).

Composition

In live coralline algae, we observed no significant change in peak frequency (Mg content) between individuals cultured in control, low pH and low, abruptly changed pH \((F_{2, 12} = 1.28, P = 0.313)\) (Fig. 2). In the low, abrupt pH treatment, live coralline algae had a significantly reduced peak frequency (lower Mg content) compared to dead thalli in the same treatment \((F_{1, 12} = 6.39, P = 0.026)\) (Fig. 2).

Net calcification, dissolution, respiration and photosynthesis in live thalli

\(Lithothamnion\) glaciale calcified during the day in both control and low, stable pH treatments (Fig. 3a). At night, \(L.\) glaciale calcified in the control treatment but dissolved in the low pH treatment (Fig. 3a). Under low, stable pH, \(L.\) glaciale calcified significantly more during the day than they dissolved at night, they also calcified marginally more in low pH than control treatments during the day \((F_{1, 54} = 50.78, P < 0.001)\) (Fig. 3a) (calcification was not measured in low but abruptly changing pH). There were no significant differences in the photosynthesis \((F_{1, 22} = 1.33, P = 0.261)\) or respiration \((F_{1, 24} = 0.60, P = 0.446)\) of \(L.\) glaciale between the control and low, stable pH treatments (Fig. 3b).
Dead thalli in low, stable pH conditions dissolved faster than control thalli (Fig. 3a) \( (F_2, 42 = 4.59, P = 0.016) \).

However, dissolution of dead thalli over the whole diel period was not significantly different from the dissolution of live thalli at night \( (F_1, 24 = 0.13, P = 0.721) \) in the low, stable pH treatment (Fig. 3a).

![Nominal 1089 cm⁻¹ peak Raman spectra characteristics](image1)

**Fig. 2** Nominal 1089 cm⁻¹ peak Raman spectra characteristics. Unsmoothed full width half maximum (FWHM, double-headed arrow) and frequencies (above peak) of the ca. 1089 cm⁻¹ peak in control (a), low, stable pH (b) and low, abrupt pH (c) treatments \( (n = 3 \pm SD) \). Grey band indicates FWHM of the control treatments for comparison.

![Algal calcification, dissolution and respiration](image2)

**Fig. 3** Algal calcification, dissolution and respiration. (a) Net calcification in the light (white) and dark (grey) in live and dead (black: light and dark) *Lithothamnion glaciale* under control and low, stable pH treatments, and (b) respiration and gross photosynthesis of live *L. glaciale* under control and low, stable pH treatments during the day (white) and night (grey). Data presented as mean \( \pm SD \).
Discussion

We observed rate-dependent pH impacts on the physical molecular-level structure of red coralline algae (L. glaciale) after exposing thalli to low pH, low but abruptly reduced pH, and control conditions for 80 days.

Bicarbonate presence

All dead thalli contained HCO$_3^-$, indicating that dissolution took place even under control pH conditions (Fig. 1). We suggest that the presence of HCO$_3^-$ indicates a chemical breakdown of the thallus skeletal material, possibly indicating the first stages of major physical breakdown. The higher percentage of individuals containing HCO$_3^-$ (Fig. 1) in the thalli exposed to the abrupt change in pH indicates that faster rates of change increase dissolution. The observed HCO$_3^-$ is unlikely to be from an external source as it was present in dead, but not live, thalli from the control treatment; an external source would have introduced the bicarbonate to both live and dead thalli.

As Raman spectroscopy was conducted on thallus calcite that was within each algal annual growth band, rather than calcite directly exposed to the sea water on the outer surface of the thallus, this suggests that damage was, at least, internal. In live algae this may be via (i) seawater ingress through low pH-damaged epithelium (Burdett et al., 2012), explaining the absence of damage in live thalli from the control treatment; or (ii) disruption of the calcification process at the centres of calcification.

Carbonate ion positional disorder

While we observed no significant differences in V$_1$ lattice mode (symmetric stretch at ca. 1089 cm$^{-1}$ frequency) between treatments within live or dead thalli, we believe it is reasonable to attribute the modest fall of FWHM in live algae exposed to low, abrupt pH to a decrease in positional disorder (Fig. 2). Less positional disorder indicates longer, and thus weaker, Mg-O bond lengths (Bischoff et al., 1985). Thus despite the differences not being statistically significant, the longer, weaker Mg-O bonds indicate that live algae exposed to abrupt reductions in pH also show evidence of weakening at the molecular bonding level, an effect not observed in thalli exposed to control or low, stable pH treatments.

Mg content

Absence of treatment induced differences in Mg content (peak frequency) suggests that, while alive, L. glaciale thalli were capable of buffering against the rate and magnitude of external carbonate chemistry changes, enabling them to continue to deposit high Mg calcite.

Reduced peak frequency of live over dead algae in the low, abrupt pH treatment indicates that live L. glaciale thalli were capable of lowering the Mg content of their calcite, making it less reactive to acidified conditions. A similar process may also occur in Neogoniolithon sp., another red coralline alga, which reduces the Mg/Ca ratio of the high Mg-calcite deposited at high pCO$_2$ concentrations (>2000 μatm) (Ries, 2011). Thus, although epithelial damage may allow dissolution of the high Mg skeleton (Fig. 3), this does not appear to preferentially dissolve out Mg. This absence of Mg dissolution in the low, abrupt pH treatment could be due to the incorporation of Mg into the calcite crystal lattice (Kamenos et al., 2009) rather than being associated organic material.

Calcification, dissolution, respiration and photosynthesis in live algae

When calcifying, L. glaciale appears to be able to compensate for OA-induced dissolution at night by increasing its calcification rate during the day. In fact, in low, stable pH conditions, L. glaciale up-regulated its calcification rates. Day-time (light) calcification rates in low pH were twice the rate required to maintain calcification in control conditions (Fig. 3). This suggests that, under OA conditions, L. glaciale may not only persist, but also increase their calcification, a phenomenon that has been observed in other algal and invertebrate species (Findlay et al., 2011, Martin et al., 2013, Rodolfo-Metalpa et al., 2011).

However, enhanced calcification in low, stable pH conditions was not supported by a change in photosynthesis (Fig. 3b), and thus an increase in available energy. While it was expected that coralline algae would obtain a photosynthetic benefit under OA (Ries et al., 2009), we observed no differences in photosynthesis or respiration of live L. glaciale between the control and low, stable pH treatments (Fig. 3b). Thus, live L. glaciale neither became metabolically challenged (which would have led to increased respiration), nor obtained photosynthetic benefits from changes in carbonate chemistry projected for 2100. We suggest that, during the day, energy usage required for enhanced calcification in thalli exposed to low, stable pH may limit photosynthetic efficiency, while at night the rate of dissolution may not be high enough to induce an increase in respiration. Over the whole diel period it is possible that the high energy requirement of day-time calcification in low pH conditions leads to energy reallocation away from night-time respiration. This may
explain the absence of an up-regulation in respiration by individuals from the low, stable pH treatment. However, respiratory CO₂ may not be exported as efficiently in low pH conditions, due to the higher CO₂ concentration in the surrounding acidified water, potentially further enhancing night-time dissolution.

Dissolution in dead algae

In dead thalli there is no protective epithelium and they cannot therefore buffer against changes in OA-associated carbonate chemistry, explaining their susceptibility to dissolution.

Wider implications of coralline algal damage

Calcifying algae are particularly important for ecosystem function in both temperate and tropical ecosystems through their roles in (i) carbon cycling; (ii) provision of habitats and associated biodiversity hotspots, (iii) association with recruitment processes; and (iv) being major structural components of coral reef systems (Nelson, 2009). Thus, any change in the three-dimensional structure and structural integrity of coralline algae may have significant effects on the ecosystem functions with which they are involved. For example live, structurally complex red coralline algae host high macro-organismal diversity (Biomaerl Team, 1999), and their ability to act as a nursery area is hierarchially controlled by the presence of a live epithelium and their complex three-dimensional skeletal heterogeneity (Kamenos et al., 2004a). Thus, while we and others (e.g., Martin et al., 2013) show evidence of increased calcification under OA, this does not necessarily result in repair of damaged skeletons, or mean that the newly calcified skeleton is as physically strong as that deposited under higher pH conditions. Weaker structural integrity may make coralline algae more prone to fragmentation, thus impacting their role in ecosystem function; this will likely vary among species and geographic locations.

When producing historical climate records from red coralline algae there is a requirement that no material is lost from the algal thallus post-deposition (Burdett et al., 2009). Such loss, for example due to grazing, significantly lowers the resolution of the reconstructed record, as the temporal constraints of material loss is unknown. In particular, this makes constructing a time series chronology particularly difficult. Our results suggest that while red coralline algae are alive, OA is not likely to affect the records they lay down at >1 day resolution due to enhanced day-time calcification.

We show that the extent of damage caused by low pH conditions is dependent on the rate of change in carbonate chemistry and live / dead status in algal biomineralisers. This is of major concern as both live and dead coralline algae are critical in service provision due to their complex three-dimensional skeletal heterogeneity (Kamenos et al., 2004a) and coral reef stabilisation / recruitment roles (Webster et al., 2013). In dead algae, the absence of post-dissolution repair mechanisms under projected OA conditions place the services they provide at significant risk. In addition, historical climate records held by dead algae within deposits are at risk from increased thallus dissolution, which is of particular concern as fossil deposits contain the longest palaeoclimatic records due to their age (Kamenos, 2010). It is thus likely that the ability of marine biomineralisers to cope with projected changes in marine carbonate chemistry and their ability to continue to provide services will be determined by the rate at which future pCO₂ enrichment occurs.

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Conflict of interest

We confirm that there are no conflicts of interest in the publication of this research.

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