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Lorenz Roth, Joachim Saur, Kurt D. Retherford, Darrell F. Strobel, John R. Spencer. Simulation of Io's Auroral Emission: Constraints on the Atmosphere in Eclipse. *Icarus*, 2011, 214 (2), pp.495. 10.1016/j.icarus.2011.05.014 . hal-00786874

HAL Id: hal-00786874

<https://hal.science/hal-00786874>

Submitted on 11 Feb 2013

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Accepted Manuscript

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PII: S0019-1035(11)00179-5
DOI: [10.1016/j.icarus.2011.05.014](https://doi.org/10.1016/j.icarus.2011.05.014)
Reference: YICAR 9821

To appear in: *Icarus*

Received Date: 18 January 2011
Revised Date: 10 May 2011
Accepted Date: 12 May 2011

Please cite this article as: Roth, L., Saur, J., Retherford, K.D., Strobel, D.F., Spencer, J.R., Simulation of Io's Auroral Emission: Constraints on the Atmosphere in Eclipse, *Icarus* (2011), doi: [10.1016/j.icarus.2011.05.014](https://doi.org/10.1016/j.icarus.2011.05.014)

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Simulation of Io's Auroral Emission: Constraints on the Atmosphere in Eclipse

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Abstract

We study the morphology of Io's aurora by comparing simulation results of a three-dimensional (3D) two-fluid plasma model to observations by the high-resolution Long-Range Reconnaissance Imager (LORRI) on-board the New Horizons spacecraft and by the Hubble Space Telescope Advanced Camera for Surveys (HST/ACS). In 2007, Io's auroral emission in eclipse has been observed simultaneously by LORRI and ACS and the observations revealed detailed features of the aurora, such as a huge glowing plume at the Tvashtar paterae close to the North pole. The auroral radiation is generated in Io's atmosphere by collisions between impinging magnetospheric electrons and various neutral gas components. We calculate the interaction of the magnetospheric plasma with Io's atmosphere-ionosphere and simulate the auroral emission. Our aurora model takes into account not only the direct influ-

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ence of the atmospheric distribution on the morphology and intensity of the emission, but also the indirect influence of the atmosphere on the plasma environment and thus on the exciting electrons. We find that the observed morphology in eclipse can be explained by a smooth (non-patchy) equatorial atmosphere with a vertical column density that corresponds to $\sim 10\%$ of the column density of the sunlit atmosphere. The atmosphere is asymmetric with two times higher density and extension on the downstream hemisphere. The auroral emission from the Tvashtar volcano enables us to constrain the plume gas content for the first time. According to our model, the observed intensity of the Tvashtar plume implies a mean column density of $\sim 5 \times 10^{15} \text{ cm}^{-2}$ for the plume region.

Key words:

Io, Jupiter, satellites, Atmospheres, structure

1. Introduction

Io probably possesses the densest and most species-rich atmosphere of the four Galilean satellites of Jupiter. Besides the main constituent sulfur dioxide various minor species, such as S, S₂, O, SO, Na, K and Cl, have already been observed in the vicinity of Io. Two possible sources for the atmosphere were discussed since its discovery. On the one hand volcanic venting as a direct source can create a neutral gas cloud around the satellite. On the other hand, sublimation of SO₂ frost from the surface by sunlight is considered a possible driver of the atmosphere (Lellouch et al., 2007). SO₂ frost abundance is found to be correlated with active volcanic regions (Douté et al., 2001). Thus, the surface distribution of the atmosphere is

related to the distribution of the volcanic regions on global scales, no matter if sublimation or direct outgassing is the main source. Prior observations of SO₂ suggested a volcanically driven atmosphere (e.g., Lellouch et al., 1992). Recent observations (Jessup et al., 2004; Moullet et al., 2010) and model results (Saur and Strobel, 2004) indicate that the sublimation driven part clearly dominates the direct outgassing.

The frequency of both the active volcanic regions and the paterae decreases with increasing latitude (Lopes-Gautier et al., 1999; Radebaugh et al., 2001) and observations of Lyman- α absorption in the atmosphere indicate that the SO₂ gas is concentrated likewise at lower latitudes (Strobel and Wolven, 2001). Feaga et al. (2009) derived a map of SO₂ column density, which shows a relatively sharp density decrease at approximately 30–45° north and south and a maximum column density on the anti-jovian hemisphere, which was first noted by Jessup et al. (2004) and Spencer et al. (2005) and later confirmed by Moullet et al. (2010). Walker et al. (2010) investigated effects of plasma heating as well as surface frost, molecular residence time and surface temperature distribution in a sophisticated SO₂ gas dynamics simulation, which includes sublimation and direct outgassing. The simulation results are compared to several observations with a backwards radiative transfer model in a companion paper (Gratiy et al., 2010) and basically confirm the previously derived atmospheric distributions assuming certain surface conditions. Considering results from modeling, millimeter, infrared and ultraviolet observations Lellouch et al. (2007) concluded that Io's atmosphere has a mean vertical column density of $\sim (1-5) \times 10^{16} \text{ cm}^{-2}$, covering 50–70% of Io's day-side hemisphere.

37 The evolution of the auroral emission, while Io passes through Jupiter's
 38 shadow, possibly provides very instructive information about the nature of
 39 the atmosphere. In eclipse the surface temperature drops and the subli-
 40 mation of SO₂ strongly decreases. Besides, no incident sunlight is reflected
 41 and thus solely the electron excited emission from the atmosphere is observ-
 42 able. Analyzing the aurora offers a possibility to determine the composition
 43 and distribution of Io's diverse atmospheric gas environment (Geissler et al.,
 44 2004).

45 Aurora (or airglow) is commonly defined as radiation caused by charged
 46 particles, which excite molecules or atoms in an atmosphere (Chamberlain
 47 and Hunten, 1987). In the case of Io, thermal electrons of the jovian magne-
 48 tosphere rotate with Jupiter's magnetic field and thus constantly flow past
 49 the slowly orbiting satellite and excite the atmospheric gas. The emission
 50 covers a broad wavelength range including ultraviolet, visible and infrared
 51 wavelengths and can be attributed to the various abundant species in Io's at-
 52 mosphere. Previous observations and analysis of Io's aurora offered insights
 53 into the satellite's atmosphere (e.g. Clarke et al., 1994; Roesler et al., 1999) .

54 Since the arrival of the Galileo probe at the jovian system, Io's auroral
 55 emission has been observed many times by ground-based and space tele-
 56 scopes (e.g., Roesler et al., 1999) as well as by on-board cameras of the
 57 Galileo (e.g., Geissler et al., 1999, 2001) and Cassini (e.g., Geissler et al.,
 58 2004) spacecrafts. Two of these observations are shown in Figure 1. Almost
 59 all observations are dominated by a key feature: bright spots close to the
 60 sub- and anti-jovian limb at low latitudes. These equatorial spots move up
 61 and down with the rocking background magnetic field of Jupiter (Retherford

et al., 2000). Geissler et al. (2001) also observed enhanced visible aurora close to volcanic plumes. Numerical simulations of the aurora by Saur et al. (2000) revealed that due to a diverted plasma flow the energy of the electrons is preferentially deposited on the flanks of Io and thus more energy reaches dense atmospheric layers around the sub- and anti-jovian points on the surface (see also aurora features in Figure 2). Therefore the bright equatorial spots are generated there. A tilt of the magnetic field leads to a tilted distribution of the magnetospheric electrons and the equatorial spots are thus correlated with the background field. Furthermore, auroral emissions appear to be brighter on the hemisphere facing the plasma torus centrifugal equator than on the other hemisphere (Retherford et al., 2003). Moore et al. (2010) found the same north/south dependency for a wake emission features simulating OI 630.0 nm emission observed by the HST Wide Field and Planetary Camera 2.

When Io enters Jupiter's shadow, two opposing effects control the evolution of this auroral radiation (Saur and Strobel, 2004). The atmospheric gas partly freezes out and thus less neutral gas can be excited by the ambient plasma. On the other hand, a decrease of atmospheric density after eclipse ingress leads to a decreasing interaction strength. The deflection and cooling of the plasma flow is lower and the streamlines of the electrons are less divergent. Saur and Strobel (2004) calculated the response of Io's electrodynamic interaction and radiation to a temporal change in the atmosphere. Depending on the total atmospheric decay, they find three qualitatively different scenarios with two of them including a transient post-eclipse brightening. Generally, the total emission intensity in eclipse was found to be lower than

87 out of eclipse . Retherford (2002) investigated the variation of key features of
 88 the eclipse aurora using HST Space Telescope Imaging Spectrograph (STIS)
 89 observations and inferred a reduction factor of ~ 1.5 to 2 for the emission close
 90 to Io during eclipse compared to the sunlit atmosphere. Geissler et al. (2004)
 91 in turn investigated the auroral emission during eclipse at different wave-
 92 lengths with Cassini filter observations. Comparing the observed brightness
 93 of an equatorial spot with modeled intensities they derived mixing ratios for
 94 various gases, which contribute to the equatorial emission, such as O, S, Na
 95 and K. Limb glows, which are attributed to minor components, indicate that
 96 O, Na and K are abundant all over the surface (Retherford, 2002; Geissler
 97 et al., 2004). Based on these eclipse observations Geissler et al. (2004) finally
 98 conclude that Io's atmosphere must be at least partly sustained directly by
 99 volcanism. Recent Monte Carlo simulations of the response to eclipse of
 100 the neutral gas by Moore et al. (2009) indicate a strong dependence on the
 101 abundance of non-condensable species during atmospheric collapse. In case
 102 of a mole-fraction of 0.35 of non-condensable gas a near-surface diffusion layer
 103 forms and the SO₂ density decreases only slowly and does not drop below
 104 0.18 of the initial column density during an eclipse event.

105 During the flyby of the New Horizons probe in 2007, the auroral emission
 106 was observed by the on-board camera LORRI (Spencer et al., 2007b) and
 107 by the Hubble Space Telescope Advanced Camera for Surveys (HST/ACS)
 108 (Retherford et al., 2007) (Figure 3a-c). The LORRI camera provides im-
 109 ages with high resolution comparable to Galileo SSI Io eclipse observations
 110 and reveals new details, such as the glowing plume of the Tvashtar volcano
 111 close to the North pole. This allows a detailed investigation of the aurora

112 morphology. With the simultaneous HST/ACS observation we are able to
 113 compare aurora morphologies from two different viewing geometries.

114 In our work we evaluate these observations regarding the morphology
 115 and the total intensity. Therefore we use a numerical model to calculate
 116 the interaction of the upstream plasma with Io's atmosphere-ionosphere and
 117 the auroral emission from the atmosphere. This method ensures that we
 118 take into account both effects, which influence the auroral response to an at-
 119 mospheric collapse during an eclipse event, as described above in this section.
 120 Investigating various atmospheric distributions, we derive one atmospheric
 121 distribution in eclipse, which is able to explain the observed intensity and
 122 morphology of the two LORRI observations and the HST/ACS image at the
 123 same time. Furthermore we infer the gas content of the Tvashtar plume. The
 124 constraints on the equatorial atmosphere and on the plume provide an esti-
 125 mation of the atmospheric collapse and a contribution of volcanic outgassing
 126 to the atmosphere.

127 2. Observations

128 In February and March 2007 the New Horizons probe passed the jovian
 129 System on its way to Pluto. The closest approach to Io occurred on 28 Feb-
 130 ruary at a distance of 2.24×10^6 km. Io was observed several times while
 131 passing through Jupiter's shadow (Retherford et al., 2007). During two of
 132 the eclipse passages LORRI, a narrow-angle high resolution camera on-board
 133 the spacecraft, took spectacular images of the auroral emission at visible and
 134 infrared wavelengths (Spencer et al., 2007a,b). According to the notation of
 135 Retherford et al. (2007) these two eclipse occasions will be denoted Ieclipse03

and Ieclipse04. During the Ieclipse03 event simultaneous spatially resolved observations of Io's far-ultraviolet aurora have been achieved by the Advanced Camera for Surveys Solar Blind Channel (ACS/SBC) on the Hubble Space Telescope. Characteristics of the eclipse occasions and the LORRI and HST/ACS observations are given in Tables 1 and 2. The geometry of the observations is depicted in Figure 2.

The NH/LORRI observations during eclipses Ieclipse04 and Ieclipse03 are shown in Figures 3a and 3b (in the order that they are discussed in the paper). The spatial resolution of LORRI is $4.96 \mu\text{rad}$ per pixel (Cheng et al., 2008). For Ieclipse03 this corresponds to a pixel size of $\sim 15 \text{ km} \times 15 \text{ km}$ (~ 200 pixels per Io diameter). During Ieclipse04 LORRI was used in the 4×4 binning mode, which reduces the spatial resolution in both directions by 4. The pixel size in that case is $\sim 56 \text{ km} \times 56 \text{ km}$, which is equivalent to ~ 50 pixels per Io diameter. The pass-band ranges from about 350 nm to 850 nm. The measured radiation originates from various emitters. In the blue range SO_2 and S_2 emission bands are the main contributors. Excited atomic oxygen, atomic sodium and potassium are the major emitters in the green, red and infra-red range (Geissler et al., 2004). Atomic sulfur contributes only a minor fraction.

The conversion from detector counts to emission brightness depends on the wavelength dependent sensitivity of the respective instrument. Since the contributing emissions in the wavelength ranges of LORRI and ACS are not clearly defined, we used Pivot wavelengths of 607.6 nm (LORRI) and 143.7 nm (ACS) to convert to Rayleighs (brightness). To validate this method we used exemplary modeled emission spectra including various emitters (SO_2 ,

161 O, S, Na and K) and calculated the brightness of the LORRI images using
 162 the sensitivity curve of Cheng et al. (2008). With this method the intensity
 163 differs less than 20% from the intensity derived with the Pivot wavelength.

164 We now describe and assign the essential features in the observation im-
 165 ages. During the Ieclipse03 observation (Figure 3b) the sub-spacecraft coor-
 166 dinates are $\sim 310^\circ\text{W}$ (west longitude, 0° towards Jupiter) and $\sim 7^\circ\text{S}$ (South-
 167 ern latitude, $\vartheta = 0^\circ$ in the orbital plane) in the Io-centric coordinate sys-
 168 tem. Hence, the sub-jovian upstream quadrant is visible during the observa-
 169 tion. The brightest small spots on the disk appear to be thermal emission.
 170 The measured intensity corresponds to a black body temperature of approx-
 171 imately 1200 K (Spencer et al., 2007b). Fainter small spots are low-altitude
 172 gas emission. On the upstream side (right-hand limb) enhanced radiation
 173 is visible only within a relatively small region, which has been attributed to
 174 a volcanic hot spot located east of the Girru paterae and therefore named
 175 “East Girru” (Spencer et al., 2007b). Considering the geometry of the back-
 176 ground magnetic field during Ieclipse03, which we discuss in detail below,
 177 the observed emission maximum around East Girru could also result from
 178 the magnetic field tilt and represent a shifted anti-jovian equatorial spot.
 179 The main part of anti-jovian spot probably is hidden behind Io. However,
 180 the East Girru emission appears to consist of one low and one high altitude
 181 emission part. Assuming that the plasma parameters do not vary discon-
 182 tinuously, the puzzling two part emission could originate from two different
 183 atmospheric species, as for example emission from SO_2 close to the surface
 184 and from atomic oxygen after dissociation of SO_2 at higher altitudes. Since
 185 a spectrally resolved observation is not available, this feature cannot be an-

186 analyzed for different emission lines. Around the downstream (left-hand) disk
 187 edge there is a diffuse emission region of approximately $500 \times 1,000 \text{ km}^2$.
 188 This region is the sub-jovian equatorial spot identified in previous observa-
 189 tions (e.g., Roesler et al., 1999). Apart from the equatorial band there is small
 190 enhanced emission just above the limb close to the North pole. The bright
 191 area can undoubtedly be assigned to the huge Tvashtar plume, which was
 192 first seen by Cassini in December 2000 (Porco et al., 2003). Unlike most of
 193 Io's volcanoes, Tvashtar is located at a high latitude near Io's North pole. As
 194 the global atmosphere is relatively thin at higher latitudes, Tvashtar creates
 195 a locally enhanced neutral atmosphere, which can be investigated separately
 196 from the equatorial atmosphere.

197 In the Ieclipse04 image (Figure 3a) the Tvashtar plume is clearly evident
 198 above the limb. The bright area is similar to the plume size derived from
 199 sunlit observations with a height of about 350 km and full width of 1100
 200 km (Spencer et al., 2007b). The geometry of this image (approx. 240°
 201 W 3° S) allows observation of both equatorial spots. While the sub-jovian
 202 spot (left-hand) is restricted to a small area, the anti-jovian spot extends
 203 further on the disk. The extent to which a few dozen lower intensity features
 204 located at known volcanic vent locations may include atmospheric emissions
 205 combined with the thermal emissions is yet to be determined. For Ieclipse04
 206 the deviation from the upstream view at 270° W, where both spots should be
 207 visible and similar in brightness, is $\sim 30^\circ$. Note that in the case of Ieclipse03
 208 the deviation from the upstream view is only $\sim 10^\circ$ more, but there is no
 209 clear anti-jovian spot observable (see description above).

210 The average of four consecutive exposures of Io in eclipse by HST/ACS

(Retherford et al., 2007) is displayed in Figure 3c. This image includes emission in the 125 nm to 190 nm band-pass, which originates mostly from excited atomic sulfur and oxygen. O and S are expected to resemble SO_2 in the region of interest for most of the analyzed emissions, although larger differences might occur at higher altitudes. The viewing angle from HST is $\sim 344^\circ$ W and $\sim 0^\circ$ N/S. Thus the complete sub-jovian side is visible, the plasma flow is directed from right (upstream) to left (downstream). The morphology is dominated by three features. The brightest area on the disk corresponds to the sub-jovian equatorial spot. The next brightest emission to the left of Io can either be attributed to the far end of the equatorial spots or originate directly in the wake of Io. And third, similar to the LORRI observation, the emission on the upstream (right-hand) side is considerably enhanced only above the East Girru region. Again, this enhancement can not definitely be assigned to a volcanic region, but might also originate from a smooth, continuous atmosphere and an inhomogeneous electron environment. As already mentioned, the NH Alice spectrograph additionally measured the total intensity of the two prominent oxygen multiplets at 130.4 nm and 135.6 nm and the sulfur multiplet at 147.9 nm during four eclipse events between 25 February and 3 March 2007 (see Table 3 and Retherford et al. (2007)).

Due to the tilt between Jupiter's dipole moment and Jupiter's rotation axis, the background field and the surrounding plasma density change while Io is moving up and down in the plasma torus during a synodic rotation period of Jupiter ($T_{syn} \approx 12.95$ h) with respect to Io. It takes about two hours for Io to pass through Jupiter's shadow, which is approximately one sixth of the

synodic period of the varying background field. All displayed observation images (Figures 3a-c) are combinations of several coaligned exposures during one eclipse event (averaging regions not contaminated with instrument scattered light), and thus include changes in the plasma environment between the exposures. Although we can not identify one exact observation geometry, we now roughly describe the plasma conditions during the LORRI and HST observations. We use an Io-centered coordinate system, where z is Jupiter (Io) North, and x is along the orbital direction of Io, i.e. approximately along the plasma flow (but rocks due to the tilt of the plasma torus). y completes the system pointing roughly (Io's orbit is slightly eccentric) towards Jupiter (Figure 2). The three components of $\vec{B}$ calculated with the model of Connerney et al. (1998) for Jupiter's internal field are listed in Table 1.

During Ieclipse03 Io is above the plasma torus and reaches the maximum distance to the torus center shortly before egress. The y component of the magnetic field vector is relatively large. The component in the direction of the orbital movement (B_x) in turn is low. The angle between the magnetic field and the polar axis (z) varies between $\sim 16^\circ$ and $\sim 19^\circ$. During Ieclipse04 Io passes through the torus center from North to South. Accordingly, the y component of $\vec{B}$ is lower and the x component somewhat larger. The tilt of $\vec{B}$ to the polar axis is around 10° . Furthermore, the ambient plasma density presumably is higher during Ieclipse04 in the torus center than during Ieclipse03, when Io is far away from the center (Bagenal, 1994). The field vectors in the xy plane during mid-eclipse are shown in Figure 2. The direction of the magnetic field for the observing geometry at mid-eclipse (Io in line with the center of the Sun and the center of Jupiter) is displayed in the

261 lower right corner in Figures 3a-c.

262 In all observations five areas are highlighted with green frames, which we
 263 will separately analyze in this paper. First, two areas of $\sim 9.2 \times 10^5 \text{ km}^2$
 264 around the equatorial spots, then a $\sim 6.4 \times 10^5 \text{ km}^2$ sized region centered at
 265 the calculated position of the Tvashtar plume. The large boxes ($\sim 3.8 \times 10^6$
 266 km^2) on the upper and lower edge of Io cover the polar areas, where the at-
 267 mosphere is expected to be less dense and thus less emission is expected. We
 268 integrate the total emission within the boxes and normalize it to the covered
 269 area. This method allows us to investigate different features of the auroral
 270 morphology quantitatively.

271

272 For an appropriate theoretical description of the formation of Io's aurora,
 273 we first need to calculate the interaction of the plasma particles with Io's
 274 environment. As explained in section 1, the density and distribution of the
 275 atmospheric gas influence this interaction, i.e. the neutral gas controls the
 276 flow pattern and the temperature and density profiles of the electrons. Since
 277 the electrons generate the auroral emission, electron temperature and den-
 278 sity are in addition to the atmospheric density the essential parameters for
 279 calculating the aurora. On the other hand, the distribution and local density
 280 of emitting gas particles is directly reflected by the intensity and morphology
 281 of the aurora. To calculate both effects self-consistently, we use a plasma
 282 model developed by Saur et al. (1999).

3. Model to interpret the observations

3.1. Plasma interaction model

The model, which we use, was developed to simulate the plasma interaction of the satellites Europa (Saur et al., 1998) and Io (Saur et al., 1999). It has undergone several improvements subsequently. The simulation results provide explanations for several observed features of Io's plasma environment such as magnetic field signatures in the wake and the rotated Alfvén wing system. Furthermore, Saur et al. (2000) were able to explain the formation of the bright equatorial spots of Io's aurora. For the full set of equations and the numerical algorithms we refer the reader to Saur et al. (1999, 2002). In the following section we explain the basics of the model and the treatment of the two plasma parameters, which are essential for aurora simulation: the density and temperature of the electrons.

The simulation is developed in the E, j approach of magnetohydrodynamics. The magnetic field is assumed to be a constant, homogeneous background field at all times. The electric conductivity parallel to the magnetic field is assumed to be infinite, so the parallel electric field vanishes. The validity of these assumptions are assessed in Neubauer (1998); Saur et al. (1999, 2002). With this model and its assumptions, Saur et al. (1999) were able to describe various aspects of the plasma interaction. In eclipse the local plasma interaction and thus the magnetic field perturbations are weaker than in the case of a sunlit atmosphere, i.e. the assumption of a homogeneous field is even more justified for our purposes. Galileo observations in sunlight revealed a magnetic field perturbation of more than $\frac{\Delta B}{B} = 0.3$ (e.g. Kivelson et al.,

1996), which were reproduced by our model (Saur et al., 2002). For the interaction in eclipse the model results indicate that the perturbation of the magnetic field is lower by a factor of 2 ($\frac{\Delta B}{B} \approx 0.15$ in eclipse). This justifies our assumption of a homogenous background magnetic field since the magnetic field environment is strongly dominated by the background field, while other plasma parameters such as velocity and electric field vary significantly due to the interaction with the atmosphere.

In the undisturbed upstream plasma the homogeneous electric field is given simply by $\vec{E}_0 = -\vec{v}_0 \times \vec{B}_0$. Inside Io's ionosphere a current system arises from the collisions between the plasma and neutral gas particles, which modify the electric field. The modified $\vec{E}$ field is calculated by a differential equation for the 2D electric potential in the plane perpendicular to $\vec{B}_0$ first derived by Wolf-Gladrow et al. (1987). The electric potential and thus the electric field around Io are calculated from the ionospheric Hall and Pedersen conductances as well as the Alfvén conductance. The undisturbed plasma velocity $\vec{v}_0$ is assumed to be perpendicular to $\vec{B}_0$. Neglecting inertia and pressure, the electron velocity $\vec{v}_e(x, y)$ in the plane perpendicular to $\vec{B}_0$ can be derived directly from the electric field. To simulate the auroral emission we need to calculate the properties of the thermal electron population in the vicinity of Io, which excite the aurora. Moving with the electron flow, the evolution of the density as well as the temperature of the electrons can be calculated as described below.

The coordinate system of the simulation corresponds to the system that we defined in section 2. The magnetic field however is constant and always anti-parallel to the z axis and the plasma flow v_0 is parallel to x . Thus, we

take into account neither changes of the inflow direction nor the rocking of the magnetic field due to the tilt of 9.6° of the jovian dipole field axis with respect to the jovian rotation axis. The geometry is discussed in detail in section 3.4.

The evolution of the electron density n_e is described by the following continuity equation:

$$\frac{d}{dt}n_e = f_{ion} n_e n_{SO_2} + k_{hee} n_{SO_2} - \alpha n_e^2 \quad (1)$$

The first term on the right hand side describes (single) ionization by electron impact on neutral gas n_{SO_2} due to thermal electrons (Sittler and Strobel, 1987) with the temperature dependent collisional ionization rate f_{ion} .

With the second term we account for ionization by kappa-distributed energetic bidirectional electrons (k_{hee}) observed in Io's vicinity (e.g., Williams et al., 1999). The electron energies in the beams span the range from keV to hundreds of keV (Williams et al., 1996; Frank and Paterson, 2002). According to the simulated electron beam morphology by Jacobsen et al. (2010) and the observations during various Galileo flybys (Williams and Thorne, 2003; Frank and Paterson, 2002) we assume a spatial distribution of the high energetic electrons given by

$$f_{hee} = f_{hee,0} [\tanh(3x + 1) + 1] \exp(-10y^6) \cdot c(x, y) \quad (2)$$

with

$$\begin{aligned} c &= 1 & \text{for } x^2 + y^2 > 1 R_{Io} \\ c &= 0.5 & \text{for } x^2 + y^2 \leq 1 R_{Io}. \end{aligned}$$

330 Thus, the beams are assumed to be relatively narrow (width in y direction:
331 $\sim 1.4 R_{Io}$) but to extend far into the wake. However, outside the atmosphere

(e.g. in the far wake), where the neutral density vanishes, no ionization can take place anyway. Using the geometrical factor $c(x, y)$ we take into account that directly above (below) Io the electron beam from the South (North) is shielded by the satellite. For the energy flux we use the derived value by Saur et al. (2002). Due to the absence of observations and models of electron beams when Io is eclipse, we note that expression (2) is derived from observations and models when Io is in sun light.

The third term on the RHS of equation (1) describes the loss due to recombination with a rate α . For the adopted parameter values we refer the reader to Saur et al. (1999, 2002).

As the electron velocity in the model is simply given by the $\vec{E} \times \vec{B}$ -drift, there is no plasma movement along $\vec{B}$ included. But, in fact plasma transport along $\vec{B}$ does occur, particularly in the wake of Io. The only electron source in the wake is the newly ionized population through the electron beams. For a low density atmosphere as in eclipse the ionization by electron beams is mainly confined to the region close to the equator, where the atmosphere is densest. Altogether, the wake region will be rather void, if movement along $\vec{B}$ is neglected. However, parallel movement due to the pressure gradient in the wake of Io possibly fills the relatively void regions in the downstream area and might thus enhance the aurora in the downstream region. To account for the parallel movement, we modified the model assuming that the plasma particles move along $\vec{B}$ with thermal velocity v_{th} .

The total time derivative on the left hand side of (1) can be written as the partial time derivative and the convective term $\vec{v}_e \cdot \nabla n_e$. Separating the flow in the xy plane from the movement along z (i.e. $\vec{v}_e \cdot \nabla n_e = \vec{v}_\perp \cdot \nabla_\perp n_e + \vec{v}_\parallel \cdot \nabla_\parallel n_e$)

we rewrite the continuity equation as

$$\vec{v}_e(x, y) \cdot \nabla_{\perp} n_e = f_{ion} n_e n_{SO_2} + k_{hee} n_{SO_2} - \alpha n_e^2 - v_{th} \nabla_{\parallel} n_e, \quad (3)$$

$\vec{v}_{\perp}$ is the electron flow $\vec{v}_e(x, y)$ and the parallel flow $\vec{v}_{\parallel}$ along z is approximated by the thermal velocity v_{th} . Thus, following the 2D electron flow $\vec{v}_e(x, y)$ a parallel flow as loss or production term depending on the density gradient along z is calculated. Numerically, we consider the net flow from or to both the grid cell above and below the current position. The flow direction is determined by the sign of the respective electron density gradient between the cells. This description is somewhat similar to a diffusive process. The thermal parallel velocity v_{th} is assumed to be the ion sound speed, as the inertia of the ions mainly determines the movement of the plasma. The ion sound speed is approximately half of the undisturbed relative flow velocity of the plasma (Kivelson et al., 2004), so the flow fills the wake at an angle of approximately $\arcsin\left(\frac{v_{th}}{v_e(x, y)}\right) \geq 30^\circ$. The propagation along $\vec{B}_0$ enables the plasma to fill the wake of Io, which is important to explain the observed auroral emission in the downstream region. Expanding into low density regions such as Io's wake the plasma can be accelerated to velocities higher than the thermal velocity (Samir et al., 1983). By the assumption of a parallel movement with v_{th} , the propagation speed into the wake might thus be underestimated.

The second important plasma quantity for aurora simulation is the thermal energy or temperature of the electrons, T_e . The temperature evolution is given by

$$\begin{aligned} \frac{3}{2} k_B n_e \frac{dT_e}{dt} = & -(\epsilon_{ion} f_{ion} + \epsilon_{dis} f_{dis} + \epsilon_{rot} f_{rot} + \epsilon_{vib} f_{vib}) \cdot n_e n_{SO_2} \\ & - \frac{3}{2} k_B T_e (f_{ion} n_e n_{SO_2} + k_{hee} n_{SO_2}) - \nabla Q_{flux}. \end{aligned} \quad (4)$$

371 We account for cooling by inelastic collisions between the magnetospheric
 372 electrons and the atmosphere, including ionization, dissociation, and rota-
 373 tional and vibrational excitation of neutral SO₂. Each process κ is described
 374 by the rate f_κ and the energy quantum ϵ_κ . The newly added electrons re-
 375 sulting from the impact ionization processes are assumed to be cold. The
 376 adjustment to magnetospheric bulk temperature of these cold electrons leads
 377 to a decrease in temperature described by the next term on the right hand
 378 side. k_B is the Boltzmann constant.

The heat flux Q_{flux} is parametrized. We take advantage of the anisotropy of the thermal conductivity and assume no heat flow perpendicular to $\vec{B}$ due to the strong background magnetic field of ~ 2000 nT (Banks and Kockarts, 1973). Parallel to $\vec{B}$ the heat conduction is extremely high. We consider the electrons in a flux tube along $\vec{B}$ outside the atmosphere to adjust to one common temperature T_{out} instantaneously. Deep inside the atmosphere the parallel thermal conductivity is lower, where the mixing ratio of neutral gas to plasma increases. When the flux tube passes through the atmosphere, we divide it in three parts, two outer parts (T_{out}) and an inner part (T_{in}). The inner part uniformly cools down due to the various collision processes described in equation 4. The heat flow from the hotter outer part to the inner part Q_{flux} is parametrized as sketched in Figure 4. Depending on the temperature dependent heat conductivity $\kappa(T_{in})$ and the temperature difference between the inner and outer parts a typical heat flow is calculated:

$$< Q_{flux} > = < \kappa(T_{in}) > \frac{T_{out} - T_{in}}{R_{typ}}. \quad (5)$$

379 The average heat transport is controlled by a typical distance R_{typ} , which is
 380 set to the scale height of the atmosphere (100 km). The electron heat conduc-

381 tivity κ in a plasma with a fraction of neutral gas as function of the ambient
 382 temperature is given by equation (22.116) of Banks and Kockarts (1973) and
 383 depends on the momentum transfer cross section for elastic collisions of elec-
 384 trons with SO_2 gas and the mixing ratio n_{SO_2}/n_e . The temperature in the
 385 outer part T_{out} decreases due to the heat flow to the cooler inner part, i.e.
 386 the outer parts serve as finite heat reservoirs and the energy of the entire flux
 387 tube is depleted. The energy capacity of these reservoirs corresponds to the
 388 electron content in the region along the magnetic field line above and below
 389 Io. For further details see Saur et al. (2002), Appendix A.

390 3.2. Atmosphere model

Based on several observations (see section 1) of Io's atmospheric distribu-
 tion, we assume a dense atmospheric ring ranging from the equator to approx-
 imately 35° North and South. At higher latitude a low density background
 n_{bg} is assumed, with a ratio of $n_{bg}/n_{eq} = 0.02$ for SO_2 . This distribution was
 calculated by Strobel and Wolven (2001) based on Lyman- α reflection ob-
 servations. Here we investigate longitudinal asymmetries, such as differences
 between sub- and anti-jovian hemispheres reported by Spencer et al. (2005),
 Feaga et al. (2009) and Moullet et al. (2010) as well as a denser downstream
 atmosphere as inferred by Saur et al. (2002). The surface density n_s is thus
 modeled by

$$n_s(\vartheta, \varphi) = n_{bg} + (n_{eq} - n_{bg}) (1 + \beta \cos(\varphi - \gamma)) e^{-\left(\frac{\vartheta}{35^\circ}\right)^6} \quad (6)$$

with the latitude ϑ and Jupiter oriented longitude φ with $\varphi = 0$ for direction
 to Jupiter and $\varphi = 90^\circ$ in Io's orbital direction. n_{eq} is the surface density at
 the equator, β and γ specify the strength and orientation of the longitudinal

inhomogeneity and are fitted to match the observations. The number density declines exponentially with increasing altitude. The vertical structure is determined by the scale height H_s . We assume the scale height to vary with φ according to the longitudinal variation of the surface density:

$$H_s(\varphi) = H_{s,0} (1 + \beta \cos(\varphi - \gamma)). \quad (7)$$

391 The simultaneous variation of scale height and surface temperature allows an
392 implementation of a larger difference in column density between two hemi-
393 spheres with smaller local gradients than choosing a strongly varying surface
394 density at a constant scale height. When the local density gradients are too
395 large, the simulation becomes unstable.

396 For modeling the plasma interaction, we assume the atmosphere to con-
397 sist solely of SO_2 , since it supposably is the by far main constituent (Lellouch
398 et al., 2007). The surface scale height $H_{s,0}$ in (7) for SO_2 is set to 100 km
399 (Saur et al., 1999). Near the surface the actual scale height is presumably
400 lower in the range of tens of km. But, for the low density eclipse atmosphere
401 the plasma likely penetrates all atmospheric layers and the model accounts
402 for integrated conductivities. At high altitudes a large scale height is ex-
403 pected. Thus, a scale height of 100 km represents an average scale height
404 for Io's atmosphere and the interaction strength still constraints the global
405 column density, since the crucial parameter for the interaction is the total
406 atmospheric gas content.

407 The longitudinal distribution of the equatorial atmosphere in eclipse,
408 which is modeled here, does not necessarily correspond to the longitudi-
409 nal distribution at daytime seen around a full Io orbit. The fractional at-
410 mospheric collapse might vary with longitude due to different atmosphere

411 and surface conditions or the position of the observer and the sun. For
 412 instance, the atmosphere may respond differently on the sub-Jupiter hemi-
 413 sphere, which is sunlit before eclipse ingress, than on the anti-jovian hemi-
 414 sphere, which is at night for several hours already before eclipse.

As the LORRI observations revealed a bright volcanic plume close to the North pole, we additionally implement a plume shaped density enhancement located at the Tvashtar paterae at 62°N and 122°W (Spencer et al., 2007b). The distribution within the plume is modeled by

$$\begin{aligned}
 n_V(h, d) = n_{V,0} & \left(\exp \left[- \left(\left(\frac{h}{H_V} \right)^2 + \left(\frac{d}{\sigma_V} \right)^2 \right)^3 \right] \right. \\
 & \left. - \exp \left[- \left(\left(\frac{h}{0.4H_V} \right)^2 + \left(\frac{d-d_0}{0.4\sigma_V} \right)^{\frac{3}{2}} \right)^3 \right] \right) \\
 & + 30 n_{V,0} \exp \left[- \left(\left(\frac{h}{0.1H_V} \right)^2 + \left(\frac{d}{0.05\sigma_V} \right)^2 \right)^3 \right]
 \end{aligned} \tag{8}$$

415 where $n_{V,0}$ is the density in the center of Tvashtar, h the vertical distance from
 416 surface and d the horizontal distance on the surface to the plume center. The
 417 height H_V and width σ_V are given by the observed plume extent in sunlight
 418 and are set to $H_V = 360$ km and $\sigma_V = 550$ km. The subtracted exponential
 419 function (with height $0.4 H_V$ and width $0.4 \sigma_V$) in the second line roughly
 420 describes the expected low density region within the plume after Zhang et al.
 421 (2003). We also include a high density region above the vent by adding the
 422 third line in the equation. The plume model of Zhang et al. (2003) also
 423 indicates the formation of a canopy shock, which we do not account for in
 424 our simulation. $n_{V,0}$ is the only free parameter of the plume. The three plume
 425 regions are marked in Figure 5, where a cross section through the volcano

density and the column density above the plume are shown.

To model the radiated emission with the simulated electron densities and temperatures all the essential atmospheric constituents are considered individually. We will derive a mixing ratio n_{comp}/n_{total} for the atmospheric components O, S, Na and K. For the simulation of the LORRI images the mixing ratio of the species is assumed to be identical in the equatorial atmosphere and in the plume. For atomic sulfur and oxygen the radial decrease is smaller than for SO₂ (Wolven et al., 2001; Summers and Strobel, 1996). The Cassini observations of Geissler et al. (2004) revealed O assigned emission at higher surface distances than the emission assigned to SO₂, see Figure 1a. This indicates a shallower decrease of O compared to SO₂. For atomic sodium, HST observations of the NaI 589 nm line revealed a very shallow drop-off with increasing height (Retherford, 2002). Therefore, we assume a slightly larger scale height for these constituents of $H_{s,0}(X) > H_{s,0}(\text{SO}_2)$ (see table 4). Moreover, we assume a higher n_{bg}/n_{eq} ratio for atomic species, as the Cassini observations (Geissler et al., 2004) as well as HST observations (Retherford et al., 2000, 2003) show a clear limb glow all around Io for emission from atomic species.

3.3. Emission simulation

Assuming equilibrium, the local intensity of the stimulated emission of an atmospheric gas is calculated by

$$i_{\lambda}(\vec{x}) = f_{\lambda}(T_e(\vec{x})) n_{gas}(\vec{x}) n_e(\vec{x}), \quad (9)$$

where f_{λ} is the gas specific rate, n_{gas} the gas density and n_e the density of the exciting electrons. Assuming spontaneous emission the emission rate cor-

responds to the collisional excitation rate, which is given as an integral over the Maxwell-Boltzmann distribution, the electron velocity and the energy-dependent cross section for the collision of the exciting electrons with the neutral species. For optically thin emission lines a 2D emission pattern is given by the line-of-sight integral over the local intensities. Not all emissions can be considered optically thin. Therefore we estimate re-absorption for Na and K. All other emission lines are assumed to be optically thin in our analysis. Due to the long lifetime or small radiative decay rate of $k_d = 0.00681 \text{ s}^{-1}$ of the $\text{O}(^1D)$ state, we include collisional quenching at low altitudes for the OI 630.0/636.4 nm emission. The lowered intensity I is calculated with the Stern-Vollmer relationship

$$I = I_0 \frac{k_d}{k_d + k_q} \quad (10)$$

where k_q is the quenching rate and I_0 the intensity without quenching. For k_q we use the estimated value of $1 \times 10^{-10} \text{ cm}^3 \text{ s}^{-1}$ from Geissler et al. (2004). Quenching and radiative decay are equally probable when the SO_2 number density is $6.8 \times 10^7 \text{ cm}^{-3}$.

The NH/LORRI observations cover wavelengths from 350 to 850 nm. Within this range emission from SO_2 bands and S_2 bands as well as from atomic oxygen (557.7 nm, 630.0 nm, 636.4 nm, 777.4 nm and 844.6 nm), atomic sulfur (772.5 nm), atomic sodium (588.9/589.6 nm) and potassium (767.0 nm) are expected. The OI 777.4 nm and OI 844.6 nm lines contribute less than 1% to the total emission and thus can be neglected (Geissler et al., 2004). The electron impact excitation cross sections that we use for atomic oxygen emission are based on the laboratory measurements of Doering and Gulcicek (1989a,b) and Doering (1992). The cross sections for the sodium

and potassium D lines were adopted from theoretical calculations by Kim (2001), which the author compares to experimental data from Enemark and Gallagher (1972). For further details on the cross sections and the associated rates see Geissler et al. (2004) and references therein. Cross sections for electron excitation of S_2 are not measured, so we can only estimate the contribution of S_2 emission roughly.

A large number of emission lines can be found in the pass-band of the HST/ACS (125–190 nm) observation, such as emission from atomic oxygen and sulfur as well as from sulfur ions (Ballester et al., 1987; Roesler et al., 1999) and also chlorine emission (Feaga et al., 2004). Since cross sections are not available for all of the lines, we solely analyze the oxygen multiplet at 135.6 nm and the sulfur multiplet at 147.9 nm, which presumably contribute most. Neglecting all the lower emissions in the pass-band, we can not compare the emission quantitatively with the ACS observation. A quantitative comparison also would necessitate an accurate analysis of optical depth of various lines. However, by simulating the OI 135.6 nm and SI 147.9 emission we are able to compare the model results to the ACS observation regarding the morphology and the relative intensities of the individually analyzed areas, since oxygen and sulfur generally are by far the main contributors in the observed wavelength range. The contribution of the OI 130.4 nm to the UV aurora is not clear yet and is therefore not considered here. For the sulfur emission at 147.9 nm we account for both the forbidden and allowed lines (Feaga et al., 2002). For both the SI147.9 nm and SI190.0 nm emission the adopted cross sections are based on the calculated collisions strengths from Zatsarinny and Tayal (2002).

483 3.4. *Simulation setup*

484 A constant magnetic field and a 2D plasma flow perpendicular to B_0 are
 485 the basic assumptions of the theoretical approach. So it is not possible to
 486 fully consider the varying plasma conditions around Io in the simulation. In
 487 the model the constant background field is parallel to the z axis and the
 488 upstream flow is in positive x direction. Moreover, the model is symmetric
 489 with respect to the xy plane given by $z = 0$. For the implementation of
 490 the volcanic plume on the North pole, we run simulations with and without
 491 plume and combine the results. The influence of the plume on the aurora in
 492 the equator region appears to be negligible ($< 1\%$). Note, the asymmetry of
 493 Io's atmosphere due to Tvashtar near Io's North pole generates also a small
 494 asymmetry in Io's plasma interaction. However, due to the small spatial
 495 extend of the plume compared to Io's diameter and the relative amount of
 496 the plume gas compared to the total gas of Io's atmosphere, the asymmetry
 497 in the plasma interaction is negligible in contrast to Enceladus' interaction,
 498 where the asymmetry plays an important role (Saur et al., 2007). We try
 499 to explain various observations with one atmospheric distribution, which is
 500 assumed to be symmetric around Io's equator due to the mentioned basic
 501 symmetry of the model around the equator. Hence, the orientation of the
 502 magnetic field and the upstream plasma flow are fixed in the model.

503 The actual plasma environment varies during a synodic rotation of Jupiter.
 504 The angle between the undisturbed plasma and the plane perpendicular to
 505 $\vec{B}_0$ varies only within -3° and 3° , which is negligible. The variation of the
 506 direction of the magnetic field with respect to the polar axis of Io is larger (up
 507 to 20°). As the simulation code is symmetric with respect to the equatorial

plane, which is always assumed to be perpendicular to the background field, we are not able to investigate a tilt between the atmospheric ring and the orbital plane. Since the displayed observations are combinations of several exposures, the geometry is also varying for the single exposures. To minimize the effect of the simplified model geometry, we only analyze larger areas around the equatorial spots and Tvashtar quantitatively. Retherford et al. (2000) showed that the inclination of the spots is somewhat lower than the tilt of the magnetic field.

We account for the variation of the electron density in the upstream plasma. For Ieclipse03 simulation, where Io is far from the torus center we assume a lower density of $n_e = 1900 \text{ cm}^{-3}$, while during Ieclipse04 (Io crosses the torus center) the electron density is presumably higher and we use $n_e = 3600 \text{ cm}^{-3}$ (Gurnett et al., 2001b,a; Bagenal, 1994). The electron temperature of the upstream plasma is $T_{e,0} = 5 \text{ eV}$ (Bagenal, 1994). For the initial velocity we assume the relative azimuthal velocity of the rotating plasma ($\vec{v}_0 = 57 \text{ km s}^{-1}$), the background field is set to $B_0 = 2050 \text{ nT}$.

The viewing geometry of the displayed 2D emission patterns is given by longitude φ and latitude ϑ , where $\vartheta = 0^\circ$, when the observer is in the orbital plane, and $\vartheta = 90^\circ$, when viewed from above the North pole.

As discussed previously in section 2, we compare the simulated images with the observations using the green frame regions in Figure 3.

4. Results

First, we derive the total gas content and its distribution, which are determined mainly by SO_2 abundance. Thereafter we briefly describe our

532 derived mixing ratios and distributions for atomic oxygen and sulfur as well
 533 as the trace elements sodium and potassium. The emission from the atomic
 534 species also contributes to the intensity and morphology of the LORRI sim-
 535 ulations, so the derivation of the abundances for SO₂ and those of the minor
 536 components are correlated to each other. Therefore, we obtain the best-fit
 537 atmosphere results from an iterative variation of both the total gas content
 538 and the mixing ratios of the minor species. Finally, we constrain the gas
 539 content of the Tvashtar plume.

540 4.1. Equatorial Atmosphere

541 The main benchmark for constraining the density of the atmospheric ring
 542 was the morphology and location of the equatorial spots. We assume the
 543 same longitudinal variations for SO₂ and the minor components.

544 First, we analyze the brightness of the equatorial spots within the green
 545 boxes of the LORRI observations. In the LORRI Ieclipse04 observation
 546 (Figure 3a) the ratio of the anti-jovian to the sub-jovian spot brightness
 547 is $I_{anti}/I_{sub} \approx 1.8$. For LORRI Ieclipse03 (Figure 3b) the anti-jovian spot
 548 can not be identified clearly. The spot might be displaced toward the East
 549 Girru feature. The ratio for the equator centered boxes in Figure 3b is
 550 $I_{anti}/I_{sub} \approx 0.2$. We get the best morphology agreement with both observa-
 551 tions for a model atmosphere with an average equatorial column density of
 552 $N_{eq,av} = 2.3 \times 10^{15} \text{ cm}^{-2}$. We find a longitudinal asymmetry in favor of the
 553 downstream side with the asymmetry parameters $\beta = \frac{1}{3}$ and $\gamma = 90^\circ$. The
 554 column density on the upstream side is thus $N_{eq,up} = 1.0 \times 10^{15} \text{ cm}^{-2}$, the
 555 downstream column density is $N_{eq,down} = 4.0 \times 10^{15} \text{ cm}^{-2}$. The simulated
 556 aurora morphologies corresponding to the LORRI images are displayed in

557 Figures 3d and 3e, where the brightness ratios are $I_{anti}/I_{sub} \approx 2.1$ (Ieclipse04)
 558 and $I_{anti}/I_{sub} \approx 0.4$ (Ieclipse03). The derived atmospheric distribution also
 559 yields a morphology that is in good agreement with the HST Ieclipse03 ob-
 560 servation on the sub-jovian side. Furthermore, the key features of previous
 561 eclipse observations by Retherford (2002) are reproduced by the simulation
 562 considering the respective observation geometries (not shown here). The de-
 563 rived eclipse column density corresponds to $\sim 10\%$ of the column density
 564 for a sunlit atmosphere summarized in Lellouch et al. (2007) ($N_{sun} = (1-$
 565 $5) \times 10^{16} \text{ cm}^{-2}$). This day-side column density also coincides with the col-
 566 umn densities derived by Saur et al. (2000) and Saur et al. (2002), where the
 567 same simulation model for a sunlit atmosphere interaction is applied. The
 568 parameters of the best-fit atmosphere are listed in Table 4. The resulting
 569 absolute simulated and measured intensities in Rayleighs (R) of the marked
 570 areas are listed in Table 5 for comparison. The relative intensities refer to
 571 the absolute values divided by the total intensity. The total intensity I_{tot} in
 572 Rayleighs (R) corresponds to the total measured emission, but averaged to
 573 the area of Io's disk.

574 In order to illustrate the formation of the aurora morphology we show
 575 results of the plasma interaction simulation for Ieclipse03 in Figure 6: the
 576 atmospheric electron temperature in the flux tubes (Fig. 6a), the electron
 577 density in the equatorial plane (Fig. 6b) and the electric current in the
 578 Northern Alfvén wing (Fig. 6c). Due to the relatively low density in the
 579 equatorial atmosphere the energy stored in the flux tubes is sufficient to
 580 keep the electron temperature in most regions between 4 and 5 eV, and
 581 above 1 eV everywhere in the interaction region. The lowest temperature

582 is found on the flanks, where the electron density is highest. These high
 583 density regions in the anti- and sub-jovian equatorial region mainly control
 584 the position of the aurora equator spots. Depending on the overall density
 585 of the equatorial atmosphere the region of the maximum ionospheric density
 586 on the flanks and thus the aurora spots are shifted along the flow direction:
 587 for a denser atmosphere the maxima form further upstream, for a lower at-
 588 mospheric density the spots move downstream. The low atmospheric density
 589 and the gas plume around Tvashtar also cause a complex current pattern in
 590 the Alfvén wing. In addition to the usual current system in the Northern
 591 wing, i.e. sub-jovian downward and anti-jovian upward currents, contrarily
 592 oriented currents evolve on both sides within the outer usual system. Because
 593 of the low neutral density around the poles and the overall low atmospheric
 594 density, only on the flanks a dense ionosphere forms. Consequently, the
 595 current system is inhomogeneously short-circuited in the atmosphere. Two
 596 separate current systems form on the anti- and sub-jovian hemisphere respec-
 597 tively. Furthermore, a small "winglet" is generated around the plume of the
 598 Tvashtar volcano, which we further discuss in Section 4.3. Simulation results
 599 for a similar model setup but a dense global atmosphere ($N = 6 \times 10^{16} \text{ cm}^{-3}$)
 600 can be found in Saur et al. (1999).

601 In the case of a lower atmospheric surface density and thus column den-
 602 sity ($N_{eq,av} < 1 \times 10^{15} \text{ cm}^{-2}$) the formation of high electron density regions
 603 (ionosphere) is weaker. The electrons also cool down less, but this plays a
 604 minor role at the given atmospheric thickness, since the cooling is already
 605 relatively low for the best-fit atmosphere. In the case of a lower density and a
 606 resulting lower ionosphere the diversion of the plasma flow decreases as well.

Thus the equatorial spots move closer to the surface and further downstream. In this case the sub-jovian (left-hand side) spot in Figure 3d disappears behind the limb. Furthermore, the total intensity of the simulated emission averaged to the disk becomes far too low compared to the observations.

In the case of a higher equatorial column density of $N_{eq,av} > 3 \times 10^{15} \text{ cm}^{-2}$ the electron density on the upstream hemisphere increases due to increasing impact ionization, the electron flow is increasingly diverted around the body, and the cooling of the electrons is stronger and less energy reaches the downstream hemisphere. The equatorial spots are centered at larger distance from Io and the emission on the upstream side gets stronger compared to downstream side, as the plasma in the wake is completely cooled down. Thus, for such a high column density the emission morphology is not in agreement with the observation.

Assuming a longitudinal asymmetry with a higher and denser atmosphere on the anti-jovian hemisphere, as inferred for the day-side atmosphere (Jessup et al., 2004; Spencer et al., 2005; Feaga et al., 2009), yields a much stronger anti-jovian spot at larger radial distance. This would imply clear visibility of the spot during Ieclipse03 in Figure 3b, where no emission on the anti-jovian limb is detected except for the East Girru region. Furthermore, during Ieclipse04 the anti-jovian spot would be brighter and at a considerably higher altitude above the surface than the sub-jovian in the LORRI observation (Figure 3a). As the anti-jovian emission maximum is only slightly higher and very close to the limb, there is no indication for an atmospheric anti-sub-jovian asymmetry in eclipse. The non-appearance of an atmospheric bulge on the anti-Jupiter side could be explained by the fact that it is nighttime

on the anti-jovian hemisphere before eclipse ingress. The response of the atmosphere to a 21-hour night compared to a 2-hour eclipse event is likely to be different and the anti-jovian atmospheric bulge might be reduced due to that. Numerical calculations by Wong and Smyth (2000) indicate that non-condensable species might even dominate the night-side atmosphere.

Instead, the simulation results imply a longitudinal asymmetry of the atmosphere with increasing scale height and surface density from upstream to downstream. Compared to a longitudinally symmetric distribution, the emission is shifted downstream, which leads to a lower intensity on the upstream hemisphere and increasing intensity in the wake. The simulation still predicts more radiation to be emitted on the upstream than on the downstream side, although the downstream emission is more located and thus peaks higher. The observing geometry of the LORRI observations does not allow a separation of emission from the two hemispheres, but the emission appears to be brighter in the downstream region. However, larger datasets containing a range of exposures observing both upstream and downstream generally confirm a higher emission on the upstream hemisphere (Oliversen et al., 2001; Retherford, 2002) in agreement with our simulation results.

In the HST/ACS eclipse observation (Figure 3c) on the upstream hemisphere (right limb) aurora is observable only around East Girru, the corresponding simulation (Figure 3f) yields clear emission on the limb of the upstream equatorial side. In this region the fluid approach and the assumption of a constant background field might not describe the behavior of the electrons with the required accuracy. A highly distorted and piled up magnetic field in this region may hinder the electrons to move along $\vec{B}$ into the

atmosphere near the equator. The electrons with pitch angles close to 90° are reflected when the $\vec{B}$ increases and they do not reach thicker atmosphere layers. Moore et al. (2010) obtained this effect in their particle simulation. However, flux tubes with large $\vec{B}$ due to the upstream pile up have a smaller cross section close to Io. As these flux tubes are connected to a larger cross section further away from Io, a smaller area is linked to a larger energy reservoir. The combination of the both effects require further investigation with kinetic models. In the downstream region, the large cavity behind Io may lead to a faster filling of the wake than we assumed. This would cause higher emission in the wake.

Nonetheless, the position and size of the sub-jovian spot and the wake feature are in good agreement for the simulation and the HST/ACS observation (Figures 3c and 3f). The so-called wake feature is composed of two components. First, emission that is stimulated by electrons, which emerge from ionization by highly energetic beams directly ‘behind’ Io. Additionally, the tail of the spots can contribute to what is observed as wake emission, depending on the exact viewing geometry. In our simulation the contribution from the flanks is stronger. HST observations with viewing longitude between $\sim 60^\circ$ and $\sim 70^\circ$ (Figure 3.7 in Retherford, 2002) also revealed brighter flanks than radiation directly in the wake.

The longitudinal positions of the simulated equatorial spots range from $\sim 5^\circ$ to $\sim 40^\circ$ downstream from the zero meridian, which is roughly sketched in Figure 2. This is also in agreement with the derived position by Retherford et al. (2000) for HST observations of the day-side atmosphere (10° – 30°). The modeled brightness ratio of the anti-jovian to the sub-jovian spot as a

function of the longitude of the observer is shown in Figure 7. Additionally, we plotted the brightness ratio of the OI 135.6 nm multiplet from observations by Retherford (2002), when Io was sunlit. These day-side observations and the eclipse simulation should not be compared directly here, but it can be pointed out that the model reproduces the basic longitudinal distribution of the aurora. The observed and simulated ratios are in agreement on the upstream hemisphere. For an observing longitude φ of 90° (into the wake) and 270° (upstream hemisphere) the ratio is ~ 1 .

On the upstream side the ratio varies roughly linearly with φ . In a low density atmosphere, e.g. during an eclipse, we would expect the spots at smaller surface distances and thus the variation to be steeper than out of eclipse. The simulated slope might therefore be too gradual due to an over-estimation of the scale height in the model, which is presumably lower than assumed in the model close to the surface (Strobel and Wolven, 2001). For a smaller scale height the spots move closer to the surface and this would imply a steeper slope of the spot ratio as a function of the longitude.

On the downstream side the relation is inverse for $80^\circ < \varphi < 100^\circ$, because the elongated spots are slightly tilted inwards. Therefore the line-of-sight integration for $\varphi \approx 85^\circ$ (resp. $\varphi \approx 95^\circ$) is approximately parallel to the anti-jovian (sub-jovian) spot and the spot appears brighter. The ratios of the wake observations by Retherford (2002) can not confirm this, but generally reveal a brighter sub-jovian spot for $\varphi < 90^\circ$ in agreement with the model results.

An absolute value for the altitude of the equatorial spots midpoints is not readily determined for the observations. First, the spot emission is elongated

over several degrees of longitude or several 100 km and thus it does not have one common altitude. Moreover, due to the two dimensional geometry and the viewing angle of the observations an altitude can not be derived without knowing the exact longitudinal extent or position. Choosing a hypothetical viewing geometry, where the observer is located right above the North pole, we can identify the height of the maximum emission on the flanks above the limb for our simulation results. Depending on the emitting species and the density of the upstream electrons, the local emission peaks at altitudes between ~ 150 km and ~ 300 km with our best-fit atmosphere. For a higher atmospheric density, the spots would appear at higher altitudes, as mentioned above, but again, the model probably overestimates the scale height and thus the spot altitude for that case.

The observed overall aurora brightness is similar between Ieclipse04 ($I_{tot} = 62.6$ kR) and Ieclipse03 ($I_{tot} = 40.9$ kR). This agrees with the simulation, in which the total intensity varies with upstream electron density. In both cases the modeled intensity differs from the observed one by a factor of ~ 2 . This discrepancy ($I_{obs}/I_{sim} \approx 2$) can be traced back either to plasma conditions or atmospheric properties. The principal factor that controls the total intensity is the electron energy, which is deposited in the atmosphere. The influence of atmospheric density on the total intensity is less significant because of the opposed effects of an increase of neutral gas abundance: More neutral gas can be excited and radiate, but stronger deflection and divergence of the plasma flow (Section 1 and Saur and Strobel (2004)). Assuming the equatorial gas ring to extend to higher latitudes (up to 45° instead of 35° , see equation 4.1)

the total intensity would be increased by $\sim 15\%$ in our simulation. However, the electron energy deposited in the atmosphere is close to maximum for our derived atmosphere. So, the difference between observation and simulation is probably caused by differences in the plasma environment, of which no simultaneous measurements are available.

For instance, during the Ieclipse04 event Io passed through the torus center and thus possibly through very high plasma density regions. An even higher upstream plasma density than we assumed, would provide more energy and generate accordingly more radiation than calculated with our input parameters. Additionally, magnetic field aligned electron fluxes of higher energies ($\sim \text{keV}$) can possibly excite auroral emission. Michael and Bhardwaj (2000) estimated a possible contribution to the aurora generation from energetic field-aligned electrons observed by the Galileo plasma instrument (Frank and Paterson, 1999). They found that the modeled emission is on the same order of brightness as the HST observations by Roesler et al. (1999), but do not consider the limited spatial distribution of the observed field aligned beams. Oliverson et al. (2001) measured short term ($\lesssim 10$ min) intensity variations in the OI 630.0 nm line and ascribed this to a time-variable energy flux of field-aligned non-thermal electrons. Another possible reason for the discrepancy of simulation and observation are additional emitters in the wavelength range of the LORRI pass-band, which we do not account for in the simulation. Emission from S_2 is expected at near-UV and blue visible wavelengths (Geissler et al., 2004), but not simulated due to the lack of exact electron impact cross sections. Since the morphology and position of the spots are to major parts determined by the absolute column density and not

by the composition of the atmosphere, the abundance of S_2 can be neglected here. In the Ieclipse03 LORRI image thermal emission around volcanic hot spots clearly contributes to the observed radiation, but is not simulated. Besides from these physical reasons, systematic uncertainties in the LORRI images might be large due to difficulties when compiling several exposures. Systematic errors such as copious instrument scattered light possibly leads to an under- or overestimation of the brightness.

4.2. *Minor components*

Since the LORRI observations do not spectrally resolve the emission, we compare our simulation results with a range of spectral observations of various emission lines resulting from oxygen, sulfur, sodium and potassium to constrain the abundance of these species. We are aware of the limitations of using observations made by different telescopes or cameras at different times, as for example long term variations and differing signal-to-noise ratios. But still, the measured intensities of the various observations differ not more than by a factor of 5 and are thus still a good way to derive abundances and check the derived atmospheric distribution with a wider range of observations. Some emission lines were frequently observed in the past, such as the oxygen multiplets OI 135.6 nm and OI 630.0 nm. In that case, we preferably used the more recent observations. For both lines, modeled intensity and morphology are in good agreement with the observations (Retherford et al., 2007; Retherford, 2002; Bouchez et al., 2000; Geissler et al., 1999; Moore et al., 2010) using a mixing ratio of 12% of atomic oxygen to sulfur dioxide at the surface and a surface scale height of 150 km. A comparison of the intensities for the analyzed emissions lines can be found in Table 3.

782 To derive a sulfur mixing ratio we used the most recent observation of the
 783 SI 147.9 multiplet taken by the NH Alice spectrograph in 2007 as benchmark.
 784 The measurements revealed a varying total intensity during three eclipse
 785 events of $I = 0.4 - 1.2$ kR. Assuming a sulfur mixing ratio of 1.5% at the
 786 surface the model calculates a total intensity $I = 0.6 - 1.1$ kR. The intensity
 787 range in the model results is due to varying upstream electron density and
 788 variation of the viewing longitude. Comparisons with previous observations
 789 (Wolven et al., 2001; Feaga et al., 2002; Ballester et al., 1987) also yield a good
 790 agreement with the derived mixing ratio. Wolven et al. (2001) determined
 791 radial profiles for the OI 135.6 and SI 147.9 emission along the spatial axis
 792 of the aperture for HST/STIS observations. For the brightness close to I_0
 793 averaged over the aperture width they find values between ~ 0.1 and ~ 0.8
 794 kR, where the emission from sulfur is found to be slightly higher than from
 795 oxygen. Depending on the sub-observer longitude, we calculate brightnesses
 796 averaged tangential to limb of 0.1–0.5 kR for OI 135.6 and 0.3–0.9 kR for
 797 SI 147.9 close to I_0 assuming the derived mixing ratios of oxygen and sulfur.
 798 The values match the observed brightness range of Wolven et al. (2001).

799 In case of the sodium emission we use the observed intensities in eclipse
 800 by Retherford (2002) and Bouchez et al. (2000) as reference. Compared to
 801 optically thin intensities the abundance of both Na and K need to be cor-
 802 rected by a factor of 3.5 and 3, respectively (Geissler et al., 2004). To achieve
 803 emission comparable to the observed intensities in the simulated sodium au-
 804 rora an abundance of sodium bearing species of 0.12 % (at the surface) has
 805 to be implemented, assuming a surface scale height of 150 km. For potas-
 806 sium there are no direct measurements of K assigned emission lines in the

close vicinity ($d < 1 R_{io}$) of Io. Using the Na/K ratio of 3.3 derived from infrared filter observations at 670–850 nm by Geissler et al. (2004) (0.04% K abundance) the total intensity of the simulated KI 767.0 nm emission line is $I = 2.2 - 4.3$ kR. This is clearly lower than the total intensity of the infrared filter observations (Geissler et al., 2004), which provide an upper limit of 10.7 kR.

All derived atmosphere parameters and the resulting equatorial column densities are summarized in Table 4. The higher scale height of atomic species is in agreement with atmospheric models (e.g. Summers and Strobel, 1996). The higher background (polar) abundance for O, S, Na and K (n_{bg}/n_{eq}) is derived from the observed limb glow and polar emission for the respective emitter. The compared total intensities for all four analyzed components are listed in Table 3.

4.3. *Tvashtar Plume*

To constrain the plume density we analyze both the morphology and the intensity of the radiation inside the framed Tvashtar regions in Figure 3. During Ieclipse04 the whole plume of Tvashtar was in the field of view of LORRI, see Figure 3a. The plume is radiating rather uniformly with a maximum intensity in the plume center ~ 220 km above the limb. The simulated absolute flux inside the framed volcano region, $I_{Tvashtar}$, is half of the observed flux from this region. Since for Ieclipse04 the simulated total intensity I_{tot} also differs approximately by factor 2 from the observation, we compared the ratio $I_{Tvashtar}/I_{tot}$ with our simulation results. For Ieclipse04 the averaged emission intensity inside the Tvashtar region is almost equal to the disk-averaged total intensity, i.e. the ratio is $I_{Tvashtar}/I_{tot} = 1.0$. Assuming a gas density

832 of $n_{V,0} = 1.7 \times 10^8 \text{ cm}^{-3}$ in the main plume or an average column density of
 833 $N_V = 5 \times 10^{15} \text{ cm}^{-2}$ over the plume region the modeled plume aurora (Figure
 834 3d) is in good agreement with the measurements ($I_{Tvashtar}/I_{tot} = 0.9$), see
 835 also Table 2. A small Alfvén wing is generated above Tvashtar due to the
 836 increased neutral and thus plasma density in the region, see Figure 6c. Yet,
 837 with the derived plume density, the electrons are still able to excite emission
 838 all over the plume (Figure 3d), since the plasma flow is diverted only mod-
 839 erately. The intensity maximum is found to be in the radial center at a limb
 840 distance of $\sim 180 \text{ km}$ and plume emission appears to be rather uniformly. In
 841 the LORRI Ieclipse03 observation the upper edge of the glowing plume is
 842 visible just above the limb. The emission in the observation and simulated
 843 aurora match very well regarding morphology and relative strength.

844 For a lower plume density of $N_V < 3 \times 10^{15} \text{ cm}^{-2}$ the total modeled
 845 emission is too weak. For both LORRI observations the intensity ratio of
 846 plume and disk-averaged emission would be lower than the observed ratio by
 847 a least a factor of 2. In the case of a higher plume density, $N_V > 7 \times 10^{15} \text{ cm}^{-2}$
 848 the current system in and above the plume (see Figure 6c) becomes stronger.
 849 This leads to a stronger diversion of the electron flow around the plume
 850 region. The electrons deposit energy mostly on the upstream side of such a
 851 dense plume and the simulated aurora appears non-uniformly with a clear
 852 maximum at the upstream edge. On the other hand, the plume averaged
 853 emission intensity hardly increases with increasing plume density. The energy
 854 deposited all over the plume is almost at maximum for our best-fit plume
 855 density $N_V = 5 \times 10^{15} \text{ cm}^{-2}$, so that a denser plume does not induce a
 856 brighter auroral emission.

Although the viewing angle of the HST/ACS Ieclipse03 image (Figure 3c) allows the whole plume to be visible by HST, only a very weak UV emission enhancement of $I_{tvashtar}/I_{poles} \approx 1.2$ was observed in the expected region. With the derived column density of $N_V = 5 \times 10^{15} \text{ cm}^{-2}$, our simulation predicts a ~ 3 times higher emission inside the Tvashtar region compared to the polar area around it. The pass-band of the HST observation includes mostly oxygen and sulfur emission lines. The lack of an emission enhancement around Tvashtar indicate a low abundance of O and S in the plume. With 3 times lower abundances of atomic sulfur ($\sim 0.5\%$) and oxygen ($\sim 4\%$) within the plume the $I_{tvashtar}/I_{poles}$ ratio matches approximately the HST/ACS observation.

As Tvashtar is considered a Pele-type plume, a high S_2 abundance is expected. Observations of the Tvashtar plume in scattered light and absorption by Jessup and Spencer (2008) revealed a spectral behavior that is consistent with previous observations of the Pele plume. For Pele various S_2/SO_2 ratios between $\sim 1\%$ and 30% have been obtained by Jessup et al. (2007). If dissociation of SO_2 and S_2 happens on time scales much larger than the average time of flight for the plume gas, the abundance of atomic sulfur and atomic oxygen in the plume will be lower than in an equatorial atmosphere, which is not solely of volcanic origin. Moses et al. (2002) show that the lifetimes for photolysis of SO_2 or S_2 producing S and O is in the range of hours, while ballistic flight times are on the order of 10 minutes. Using a thermodynamic model Fegley and Zolotov (2000) infer mixing ratios of atomic sulfur and oxygen to SO_2 and S_2 of the order of 10^{-2} and below. This implies a low abundance of S and O and would thus explain the low intensity in the UV

range. On the other hand, with a low abundance of O, which is also an essential emitter at visible wavelengths, the simulated plume intensity for the LORRI images would be reduced. Thus, the abundance of other species emitting in the LORRI wavelength must be higher to sustain the plume radiation. Another explanation for the non-visibility of the Tvashtar plume is the low resolution of the HST image, which possibly masks the local volcanic emission enhancement. A high resolution image in this UV range would offer a possibility to determine the exact abundance of atomic oxygen and sulfur in the plume.

The local density for a cut through the best-fit plume for the LORRI observations and the vertical column density as a function of the distance to the plume center (d in (8)) are shown in Figure 5. Applying the model of a volcanic system developed by Kieffer (1982) and later adapted by Strobel and Wolven (2001) to the derived average column density of $N_V = 5 \times 10^{15} \text{ cm}^{-2}$ and the implemented plume size, Tvashtar appears to be a high temperature, low pressure volcano. The large size of the plume indicates relatively high particles velocities and thus a high temperature at the volcanic crater, which results from a high temperature ($T_0 \approx 800 \text{ K}$) reservoir in the model system of Kieffer (1982). Our derived average column density is low compared to common plume models (e.g. Zhang et al., 2003) and would imply a low mean plume pressure of 0.1 nbar and correspond to the low pressure case (small plume size) discussed in Strobel and Wolven (2001).

Measurements and model calculations (e.g. by McGrath et al., 2000; Zhang et al., 2003) generally yield comparatively denser plumes for volcanoes close to the equator. Thus, our derived plume density and pressure appears

to be relatively low. However, our method is hardly able to determine the plume density close to the surface accurately as the observed and simulated radiation is mostly emitted at a distance > 50 km from the ground. A higher localized density at very low altitudes might also not influence the plasma interaction strongly. If the density below 50 km increased exponentially over the whole plume area with an atmospheric scale height on the order of tens of km (as inferred by Strobel et al. (1994) for the equatorial atmosphere near the surface), the plume column density would go up by at least a factor of 2. For instance, a very low sticking coefficient for molecules that contact the surface could lead to a density increase at low altitudes.

The total content of the equatorial atmosphere, which we derived in section 4.1, equals ~ 5 – 10 times the total content of a volcanic plume with the plume density derived for Tvashtar. In other words, if we ruled out sublimation as possible atmospheric source, about 5–10 active volcanoes of the size of Tvashtar would be necessary to sustain the eclipse atmosphere. Pele-type plumes, such as Tvashtar, are the largest observed plumes and not more than 16 plumes have been observed so far (Geissler and Goldstein, 2007). Therefore it is unlikely that an atmosphere of the derived column density can be sustained solely by direct volcanic outgassing in eclipse, but there has to be an essential amount of gas species that do not condense during eclipse. The atmospheric density in sunlight is approximately ten times higher (Lelouch et al., 2007) than our derived equatorial atmosphere, so there would have to be ~ 50 – 100 Tvashtar-sized active volcanoes to create such a dense atmosphere without sublimation. Although already more than 150 active regions have been discovered, above most regions no plumes have been ob-

served (Lopes et al., 2004). Hence, direct volcanic outgassing is possibly able to sustain at least parts of the eclipse atmosphere, but cannot be considered as an essential source for the sunlit atmosphere.

5. Summary and conclusions

We modeled the auroral emission from the atmosphere while Io was in eclipse of Jupiter and investigated the effect of various atmospheric distributions on the intensity and morphology of the aurora. Our model results imply an atmospheric column density of $N_{eclipse} = (1-4) \times 10^{15} \text{ cm}^{-2}$, i.e. when Io moves into Jupiter's shadow, its atmospheric density decreases down to $\sim 10\%$ of the sunlit atmosphere assuming $N_{sun} = (1-5) \times 10^{16} \text{ cm}^{-2}$ (Lellouch et al., 2007; Saur et al., 2000; Saur et al., 2002). Despite the density decrease, the atmosphere probably still covers most of the surface around the equator up to $\sim 35^\circ$ of latitude at these densities. With a smooth atmospheric ring and a low density (2% of equatorial density) at high latitudes we were able to reproduce the main features of the auroral emission, observed in various HST and spacecraft observations. Independently from the exact location of active volcanoes, sub-jovian and anti-jovian bright spots arise due to the diverted plasma flow. The various observed spot morphologies primarily result from the respective viewing geometry of the observer and the resulting visibility of the aurora features, such as the spots.

Compared to detailed atmospheric models (e.g. Walker et al., 2010), the inferred neutral distribution appears to be simplified. However, since the advantage of our model is the inclusion of the influence of a chosen atmospheric distribution on the plasma interaction and thus on the aurora generation, we

956 did not consider the small scale variations of the equatorial atmosphere for
957 numerical reasons but focused on the global structure.

958 Strong emission appearing in the downstream region of Io, which often
959 is denoted as “wake emission”, can be explained for major parts by flank
960 emission extending far downstream. The derived variation of the brightness
961 ratio between anti-jovian and sub-jovian spots in eclipse coincides with the
962 observed variation in sunlight qualitatively. This implies that the aurora
963 morphology is controlled by the plasma interaction also in eclipse and not
964 only by the exact distribution of the atmospheric gas and/or the locations of
965 volcanoes.

966 Analyzing various monochromatic observations we derived mixing ratios
967 for minor components in the atmosphere: 12% atomic oxygen, 1.5% atomic
968 sulfur and 0.12% sodium. The derived mixing ratios and the resulting column
969 densities are in general agreement with previous observations and model
970 results (Geissler et al., 2004; Wong and Smyth, 2000; Summers and Strobel,
971 1996), although our values are comparatively high.

972 For the Tvashtar plume we find a column density of $N_V = 5 \times 10^{15} \text{ cm}^{-2}$.
973 This relatively low plume content supports the idea of an atmosphere that
974 is sustained almost solely by sublimation. If a larger number of plumes of
975 the Tvashtar size were active in the last decades, they likely would have
976 been observed already. So far, 16 mostly smaller plumes have been observed
977 (Geissler and Goldstein, 2007) and thus volcanic outgassing probably is not
978 able to sustain a dense atmosphere as measured in sunlight.

979 Spectrally resolved observations with a resolution comparable to the
980 LORRI images would provide a possibility to determine absolute values of

the abundance of the various emitters. As the electron energy is distributed very inhomogeneously around Io, local inhomogeneities in the neutral gas abundance possibly influence the total intensity strongly. Spectral observations with a high spatial resolution like the LORRI images would allow a detailed analysis of various regions, taking into account the local electron parameters as well as the local neutral gas density and composition. For example, spectrally resolved observations of the Tvashtar plume emission could offer detailed information about the abundant species in this huge and outstanding plume.

Acknowledgments

L.R. and J.S. acknowledge support from BMWi and DLR, grant FKZ 50 OR 0805. K.D.R. appreciates support from NASA grant NNX09AU28G. D.F.S. acknowledges support from the New Horizons Mission through SWRI Contract No. 277043Q and NASA grant NNX10AB846.

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Event	Date	Moment	UTC (hh:mm)	Sys-III long. ^a	Centrif. lat. ^a	Background field ^b (nT)		
						B_x	B_y	B_z
Ieclipse03	2/27/2007	Ingress	14:21	154°	4.3°	308	-580	-2145
		Mid-eclipse	15:24	183°	6.0°	79	-788	-2141
		Egress	16:27	212°	6.3°	-142	-715	-2108
Ieclipse04	3/1/2007	Ingress	08:50	255°	3.9°	-300	-321	-2022
		Mid-eclipse	09:53	284°	0.9°	-308	-11	-1966
		Egress	10:56	313°	-2.3°	-261	264	-1938

Table 1: Eclipse occasions during LORRI and HST/ACS observations. ^a System-III longitude and the angle between Io's position and the centrifugal torus center. ^b calculated with the magnetic field model of Connerney et al. (1998) for the internal field of Jupiter.

Observation	Bandpass (nm)	Resol. (km/pixel)	Viewing geometry ^a	Figure
LORRI Ieclipse04	350-850	~56	3° S 240° W	3a
LORRI Ieclipse03	350-850	~15	7° S 310° W	3b
HST/ACS Ieclipse03	125-190	~135	~0° 344° W	3c

Table 2: Summary of the 2007 NH/LORRI and HST/ACS eclipse observations. The observations are displayed in Figures 3a-c. All observations are combinations of several exposures during one eclipse event, as reported in Spencer et al. (2007b). ^a The viewing geometry of the respective observer (New Horizons, HST) refers to the mid-eclipse time and can thus differ for single exposures.

Emitter	Sunlight observations I_{sun} (kR)	Eclipse observations I_{ecl} (kR)	Eclipse simulation $I_{ecl,sim}$ (kR)
OI 135.6	$0.48 - 0.88^a$	$0.38 - 0.80^a$	$0.4 - 0.7$
OI 630.0/636.4	–	$6.8^f / 14.2^b / 23.7^c$	$6.2 - 11.8$
SI 147.9	$0.60 - 1.17^a$	$0.43 - 1.20^a$	$0.6 - 1.1$
SI 190.0	$1.68 - 2.33^d$	–	$0.5 - 0.9$
NaI 588.9/589.6	–	$4^b / 6.9^c$	$3.4 - 6.6$
KI 767.0	–	$< 10.35^e$	$2.2 - 4.3$

Table 3: Observed and simulated total intensities of the investigated emission lines. The variation of the simulated intensity results from changes in the viewing geometry as well as various upstream electron densities ($n_e = 1900 - 3600 \text{ cm}^{-3}$). ^aRetherford et al. (2007), ^bRetherford (2002), ^cBouchez et al. (2000), ^dBallester et al. (1987), ^eGeissler et al. (2004), ^fGeissler et al. (1999)

Atmosp.	Equ. surface mixing ratio	<i>Geissler et al., 2004</i>	Sc.height H (km)	backgr. n_{bg}/n_{eq}	Long.asym. γ , β	Vert. col. dens. upstream	downstr.	Tang.col.dens. (cm^{-2})
All	100 %	-	-	-	90° , $1/3$	1.0×10^{15}	4.0×10^{15}	1.8×10^{16}
SO ₂	80 %	82 %	100	0.02	90° , $1/3$	8.0×10^{14}	3.3×10^{15}	1.5×10^{16}
O	12 %	5 %	150	0.15	90° , $1/3$	1.8×10^{14}	7.3×10^{14}	3.2×10^{15}
S	1.5 %	2.5 %	150	0.15	90° , $1/3$	1.9×10^{13}	7.8×10^{13}	4.1×10^{14}
Na	0.12 %	0.12 %	150	0.5	90° , $1/3$	1.8×10^{12}	7.3×10^{12}	6.5×10^{13}
K	0.04 %	0.036 %	150	0.5	90° , $1/3$	6.0×10^{11}	2.4×10^{12}	1.1×10^{13}
S ₂ ,O ₂ ,SO	~5%	~10%						
<hr/>								
	Height h (km)	Width σ (km)	Mean col. dens. (cm^{-2})		Max. col. dens. (cm^{-2})			
Tvashtar	360	1100	3×10^{15}		7×10^{15}			

Table 4: Best-fit parameters of model atmosphere.

Fig.	Total	Equ. spots		Ratio	North			South			Tvashtar		
		I_{tot}	$I_{sub}(1)$		$I_{anti}(2)$	I_{anti}/I_{sub}	$I_N(3)$	I_N/I_{tot}	$I_S(4)$	I_S/I_{tot}	$I_t(5)$	I_T/I_{tot}	
Observation													
LORRI Iec104	3a	62.6	53.6	97.3	1.8		21.3	0.3		20.5	0.3	62.6	1.0
LORRI Iec103	3b	40.9	70.6	22.2	0.3		6.5	0.2		11.2	0.3	12.1	0.3
HST/ACS Iec103	3c	1.8	0.7 ^a	1.6 ^a	0.4 ^a		0.58	0.4		0.7	0.4	0.7	0.4
Simulation													
LORRI Iec104	3d	29.0	29.3	60.4	2.1		9.1	0.3		5.2	0.2	23.6	0.9
LORRI Iec103	3e	17.1	38.2	13.7	0.4		4.0	0.2		3.1	0.1	8.9	0.5
HST/ACS Iec103	3f	1.4	0.9 ^a	1.9 ^a	0.5 ^a		0.4	0.3		0.3	0.2	1.1	0.8

Table 5: Comparison between the emission intensities of the observations (Figures 3a-c) and the simulation results (Figures 3d-f). Absolute intensities are in kRayleigh. The measured and calculated intensities of the framed regions in Figures 3a-f are all averaged to the box areas, the total intensity is averaged to the area of Io's disk. The numbers in parentheses (1-5) refer to the green box numbers. ^a Direct comparison of absolute intensities of the HST/ACS observations with the simulation is limited, since the simulation does not take into account all emission lines in the observed wavelength range. SI 147.9 nm emission is assumed to be optically thin.

Figure 1: Observations of Io's auroral emission. (a) Multispectral eclipse image of Io taken by Cassini on January 1, 2001 (Geissler et al., 2004). Near-UV emission displayed in blue is attributed to SO₂, emission in the red visible wavelength range (red in image) is primary from atomic oxygen (OI630.0 nm). (b) HST/STIS OI 135.6 nm image of October 14, 1997 (Roesler et al., 1999).

Figure 2: Geometry of the Ieclipse03 and Ieclipse04 observations in the planetocentric coordinate system of Io. The black arrows point out the viewing longitude of the New Horizons probe and the Hubble Space Telescope, respectively. The jovian background field vectors in the xy (orbital) plane at mid-eclipse are shown with purple arrows. Up to the right a B field vector of 500 nT is displayed for comparison ($B_z \approx 2000$ nT). In our simulation model the undisturbed plasma flow is directed in positive x direction and the background field is assumed constant on negative z direction (perpendicular to the displayed xy plane). The expected key features of the aurora (very simplified) are marked with the red areas.

Figure 3: Aurora observations and the respective simulated emission patterns. Properties of the New Horizons LORRI images (a and b) (Spencer et al., 2007a,b) and the HST/ACS image (c) (Retherford et al., 2007) are listed in Table 2. The displayed LORRI images have undergone some corrections to remove blemish emission. The corresponding simulated emission morphologies (d, e and f) are displayed with the respective viewing geometry. The color scale and the contour lines differ by a factor of 2 between observation and simulation for the LORRI images. Contours are 150 kR (a), 100 kR (b), 75 kR (d) and 50 kR (e). No contours in (c) and (f). Note that the HST/ACS observation (c) covers wavelengths from 125 – 190 nm, whereas the corresponding simulation (f) takes into account only emission from OI 135.6 nm and SI 147.9 nm. The orbital trailing (270°) longitudes are indicated with dashed meridians, the sub-jovian (0°) and anti-jovian (180°) meridians are displayed in plain bold. The emission within the green framed areas is investigated quantitatively. The total and relative intensities of the boxes are shown in Table 5.

Figure 4: Evolution of the electron temperature in the simulation model. The flux-tubes parallel to $\vec{B}_0$ ($\parallel z$) convect through the Io's interaction region given by the 2D plasma flow profile in the xy plane. The heat conduction perpendicular to $\vec{B}$ is neglected, the heat conduction parallel to $\vec{B}$ is infinite outside the atmosphere. Inside the atmosphere the plasma cools down to T_{in} and the heat flow from outside (T_{out}) is limited due to a finite heat conductivity along $\vec{B}$. The heat flow from the outer tube part to the inner is parametrized and depends on the temperature gradient between the two parts ($\sim (T_{out} - T_{in})/R_{typ}$).

Figure 5: The density distribution model for the Tvashtar plume displayed in a vertical plane through the center. The local neutral gas density is color-coded, the dashed line represents the derived vertical column density above the plume. The plume density is given by equation 8 and consists of three components: (1) the main plume, (2) a low density ring as found by ballistic models and (3) a high density region above the crater. The shape replicates roughly the modeled Pele-like plume by Zhang et al. (2003) with a peak column density over the plume center and a shallow decrease of the column density from ~ 200 to ~ 500 km distance from the center. The plume-averaged column density is determined within a distance to the plume center of < 700 km (dashed-dotted vertical lines).

Figure 6: Results of the plasma interaction simulation: (a) Electron temperature within the atmosphere in xy plane. Isolines for 1 to 5 eV. (b) Electron density in units of torus density (1900 cm^{-3}) in the equatorial plane at $z = 0$. (c) Isolines of electric currents in 10^{-7} A m^{-2} in the northern Alfvén wing.

Figure 7: Solid line: Modeled brightness ratio of the anti-jovian/sub-jovian spots. The brightnesses refer to the average emission intensity within a $\sim 800 \times 1100 \text{ km}^2$ box located at the equatorial limb (see Figures 3). A spot brightness ratio derived from UV observations in sunlight by Retherford (2002) is shown (+) for observing longitudes between 59° and 74° and between 241° and 297° . Dotted: The longitudinal variation of the total intensity of OI 135.6 emission normalized to its minimum.

Figure 2

Figure 3

Figure 4

Figure 5

Figure 6

Figure 7