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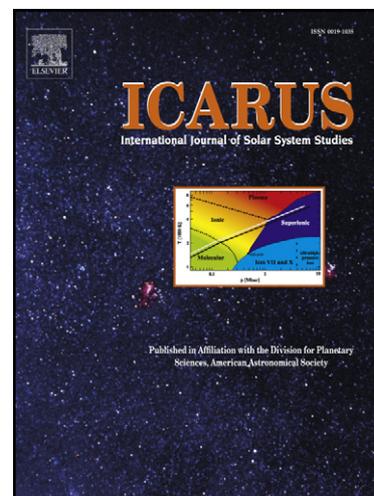
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Simulation of Io's Auroral Emission: Constraints on the Atmosphere in Eclipse

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Abstract

We study the morphology of Io's aurora by comparing simulation results of a three-dimensional (3D) two-fluid plasma model to observations by the high-resolution Long-Range Reconnaissance Imager (LORRI) on-board the New Horizons spacecraft and by the Hubble Space Telescope Advanced Camera for Surveys (HST/ACS). In 2007, Io's auroral emission in eclipse has been observed simultaneously by LORRI and ACS and the observations revealed detailed features of the aurora, such as a huge glowing plume at the Tvashtar paterae close to the North pole. The auroral radiation is generated in Io's atmosphere by collisions between impinging magnetospheric electrons and various neutral gas components. We calculate the interaction of the magnetospheric plasma with Io's atmosphere-ionosphere and simulate the auroral emission. Our aurora model takes into account not only the direct influ-

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ence of the atmospheric distribution on the morphology and intensity of the emission, but also the indirect influence of the atmosphere on the plasma environment and thus on the exciting electrons. We find that the observed morphology in eclipse can be explained by a smooth (non-patchy) equatorial atmosphere with a vertical column density that corresponds to $\sim 10\%$ of the column density of the sunlit atmosphere. The atmosphere is asymmetric with two times higher density and extension on the downstream hemisphere. The auroral emission from the Tvashtar volcano enables us to constrain the plume gas content for the first time. According to our model, the observed intensity of the Tvashtar plume implies a mean column density of $\sim 5 \times 10^{15} \text{ cm}^{-2}$ for the plume region.

Key words:

Io, Jupiter, satellites, Atmospheres, structure

1. Introduction

Io probably possesses the densest and most species-rich atmosphere of the four Galilean satellites of Jupiter. Besides the main constituent sulfur dioxide various minor species, such as S, S₂, O, SO, Na, K and Cl, have already been observed in the vicinity of Io. Two possible sources for the atmosphere were discussed since its discovery. On the one hand volcanic venting as a direct source can create a neutral gas cloud around the satellite. On the other hand, sublimation of SO₂ frost from the surface by sunlight is considered a possible driver of the atmosphere (Lellouch et al., 2007). SO₂ frost abundance is found to be correlated with active volcanic regions (Douté et al., 2001). Thus, the surface distribution of the atmosphere is

12 related to the distribution of the volcanic regions on global scales, no matter
13 if sublimation or direct outgassing is the main source. Prior observations of
14 SO_2 suggested a volcanically driven atmosphere (e.g., Lellouch et al., 1992).
15 Recent observations (Jessup et al., 2004; Moullet et al., 2010) and model
16 results (Saur and Strobel, 2004) indicate that the sublimation driven part
17 clearly dominates the direct outgassing.

18 The frequency of both the active volcanic regions and the paterae de-
19 creases with increasing latitude (Lopes-Gautier et al., 1999; Radebaugh et al.,
20 2001) and observations of Lyman- α absorption in the atmosphere indicate
21 that the SO_2 gas is concentrated likewise at lower latitudes (Strobel and Wol-
22 ven, 2001). Feaga et al. (2009) derived a map of SO_2 column density, which
23 shows a relatively sharp density decrease at approximately $30\text{--}45^\circ$ north and
24 south and a maximum column density on the anti-jovian hemisphere, which
25 was first noted by Jessup et al. (2004) and Spencer et al. (2005) and later
26 confirmed by Moullet et al. (2010). Walker et al. (2010) investigated effects of
27 plasma heating as well as surface frost, molecular residence time and surface
28 temperature distribution in a sophisticated SO_2 gas dynamics simulation,
29 which includes sublimation and direct outgassing. The simulation results are
30 compared to several observations with a backwards radiative transfer model
31 in a companion paper (Gratiy et al., 2010) and basically confirm the previ-
32 ously derived atmospheric distributions assuming certain surface conditions.
33 Considering results from modeling, millimeter, infrared and ultraviolet ob-
34 servations Lellouch et al. (2007) concluded that Io's atmosphere has a mean
35 vertical column density of $\sim (1\text{--}5) \times 10^{16} \text{ cm}^{-2}$, covering 50–70% of Io's
36 day-side hemisphere.

37 The evolution of the auroral emission, while Io passes through Jupiter's
38 shadow, possibly provides very instructive information about the nature of
39 the atmosphere. In eclipse the surface temperature drops and the subli-
40 mation of SO_2 strongly decreases. Besides, no incident sunlight is reflected
41 and thus solely the electron excited emission from the atmosphere is observ-
42 able. Analyzing the aurora offers a possibility to determine the composition
43 and distribution of Io's diverse atmospheric gas environment (Geissler et al.,
44 2004).

45 Aurora (or airglow) is commonly defined as radiation caused by charged
46 particles, which excite molecules or atoms in an atmosphere (Chamberlain
47 and Hunten, 1987). In the case of Io, thermal electrons of the jovian magne-
48 tosphere rotate with Jupiter's magnetic field and thus constantly flow past
49 the slowly orbiting satellite and excite the atmospheric gas. The emission
50 covers a broad wavelength range including ultraviolet, visible and infrared
51 wavelengths and can be attributed to the various abundant species in Io's at-
52 mosphere. Previous observations and analysis of Io's aurora offered insights
53 into the satellite's atmosphere (e.g. Clarke et al., 1994; Roesler et al., 1999) .

54 Since the arrival of the Galileo probe at the jovian system, Io's auroral
55 emission has been observed many times by ground-based and space tele-
56 scopes (e.g., Roesler et al., 1999) as well as by on-board cameras of the
57 Galileo (e.g., Geissler et al., 1999, 2001) and Cassini (e.g., Geissler et al.,
58 2004) spacecrafts. Two of these observations are shown in Figure 1. Almost
59 all observations are dominated by a key feature: bright spots close to the
60 sub- and anti-jovian limb at low latitudes. These equatorial spots move up
61 and down with the rocking background magnetic field of Jupiter (Retherford

62 et al., 2000). Geissler et al. (2001) also observed enhanced visible aurora
63 close to volcanic plumes. Numerical simulations of the aurora by Saur et al.
64 (2000) revealed that due to a diverted plasma flow the energy of the electrons
65 is preferentially deposited on the flanks of Io and thus more energy reaches
66 dense atmospheric layers around the sub- and anti-jovian points on the sur-
67 face (see also aurora features in Figure 2). Therefore the bright equatorial
68 spots are generated there. A tilt of the magnetic field leads to a tilted dis-
69 tribution of the magnetospheric electrons and the equatorial spots are thus
70 correlated with the background field. Furthermore, auroral emissions appear
71 to be brighter on the hemisphere facing the plasma torus centrifugal equator
72 than on the other hemisphere (Retherford et al., 2003). Moore et al. (2010)
73 found the same north/south dependency for a wake emission features simu-
74 lating OI 630.0 nm emission observed by the HST Wide Field and Planetary
75 Camera 2.

76 When Io enters Jupiter's shadow, two opposing effects control the evolu-
77 tion of this auroral radiation (Saur and Strobel, 2004). The atmospheric gas
78 partly freezes out and thus less neutral gas can be excited by the ambient
79 plasma. On the other hand, a decrease of atmospheric density after eclipse
80 ingress leads to a decreasing interaction strength. The deflection and cooling
81 of the plasma flow is lower and the streamlines of the electrons are less diver-
82 gent. Saur and Strobel (2004) calculated the response of Io's electrodynamic
83 interaction and radiation to a temporal change in the atmosphere. Depend-
84 ing on the total atmospheric decay, they find three qualitatively different
85 scenarios with two of them including a transient post-eclipse brightening.
86 Generally, the total emission intensity in eclipse was found to be lower than

87 out of eclipse . Retherford (2002) investigated the variation of key features of
88 the eclipse aurora using HST Space Telescope Imaging Spectrograph (STIS)
89 observations and inferred a reduction factor of ~ 1.5 to 2 for the emission close
90 to Io during eclipse compared to the sunlit atmosphere. Geissler et al. (2004)
91 in turn investigated the auroral emission during eclipse at different wave-
92 lengths with Cassini filter observations. Comparing the observed brightness
93 of an equatorial spot with modeled intensities they derived mixing ratios for
94 various gases, which contribute to the equatorial emission, such as O, S, Na
95 and K. Limb glows, which are attributed to minor components, indicate that
96 O, Na and K are abundant all over the surface (Retherford, 2002; Geissler
97 et al., 2004). Based on these eclipse observations Geissler et al. (2004) finally
98 conclude that Io's atmosphere must be at least partly sustained directly by
99 volcanism. Recent Monte Carlo simulations of the response to eclipse of
100 the neutral gas by Moore et al. (2009) indicate a strong dependence on the
101 abundance of non-condensable species during atmospheric collapse. In case
102 of a mole-fraction of 0.35 of non-condensable gas a near-surface diffusion layer
103 forms and the SO₂ density decreases only slowly and does not drop below
104 0.18 of the initial column density during an eclipse event.

105 During the flyby of the New Horizons probe in 2007, the auroral emission
106 was observed by the on-board camera LORRI (Spencer et al., 2007b) and
107 by the Hubble Space Telescope Advanced Camera for Surveys (HST/ACS)
108 (Retherford et al., 2007) (Figure 3a-c). The LORRI camera provides im-
109 ages with high resolution comparable to Galileo SSI Io eclipse observations
110 and reveals new details, such as the glowing plume of the Tvashtar volcano
111 close to the North pole. This allows a detailed investigation of the aurora

112 morphology. With the simultaneous HST/ACS observation we are able to
113 compare aurora morphologies from two different viewing geometries.

114 In our work we evaluate these observations regarding the morphology
115 and the total intensity. Therefore we use a numerical model to calculate
116 the interaction of the upstream plasma with Io's atmosphere-ionosphere and
117 the auroral emission from the atmosphere. This method ensures that we
118 take into account both effects, which influence the auroral response to an at-
119 mospheric collapse during an eclipse event, as described above in this section.
120 Investigating various atmospheric distributions, we derive one atmospheric
121 distribution in eclipse, which is able to explain the observed intensity and
122 morphology of the two LORRI observations and the HST/ACS image at the
123 same time. Furthermore we infer the gas content of the Tvashtar plume. The
124 constraints on the equatorial atmosphere and on the plume provide an esti-
125 mation of the atmospheric collapse and a contribution of volcanic outgassing
126 to the atmosphere.

127 **2. Observations**

128 In February and March 2007 the New Horizons probe passed the jovian
129 System on its way to Pluto. The closest approach to Io occurred on 28 Feb-
130 ruary at a distance of 2.24×10^6 km. Io was observed several times while
131 passing through Jupiter's shadow (Retherford et al., 2007). During two of
132 the eclipse passages LORRI, a narrow-angle high resolution camera on-board
133 the spacecraft, took spectacular images of the auroral emission at visible and
134 infrared wavelengths (Spencer et al., 2007a,b). According to the notation of
135 Retherford et al. (2007) these two eclipse occasions will be denoted Ieclipse03

136 and Ieclipse04. During the Ieclipse03 event simultaneous spatially resolved
137 observations of Io's far-ultraviolet aurora have been achieved by the Ad-
138 vanced Camera for Surveys Solar Blind Channel (ACS/SBC) on the Hubble
139 Space Telescope. Characteristics of the eclipse occasions and the LORRI and
140 HST/ACS observations are given in Tables 1 and 2. The geometry of the
141 observations is depicted in Figure 2.

142 The NH/LORRI observations during eclipses Ieclipse04 and Ieclipse03
143 are shown in Figures 3a and 3b (in the order that they are discussed in the
144 paper). The spatial resolution of LORRI is $4.96 \mu\text{rad}$ per pixel (Cheng et al.,
145 2008). For Ieclipse03 this corresponds to a pixel size of $\sim 15 \text{ km} \times 15 \text{ km}$
146 (~ 200 pixels per Io diameter). During Ieclipse04 LORRI was used in the
147 4×4 binning mode, which reduces the spatial resolution in both directions
148 by 4. The pixel size in that case is $\sim 56 \text{ km} \times 56 \text{ km}$, which is equivalent
149 to ~ 50 pixels per Io diameter. The pass-band ranges from about 350 nm to
150 850 nm. The measured radiation originates from various emitters. In the blue
151 range SO_2 and S_2 emission bands are the main contributors. Excited atomic
152 oxygen, atomic sodium and potassium are the major emitters in the green,
153 red and infra-red range (Geissler et al., 2004). Atomic sulfur contributes only
154 a minor fraction.

155 The conversion from detector counts to emission brightness depends on
156 the wavelength dependent sensitivity of the respective instrument. Since the
157 contributing emissions in the wavelength ranges of LORRI and ACS are not
158 clearly defined, we used Pivot wavelengths of 607.6 nm (LORRI) and 143.7
159 nm (ACS) to convert to Rayleighs (brightness). To validate this method we
160 used exemplary modeled emission spectra including various emitters (SO_2 ,

161 O, S, Na and K) and calculated the brightness of the LORRI images using
162 the sensitivity curve of Cheng et al. (2008). With this method the intensity
163 differs less than 20% from the intensity derived with the Pivot wavelength.

164 We now describe and assign the essential features in the observation im-
165 ages. During the Ieclipse03 observation (Figure 3b) the sub-spacecraft coor-
166 dinates are $\sim 310^\circ\text{W}$ (west longitude, 0° towards Jupiter) and $\sim 7^\circ\text{S}$ (South-
167 ern latitude, $\vartheta = 0^\circ$ in the orbital plane) in the Io-centric coordinate sys-
168 tem. Hence, the sub-jovian upstream quadrant is visible during the observa-
169 tion. The brightest small spots on the disk appear to be thermal emission.
170 The measured intensity corresponds to a black body temperature of approx-
171 imately 1200 K (Spencer et al., 2007b). Fainter small spots are low-altitude
172 gas emission. On the upstream side (right-hand limb) enhanced radiation
173 is visible only within a relatively small region, which has been attributed to
174 a volcanic hot spot located east of the Girru paterae and therefore named
175 “East Girru“ (Spencer et al., 2007b). Considering the geometry of the back-
176 ground magnetic field during Ieclipse03, which we discuss in detail below,
177 the observed emission maximum around East Girru could also result from
178 the magnetic field tilt and represent a shifted anti-jovian equatorial spot.
179 The main part of anti-jovian spot probably is hidden behind Io. However,
180 the East Girru emission appears to consist of one low and one high altitude
181 emission part. Assuming that the plasma parameters do not vary discon-
182 tinuously, the puzzling two part emission could originate from two different
183 atmospheric species, as for example emission from SO_2 close to the surface
184 and from atomic oxygen after dissociation of SO_2 at higher altitudes. Since
185 a spectrally resolved observation is not available, this feature cannot be an-

186 analyzed for different emission lines. Around the downstream (left-hand) disk
187 edge there is a diffuse emission region of approximately $500 \times 1,000 \text{ km}^2$.
188 This region is the sub-jovian equatorial spot identified in previous observa-
189 tions (e.g., Roesler et al., 1999). Apart from the equatorial band there is small
190 enhanced emission just above the limb close to the North pole. The bright
191 area can undoubtedly be assigned to the huge Tvashtar plume, which was
192 first seen by Cassini in December 2000 (Porco et al., 2003). Unlike most of
193 Io's volcanoes, Tvashtar is located at a high latitude near Io's North pole. As
194 the global atmosphere is relatively thin at higher latitudes, Tvashtar creates
195 a locally enhanced neutral atmosphere, which can be investigated separately
196 from the equatorial atmosphere.

197 In the Ieclipse04 image (Figure 3a) the Tvashtar plume is clearly evident
198 above the limb. The bright area is similar to the plume size derived from
199 sunlit observations with a height of about 350 km and full width of 1100
200 km (Spencer et al., 2007b). The geometry of this image (approx. 240°
201 W 3° S) allows observation of both equatorial spots. While the sub-jovian
202 spot (left-hand) is restricted to a small area, the anti-jovian spot extends
203 further on the disk. The extent to which a few dozen lower intensity features
204 located at known volcanic vent locations may include atmospheric emissions
205 combined with the thermal emissions is yet to be determined. For Ieclipse04
206 the deviation from the upstream view at 270° W, where both spots should be
207 visible and similar in brightness, is $\sim 30^\circ$. Note that in the case of Ieclipse03
208 the deviation from the upstream view is only $\sim 10^\circ$ more, but there is no
209 clear anti-jovian spot observable (see description above).

210 The average of four consecutive exposures of Io in eclipse by HST/ACS

211 (Retherford et al., 2007) is displayed in Figure 3c. This image includes emis-
212 sion in the 125 nm to 190 nm band-pass, which originates mostly from excited
213 atomic sulfur and oxygen. O and S are expected to resemble SO_2 in the re-
214 gion of interest for most of the analyzed emissions, although larger differences
215 might occur at higher altitudes. The viewing angle from HST is $\sim 344^\circ$ W
216 and $\sim 0^\circ$ N/S. Thus the complete sub-jovian side is visible, the plasma flow
217 is directed from right (upstream) to left (downstream). The morphology is
218 dominated by three features. The brightest area on the disk corresponds to
219 the sub-jovian equatorial spot. The next brightest emission to the left of Io
220 can either be attributed to the far end of the equatorial spots or originate
221 directly in the wake of Io. And third, similar to the LORRI observation, the
222 emission on the upstream (right-hand) side is considerably enhanced only
223 above the East Girru region. Again, this enhancement can not definitely
224 be assigned to a volcanic region, but might also originate from a smooth,
225 continuous atmosphere and an inhomogeneous electron environment. As al-
226 ready mentioned, the NH Alice spectrograph additionally measured the total
227 intensity of the two prominent oxygen multiplets at 130.4 nm and 135.6 nm
228 and the sulfur multiplet at 147.9 nm during four eclipse events between 25
229 February and 3 March 2007 (see Table 3 and Retherford et al. (2007)).

230

231 Due to the tilt between Jupiter's dipole moment and Jupiter's rotation axis,
232 the background field and the surrounding plasma density change while Io is
233 moving up and down in the plasma torus during a synodic rotation period
234 of Jupiter ($T_{syn} \approx 12.95$ h) with respect to Io. It takes about two hours for
235 Io to pass through Jupiter's shadow, which is approximately one sixth of the

236 synodic period of the varying background field. All displayed observation
 237 images (Figures 3a-c) are combinations of several coaligned exposures during
 238 one eclipse event (averaging regions not contaminated with instrument scat-
 239 tered light), and thus include changes in the plasma environment between
 240 the exposures. Although we can not identify one exact observation geometry,
 241 we now roughly describe the plasma conditions during the LORRI and HST
 242 observations. We use an Io-centered coordinate system, where z is Jupiter
 243 (Io) North, and x is along the orbital direction of Io, i.e. approximately
 244 along the plasma flow (but rocks due to the tilt of the plasma torus). y com-
 245 pletes the system pointing roughly (Io's orbit is slightly eccentric) towards
 246 Jupiter (Figure 2). The three components of \vec{B} calculated with the model of
 247 Connerney et al. (1998) for Jupiter's internal field are listed in Table 1.

248 During Ieclipse03 Io is above the plasma torus and reaches the maximum
 249 distance to the torus center shortly before egress. The y component of the
 250 magnetic field vector is relatively large. The component in the direction of
 251 the orbital movement (B_x) in turn is low. The angle between the magnetic
 252 field and the polar axis (z) varies between $\sim 16^\circ$ and $\sim 19^\circ$. During Ieclipse04
 253 Io passes through the torus center from North to South. Accordingly, the y
 254 component of \vec{B} is lower and the x component somewhat larger. The tilt of
 255 \vec{B} to the polar axis is around 10° . Furthermore, the ambient plasma den-
 256 sity presumably is higher during Ieclipse04 in the torus center than during
 257 Ieclipse03, when Io is far away from the center (Bagenal, 1994). The field
 258 vectors in the xy plane during mid-eclipse are shown in Figure 2. The di-
 259 rection of the magnetic field for the observing geometry at mid-eclipse (Io in
 260 line with the center of the Sun and the center of Jupiter) is displayed in the

261 lower right corner in Figures 3a-c.

262 In all observations five areas are highlighted with green frames, which we
263 will separately analyze in this paper. First, two areas of $\sim 9.2 \times 10^5 \text{ km}^2$
264 around the equatorial spots, then a $\sim 6.4 \times 10^5 \text{ km}^2$ sized region centered at
265 the calculated position of the Tvashtar plume. The large boxes ($\sim 3.8 \times 10^6$
266 km^2) on the upper and lower edge of Io cover the polar areas, where the at-
267 mosphere is expected to be less dense and thus less emission is expected. We
268 integrate the total emission within the boxes and normalize it to the covered
269 area. This method allows us to investigate different features of the auroral
270 morphology quantitatively.

271

272 For an appropriate theoretical description of the formation of Io's aurora,
273 we first need to calculate the interaction of the plasma particles with Io's
274 environment. As explained in section 1, the density and distribution of the
275 atmospheric gas influence this interaction, i.e. the neutral gas controls the
276 flow pattern and the temperature and density profiles of the electrons. Since
277 the electrons generate the auroral emission, electron temperature and den-
278 sity are in addition to the atmospheric density the essential parameters for
279 calculating the aurora. On the other hand, the distribution and local density
280 of emitting gas particles is directly reflected by the intensity and morphology
281 of the aurora. To calculate both effects self-consistently, we use a plasma
282 model developed by Saur et al. (1999).

283 3. Model to interpret the observations

284 3.1. Plasma interaction model

285 The model, which we use, was developed to simulate the plasma inter-
286 action of the satellites Europa (Saur et al., 1998) and Io (Saur et al., 1999).
287 It has undergone several improvements subsequently. The simulation results
288 provide explanations for several observed features of Io's plasma environment
289 such as magnetic field signatures in the wake and the rotated Alfvén wing
290 system. Furthermore, Saur et al. (2000) were able to explain the formation
291 of the bright equatorial spots of Io's aurora. For the full set of equations and
292 the numerical algorithms we refer the reader to Saur et al. (1999, 2002). In
293 the following section we explain the basics of the model and the treatment
294 of the two plasma parameters, which are essential for aurora simulation: the
295 density and temperature of the electrons.

296
297 The simulation is developed in the E, j approach of magnetohydrodynamics.
298 The magnetic field is assumed to be a constant, homogeneous background
299 field at all times. The electric conductivity parallel to the magnetic field is
300 assumed to be infinite, so the parallel electric field vanishes. The validity of
301 these assumptions are assessed in Neubauer (1998); Saur et al. (1999, 2002).
302 With this model and its assumptions, Saur et al. (1999) were able to describe
303 various aspects of the plasma interaction. In eclipse the local plasma inter-
304 action and thus the magnetic field perturbations are weaker than in the case
305 of a sunlit atmosphere, i.e. the assumption of a homogeneous field is even
306 more justified for our purposes. Galileo observations in sunlight revealed
307 a magnetic field perturbation of more than $\frac{\Delta B}{B} = 0.3$ (e.g. Kivelson et al.,

1996), which were reproduced by our model (Saur et al., 2002). For the interaction in eclipse the model results indicate that the perturbation of the magnetic field is lower by a factor of 2 ($\frac{\Delta B}{B} \approx 0.15$ in eclipse). This justifies our assumption of a homogenous background magnetic field since the magnetic field environment is strongly dominated by the background field, while other plasma parameters such as velocity and electric field vary significantly due to the interaction with the atmosphere.

In the undisturbed upstream plasma the homogeneous electric field is given simply by $\vec{E}_0 = -\vec{v}_0 \times \vec{B}_0$. Inside Io's ionosphere a current system arises from the collisions between the plasma and neutral gas particles, which modify the electric field. The modified \vec{E} field is calculated by a differential equation for the 2D electric potential in the plane perpendicular to \vec{B}_0 first derived by Wolf-Gladrow et al. (1987). The electric potential and thus the electric field around Io are calculated from the ionospheric Hall and Pedersen conductances as well as the Alfvén conductance. The undisturbed plasma velocity \vec{v}_0 is assumed to be perpendicular to \vec{B}_0 . Neglecting inertia and pressure, the electron velocity $\vec{v}_e(x, y)$ in the plane perpendicular to \vec{B}_0 can be derived directly from the electric field. To simulate the auroral emission we need to calculate the properties of the thermal electron population in the vicinity of Io, which excite the aurora. Moving with the electron flow, the evolution of the density as well as the temperature of the electrons can be calculated as described below.

The coordinate system of the simulation corresponds to the system that we defined in section 2. The magnetic field however is constant and always anti-parallel to the z axis and the plasma flow v_0 is parallel to x . Thus, we

take into account neither changes of the inflow direction nor the rocking of the magnetic field due to the tilt of 9.6° of the jovian dipole field axis with respect to the jovian rotation axis. The geometry is discussed in detail in section 3.4.

The evolution of the electron density n_e is described by the following continuity equation:

$$\frac{d}{dt}n_e = f_{ion} n_e n_{SO_2} + k_{hee} n_{SO_2} - \alpha n_e^2 \quad (1)$$

The first term on the right hand side describes (single) ionization by electron impact on neutral gas n_{SO_2} due to thermal electrons (Sittler and Strobel, 1987) with the temperature dependent collisional ionization rate f_{ion} .

With the second term we account for ionization by kappa-distributed energetic bidirectional electrons (k_{hee}) observed in Io's vicinity (e.g., Williams et al., 1999). The electron energies in the beams span the range from keV to hundreds of keV (Williams et al., 1996; Frank and Paterson, 2002). According to the simulated electron beam morphology by Jacobsen et al. (2010) and the observations during various Galileo flybys (Williams and Thorne, 2003; Frank and Paterson, 2002) we assume a spatial distribution of the high energetic electrons given by

$$f_{hee} = f_{hee,0} [\tanh(3x + 1) + 1] \exp(-10y^6) \cdot c(x, y) \quad (2)$$

with

$$\begin{aligned} c &= 1 & \text{for } x^2 + y^2 > 1 R_{Io} \\ c &= 0.5 & \text{for } x^2 + y^2 \leq 1 R_{Io}. \end{aligned}$$

330 Thus, the beams are assumed to be relatively narrow (width in y direction:
331 $\sim 1.4 R_{Io}$) but to extend far into the wake. However, outside the atmosphere

332 (e.g. in the far wake), where the neutral density vanishes, no ionization can
 333 take place anyway. Using the geometrical factor $c(x, y)$ we take into account
 334 that directly above (below) Io the electron beam from the South (North) is
 335 shielded by the satellite. For the energy flux we use the derived value by
 336 Saur et al. (2002). Due to the absence of observations and models of elec-
 337 tron beams when Io is eclipse, we note that expression (2) is derived from
 338 observations and models when Io is in sun light.

339 The third term on the RHS of equation (1) describes the loss due to recom-
 340 bination with a rate α . For the adopted parameter values we refer the reader
 341 to Saur et al. (1999, 2002).

342 As the electron velocity in the model is simply given by the $\vec{E} \times \vec{B}$ -drift,
 343 there is no plasma movement along \vec{B} included. But, in fact plasma transport
 344 along \vec{B} does occur, particularly in the wake of Io. The only electron source
 345 in the wake is the newly ionized population through the electron beams. For
 346 a low density atmosphere as in eclipse the ionization by electron beams is
 347 mainly confined to the region close to the equator, where the atmosphere is
 348 densest. Altogether, the wake region will be rather void, if movement along
 349 \vec{B} is neglected. However, parallel movement due to the pressure gradient in
 350 the wake of Io possibly fills the relatively void regions in the downstream area
 351 and might thus enhance the aurora in the downstream region. To account
 352 for the parallel movement, we modified the model assuming that the plasma
 353 particles move along \vec{B} with thermal velocity v_{th} .

The total time derivative on the left hand side of (1) can be written as the
 partial time derivative and the convective term $\vec{v}_e \cdot \nabla n_e$. Separating the flow in
 the xy plane from the movement along z (i.e. $\vec{v}_e \cdot \nabla n_e = \vec{v}_\perp \cdot \nabla_\perp n_e + \vec{v}_\parallel \cdot \nabla_\parallel n_e$)

we rewrite the continuity equation as

$$\vec{v}_e(x, y) \cdot \nabla_{\perp} n_e = f_{ion} n_e n_{SO_2} + k_{hee} n_{SO_2} - \alpha n_e^2 - v_{th} \nabla_{\parallel} n_e, \quad (3)$$

354 \vec{v}_{\perp} is the electron flow $\vec{v}_e(x, y)$ and the parallel flow \vec{v}_{\parallel} along z is approximated
 355 by the thermal velocity v_{th} . Thus, following the 2D electron flow $\vec{v}_e(x, y)$ a
 356 parallel flow as loss or production term depending on the density gradient
 357 along z is calculated. Numerically, we consider the net flow from or to both
 358 the grid cell above and below the current position. The flow direction is
 359 determined by the sign of the respective electron density gradient between the
 360 cells. This description is somewhat similar to a diffusive process. The thermal
 361 parallel velocity v_{th} is assumed to be the ion sound speed, as the inertia of the
 362 ions mainly determines the movement of the plasma. The ion sound speed
 363 is approximately half of the undisturbed relative flow velocity of the plasma
 364 (Kivelson et al., 2004), so the flow fills the wake at an angle of approximately
 365 $\arcsin\left(\frac{v_{th}}{v_e(x, y)}\right) \geq 30^{\circ}$. The propagation along \vec{B}_0 enables the plasma to fill
 366 the wake of Io, which is important to explain the observed auroral emission
 367 in the downstream region. Expanding into low density regions such as Io's
 368 wake the plasma can be accelerated to velocities higher than the thermal
 369 velocity (Samir et al., 1983). By the assumption of a parallel movement with
 370 v_{th} , the propagation speed into the wake might thus be underestimated.

The second important plasma quantity for aurora simulation is the thermal energy or temperature of the electrons, T_e . The temperature evolution is given by

$$\begin{aligned} \frac{3}{2} k_B n_e \frac{dT_e}{dt} = & -(\epsilon_{ion} f_{ion} + \epsilon_{dis} f_{dis} + \epsilon_{rot} f_{rot} + \epsilon_{vib} f_{vib}) \cdot n_e n_{SO_2} \\ & - \frac{3}{2} k_B T_e (f_{ion} n_e n_{SO_2} + k_{hee} n_{SO_2}) - \nabla Q_{flux}. \end{aligned} \quad (4)$$

371 We account for cooling by inelastic collisions between the magnetospheric
 372 electrons and the atmosphere, including ionization, dissociation, and rota-
 373 tional and vibrational excitation of neutral SO_2 . Each process κ is described
 374 by the rate f_κ and the energy quantum ϵ_κ . The newly added electrons re-
 375 sulting from the impact ionization processes are assumed to be cold. The
 376 adjustment to magnetospheric bulk temperature of these cold electrons leads
 377 to a decrease in temperature described by the next term on the right hand
 378 side. k_B is the Boltzmann constant.

The heat flux Q_{flux} is parametrized. We take advantage of the anisotropy
 of the thermal conductivity and assume no heat flow perpendicular to \vec{B} due
 to the strong background magnetic field of ~ 2000 nT (Banks and Kockarts,
 1973). Parallel to \vec{B} the heat conduction is extremely high. We consider
 the electrons in a flux tube along \vec{B} outside the atmosphere to adjust to one
 common temperature T_{out} instantaneously. Deep inside the atmosphere the
 parallel thermal conductivity is lower, where the mixing ratio of neutral gas
 to plasma increases. When the flux tube passes through the atmosphere,
 we divide it in three parts, two outer parts (T_{out}) and an inner part (T_{in}).
 The inner part uniformly cools down due to the various collision processes
 described in equation 4. The heat flow from the hotter outer part to the
 inner part Q_{flux} is parametrized as sketched in Figure 4. Depending on
 the temperature dependent heat conductivity $\kappa(T_{in})$ and the temperature
 difference between the inner and outer parts a typical heat flow is calculated:

$$\langle Q_{flux} \rangle = \langle \kappa(T_{in}) \rangle \frac{T_{out} - T_{in}}{R_{typ}}. \quad (5)$$

379 The average heat transport is controlled by a typical distance R_{typ} , which is
 380 set to the scale height of the atmosphere (100 km). The electron heat conduc-

381 tivity κ in a plasma with a fraction of neutral gas as function of the ambient
 382 temperature is given by equation (22.116) of Banks and Kockarts (1973) and
 383 depends on the momentum transfer cross section for elastic collisions of elec-
 384 trons with SO_2 gas and the mixing ratio n_{SO_2}/n_e . The temperature in the
 385 outer part T_{out} decreases due to the heat flow to the cooler inner part, i.e.
 386 the outer parts serve as finite heat reservoirs and the energy of the entire flux
 387 tube is depleted. The energy capacity of these reservoirs corresponds to the
 388 electron content in the region along the magnetic field line above and below
 389 Io. For further details see Saur et al. (2002), Appendix A.

390 3.2. Atmosphere model

Based on several observations (see section 1) of Io's atmospheric distribu-
 tion, we assume a dense atmospheric ring ranging from the equator to approx-
 imately 35° North and South. At higher latitude a low density background
 n_{bg} is assumed, with a ratio of $n_{bg}/n_{eq} = 0.02$ for SO_2 . This distribution was
 calculated by Strobel and Wolven (2001) based on Lyman- α reflection ob-
 servations. Here we investigate longitudinal asymmetries, such as differences
 between sub- and anti-jovian hemispheres reported by Spencer et al. (2005),
 Feaga et al. (2009) and Moullet et al. (2010) as well as a denser downstream
 atmosphere as inferred by Saur et al. (2002). The surface density n_s is thus
 modeled by

$$n_s(\vartheta, \varphi) = n_{bg} + (n_{eq} - n_{bg}) (1 + \beta \cos(\varphi - \gamma)) e^{-\left(\frac{\vartheta}{35^\circ}\right)^6} \quad (6)$$

with the latitude ϑ and Jupiter oriented longitude φ with $\varphi = 0$ for direction
 to Jupiter and $\varphi = 90^\circ$ in Io's orbital direction. n_{eq} is the surface density at
 the equator, β and γ specify the strength and orientation of the longitudinal

inhomogeneity and are fitted to match the observations. The number density declines exponentially with increasing altitude. The vertical structure is determined by the scale height H_s . We assume the scale height to vary with φ according to the longitudinal variation of the surface density:

$$H_s(\varphi) = H_{s,0} (1 + \beta \cos(\varphi - \gamma)). \quad (7)$$

391 The simultaneous variation of scale height and surface temperature allows an
 392 implementation of a larger difference in column density between two hemi-
 393 spheres with smaller local gradients than choosing a strongly varying surface
 394 density at a constant scale height. When the local density gradients are too
 395 large, the simulation becomes unstable.

396 For modeling the plasma interaction, we assume the atmosphere to con-
 397 sist solely of SO_2 , since it supposedly is the by far main constituent (Lellouch
 398 et al., 2007). The surface scale height $H_{s,0}$ in (7) for SO_2 is set to 100 km
 399 (Saur et al., 1999). Near the surface the actual scale height is presumably
 400 lower in the range of tens of km. But, for the low density eclipse atmosphere
 401 the plasma likely penetrates all atmospheric layers and the model accounts
 402 for integrated conductivities. At high altitudes a large scale height is ex-
 403 pected. Thus, a scale height of 100 km represents an average scale height
 404 for Io's atmosphere and the interaction strength still constraints the global
 405 column density, since the crucial parameter for the interaction is the total
 406 atmospheric gas content.

407 The longitudinal distribution of the equatorial atmosphere in eclipse,
 408 which is modeled here, does not necessarily correspond to the longitudi-
 409 nal distribution at daytime seen around a full Io orbit. The fractional at-
 410 mospheric collapse might vary with longitude due to different atmosphere

411 and surface conditions or the position of the observer and the sun. For
 412 instance, the atmosphere may respond differently on the sub-Jupiter hemi-
 413 sphere, which is sunlit before eclipse ingress, than on the anti-jovian hemi-
 414 sphere, which is at night for several hours already before eclipse.

As the LORRI observations revealed a bright volcanic plume close to the North pole, we additionally implement a plume shaped density enhancement located at the Tvashtar paterae at 62°N and 122°W (Spencer et al., 2007b). The distribution within the plume is modeled by

$$\begin{aligned}
 n_V(h, d) = n_{V,0} & \left(\exp \left[- \left(\left(\frac{h}{H_V} \right)^2 + \left(\frac{d}{\sigma_V} \right)^2 \right)^3 \right] \right. \\
 & \left. - \exp \left[- \left(\left(\frac{h}{0.4H_V} \right)^2 + \left(\frac{d-d_0}{0.4\sigma_V} \right)^{\frac{3}{2}} \right)^3 \right] \right) \\
 & + 30 n_{V,0} \exp \left[- \left(\left(\frac{h}{0.1H_V} \right)^2 + \left(\frac{d}{0.05\sigma_V} \right)^2 \right)^3 \right]
 \end{aligned} \tag{8}$$

415 where $n_{V,0}$ is the density in the center of Tvashtar, h the vertical distance from
 416 surface and d the horizontal distance on the surface to the plume center. The
 417 height H_V and width σ_V are given by the observed plume extent in sunlight
 418 and are set to $H_V = 360$ km and $\sigma_V = 550$ km. The subtracted exponential
 419 function (with height $0.4 H_V$ and width $0.4 \sigma_V$) in the second line roughly
 420 describes the expected low density region within the plume after Zhang et al.
 421 (2003). We also include a high density region above the vent by adding the
 422 third line in the equation. The plume model of Zhang et al. (2003) also
 423 indicates the formation of a canopy shock, which we do not account for in
 424 our simulation. $n_{V,0}$ is the only free parameter of the plume. The three plume
 425 regions are marked in Figure 5, where a cross section through the volcano

426 density and the column density above the plume are shown.

427 To model the radiated emission with the simulated electron densities and
 428 temperatures all the essential atmospheric constituents are considered in-
 429 dividually. We will derive a mixing ratio n_{comp}/n_{total} for the atmospheric
 430 components O, S, Na and K. For the simulation of the LORRI images the
 431 mixing ratio of the species is assumed to be identical in the equatorial at-
 432 mosphere and in the plume. For atomic sulfur and oxygen the radial decrease
 433 is smaller than for SO₂ (Wolven et al., 2001; Summers and Strobel, 1996).
 434 The Cassini observations of Geissler et al. (2004) revealed O assigned emis-
 435 sion at higher surface distances than the emission assigned to SO₂, see Figure
 436 1a. This indicates a shallower decrease of O compared to SO₂. For atomic
 437 sodium, HST observations of the NaI 589 nm line revealed a very shallow
 438 drop-off with increasing height (Retherford, 2002). Therefore, we assume a
 439 slightly larger scale height for these constituents of $H_{s,0}(X) > H_{s,0}(\text{SO}_2)$
 440 (see table 4). Moreover, we assume a higher n_{bg}/n_{eq} ratio for atomic species,
 441 as the Cassini observations (Geissler et al., 2004) as well as HST observa-
 442 tions (Retherford et al., 2000, 2003) show a clear limb glow all around Io for
 443 emission from atomic species.

444 3.3. Emission simulation

Assuming equilibrium, the local intensity of the stimulated emission of
 an atmospheric gas is calculated by

$$i_{\lambda}(\vec{x}) = f_{\lambda}(T_e(\vec{x})) n_{gas}(\vec{x}) n_e(\vec{x}), \quad (9)$$

where f_{λ} is the gas specific rate, n_{gas} the gas density and n_e the density of
 the exciting electrons. Assuming spontaneous emission the emission rate cor-

responds to the collisional excitation rate, which is given as an integral over the Maxwell-Boltzmann distribution, the electron velocity and the energy-dependent cross section for the collision of the exciting electrons with the neutral species. For optically thin emission lines a 2D emission pattern is given by the line-of-sight integral over the local intensities. Not all emissions can be considered optically thin. Therefore we estimate re-absorption for Na and K. All other emission lines are assumed to be optically thin in our analysis. Due to the long lifetime or small radiative decay rate of $k_d = 0.00681 \text{ s}^{-1}$ of the $O(^1D)$ state, we include collisional quenching at low altitudes for the OI 630.0/636.4 nm emission. The lowered intensity I is calculated with the Stern-Vollmer relationship

$$I = I_0 \frac{k_d}{k_d + k_q} \quad (10)$$

445 where k_q is the quenching rate and I_0 the intensity without quenching. For
 446 k_q we use the estimated value of $1 \times 10^{-10} \text{ cm}^3\text{s}^{-1}$ from Geissler et al. (2004).
 447 Quenching and radiative decay are equally probable when the SO_2 number
 448 density is $6.8 \times 10^7 \text{ cm}^{-3}$.

449 The NH/LORRI observations cover wavelengths from 350 to 850 nm.
 450 Within this range emission from SO_2 bands and S_2 bands as well as from
 451 atomic oxygen (557.7 nm, 630.0 nm, 636.4 nm, 777.4 nm and 844.6 nm),
 452 atomic sulfur (772.5 nm), atomic sodium (588.9/589.6 nm) and potassium
 453 (767.0 nm) are expected. The OI 777.4 nm and OI 844.6 nm lines contribute
 454 less than 1% to the total emission and thus can be neglected (Geissler et al.,
 455 2004). The electron impact excitation cross sections that we use for atomic
 456 oxygen emission are based on the laboratory measurements of Doering and
 457 Gulcicek (1989a,b) and Doering (1992). The cross sections for the sodium

458 and potassium D lines were adopted from theoretical calculations by Kim
459 (2001), which the author compares to experimental data from Enemark and
460 Gallagher (1972). For further details on the cross sections and the associ-
461 ated rates see Geissler et al. (2004) and references therein. Cross sections
462 for electron excitation of S_2 are not measured, so we can only estimate the
463 contribution of S_2 emission roughly.

464 A large number of emission lines can be found in the pass-band of the
465 HST/ACS (125–190 nm) observation, such as emission from atomic oxygen
466 and sulfur as well as from sulfur ions (Ballester et al., 1987; Roesler et al.,
467 1999) and also chlorine emission (Feaga et al., 2004). Since cross sections are
468 not available for all of the lines, we solely analyze the oxygen multiplet at
469 135.6 nm and the sulfur multiplet at 147.9 nm, which presumably contribute
470 most. Neglecting all the lower emissions in the pass-band, we can not com-
471 pare the emission quantitatively with the ACS observation. A quantitative
472 comparison also would necessitate an accurate analysis of optical depth of
473 various lines. However, by simulating the OI 135.6 nm and SI 147.9 emission
474 we are able to compare the model results to the ACS observation regarding
475 the morphology and the relative intensities of the individually analyzed ar-
476 eas, since oxygen and sulfur generally are by far the main contributors in the
477 observed wavelength range. The contribution of the OI 130.4 nm to the UV
478 aurora is not clear yet and is therefore not considered here. For the sulfur
479 emission at 147.9 nm we account for both the forbidden and allowed lines
480 (Feaga et al., 2002). For both the SI147.9 nm and SI190.0 nm emission the
481 adopted cross sections are based on the calculated collisions strengths from
482 Zatsarinny and Tayal (2002).

483 *3.4. Simulation setup*

484 A constant magnetic field and a 2D plasma flow perpendicular to B_0 are
485 the basic assumptions of the theoretical approach. So it is not possible to
486 fully consider the varying plasma conditions around Io in the simulation. In
487 the model the constant background field is parallel to the z axis and the
488 upstream flow is in positive x direction. Moreover, the model is symmetric
489 with respect to the xy plane given by $z = 0$. For the implementation of
490 the volcanic plume on the North pole, we run simulations with and without
491 plume and combine the results. The influence of the plume on the aurora in
492 the equator region appears to be negligible ($< 1\%$). Note, the asymmetry of
493 Io's atmosphere due to Tvashtar near Io's North pole generates also a small
494 asymmetry in Io's plasma interaction. However, due to the small spatial
495 extend of the plume compared to Io's diameter and the relative amount of
496 the plume gas compared to the total gas of Io's atmosphere, the asymmetry
497 in the plasma interaction is negligible in contrast to Enceladus' interaction,
498 where the asymmetry plays an important role (Saur et al., 2007). We try
499 to explain various observations with one atmospheric distribution, which is
500 assumed to be symmetric around Io's equator due to the mentioned basic
501 symmetry of the model around the equator. Hence, the orientation of the
502 magnetic field and the upstream plasma flow are fixed in the model.

503 The actual plasma environment varies during a synodic rotation of Jupiter.
504 The angle between the undisturbed plasma and the plane perpendicular to
505 \vec{B}_0 varies only within -3° and 3° , which is negligible. The variation of the
506 direction of the magnetic field with respect to the polar axis of Io is larger (up
507 to 20°). As the simulation code is symmetric with respect to the equatorial

508 plane, which is always assumed to be perpendicular to the background field,
 509 we are not able to investigate a tilt between the atmospheric ring and the
 510 orbital plane. Since the displayed observations are combinations of several
 511 exposures, the geometry is also varying for the single exposures. To mini-
 512 mize the effect of the simplified model geometry, we only analyze larger areas
 513 around the equatorial spots and Tvashtar quantitatively. Retherford et al.
 514 (2000) showed that the inclination of the spots is somewhat lower than the
 515 tilt of the magnetic field.

516 We account for the variation of the electron density in the upstream
 517 plasma. For Ieclipse03 simulation, where Io is far from the torus center
 518 we assume a lower density of $n_e = 1900 \text{ cm}^{-3}$, while during Ieclipse04 (Io
 519 crosses the torus center) the electron density is presumably higher and we
 520 use $n_e = 3600 \text{ cm}^{-3}$ (Gurnett et al., 2001b,a; Bagenal, 1994). The electron
 521 temperature of the upstream plasma is $T_{e,0} = 5 \text{ eV}$ (Bagenal, 1994). For
 522 the initial velocity we assume the relative azimuthal velocity of the rotating
 523 plasma ($\vec{v}_0 = 57 \text{ km s}^{-1}$), the background field is set to $B_0 = 2050 \text{ nT}$.

524 The viewing geometry of the displayed 2D emission patterns is given by
 525 longitude φ and latitude ϑ , where $\vartheta = 0^\circ$, when the observer is in the orbital
 526 plane, and $\vartheta = 90^\circ$, when viewed from above the North pole.

527 As discussed previously in section 2, we compare the simulated images with
 528 the observations using the green frame regions in Figure 3.

529 4. Results

530 First, we derive the total gas content and its distribution, which are de-
 531 termined mainly by SO_2 abundance. Thereafter we briefly describe our

532 derived mixing ratios and distributions for atomic oxygen and sulfur as well
 533 as the trace elements sodium and potassium. The emission from the atomic
 534 species also contributes to the intensity and morphology of the LORRI sim-
 535 ulations, so the derivation of the abundances for SO₂ and those of the minor
 536 components are correlated to each other. Therefore, we obtain the best-fit
 537 atmosphere results from an iterative variation of both the total gas content
 538 and the mixing ratios of the minor species. Finally, we constrain the gas
 539 content of the Tvashtar plume.

540 4.1. Equatorial Atmosphere

541 The main benchmark for constraining the density of the atmospheric ring
 542 was the morphology and location of the equatorial spots. We assume the
 543 same longitudinal variations for SO₂ and the minor components.

544 First, we analyze the brightness of the equatorial spots within the green
 545 boxes of the LORRI observations. In the LORRI Ieclipse04 observation
 546 (Figure 3a) the ratio of the anti-jovian to the sub-jovian spot brightness
 547 is $I_{anti}/I_{sub} \approx 1.8$. For LORRI Ieclipse03 (Figure 3b) the anti-jovian spot
 548 can not be identified clearly. The spot might be displaced toward the East
 549 Girru feature. The ratio for the equator centered boxes in Figure 3b is
 550 $I_{anti}/I_{sub} \approx 0.2$. We get the best morphology agreement with both observa-
 551 tions for a model atmosphere with an average equatorial column density of
 552 $N_{eq,av} = 2.3 \times 10^{15} \text{ cm}^{-2}$. We find a longitudinal asymmetry in favor of the
 553 downstream side with the asymmetry parameters $\beta = \frac{1}{3}$ and $\gamma = 90^\circ$. The
 554 column density on the upstream side is thus $N_{eq,up} = 1.0 \times 10^{15} \text{ cm}^{-2}$, the
 555 downstream column density is $N_{eq,down} = 4.0 \times 10^{15} \text{ cm}^{-2}$. The simulated
 556 aurora morphologies corresponding to the LORRI images are displayed in

557 Figures 3d and 3e, where the brightness ratios are $I_{anti}/I_{sub} \approx 2.1$ (Ieclipse04)
 558 and $I_{anti}/I_{sub} \approx 0.4$ (Ieclipse03). The derived atmospheric distribution also
 559 yields a morphology that is in good agreement with the HST Ieclipse03 ob-
 560 servation on the sub-jovian side. Furthermore, the key features of previous
 561 eclipse observations by Retherford (2002) are reproduced by the simulation
 562 considering the respective observation geometries (not shown here). The de-
 563 rived eclipse column density corresponds to $\sim 10\%$ of the column density
 564 for a sunlit atmosphere summarized in Lellouch et al. (2007) ($N_{sun} = (1-$
 565 $5) \times 10^{16} \text{ cm}^{-2}$). This day-side column density also coincides with the col-
 566 umn densities derived by Saur et al. (2000) and Saur et al. (2002), where the
 567 same simulation model for a sunlit atmosphere interaction is applied. The
 568 parameters of the best-fit atmosphere are listed in Table 4. The resulting
 569 absolute simulated and measured intensities in Rayleighs (R) of the marked
 570 areas are listed in Table 5 for comparison. The relative intensities refer to
 571 the absolute values divided by the total intensity. The total intensity I_{tot} in
 572 Rayleighs (R) corresponds to the total measured emission, but averaged to
 573 the area of Io's disk.

574 In order to illustrate the formation of the aurora morphology we show
 575 results of the plasma interaction simulation for Ieclipse03 in Figure 6: the
 576 atmospheric electron temperature in the flux tubes (Fig. 6a), the electron
 577 density in the equatorial plane (Fig. 6b) and the electric current in the
 578 Northern Alfvén wing (Fig. 6c). Due to the relatively low density in the
 579 equatorial atmosphere the energy stored in the flux tubes is sufficient to
 580 keep the electron temperature in most regions between 4 and 5 eV, and
 581 above 1 eV everywhere in the interaction region. The lowest temperature

582 is found on the flanks, where the electron density is highest. These high
583 density regions in the anti- and sub-jovian equatorial region mainly control
584 the position of the aurora equator spots. Depending on the overall density
585 of the equatorial atmosphere the region of the maximum ionospheric density
586 on the flanks and thus the aurora spots are shifted along the flow direction:
587 for a denser atmosphere the maxima form further upstream, for a lower at-
588 mospheric density the spots move downstream. The low atmospheric density
589 and the gas plume around Tvashtar also cause a complex current pattern in
590 the Alfvén wing. In addition to the usual current system in the Northern
591 wing, i.e. sub-jovian downward and anti-jovian upward currents, contrarily
592 oriented currents evolve on both sides within the outer usual system. Because
593 of the low neutral density around the poles and the overall low atmospheric
594 density, only on the flanks a dense ionosphere forms. Consequently, the
595 current system is inhomogeneously short-circuited in the atmosphere. Two
596 separate current systems form on the anti- and sub-jovian hemisphere respec-
597 tively. Furthermore, a small "winglet" is generated around the plume of the
598 Tvashtar volcano, which we further discuss in Section 4.3. Simulation results
599 for a similar model setup but a dense global atmosphere ($N = 6 \times 10^{16} \text{ cm}^{-3}$)
600 can be found in Saur et al. (1999).

601 In the case of a lower atmospheric surface density and thus column den-
602 sity ($N_{eq,av} < 1 \times 10^{15} \text{ cm}^{-2}$) the formation of high electron density regions
603 (ionosphere) is weaker. The electrons also cool down less, but this plays a
604 minor role at the given atmospheric thickness, since the cooling is already
605 relatively low for the best-fit atmosphere. In the case of a lower density and a
606 resulting lower ionosphere the diversion of the plasma flow decreases as well.

607 Thus the equatorial spots move closer to the surface and further downstream.
608 In this case the sub-jovian (left-hand side) spot in Figure 3d disappears be-
609 hind the limb. Furthermore, the total intensity of the simulated emission
610 averaged to the disk becomes far too low compared to the observations.

611 In the case of a higher equatorial column density of $N_{eq,av} > 3 \times 10^{15} \text{ cm}^{-2}$
612 the electron density on the upstream hemisphere increases due to increasing
613 impact ionization, the electron flow is increasingly diverted around the body,
614 and the cooling of the electrons is stronger and less energy reaches the down-
615 stream hemisphere. The equatorial spots are centered at larger distance from
616 Io and the emission on the upstream side gets stronger compared to down-
617 stream side, as the plasma in the wake is completely cooled down. Thus,
618 for such a high column density the emission morphology is not in agreement
619 with the observation.

620 Assuming a longitudinal asymmetry with a higher and denser atmosphere
621 on the anti-jovian hemisphere, as inferred for the day-side atmosphere (Jessup
622 et al., 2004; Spencer et al., 2005; Feaga et al., 2009), yields a much stronger
623 anti-jovian spot at larger radial distance. This would imply clear visibility
624 of the spot during Ieclipse03 in Figure 3b, where no emission on the anti-
625 jovian limb is detected except for the East Girru region. Furthermore, during
626 Ieclipse04 the anti-jovian spot would be brighter and at a considerably higher
627 altitude above the surface than the sub-jovian in the LORRI observation
628 (Figure3a). As the anti-jovian emission maximum is only slightly higher and
629 very close to the limb, there is no indication for an atmospheric anti-sub-
630 jovian asymmetry in eclipse. The non-appearance of an atmospheric bulge
631 on the anti-Jupiter side could be explained by the fact that it is nighttime

632 on the anti-jovian hemisphere before eclipse ingress. The response of the
633 atmosphere to a 21-hour night compared to a 2-hour eclipse event is likely
634 to be different and the anti-jovian atmospheric bulge might be reduced due
635 to that. Numerical calculations by Wong and Smyth (2000) indicate that
636 non-condensable species might even dominate the night-side atmosphere.

637 Instead, the simulation results imply a longitudinal asymmetry of the
638 atmosphere with increasing scale height and surface density from upstream
639 to downstream. Compared to a longitudinally symmetric distribution, the
640 emission is shifted downstream, which leads to a lower intensity on the up-
641 stream hemisphere and increasing intensity in the wake. The simulation still
642 predicts more radiation to be emitted on the upstream than on the down-
643 stream side, although the downstream emission is more located and thus
644 peaks higher. The observing geometry of the LORRI observations does not
645 allow a separation of emission from the two hemispheres, but the emission
646 appears to be brighter in the downstream region. However, larger datasets
647 containing a range of exposures observing both upstream and downstream
648 generally confirm a higher emission on the upstream hemisphere (Oliver-
649 et al., 2001; Retherford, 2002) in agreement with our simulation results.

650 In the HST/ACS eclipse observation (Figure 3c) on the upstream hemi-
651 sphere (right limb) aurora is observable only around East Girru, the cor-
652 responding simulation (Figure 3f) yields clear emission on the limb of the
653 upstream equatorial side. In this region the fluid approach and the assump-
654 tion of a constant background field might not describe the behavior of the
655 electrons with the required accuracy. A highly distorted and piled up mag-
656 netic field in this region may hinder the electrons to move along \vec{B} into the

657 atmosphere near the equator. The electrons with pitch angles close to 90°
658 are reflected when the \vec{B} increases and they do not reach thicker atmosphere
659 layers. Moore et al. (2010) obtained this effect in their particle simulation.
660 However, flux tubes with large \vec{B} due to the upstream pile up have a smaller
661 cross section close to Io. As these flux tubes are connected to a larger cross
662 section further away from Io, a smaller area is linked to a larger energy reser-
663 voir. The combination of the both effects require further investigation with
664 kinetic models. In the downstream region, the large cavity behind Io may
665 lead to a faster filling of the wake than we assumed. This would cause higher
666 emission in the wake.

667 Nonetheless, the position and size of the sub-jovian spot and the wake
668 feature are in good agreement for the simulation and the HST/ACS obser-
669 vation (Figures 3c and 3f). The so-called wake feature is composed of two
670 components. First, emission that is stimulated by electrons, which emerge
671 from ionization by highly energetic beams directly ‘behind’ Io. Additionally,
672 the tail of the spots can contribute to what is observed as wake emission,
673 depending on the exact viewing geometry. In our simulation the contribu-
674 tion from the flanks is stronger. HST observations with viewing longitude
675 between $\sim 60^\circ$ and $\sim 70^\circ$ (Figure 3.7 in Retherford, 2002) also revealed
676 brighter flanks than radiation directly in the wake.

677 The longitudinal positions of the simulated equatorial spots range from
678 $\sim 5^\circ$ to $\sim 40^\circ$ downstream from the zero meridian, which is roughly sketched in
679 Figure 2. This is also in agreement with the derived position by Retherford
680 et al. (2000) for HST observations of the day-side atmosphere (10° – 30°).
681 The modeled brightness ratio of the anti-jovian to the sub-jovian spot as a

682 function of the longitude of the observer is shown in Figure 7. Additionally,
683 we plotted the brightness ratio of the OI 135.6 nm multiplet from observations
684 by Retherford (2002), when Io was sunlit. These day-side observations and
685 the eclipse simulation should not be compared directly here, but it can be
686 pointed out that the model reproduces the basic longitudinal distribution
687 of the aurora. The observed and simulated ratios are in agreement on the
688 upstream hemisphere. For an observing longitude φ of 90° (into the wake)
689 and 270° (upstream hemisphere) the ratio is ~ 1 .

690 On the upstream side the ratio varies roughly linearly with φ . In a low
691 density atmosphere, e.g. during an eclipse, we would expect the spots at
692 smaller surface distances and thus the variation to be steeper than out of
693 eclipse. The simulated slope might therefore be too gradual due to an over-
694 estimation of the scale height in the model, which is presumably lower than
695 assumed in the model close to the surface (Strobel and Wolven, 2001). For
696 a smaller scale height the spots move closer to the surface and this would
697 imply a steeper slope of the spot ratio as a function of the longitude.

698 On the downstream side the relation is inverse for $80^\circ < \varphi < 100^\circ$,
699 because the elongated spots are slightly tilted inwards. Therefore the line-of-
700 sight integration for $\varphi \approx 85^\circ$ (resp. $\varphi \approx 95^\circ$) is approximately parallel to the
701 anti-jovian (sub-jovian) spot and the spot appears brighter. The ratios of the
702 wake observations by Retherford (2002) can not confirm this, but generally
703 reveal a brighter sub-jovian spot for $\varphi < 90^\circ$ in agreement with the model
704 results.

705 An absolute value for the altitude of the equatorial spots midpoints is not
706 readily determined for the observations. First, the spot emission is elongated

707 over several degrees of longitude or several 100 km and thus it does not have
708 one common altitude. Moreover, due to the two dimensional geometry and
709 the viewing angle of the observations an altitude can not be derived without
710 knowing the exact longitudinal extent or position. Choosing a hypothetical
711 viewing geometry, where the observer is located right above the North pole,
712 we can identify the height of the maximum emission on the flanks above
713 the limb for our simulation results. Depending on the emitting species and
714 the density of the upstream electrons, the local emission peaks at altitudes
715 between ~ 150 km and ~ 300 km with our best-fit atmosphere. For a higher
716 atmospheric density, the spots would appear at higher altitudes, as men-
717 tioned above, but again, the model probably overestimates the scale height
718 and thus the spot altitude for that case.

719

720 The observed overall aurora brightness is similar between Ieclipse04 ($I_{tot} =$
721 62.6 kR) and Ieclipse03 ($I_{tot} = 40.9$ kR). This agrees with the simulation, in
722 which the total intensity varies with upstream electron density. In both cases
723 the modeled intensity differs from the observed one by a factor of ~ 2 . This
724 discrepancy ($I_{obs}/I_{sim} \approx 2$) can be traced back either to plasma conditions or
725 atmospheric properties. The principal factor that controls the total intensity
726 is the electron energy, which is deposited in the atmosphere. The influence
727 of atmospheric density on the total intensity is less significant because of the
728 opposed effects of an increase of neutral gas abundance: More neutral gas can
729 be excited and radiate, but stronger deflection and divergence of the plasma
730 flow (Section 1 and Saur and Strobel (2004)). Assuming the equatorial gas
731 ring to extend to higher latitudes (up to 45° instead of 35° , see equation 4.1)

732 the total intensity would be increased by $\sim 15\%$ in our simulation. However,
733 the electron energy deposited in the atmosphere is close to maximum for our
734 derived atmosphere. So, the difference between observation and simulation
735 is probably caused by differences in the plasma environment, of which no
736 simultaneous measurements are available.

737 For instance, during the Ieclipse04 event Io passed through the torus cen-
738 ter and thus possibly through very high plasma density regions. An even
739 higher upstream plasma density than we assumed, would provide more en-
740 ergy and generate accordingly more radiation than calculated with our input
741 parameters. Additionally, magnetic field aligned electron fluxes of higher
742 energies ($\sim\text{keV}$) can possibly excite auroral emission. Michael and Bhard-
743 waj (2000) estimated a possible contribution to the aurora generation from
744 energetic field-aligned electrons observed by the Galileo plasma instrument
745 (Frank and Paterson, 1999). They found that the modeled emission is on the
746 same order of brightness as the HST observations by Roesler et al. (1999),
747 but do not consider the limited spatial distribution of the observed field
748 aligned beams. Oliverson et al. (2001) measured short term ($\lesssim 10$ min) in-
749 tensity variations in the OI 630.0 nm line and ascribed this to a time-variable
750 energy flux of field-aligned non-thermal electrons. Another possible reason
751 for the discrepancy of simulation and observation are additional emitters in
752 the wavelength range of the LORRI pass-band, which we do not account for
753 in the simulation. Emission from S_2 is expected at near-UV and blue visible
754 wavelengths (Geissler et al., 2004), but not simulated due to the lack of ex-
755 act electron impact cross sections. Since the morphology and position of the
756 spots are to major parts determined by the absolute column density and not

757 by the composition of the atmosphere, the abundance of S_2 can be neglected
758 here. In the Ieclipse03 LORRI image thermal emission around volcanic hot
759 spots clearly contributes to the observed radiation, but is not simulated.
760 Besides from these physical reasons, systematic uncertainties in the LORRI
761 images might be large due to difficulties when compiling several exposures.
762 Systematic errors such as copious instrument scattered light possibly leads
763 to an under- or overestimation of the brightness.

764 *4.2. Minor components*

765 Since the LORRI observations do not spectrally resolve the emission, we
766 compare our simulation results with a range of spectral observations of var-
767 ious emission lines resulting from oxygen, sulfur, sodium and potassium to
768 constrain the abundance of these species. We are aware of the limitations
769 of using observations made by different telescopes or cameras at different
770 times, as for example long term variations and differing signal-to-noise ratios.
771 But still, the measured intensities of the various observations differ not more
772 than by a factor of 5 and are thus still a good way to derive abundances and
773 check the derived atmospheric distribution with a wider range of observa-
774 tions. Some emission lines were frequently observed in the past, such as the
775 oxygen multiplets OI 135.6 nm and OI 630.0 nm. In that case, we preferably
776 used the more recent observations. For both lines, modeled intensity and
777 morphology are in good agreement with the observations (Retherford et al.,
778 2007; Retherford, 2002; Bouchez et al., 2000; Geissler et al., 1999; Moore
779 et al., 2010) using a mixing ratio of 12% of atomic oxygen to sulfur dioxide
780 at the surface and a surface scale height of 150 km. A comparison of the
781 intensities for the analyzed emissions lines can be found in Table 3.

782 To derive a sulfur mixing ratio we used the most recent observation of the
783 SI 147.9 multiplet taken by the NH Alice spectrograph in 2007 as benchmark.
784 The measurements revealed a varying total intensity during three eclipse
785 events of $I = 0.4 - 1.2$ kR. Assuming a sulfur mixing ratio of 1.5% at the
786 surface the model calculates a total intensity $I = 0.6 - 1.1$ kR. The intensity
787 range in the model results is due to varying upstream electron density and
788 variation of the viewing longitude. Comparisons with previous observations
789 (Wolven et al., 2001; Feaga et al., 2002; Ballester et al., 1987) also yield a good
790 agreement with the derived mixing ratio. Wolven et al. (2001) determined
791 radial profiles for the OI 135.6 and SI 147.9 emission along the spatial axis
792 of the aperture for HST/STIS observations. For the brightness close to I_0
793 averaged over the aperture width they find values between ~ 0.1 and ~ 0.8
794 kR, where the emission from sulfur is found to be slightly higher than from
795 oxygen. Depending on the sub-observer longitude, we calculate brightnesses
796 averaged tangential to limb of 0.1–0.5 kR for OI 135.6 and 0.3–0.9 kR for
797 SI 147.9 close to I_0 assuming the derived mixing ratios of oxygen and sulfur.
798 The values match the observed brightness range of Wolven et al. (2001).

799 In case of the sodium emission we use the observed intensities in eclipse
800 by Retherford (2002) and Bouchez et al. (2000) as reference. Compared to
801 optically thin intensities the abundance of both Na and K need to be cor-
802 rected by a factor of 3.5 and 3, respectively (Geissler et al., 2004). To achieve
803 emission comparable to the observed intensities in the simulated sodium au-
804 rora an abundance of sodium bearing species of 0.12 % (at the surface) has
805 to be implemented, assuming a surface scale height of 150 km. For potas-
806 sium there are no direct measurements of K assigned emission lines in the

807 close vicinity ($d < 1 R_{io}$) of Io. Using the Na/K ratio of 3.3 derived from
 808 infrared filter observations at 670–850 nm by Geissler et al. (2004) (0.04% K
 809 abundance) the total intensity of the simulated KI 767.0 nm emission line is
 810 $I = 2.2 - 4.3$ kR. This is clearly lower than the total intensity of the infrared
 811 filter observations (Geissler et al., 2004), which provide an upper limit of 10.7
 812 kR.

813 All derived atmosphere parameters and the resulting equatorial column
 814 densities are summarized in Table 4. The higher scale height of atomic species
 815 is in agreement with atmospheric models (e.g. Summers and Strobel, 1996).
 816 The higher background (polar) abundance for O, S, Na and K (n_{bg}/n_{eq}) is
 817 derived from the observed limb glow and polar emission for the respective
 818 emitter. The compared total intensities for all four analyzed components are
 819 listed in Table 3.

820 4.3. *Tvashtar Plume*

821 To constrain the plume density we analyze both the morphology and the
 822 intensity of the radiation inside the framed Tvashtar regions in Figure 3. Dur-
 823 ing Ieclipse04 the whole plume of Tvashtar was in the field of view of LORRI,
 824 see Figure 3a. The plume is radiating rather uniformly with a maximum in-
 825 tensity in the plume center ~ 220 km above the limb. The simulated absolute
 826 flux inside the framed volcano region, $I_{Tvashtar}$, is half of the observed flux
 827 from this region. Since for Ieclipse04 the simulated total intensity I_{tot} also
 828 differs approximately by factor 2 from the observation, we compared the ratio
 829 $I_{Tvashtar}/I_{tot}$ with our simulation results. For Ieclipse04 the averaged emis-
 830 sion intensity inside the Tvashtar region is almost equal to the disk-averaged
 831 total intensity, i.e. the ratio is $I_{Tvashtar}/I_{tot} = 1.0$. Assuming a gas density

832 of $n_{V,0} = 1.7 \times 10^8 \text{ cm}^{-3}$ in the main plume or an average column density of
 833 $N_V = 5 \times 10^{15} \text{ cm}^{-2}$ over the plume region the modeled plume aurora (Figure
 834 3d) is in good agreement with the measurements ($I_{Tvashtar}/I_{tot} = 0.9$), see
 835 also Table 2. A small Alfvén wing is generated above Tvashtar due to the
 836 increased neutral and thus plasma density in the region, see Figure 6c. Yet,
 837 with the derived plume density, the electrons are still able to excite emission
 838 all over the plume (Figure 3d), since the plasma flow is diverted only mod-
 839 erately. The intensity maximum is found to be in the radial center at a limb
 840 distance of $\sim 180 \text{ km}$ and plume emission appears to be rather uniformly. In
 841 the LORRI Ieclipse03 observation the upper edge of the glowing plume is
 842 visible just above the limb. The emission in the observation and simulated
 843 aurora match very well regarding morphology and relative strength.

844 For a lower plume density of $N_V < 3 \times 10^{15} \text{ cm}^{-2}$ the total modeled
 845 emission is too weak. For both LORRI observations the intensity ratio of
 846 plume and disk-averaged emission would be lower than the observed ratio by
 847 a least a factor of 2. In the case of a higher plume density, $N_V > 7 \times 10^{15} \text{ cm}^{-2}$
 848 the current system in and above the plume (see Figure 6c) becomes stronger.
 849 This leads to a stronger diversion of the electron flow around the plume
 850 region. The electrons deposit energy mostly on the upstream side of such a
 851 dense plume and the simulated aurora appears non-uniformly with a clear
 852 maximum at the upstream edge. On the other hand, the plume averaged
 853 emission intensity hardly increases with increasing plume density. The energy
 854 deposited all over the plume is almost at maximum for our best-fit plume
 855 density $N_V = 5 \times 10^{15} \text{ cm}^{-2}$, so that a denser plume does not induce a
 856 brighter auroral emission.

857 Although the viewing angle of the HST/ACS Ieclipse03 image (Figure
858 3c) allows the whole plume to be visible by HST, only a very weak UV
859 emission enhancement of $I_{tvashtar}/I_{poles} \approx 1.2$ was observed in the expected
860 region. With the derived column density of $N_V = 5 \times 10^{15} \text{ cm}^{-2}$, our simu-
861 lation predicts a ~ 3 times higher emission inside the Tvashtar region com-
862 pared to the polar area around it. The pass-band of the HST observation
863 includes mostly oxygen and sulfur emission lines. The lack of an emission
864 enhancement around Tvashtar indicate a low abundance of O and S in the
865 plume. With 3 times lower abundances of atomic sulfur ($\sim 0.5\%$) and oxygen
866 ($\sim 4\%$) within the plume the $I_{tvashtar}/I_{poles}$ ratio matches approximately the
867 HST/ACS observation.

868 As Tvashtar is considered a Pele-type plume, a high S_2 abundance is ex-
869 pected. Observations of the Tvashtar plume in scattered light and absorption
870 by Jessup and Spencer (2008) revealed a spectral behavior that is consistent
871 with previous observations of the Pele plume. For Pele various S_2/SO_2 ratios
872 between $\sim 1\%$ and 30% have been obtained by Jessup et al. (2007). If disso-
873 ciation of SO_2 and S_2 happens on time scales much larger than the average
874 time of flight for the plume gas, the abundance of atomic sulfur and atomic
875 oxygen in the plume will be lower than in an equatorial atmosphere, which
876 is not solely of volcanic origin. Moses et al. (2002) show that the lifetimes
877 for photolysis of SO_2 or S_2 producing S and O is in the range of hours, while
878 ballistic flight times are on the order of 10 minutes. Using a thermodynamic
879 model Fegley and Zolotov (2000) infer mixing ratios of atomic sulfur and
880 oxygen to SO_2 and S_2 of the order of 10^{-2} and below. This implies a low
881 abundance of S and O and would thus explain the low intensity in the UV

882 range. On the other hand, with a low abundance of O, which is also an
883 essential emitter at visible wavelengths, the simulated plume intensity for
884 the LORRI images would be reduced. Thus, the abundance of other species
885 emitting in the LORRI wavelength must be higher to sustain the plume ra-
886 diation. Another explanation for the non-visibility of the Tvashtar plume is
887 the low resolution of the HST image, which possibly masks the local volcanic
888 emission enhancement. A high resolution image in this UV range would offer
889 a possibility to determine the exact abundance of atomic oxygen and sulfur
890 in the plume.

891 The local density for a cut through the best-fit plume for the LORRI
892 observations and the vertical column density as a function of the distance to
893 the plume center (d in (8)) are shown in Figure 5. Applying the model of a
894 volcanic system developed by Kieffer (1982) and later adapted by Strobel and
895 Wolven (2001) to the derived average column density of $N_V = 5 \times 10^{15} \text{ cm}^{-2}$
896 and the implemented plume size, Tvashtar appears to be a high temperature,
897 low pressure volcano. The large size of the plume indicates relatively high
898 particles velocities and thus a high temperature at the volcanic crater, which
899 results from a high temperature ($T_0 \approx 800 \text{ K}$) reservoir in the model system
900 of Kieffer (1982). Our derived average column density is low compared to
901 common plume models (e.g. Zhang et al., 2003) and would imply a low mean
902 plume pressure of 0.1 nbar and correspond to the low pressure case (small
903 plume size) discussed in Strobel and Wolven (2001).

904 Measurements and model calculations (e.g. by McGrath et al., 2000;
905 Zhang et al., 2003) generally yield comparatively denser plumes for volcanoes
906 close to the equator. Thus, our derived plume density and pressure appears

907 to be relatively low. However, our method is hardly able to determine the
908 plume density close to the surface accurately as the observed and simulated
909 radiation is mostly emitted at a distance > 50 km from the ground. A higher
910 localized density at very low altitudes might also not influence the plasma
911 interaction strongly. If the density below 50 km increased exponentially over
912 the whole plume area with an atmospheric scale height on the order of tens
913 of km (as inferred by Strobel et al. (1994) for the equatorial atmosphere near
914 the surface), the plume column density would go up by at least a factor of
915 2. For instance, a very low sticking coefficient for molecules that contact the
916 surface could lead to a density increase at low altitudes.

917 The total content of the equatorial atmosphere, which we derived in sec-
918 tion 4.1, equals ~ 5 – 10 times the total content of a volcanic plume with the
919 plume density derived for Tvashtar. In other words, if we ruled out sublima-
920 tion as possible atmospheric source, about 5–10 active volcanoes of the size
921 of Tvashtar would be necessary to sustain the eclipse atmosphere. Pele-type
922 plumes, such as Tvashtar, are the largest observed plumes and not more
923 than 16 plumes have been observed so far (Geissler and Goldstein, 2007).
924 Therefore it is unlikely that an atmosphere of the derived column density
925 can be sustained solely by direct volcanic outgassing in eclipse, but there has
926 to be an essential amount of gas species that do not condense during eclipse.
927 The atmospheric density in sunlight is approximately ten times higher (Lel-
928 louch et al., 2007) than our derived equatorial atmosphere, so there would
929 have to be ~ 50 – 100 Tvashtar-sized active volcanoes to create such a dense
930 atmosphere without sublimation. Although already more than 150 active
931 regions have been discovered, above most regions no plumes have been ob-

932 served (Lopes et al., 2004). Hence, direct volcanic outgassing is possibly able
933 to sustain at least parts of the eclipse atmosphere, but cannot be considered
934 as an essential source for the sunlit atmosphere.

935 5. Summary and conclusions

936 We modeled the auroral emission from the atmosphere while Io was in
937 eclipse of Jupiter and investigated the effect of various atmospheric distribu-
938 tions on the intensity and morphology of the aurora. Our model results im-
939 ply an atmospheric column density of $N_{eclipse} = (1-4) \times 10^{15} \text{ cm}^{-2}$, i.e. when
940 Io moves into Jupiter's shadow, its atmospheric density decreases down to
941 $\sim 10\%$ of the sunlit atmosphere assuming $N_{sun} = (1-5) \times 10^{16} \text{ cm}^{-2}$ (Lellouch
942 et al., 2007; Saur et al., 2000; Saur et al., 2002). Despite the density decrease,
943 the atmosphere probably still covers most of the surface around the equator
944 up to $\sim 35^\circ$ of latitude at these densities. With a smooth atmospheric ring
945 and a low density (2% of equatorial density) at high latitudes we were able to
946 reproduce the main features of the auroral emission, observed in various HST
947 and spacecraft observations. Independently from the exact location of active
948 volcanoes, sub-jovian and anti-jovian bright spots arise due to the diverted
949 plasma flow. The various observed spot morphologies primarily result from
950 the respective viewing geometry of the observer and the resulting visibility
951 of the aurora features, such as the spots.

952 Compared to detailed atmospheric models (e.g. Walker et al., 2010), the
953 inferred neutral distribution appears to be simplified. However, since the ad-
954 vantage of our model is the inclusion of the influence of a chosen atmospheric
955 distribution on the plasma interaction and thus on the aurora generation, we

956 did not consider the small scale variations of the equatorial atmosphere for
957 numerical reasons but focused on the global structure.

958 Strong emission appearing in the downstream region of Io, which often
959 is denoted as “wake emission”, can be explained for major parts by flank
960 emission extending far downstream. The derived variation of the brightness
961 ratio between anti-jovian and sub-jovian spots in eclipse coincides with the
962 observed variation in sunlight qualitatively. This implies that the aurora
963 morphology is controlled by the plasma interaction also in eclipse and not
964 only by the exact distribution of the atmospheric gas and/or the locations of
965 volcanoes.

966 Analyzing various monochromatic observations we derived mixing ratios
967 for minor components in the atmosphere: 12% atomic oxygen, 1.5% atomic
968 sulfur and 0.12% sodium. The derived mixing ratios and the resulting column
969 densities are in general agreement with previous observations and model
970 results (Geissler et al., 2004; Wong and Smyth, 2000; Summers and Strobel,
971 1996), although our values are comparatively high.

972 For the Tvashtar plume we find a column density of $N_V = 5 \times 10^{15} \text{ cm}^{-2}$.
973 This relatively low plume content supports the idea of an atmosphere that
974 is sustained almost solely by sublimation. If a larger number of plumes of
975 the Tvashtar size were active in the last decades, they likely would have
976 been observed already. So far, 16 mostly smaller plumes have been observed
977 (Geissler and Goldstein, 2007) and thus volcanic outgassing probably is not
978 able to sustain a dense atmosphere as measured in sunlight.

979 Spectrally resolved observations with a resolution comparable to the
980 LORRI images would provide a possibility to determine absolute values of

981 the abundance of the various emitters. As the electron energy is distributed
982 very inhomogeneously around Io, local inhomogeneities in the neutral gas
983 abundance possibly influence the total intensity strongly. Spectral observa-
984 tions with a high spatial resolution like the LORRI images would allow a
985 detailed analysis of various regions, taking into account the local electron
986 parameters as well as the local neutral gas density and composition. For
987 example, spectrally resolved observations of the Tvashtar plume emission
988 could offer detailed information about the abundant species in this huge and
989 outstanding plume.

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Event	Date	Moment	UTC (hh:mm)	Sys-III long. ^a	Centrif. lat. ^a	Background field ^b (nT)		
						B_x	B_y	B_z
Ieclipse03	2/27/2007	Ingress	14:21	154°	4.3°	308	-580	-2145
		Mid-eclipse	15:24	183°	6.0°	79	-788	-2141
		Egress	16:27	212°	6.3°	-142	-715	-2108
Ieclipse04	3/1/2007	Ingress	08:50	255°	3.9°	-300	-321	-2022
		Mid-eclipse	09:53	284°	0.9°	-308	-11	-1966
		Egress	10:56	313°	-2.3°	-261	264	-1938

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Table 1: Eclipse occasions during LORRI and HST/ACS observations. ^a System-III longitude and the angle between Io's position and the centrifugal torus center. ^b calculated with the magnetic field model of Connerney et al. (1998) for the internal field of Jupiter.

Observation	Bandpass (nm)	Resol. (km/pixel)	Viewing geometry ^a	Figure
LORRI Ieclipse04	350-850	~56	3° S 240° W	3a
LORRI Ieclipse03	350-850	~15	7° S 310° W	3b
HST/ACS Ieclipse03	125-190	~135	~0° 344° W	3c

Table 2: Summary of the 2007 NH/LORRI and HST/ACS eclipse observations. The observations are displayed in Figures 3a-c. All observations are combinations of several exposures during one eclipse event, as reported in Spencer et al. (2007b). ^a The viewing geometry of the respective observer (New Horizons, HST) refers to the mid-eclipse time and can thus differ for single exposures.

Emitter	Sunlight observations I_{sun} (kR)	Eclipse observations I_{ecl} (kR)	Eclipse simulation $I_{ecl,sim}$ (kR)
OI 135.6	0.48 – 0.88 ^a	0.38 – 0.80 ^a	0.4 – 0.7
OI 630.0/636.4	–	6.8 ^f /14.2 ^b /23.7 ^c	6.2 – 11.8
SI 147.9	0.60 – 1.17 ^a	0.43 – 1.20 ^a	0.6 – 1.1
SI 190.0	1.68 – 2.33 ^d	–	0.5 – 0.9
NaI 588.9/589.6	–	4 ^b / 6.9 ^c	3.4 – 6.6
KI 767.0	–	< 10.35 ^e	2.2 – 4.3

Table 3: Observed and simulated total intensities of the investigated emission lines. The variation of the simulated intensity results from changes in the viewing geometry as well as various upstream electron densities ($n_e = 1900 - 3600 \text{ cm}^{-3}$). ^aRetherford et al. (2007), ^bRetherford (2002), ^cBouchez et al. (2000), ^dBallester et al. (1987), ^eGeissler et al. (2004), ^fGeissler et al. (1999)

Atmosp.	Equ. surface mixing ratio	<i>Geissler et al., 2004</i>	Sc.height H (km)	backgr. n_{bg}/n_{eq}	Long.asym. γ, β	Vert. col. dens. upstream	downstr.	Tang.col.dens. (cm^{-2})
All	100 %	-	-	-	$90^\circ, 1/3$	1.0×10^{15}	4.0×10^{15}	1.8×10^{16}
SO ₂	80 %	82 %	100	0.02	$90^\circ, 1/3$	8.0×10^{14}	3.3×10^{15}	1.5×10^{16}
O	12 %	5 %	150	0.15	$90^\circ, 1/3$	1.8×10^{14}	7.3×10^{14}	3.2×10^{15}
S	1.5 %	2.5 %	150	0.15	$90^\circ, 1/3$	1.9×10^{13}	7.8×10^{13}	4.1×10^{14}
Na	0.12 %	0.12 %	150	0.5	$90^\circ, 1/3$	1.8×10^{12}	7.3×10^{12}	6.5×10^{13}
K	0.04 %	0.036 %	150	0.5	$90^\circ, 1/3$	6.0×10^{11}	2.4×10^{12}	1.1×10^{13}
S ₂ ,O ₂ ,SO	~5%	~10%						
	Height h (km)		Width σ (km)		Mean col. dens. (cm^{-2})		Max. col. dens. (cm^{-2})	
Tvashtar	360		1100		3×10^{15}		7×10^{15}	

Table 4: Best-fit parameters of model atmosphere.

Fig.	Total I_{tot}	Equ. spots		Ratio I_{anti}/I_{sub}	North		South		Tvashtar		
		$I_{sub}(1)$	$I_{anti}(2)$		$I_N(3)$	I_N/I_{tot}	$I_S(4)$	I_S/I_{tot}	$I_t(5)$	I_T/I_{tot}	
Observation											
LORRI Iec104	3a	62.6	53.6	97.3	1.8	21.3	0.3	20.5	0.3	62.6	1.0
LORRI Iec103	3b	40.9	70.6	22.2	0.3	6.5	0.2	11.2	0.3	12.1	0.3
HST/ACS Iec103	3c	1.8	0.7 ^a	1.6 ^a	0.4 ^a	0.58	0.4	0.7	0.4	0.7	0.4
Simulation											
LORRI Iec104	3d	29.0	29.3	60.4	2.1	9.1	0.3	5.2	0.2	23.6	0.9
LORRI Iec103	3e	17.1	38.2	13.7	0.4	4.0	0.2	3.1	0.1	8.9	0.5
HST/ACS Iec103	3f	1.4	0.9 ^a	1.9 ^a	0.5 ^a	0.4	0.3	0.3	0.2	1.1	0.8

Table 5: Comparison between the emission intensities of the observations (Figures 3a-c) and the simulation results (Figures 3d-f). Absolute intensities are in kRayleigh. The measured and calculated intensities of the framed regions in Figures 3a-f are all averaged to the box areas, the total intensity is averaged to the area of Io's disk. The numbers in parentheses (1-5) refer to the green box numbers. ^a Direct comparison of absolute intensities of the HST/ACS observations with the simulation is limited, since the simulation does not take into account all emission lines in the observed wavelength range. SI 147.9 nm emission is assumed to be optically thin.

Figure 1: Observations of Io's auroral emission. (a) Multispectral eclipse image of Io taken by Cassini on January 1, 2001 (Geissler et al., 2004). Near-UV emission displayed in blue is attributed to SO₂, emission in the red visible wavelength range (red in image) is primary from atomic oxygen (OI630.0 nm). (b) HST/STIS OI 135.6 nm image of October 14, 1997 (Roesler et al., 1999).

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Figure 2: Geometry of the Ieclipse03 and Ieclipse04 observations in the planetocentric coordinate system of Io. The black arrows point out the viewing longitude of the New Horizons probe and the Hubble Space Telescope, respectively. The jovian background field vectors in the xy (orbital) plane at mid-eclipse are shown with purple arrows. Up to the right a B field vector of 500 nT is displayed for comparison ($B_z \approx 2000$ nT). In our simulation model the undisturbed plasma flow is directed in positive x direction and the background field is assumed constant on negative z direction (perpendicular to the displayed xy plane). The expected key features of the aurora (very simplified) are marked with the red areas.

Figure 3: Aurora observations and the respective simulated emission patterns. Properties of the New Horizons LORRI images (a and b) (Spencer et al., 2007a,b) and the HST/ACS image (c) (Retherford et al., 2007) are listed in Table 2. The displayed LORRI images have undergone some corrections to remove blemish emission. The corresponding simulated emission morphologies (d, e and f) are displayed with the respective viewing geometry. The color scale and the contour lines differ by a factor of 2 between observation and simulation for the LORRI images. Contours are 150 kR (a), 100 kR (b), 75 kR (d) and 50 kR (e). No contours in (c) and (f). Note that the HST/ACS observation (c) covers wavelengths from 125 – 190 nm, whereas the corresponding simulation (f) takes into account only emission from OI 135.6 nm and SI 147.9 nm. The orbital trailing (270°) longitudes are indicated with dashed meridians, the sub-jovian (0°) and anti-jovian (180°) meridians are displayed in plain bold. The emission within the green framed areas is investigated quantitatively. The total and relative intensities of the boxes are shown in Table 5.

Figure 4: Evolution of the electron temperature in the simulation model. The flux-tubes parallel to \vec{B}_0 ($\parallel z$) convect through the Io's interaction region given by the 2D plasma flow profile in the xy plane. The heat conduction perpendicular to \vec{B} is neglected, the heat conduction parallel to \vec{B} is infinite outside the atmosphere. Inside the atmosphere the plasma cools down to T_{in} and the heat flow from outside (T_{out}) is limited due to a finite heat conductivity along \vec{B} . The heat flow from the outer tube part to the inner is parametrized and depends on the temperature gradient between the two parts ($\sim (T_{out} - T_{in})/R_{typ}$).

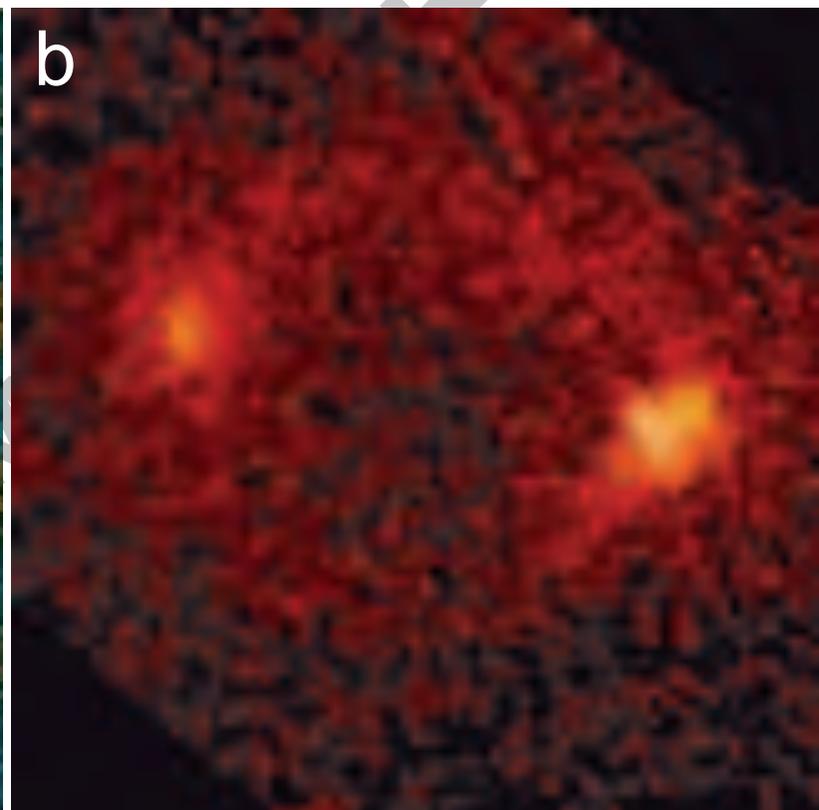
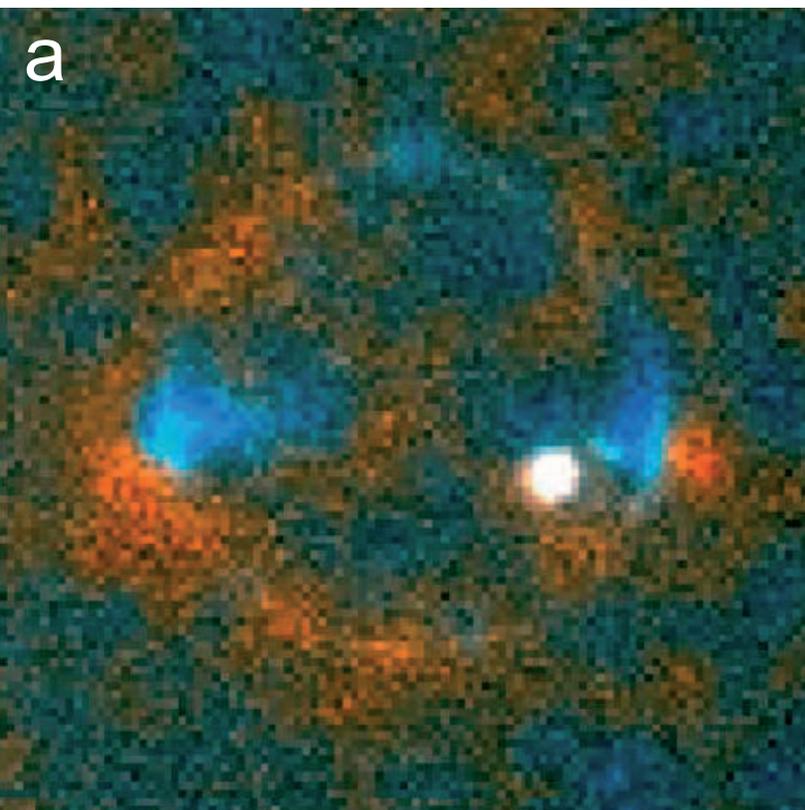
Figure 5: The density distribution model for the Tvashtar plume displayed in a vertical plane through the center. The local neutral gas density is color-coded, the dashed line represents the derived vertical column density above the plume. The plume density is given by equation 8 and consists of three components: (1) the main plume, (2) a low density ring as found by ballistic models and (3) a high density region above the crater. The shape replicates roughly the modeled Pele-like plume by Zhang et al. (2003) with a peak column density over the plume center and a shallow decrease of the column density from ~ 200 to ~ 500 km distance from the center. The plume-averaged column density is determined within a distance to the plume center of < 700 km (dashed-dotted vertical lines).

Figure 6: Results of the plasma interaction simulation: (a) Electron temperature within the atmosphere in xy plane. Isolines for 1 to 5 eV. (b) Electron density in units of torus density (1900 cm^{-3}) in the equatorial plane at $z = 0$. (c) Isolines of electric currents in 10^{-7} A m^{-2} in the northern Alfvén wing.

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Figure 7: Solid line: Modeled brightness ratio of the anti-jovian/sub-jovian spots. The brightnesses refer to the average emission intensity within a $\sim 800 \times 1100 \text{ km}^2$ box located at the equatorial limb (see Figures 3). A spot brightness ratio derived from UV observations in sunlight by Retherford (2002) is shown (+) for observing longitudes between 59° and 74° and between 241° and 297° . Dotted: The longitudinal variation of the total intensity of OI 135.6 emission normalized to its minimum.

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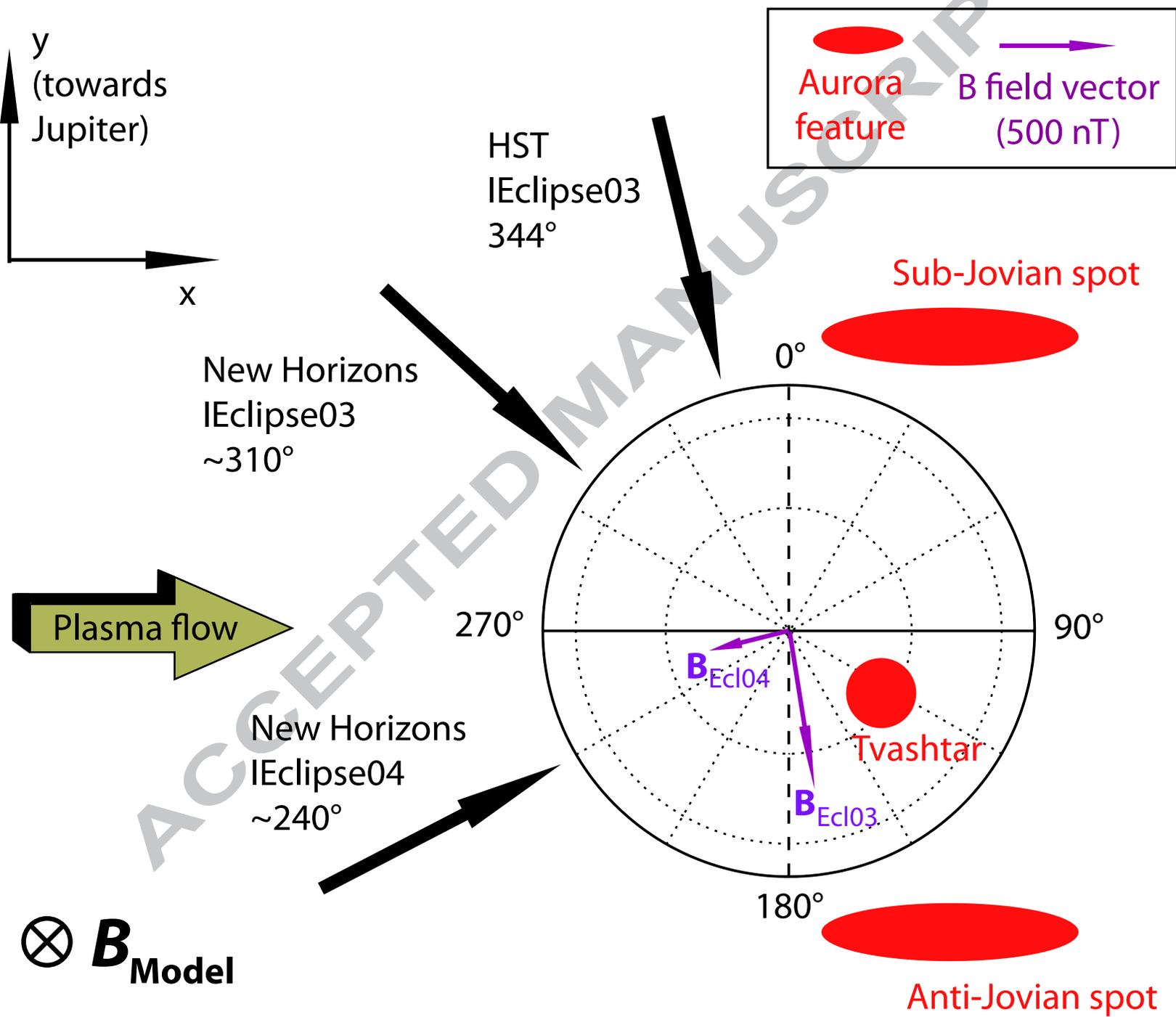


Figure 3

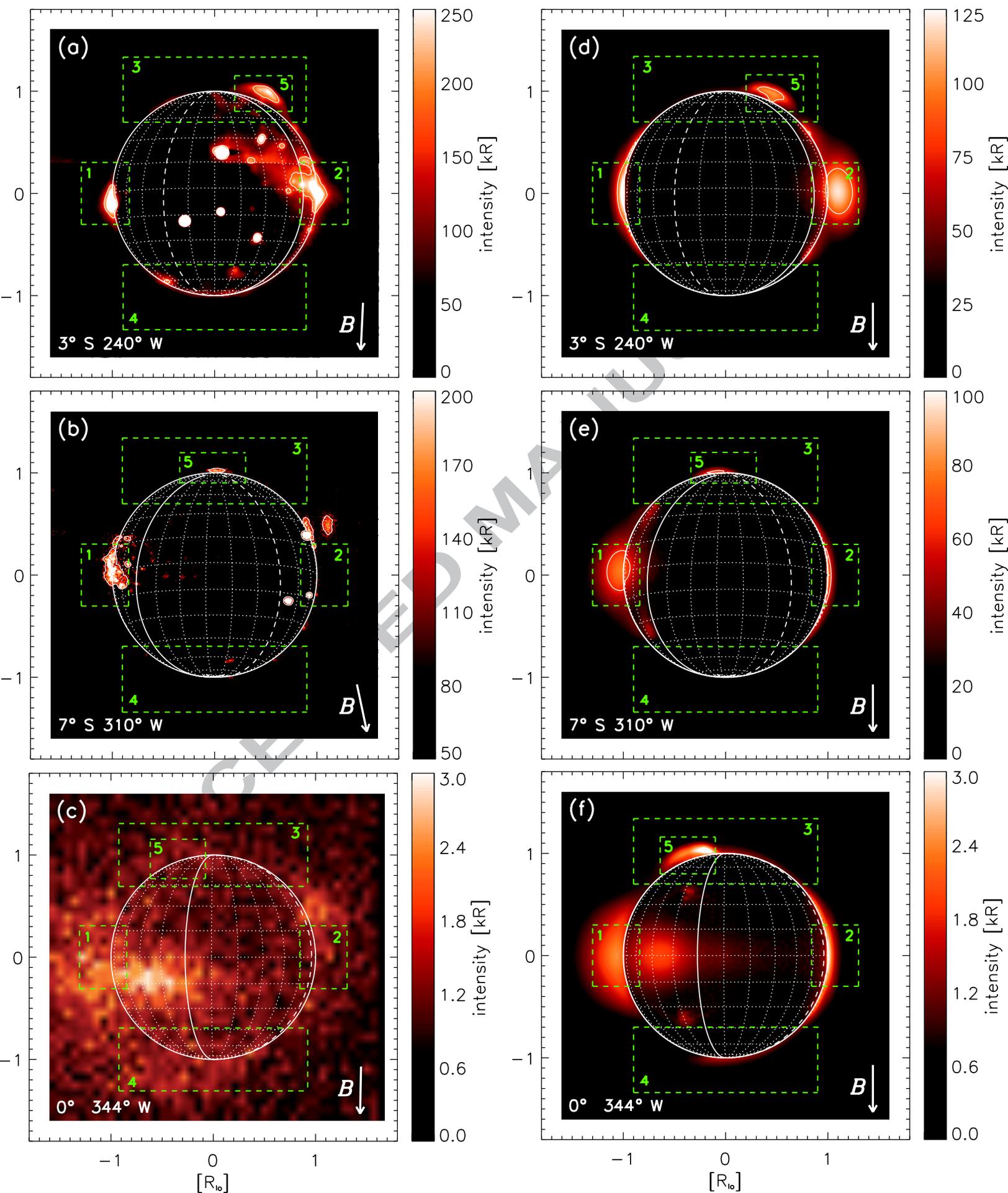


Figure 4

