MIPAS reference atmospheres and comparisons to V4.61/V4.62 MIPAS level 2 geophysical data sets

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Reliable reference profiles and estimates of variability are a necessity for a variety of processes relating to ENVISAT including the development of key aspects and inputs for the operational processor for the Michelson Interferometer for Passive Atmospheric Sounding. MIPAS reference atmospheres have therefore been produced in two forms, namely standard atmospheres for modelling and error analysis for typical atmospheric situations and the IG2 seasonal climatologies for initial guess profiles used as part of the operational processing. The reference states cover 36 species on a common altitude, pressure, and temperature grid from 0 to 120 km, and include both means and estimates of variability (maximum, minimum and one sigma values). This paper describes V3.1 of the standard atmospheres and V4.0 of the IG2 atmospheres which are the current versions of the reference atmospheres. Particular attention is paid to the MIPAS operational geophysical products (pressure/temperature, H₂O, O₃, CH₄, N₂O, HNO₃ and NO₂) and to CO₂ whose mixing ratio is required for the retrieval of pressure and temperature. A dynamic representation of CO₂ is presented which shows the presence of CO₂ gradients in the troposphere and the lower stratosphere.

Since these atmospheres have been produced independently of MIPAS data, it is also possible to compare the data to the MIPAS operational products and derive valuable information on both the reference atmospheres and on MIPAS data products themselves. This process has been performed for V4.61/V4.62 data from the year 2003 as part of the MIPAS validation activity. It is demonstrated that the agreement between the MIPAS mean data and the reference atmospheres is very good in mid-latitudes and the tropics, verifying these data to first order. There is also reasonable agreement in standard deviations between the IG2 atmospheres and the corresponding sigmas calculated from the MIPAS data. Knowledge of tropospheric concentrations of CH₄ and N₂O is used to examine the accuracy of the MIPAS data and their susceptibility to cloud effects. It is shown that for the highest accuracy, MIPAS data should be filtered with cloud index values of 2.5 for N₂O and 3.5 for CH₄. Once such filtering has

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been performed, the MIPAS data for these species appear to be accurate to within 10% in the upper troposphere. The use of cloud index data in combination with MIPAS data is recommended for studies of the polar winter stratosphere and the upper troposphere/lower stratosphere.

1 Introduction

Reference atmospheres for atmospheric gases can deliver a simple, accessible but quantitative description of the mean state of the atmosphere and likely deviations from that state. As such they are useful for a variety of purposes from inputs to radiative transfer simulations and inversion algorithms, to initial states and first order verifications for atmospheric measurements and models. The reference atmospheres described here were constructed specifically within the frame of the Michelson Interferometer for Passive Atmospheric Sounding (MIPAS) to provide well-characterized situations for development of key aspects of the retrieval process for pressure/temperature and trace gases, to provide initial guess/contaminant profile data for the operational Level 2 processing of infra-red radiance to geophysical data products and to provide atmospheric concentrations for evaluation of error budgets for the data products.

Prior to MIPAS launch, existing reference atmospheres consisted of standard profiles covering many species for typical atmosphere situations, e.g., [U.S. Standard Atmosphere \(1976\)](#), [Anderson et al. \(1986\)](#), or more sophisticated zonal mean climatologies but for restricted numbers of species and altitude ranges, e.g. [CIRA \(1986\)](#), or advanced climatologies with error/variability estimates for specific species, e.g., [Fortuin et al. \(1998\)](#) for ozone, [Chiou et al. \(1997\)](#) for water vapour. For MIPAS purposes, it was desirable to include information on many infra-red active species for both standard states and seasonal climatologies on a grid from the surface to 120 km, with estimates of variability in order to cover the varying needs of the project. An initial study by [Echle et al. \(1992\)](#) had generated typical profiles for MIPAS but the study reported here was intended to be more detailed in estimation of seasonal climatologies and standard de-

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viations. In addition, it was anticipated that the MIPAS reference atmospheres should also be more representative of the information available from the Upper Atmosphere Research Satellite (UARS) and from tropospheric monitoring of changes in source gases, particularly the anthropogenic greenhouse gases in the 1990s. With these considerations, it was decided that what was required for MIPAS reference atmospheres were two sets of reference databases: 1) a set of standard atmospheres covering five atmospheric states encompassing broad latitudinal and seasonal variability (tropical, mid-latitude day/night, polar summer and polar winter); 2) a seasonal climatology comprising a four season, six latitude climatology database (the “IG2 climatology or Initial Guess 2 climatology”).

The generation of reference descriptions of the atmosphere necessarily implies that some prior knowledge of the expected state is contained within the databases. Hence the availability of these detailed MIPAS reference atmosphere databases also provides an opportunity to intercompare with MIPAS data products since the reference databases are independent of MIPAS data. These enable both the reference atmospheres to be considered for their appropriateness and the large-scale mean and statistical variations of MIPAS data products to be verified. This exercise is performed in this paper for the MIPAS operational products and enables identification of mean behaviour, statistical variability, and cloud effects. In keeping with the MIPAS validation papers in this special issue, this paper reports on results for MIPAS V4.61/V4.62 data for the 2003 period.

2 The MIPAS instrument

The MIPAS instrument (Fischer and Oelhaf, 1996; Fischer et al., 2000) is a Fourier transform spectrometer (FTS) on ENVISAT employing limb sounding to obtain vertically resolved profiles of atmospheric infra-red emission spectra between 685 cm^{-1} ($14.60\text{ }\mu\text{m}$) and 2410 cm^{-1} ($4.15\text{ }\mu\text{m}$). Key characteristics of MIPAS include its high spectral resolution (in full spectral resolution mode), the acquisition of spectra from

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the mid-troposphere (or lower) to the mesopause (and into the mesosphere in some modes) at moderately high vertical resolution, and its global coverage with true pole to pole, latitudinal sampling. Between launch and March 2004, the nominal full spectral resolution mode of operation for the majority of acquisitions allowed for a spectral resolution of 0.025 cm^{-1} (unapodized) over the entire spectral range with only narrow gaps between the five spectral channels of the instrument. Vertical profiles in this nominal mode were scanned from 6 km to 68 km with a vertical spacing of 3 km in the UT and lower stratosphere commensurate with the field-of-view of the instrument. Each profile required approximately 76 s corresponding to an along-track distance of 500 km between vertical scans.

The essential steps in subsequent treatment of the data recorded by the instrument are transformation of the recorded interferograms to atmospheric spectra, calibration and inversion to geophysical data sets. The processing of the data to produce geolocated, radiometrically and spectrally corrected spectra at level 1b is described in Kleinert et al. (2007). The quality of the level 1b spectra produced operationally by the European Space Agency (ESA) are excellent with offsets in the calibrated spectra believed to be less than $2\text{ nW/cm}^2\text{ sr cm}^{-1}$ (Spang et al., 2005). Typical noise values for a single spectrum vary between $40\text{ nW/cm}^2\text{ sr cm}^{-1}$ at 800 cm^{-1} to less than half that value at 1200 cm^{-1} and to $3\text{ nW/cm}^2\text{ sr cm}^{-1}$ at 2100 cm^{-1} .

Within these spectra, the spectral signatures of more than thirty trace gases can be distinguished from the immensely rich information content of atmospheric emission observed at high spectral resolution. The ESA operational processing of the data to level 2 geophysical products focusses on pressure/temperature, water vapour, ozone, methane, nitrous oxide, nitric acid and nitrogen dioxide in the altitude range from 6 to 68 km. The retrieval methodologies for the operational processor are described by Raspollini et al., (2006). It is V4.61/V4.62 data produced by this processor which is the subject of the comparisons in this paper and hence a particular focus also for the reference atmosphere descriptions. Nonetheless, in the generation of the reference atmospheres for MIPAS all the major relevant species emitting in this spectral range

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had to be considered and estimates made for their concentrations.

3 Reference atmospheres

The construction of the MIPAS reference atmospheres was driven by the desire to have an accessible, useful description of expected atmospheric states which could be employed either directly in the MIPAS operational processor or used indirectly in off-line calculations related to auxiliary data or information for the processor. Standard atmospheres representing typical atmospheric states and their latitudinal variation were more appropriate for computationally expensive operations such as microwindow optimisation and selection, level 2 geophysical error estimation and for forward model optimisation. For initial guess and contaminant profiles within the processing itself, a more climatological database was necessary which encapsulated both seasonal and latitudinal variations (the IG2 database). For both types of atmospheres, estimates of variability were an important addition to climatological data sets. Both the standard atmospheres and the IG2 seasonal climatologies were constructed to cover pressure and temperature, and concentrations of thirty six species: N₂, O₂, C₂H₂, C₂H₆, CO₂, O₃, H₂O, CH₄, N₂O, CFC-11, CFC-12, CFC-13, CFC-14, CFC-21, HCFC-22, CFC-113, CFC-114, CFC-115, CH₃Cl, CCl₄, HCN, NH₃, SF₆, HNO₃, HNO₄, NO, NO₂, SO₂, CO, HOCl, ClO, H₂O₂, N₂O₅, OCS, ClONO₂, COF₂.

3.1 Sources of data

The compilation of the reference atmospheres for a large number of species was able to benefit from the efforts of a number of observational and modelling studies that delivered valuable data sets prior to the launch of ENVISAT in 2002. The zonal mean climatology for temperature was taken from the CIRA (1986) which ensured hydrostatic equilibrium. Other single product climatologies utilised included the climatology of Fortuin et al. (1998) for ozone and Chiou et al. (1997) for water vapour which were

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very useful for the troposphere and upper troposphere/lower stratosphere respectively. As part of this work also, an average tropospheric climatology of water vapour was calculated for five years of ECMWF re-analysis data (1979–1983).

Data from the UARS Reference Atmosphere Project (URAP), co-ordinated by the lead author of this publication, constituted a large element of the stratospheric observational data set. In this project, climatologies of major stratospheric constituents were developed from observations by instruments on the UARS launched in 1991; data from UARS instruments provided a significant observational advance in global knowledge of stratospheric constituents. The zonal climatologies were mostly calculated on both log(pressure) and theta surfaces; the theta data were mapped with equivalent latitude also. For the log(pressure) case, standard deviations were also estimated for the monthly data sets. “Baseline” datasets were calculated for the period April 1992 to March 1993, since this provided the maximum species coverage by UARS instruments and minimised the effects of aerosol loading from the Mt. Pinatubo eruption. However, where possible, the URAP project also allowed for “extended” atmospheres which were averaged over a longer period of time, and additional atmospheres which were averages of data not falling within the baseline period. The extended atmospheres have been used in this study where possible. Further information on some of the URAP climatologies can be found in Wang et al. (1999) and Randel et al. (1999).

Ground-based networks played a particularly important role in providing surface data for tropospheric source gases such as the anthropogenic greenhouse gases including the halocarbons, and sulphur compounds. Data from the NOAA-related networks of carbon dioxide, including the GLOBALVIEW-CO₂ (2006) system have been incorporated as have reported observations from the AGAGE network (Prinn et al., 2000). At the time of the major part of the construction of the base information for the reference atmospheres, much of the data were available only up to 2000 although the IG2 carbon dioxide dataset reported here is dynamic and is updated on an annual basis. Ground-based observations have been complemented by observations of mixing ratios from aircraft measurements, e.g. as part of the NOAA Globalview project or from cam-

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paigns, and have been useful also in checking on maximum/minimum limits of mixing ratios.

Model data were also input to the study and were used particularly for the seasonal IG2 climatologies. Stratospheric trace gas distributions were calculated using the SLIMCAT 3-D off-line chemical transport model (CTM) (Chipperfield, 1999). The model uses an isentropic vertical coordinate with horizontal winds and temperatures are specified from meteorological analyses, while the vertical (diabatic) motion is calculated from a radiation scheme within the model. The model is coupled to a detailed stratospheric chemistry scheme, which contains a treatment of gas phase and heterogeneous reactions. More details are given in Chipperfield (1999). For the results used here the model was integrated from October 1991 until March 2000 at the horizontal resolution of 7.5×7.5 degrees and 18 isentropic levels from 330 K to 3000 K. The model was forced using UK Met. Office analyses (Swinbank and O'Neill, 1994). Monthly mean data and standard deviations were computed from the model 5 day output between 23 April 1992 and 7 March 2000 with model fields interpolated (in $\log(p)$) from the model isentropic levels to standard UARS pressure levels.

Tropospheric model climatologies were generated based on a complete year of data from the Model for Ozone And Related Chemical Tracers (MOZART) global model. This is a three-dimensional Chemical Transport Model (CTM) of the global troposphere described and evaluated by Brasseur et al. (1998) and Hauglustaine et al. (1998). The model was originally developed in the framework of the National Center for Atmospheric Research (NCAR) Community Climate Model (CCM). In the version of MOZART used here (version 1), the time history of 56 chemical species is calculated on the global scale from the surface to the mid-stratosphere. The model accounts for surface emissions of chemical compounds (N_2O , CH_4 , NMHCs, CO , NO_x , CH_2O , and acetone), advective transport, convective transport, diffusive exchanges in the boundary layer, chemical and photochemical reactions, wet deposition of 11 soluble species, and surface dry deposition. This version of the model chemical scheme included 140 chemical and photochemical reactions and considered the photochemical oxidation

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schemes of methane (CH₄), ethane (C₂H₆), propane (C₃H₈), ethylene (C₂H₄), propylene (C₃H₆), isoprene (C₅H₈), terpenes (α -pinene, C₁₀H₁₆), and a lumped compound n-butane (C₄H₁₀) used as a surrogate for heavier hydrocarbons. The evolution of species was calculated with a numerical time step of 20 min for both chemistry and transport processes. The model was run with a horizontal resolution which is identical to that of CCM (triangular truncation at 42 waves, T42) corresponding to about 2.8 degrees in both latitude and longitude. In the vertical, the model uses hybrid sigma-pressure coordinates with 25 levels extending from the surface to the level of 3 mb.

A number of standard atmosphere profiles were also available (Echle et al., 1992; Anderson et al., 1986) based on previous literature and these were very useful for species, such as NH₃ for which relatively little is known about the height profile.

3.2 MIPAS standard atmospheres

The MIPAS standard atmospheres were intended to provide a updated set of profiles, pre-launch, for characteristic atmospheric states such as those derived by the studies of Echle et al. (1992) and Anderson et al. (1986). Its main objective was to provide reference states for computationally expensive calculations in the development of the MIPAS operational processor including the selection of microwindows and occupation matrices (MW/OM) selection for the operational processor, the generation of look-up tables (LUTs) for forward model calculations and error estimation. However, the atmospheres are also useful for a range of investigations from retrieval algorithm development and specification to instrument design. Since the MIPAS operational processor is based on a latitudinal switching of associated files such as microwindows and LUTs, the standard atmosphere states were also chosen to be latitudinal in form with different conditions for the poles (polar summer and polar winter) and a day/night form for mid-latitudes although the processor itself does not allow day/night switching of auxiliary data. This resulted in five atmospheric states being represented on a 0–120 km grid with 1 km spacing: tropical, mid-latitude (day/night), polar winter and polar summer (also sigma, maximum and minimum profiles). Current

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data versions are V3.1 for the MIPAS standard atmospheres (available for download at http://www.atm.ox.ac.uk/RFM/rfm_downloads.html). To allow for variation in the states, maximum and minimum values (extreme values) were calculated at each altitude. From these sigma data were also produced although the mean, maximum and minimum values are the primary products. The sigma data were extracted assuming maximum and minimum profiles were equivalent to 3 sigma of a Gaussian distribution.

It was decided that the MIPAS Standard atmospheres should primarily be based on observations where possible but clearly model data was required for the troposphere for virtually all species and in the stratosphere for less abundant species. For the major species, such as the operational products for MIPAS, the standard atmospheres could be built from data sets with much information so that variability should be estimated reasonably well. Temperature climatologies are well defined by the CIRA (1986) data with variabilities estimated from variabilities in Cryogenic Limb Etalon Array Spectrometer (CLAES) V8 data. For H₂O, data were combined from the ECMWF averages, the SAGE II climatology (particularly for the UTLS and lowermost stratosphere or so-called “middle world”) and the URAP extended atmosphere (Randel et al., 1999). For O₃ and CH₄, the extended URAP climatologies were combined with the tropospheric MOZART model and Fortuin et al. (1998) data sets (for O₃). For N₂O and HNO₃, the process used the URAP baseline monthly means (J. Gille, personal communication) produced from CLAES V8 data sets for these two molecules. The N₂O values were merged with tropospheric concentrations (see next paragraph) using profiles from the Echle et al. (1992) study.

For “well-mixed” gases such as N₂O and halocarbons, tropospheric values were in-filled assuming that surface observations provided an adequate climatological representation for such “well mixed” gases. A similar approach was used for other source gases of a similar type, i.e., CCl₄, COF₂, CH₃Cl, SF₆, and OCS; although this could have been assumed for CH₄, the gradients in the troposphere and the availability of MOZART model data suggested the use of model data instead. Table 1 shows the mean global concentrations of greenhouse gas from the standard atmospheres com-

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pared to 2000 and 2003 estimates where these are available through the IPCC 2001 and 2003 reports, and the NOAA CMDL and AGAGE networks. The question of CO₂, which is slightly underestimated in V3.1 (2%), is returned to in the next section but CH₄, N₂O, CFC-12, CCl₄ and CFC-11 (<5% change) all show reasonable values compared to the 2003 period. Major changes have occurred for HCFC-22, CFC-113, CFC-115 and CH₃Cl where the first three have risen by between 10% and a factor of four (CH₃Cl is estimated to be 25% smaller than the standard atmosphere values).

Gases for which there was much less global data at the time have been assigned constant tropospheric profiles, i.e., HCN, NH₃, SO₂, OCS, SF₆, and the less abundant halocarbons. The uncertainties in the profiles for these gases is therefore high. Finally for CO₂, the profile is also the same for all five atmospheres as its variability was not believed to be significant compared to MIPAS requirements (see next section). Essentially, the CO₂ profile was based on the estimated surface value for 2000 with a five year time lag for the stratosphere and a modified profile above 80 km consistent with rocket soundings and the ATMOS results of Rinsland et al. (1992). Typical reference profiles for CO₂ and the operational species for MIPAS are shown in successive figures. Figure 1 therefore shows the standard atmosphere profiles for mid-latitudes for all remaining species (except N₂ and O₂) as examples of the standard atmosphere profiles.

3.3 IG2 seasonal climatology

Although the standard atmospheres provide a good basic set of reference profiles, a zonal mean climatology is a more powerful means of encapsulating a greater range of atmospheric states in a regular manner whilst retaining a simple representation. In the case of MIPAS, this need for a simple climatology was driven by constraints imposed by the desire to constrain auxiliary data changes in the operational processor. This had the beneficial impact of retaining a well-defined set of profiles for initial guess and contaminant profiles at the expense of less accuracy for contaminant profiles for a particular retrieval at a given location and time. The IG2 climatology developed in

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this study represents the seasonal average atmosphere through a four season, six latitude band set of states. The four seasons are centred on January, April, July and October. The six latitude bands are $\pm(90-65^\circ)$, $\pm(65-20^\circ)$, and $\pm(20-0^\circ)$. A common altitude grid of 0–120 km is employed for the seasonal climatologies as for the five standard atmospheres. The current version of the IG2 seasonal climatology is V4.0 and is the version used here for comparison to the operational MIPAS data. The IG2 data are downloadable as part of the auxiliary data for the MIPAS processing but are also available from the lead author of this paper.

In contrast to the standard atmospheres, the regular grid of the climatology allowed for and required a more automated processing. Therefore, the IG2 climatologies were based on five sets of climatological data: the URAP data sets, the MOZART and SLIMCAT models, the ECMWF analysis and the CIRA climatology. Where no appropriate data sets are available, the IG2 climatology defaults to the most recent standard atmospheres, in this case the V3.1 data. An input grid was defined mapping each of the 24 IG2 conditions to one of the standard atmospheres. A priority order of data utilisation was employed with defined ranges of validity for each of the input data sets with essentially the troposphere for MOZART and ECMWF H₂O data and the URAP data and SLIMCAT model data for the stratosphere; the URAP data has priority in the stratosphere.

One of the major updates to the IG2 climatology compared to the standard atmospheres concerns carbon dioxide. Since the MIPAS pressure/temperature retrieval depends on knowledge of the CO₂ mixing ratio, knowledge of this gas was felt to be particularly important. Therefore, a dynamic climatology was developed compared to the standard CO₂ representation to account for trends in the mixing ratios. Clearly, the concentrations of greenhouse gases in the climatologies need continuous updates due to trends in their concentrations. However, it is also the case that for carbon dioxide there is some horizontal and vertical structure to the concentrations in the troposphere and stratosphere due to the combined effects of source distributions, the seasonal cycles in CO₂ due to vegetative uptake, and the nature of transport into the stratosphere

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with a slow transport of air to the poles. The IG2 V4.0 CO₂ seasonal IG2 climatology includes these effects. The process starts with ground-based and aircraft measurements from the Globalview data set, which are advected into the stratosphere according to age of air relationships as a function of N₂O (Andrews et al., 1999); N₂O values in the stratosphere have been derived using compact relationships derived from ATMOS data (Michelsen et al., 1998) because comparisons with equivalent relationships for CLAES data showed a clear bias in equatorial latitudes, presumably resulting from the residual effects of Mt. Pinatubo aerosol on the retrievals.

The efficacy of the age of air process is illustrated in Fig. 2 where tropical retrievals of CFC-12 from MIPAS data for July 2003, performed off-line using the OPERA scheme at Leicester (Moore et al., 2006), are compared to V3 data from the ATMOS instrument updated for tropospheric trends (Cunnold et al., 2002). The agreement is very good as might be expected in the tropics where the age of air is younger than in the poles. The IG2 and tropical standard atmosphere are biased high due to the Pinatubo effect on CLAES data. Uncertainty in the age of air will increase towards the poles but variations in CO₂ should be close to the expected one sigma of 2 ppmv. To be conservative above 25 km, where the age of air becomes more uncertain, constant values are assumed for CO₂ due to the tropospheric source as represented at 25 km and only a small methane oxidation term changes the vertical profile in the upper stratosphere. Above 80 km, upper atmosphere CO₂ has been further modified to follow Lopez Puertas et al. (2000). The atmosphere is updated on an annual basis using the latest CO₂ Globalview data sets.

Figure 3 shows the final V4.0 latitude dependent profiles produced for the 2003 annual average. On a full scale (left hand plot), there is clearly little variation of CO₂ in the troposphere and stratosphere. However, a finer scale (right hand plot) highlights the structure present in the CO₂ profiles, particularly in the troposphere but also in the vicinity of the tropopause. In the lower troposphere, there is a clear difference between the polluted northern hemisphere and the cleaner southern hemisphere. There is also considerable variability in the upper troposphere with potential for strong gradients into

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the lower stratosphere. In the northern hemisphere this gradient may be particularly strong because of the strong tropospheric seasonal cycle and anthropogenic inputs coupled to a long age of air for the lower stratospheric layers overlying the troposphere. Differing values in the stratosphere reflect age of air variations with latitude as expected.

Similarly it will be important in the future to re-visit the MIPAS reference atmospheres for all tropospheric species to consider in more detail the variability of key species. Currently the IG2 seasonal climatology already contains variations, for example of tropospheric ozone. Nonetheless, scrutiny of the climatologies and updates for the range of satellite data available will provide insight into the climatological representation.

4 Comparisons with MIPAS operational geophysical products

The seven operational MIPAS products are profiles of pressure/temperature, H_2O , O_3 , HNO_3 , CH_4 , N_2O and NO_2 retrieved along an orbit-track,. As for the validation papers in this MIPAS special issue, the aim of this study is to describe the chief characteristics of V4.61/V4.62 of the MIPAS operational product. The architecture of the operational retrieval and optimised forward model, and the strategy for its implementation, are described in detail in Raspollini et al. (2006). Briefly, the retrieval is based on a global, non-linear least-squares fit of each entire limb sequence of spectra, with products being retrieved sequentially in the order in which they are listed above. The retrievals are performed in optimised microwindows extracted from calibrated, geo-located and cloud flagged level 1b spectra and utilise a dedicated spectroscopic database. Estimated errors for the products are described in Raspollini et al. (2006) and can be found in more detail at <http://www-atm.atm.ox.ac.uk/group/mipas/err/>.

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4.1 Cloud flagging

Cloud flagging is of some importance for MIPAS data quality. In MIPAS, cloud flagging is implemented in the operational processor, as described in [Spang et al. \(2004\)](#), using pairs of microwindows in the MIPAS spectral channels A, B and D in order of priority.

5 The band A pair of microwindows is therefore of the highest priority and is used to flag the majority of MIPAS data as cloudy or non-cloudy; its use depends only on the validity of the MIPAS A spectral data recorded for any limb observation (or sweep). The cloud index for channel A (CI-A) is defined as the ratio of the signals in two designated cloud microwindows (MIPAS band A), 788.2 to 799.25 cm^{-1} and 832.3 to 834.4 cm^{-1} , such
10 that values below a threshold value are used to indentify particularly cloudy spectra which are not used in the geophysical product retrievals. The V4.61/V.62 processor employed a value of 1.8 for the CI-A threshold. In fact optically thinner clouds (cloud index greater than 1.8) may also cause residual errors in the trace gas retrievals and the definition of threshold is dependent to some extent on the degree of systematic
15 error that can be tolerated for a given application. For example, [Waterfall et al. \(2004\)](#) with respect to all gases and [Glatthor et al. \(2006\)](#) with respect to ozone have indicated values of 2.2 and 4.0 respectively for more rigorous flagging than the current operational processing. The comparisons to reference atmospheres in this study afford some evidence for better definition of the CI-A threshold, in particular for CH_4 and N_2O . All
20 MIPAS geophysical data examined in this study are only utilised if a valid CI-A cloud flag can be computed and its value examined.

4.2 Results for MIPAS operational data

In this study, each of the MIPAS geophysical products have been compared to the MIPAS reference atmospheres on a monthly basis for each month in 2003. However most
25 of the results in a climatological sense are consistent across the year and hence the plots in Figs. 4 to 18 are based on the month of July 2003, except for those for O_3 . Each of the main plots is composed of two panels. In the left hand panel, labelled (a), the

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objective is to examine mean profiles and compare the variability of MIPAS data with that expected. Individual MIPAS data points are represented by small crosses, with black indicating cloud-free measurements, and red crosses indicating measurements with a CI-A of less than 2.5 (see comments below for CH₄ and N₂O). All data plots plotted are those with a valid CI-A flag and with valid convergence criteria. Ranges for the plots are chosen so as to show the chief features of the profiles and not all MIPAS points may appear on them. In particular, it was observed in this study that all MIPAS trace gas products show values of 10⁻¹⁰ occasionally, these values being the result of retrieval instabilities (see Fig. 17 for NO₂). Mean profiles from the MIPAS data are represented by the larger green (mean of data with CI-A >2.5) and red (all data) crosses. The reference atmosphere profiles plotted are the corresponding IG2 climatologies (orange line) together with the associated three-sigma values (dotted orange lines) and the appropriate standard atmosphere with maximum and minimum thresholds (blue lines). Since the standard atmospheres are only derived for five atmospheric states, one would expect that these would show the larger discrepancies from the observed data. The right hand panel plots illustrate the variability in a different manner, i.e. via a calculated sigma for the various data sources: one-sigma values from the IG2 climatologies (light orange dotted lines) compared to sigma uncertainties in the standard atmospheres ((max-min)/6: the dark orange dashed lines) and the standard deviation of MIPAS measurements for a given month for cloud-free scenes (green diamonds) and cloud-flagged scenes (black asterisks). Sigmas derived from maximum/minimum profiles will not necessarily agree with the IG2 sigmas and the latter should be a better estimate. An upper limit of 0.01 mb is used in the plots based on the upper altitude range of MIPAS data.

Temperature data shown in Fig. 4 indicate the general agreement with MIPAS data in mid-latitudes and tropics. Monthly mean data in 2003 typically agree very well away from the polar regions. For temperature, the standard deviations in the IG2 atmospheres are derived from the standard atmospheres directly hence only one variability line on the plots. The agreement shown is also typical with MIPAS data tending to

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show a variability of approximately 3 K in the middle stratosphere, away from the winter pole, whereas the reference sigmas tend to vary between 3 and 5 K. In the polar regions, there are larger discrepancies between MIPAS data and the IG2 climatology in either the lower stratosphere, e.g. in the Antarctic winter, or in the upper stratosphere in the summer pole. However, there is considerable variability in temperatures in this region and further study is required to investigate whether the effects are interannual variability or problems with either the instrument data or the reference atmospheres.

Figures 5 and 6 show comparisons for H₂O which essentially cover two main regimes, the dry stratosphere and the increasingly moist troposphere. In the troposphere, mean MIPAS profiles and reference atmosphere profiles agree well in shape but the mean MIPAS data (cloud free) are lower than both reference atmosphere profiles, a result which is typical of comparisons for 2003. This apparent discrepancy is likely to be a consequence of the fact that the reference atmospheres are expected to be typical of all conditions whereas the MIPAS data are biased towards thin cloud and clear sky. However, the estimated systematic uncertainty in MIPAS observations is also large in this region and of the order of 30%. It is also noticeable that whilst there are considerable positive discrepancies in the cloudier flagged data (mean of all data greater than mean for CI-A <2.5), there are also a number of very low MIPAS data points particularly near the cold tropopause. Although there is a possibility of low H₂O mixing ratios, evidence from retrieval simulations suggest that the lowest MIPAS values can arise from two sources: (a) limitations in the MIPAS operational processor to retrieve accurately in this situation; (b) a tendency of the H₂O retrieval to assign anomalously low values at heights immediately above unflagged clouds near the CI-A threshold. The dry constrained nature of stratospheric H₂O, with values limited to a few ppmv is evident in Fig. 5 for the tropics. Mean values agree particularly well between 20 mb and 5 mb with the MIPAS mean data higher than the reference atmosphere data. At pressures of greater than 20 mb, some tendency towards an oscillating mean profile can be seen. Individual data points sit very well within the max/min profiles and standard deviations are in reasonable agreement indicating random noise and atmospheric

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variability of 10% or less 1 sigma; expected random errors vary between 5 and 10% in the middle stratosphere. Similarly good comparisons are generally observed between MIPAS and the IG2 data in the southern hemisphere summer but the sigmas are larger for MIPAS data in the northern hemisphere winter.

For ozone, data from October 2003 illustrates the main features very well (Figs. 7 and 8). In the tropical stratosphere, there is excellent agreement between mean MIPAS profiles and the reference atmosphere profiles, demonstrating the well-constrained and characterised stratospheric profile. As for H₂O in the stratosphere, climatological representations tend to do a good job of representing most of the stratospheric behaviour. Relative to MIPAS data, the minimum profile tends to be rather low but it is also noticeable that the one sigma standard profiles tend to be smaller than the IG2 and MIPAS one sigma profiles. This implies that distributions are somewhat narrower than would be expected from maximum and minimum profiles as representing three sigma extremes. In the southern polar spring, much better agreement between sigma profiles is obtained. The low ozone values for the mean profile and the data points are most likely real with the neither reference atmosphere particularly capturing the lowness of the Antarctic ozone hole in this year. However, the IG2 data do a better job of capturing the variability represented by the individual MIPAS observations than do the standard atmosphere one sigmas. Both plots show high and low ozone values potentially associated with CI-A less than 2.5 in the troposphere, with more variability clear in the Northern high latitudes. Tropospheric concentrations from MIPAS are lower than the reference atmosphere profiles, particularly in equatorial regions and mid-latitudes (not shown).

Methane and nitrous oxide are dynamical tracers which are “well-mixed” in the troposphere with faster sinks in the stratosphere. Their observed distributions therefore provide a good test of data quality. Figures 9 and 10 show results for CH₄ in southern mid-latitudes and northern polar summer respectively. The agreement between the MIPAS data and the reference atmosphere is generally very good in both mean and variability compared to standard max/min and IG2 sigma. In Fig. 9, at altitudes near

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1 mb, there is a lower CH₄ concentration probably due to descent since it becomes larger near the southern pole. Near 10.0 mb, an increased variability is observed which is under investigation. In polar summer, the agreement between the mean data sets is excellent for the entire profile. The IG2 sigmas appear to be considerably overestimated as do the max/min profiles. In the polar summer mean plot, there is a noticeable set of red points which deviate considerably from the expected tropospheric concentrations in both a negative and positive sense. The distributions of CH₄ are well known in the troposphere with small variances of less than 10% expected as can be seen from the reference atmosphere variability profiles. Here it is the MIPAS data that are in error due to residual clouds effects with somewhat persistent polar cirrus likely to be present. Similar effects can be seen for the southern mid-latitudes plots although not as clearly. The red points at pressures of less than 10 mb are not clouds but are due to noise in the CI-A test at 2.5; recall that the operational processor flags the data at 1.8 which is a lower threshold.

The cloud effects are examined further in Fig. 11 where the average and standard deviation of MIPAS data have been computed as a function of CI-A (bins of 0.5 in CI-A) for each hemisphere at latitudes of less than 50° and pressures greater than 150 mb. For low CI-A values, indicating thicker clouds/aerosol particle concentrations, it is clear that both the mean and the standard deviations of the measurements increase considerably for CI-A less than 3.0 or even 3.5 (northern hemisphere plot). Using this test, it is not possible to be exact about the expected behaviour since CH₄ in the upper troposphere will deviate slightly from constant volume mixing ratio if only due to transport and mixing of tropospheric and stratospheric air. In addition, there is a hemispheric gradient of approximately 0.1 ppmv as well as possible variability due to fast transport of surface concentrations. So for example, the gradient observed for the southern hemisphere with CI-A may be entirely due to the gradient of CH₄ with height rather than particle effects on the retrieval. Expected values of tropospheric surface CH₄ have been estimated from values observed by the AGAGE network (Prinn et al., 2000; Cunold et al., 2002) as 1.829 and 1.732 ppbv for the northern hemisphere and

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southern hemispheres respectively. Applying a threshold of CI-A of 3.5 would then suggest that more stringent cloud flagging results in “clear sky” CH₄ values which are less than 10% high and that the random error in the retrieval (noise) is of the order of 10% also. These error estimates are broadly consistent with those expected from the MIPAS estimated error budget for the upper troposphere, with the limiting factor being temperature uncertainties.

Similar evaluations have been performed for nitrous oxide. The comparisons here are shown on the usual altitude scale for southern mid-latitudes and an expanded altitude scale plot for the tropics to illustrate effects on N₂O in more detail. In Fig. 13, the mean profile and standard deviations are in excellent agreement in the stratosphere with a considerably lower MIPAS mean profile in the upper stratosphere. Again there is an effect of increased variability at approximately 10 mb as for CH₄. The results shown in Fig. 12 illustrate how the reference atmosphere data encapsulate both mean and standard deviations rather well in the lower stratosphere. However from 100 mb and below, higher deviations can be seen which are most likely cloud effects and also the data are perhaps more variable indicating that single profile noise is higher than expected atmospheric variability. Figure 14 shows plots for N₂O mean concentrations versus CI-A, as for CH₄. Similar results are found, with high sensitivity to CI-A in mean and standard deviation for CI less than 2.5. For higher CI-A thresholds, residual uncertainties might be just over 10% systematic and 10% random error, reducing a little as the CI-A threshold is increased. It is possible that the actual total systematic error in MIPAS data from non-cloud sources is less than 5% but it is not possible to be definitive about this.

For the remaining two gases, nitric acid and nitrogen dioxide, it is not possible to undertake the same types of cloud tests. Figures 15 and 16 show data for nitric acid for northern polar summer and for southern mid-latitudes respectively representing contrasting views of nitric acid. Considering the former first, excellent agreement is seen in profile shape and magnitude between 100 mb and 7 mb. At higher altitudes, the random error increases considerably but there is also a significant difference in the

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nitric acid data. At lower altitudes, cloud effects become significant and the mean HNO_3 profile, even for the green crosses, is considerably larger than the reference state in the upper troposphere at altitudes below a pressure of 150 mb. It is not possible in this study to say whether this is due to cloud effects on retrievals or interactions in the atmosphere between cloud particles and nitric acid.

Figure 17 shows nitrogen dioxide data comparisons and highlights some of the current issues with the operational NO_2 product. The MIPAS data for NO_2 are retrieved in a restricted altitude range within which the mean and sigma values for NO_2 are well characterised in the reference atmospheres and match well with MIPAS data. This is particularly true of the standard deviations in all latitudes and altitudes except for the southern polar regions (not shown). In the mean profile, the comparisons are very good except at the high altitudes where the diurnal variability of NO_2 complicates interpretation of the plots. However, clear separation can be seen in the southern midlatitude plot for regions above 1 mb, with part of the data set clearly converging near the IG2 climatological profiles, and part of the data set retrieving significantly lower concentrations closer to the nighttime standard atmosphere. Erroneous values of 10^{-10} are also present in the operation product, and shown here. These are believed to be instabilities in the operational retrieval processor within the limits of convergence allowed for computing time.

5 Summary

In this paper, the MIPAS reference atmospheres have been introduced and descriptions given of the current five standard atmospheres (V3.1) and the IG2 seasonal climatology (V4.0). These atmospheres include not only mean concentrations for 36 species from 0 to 120 km, as well as pressure and temperature, but also estimates of variability. For the standard atmospheres, the aim was to capture the maximum/minimum values of each species at each altitude, i.e., the extreme values, in order to assess the full range of possible atmospheric distributions within practical constraints. One sigma

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standard deviations can be estimated from the maximum/minimum profiles, assuming Gaussian statistics but this clearly is not a good assumption for many trace species. For sigma values, it is probably better to use the IG2 data for which these values have been calculated explicitly. A significant aspect of the IG2 climatology is the addition of a dynamic, yearly dependent update through the development of a time varying CO₂ climatology. For other “well-mixed” gases in the troposphere, changes were not found to be significant for most species but HCFC-22, CFC-113 and CFC-115 clearly require further updates.

The MIPAS reference atmospheres have been compared to MIPAS observations for 2003 (full spectral resolution mode) and are revealing. It is found that the MIPAS data agree very well with the reference atmospheres in mean behaviour at most altitudes, verifying the consistency of the reference states and the independent MIPAS data. Agreements are particularly good for the middle stratosphere data. Discrepancies due to clouds and noise have been noted at the upper and lower limits of the MIPAS retrieval range. For methane and nitrous oxide, it has been demonstrated that cloud influences on the retrieval are important, and are presumably so for other species. The results presented here suggest that for the highest accuracy, MIPAS data should be filtered with cloud index values of 2.5 for N₂O and 3.5 for CH₄. Once such filtering has been performed, the MIPAS data for these species appear to be accurate to within 10% in the upper troposphere. The presence of data with values of 10⁻¹⁰ has also been identified in all the products despite the use of MIPAS convergence flags. It is believed that these arise from retrieval instabilities. There are also some regions where MIPAS data are not in good agreement with the mean climatological profiles particularly in polar regions in the lower stratosphere and troposphere. These will be investigated further. The results presented here illustrate the significant utility of comparisons of observed data with reference climatological atmospheric states. The use of cloud index data in combination with MIPAS data is recommended for studies of the polar winter stratosphere and the upper troposphere/lower stratosphere.

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Table 1. Comparison of MIPAS standard atmosphere mean tropospheric values for well-mixed gases compared to IPCC estimates (IPCC, 2001, 2005). Values for CH₃Cl are derived from Mace Head(a) and Cape Grim(b) observations (O'Doherty et al., 2001).

Gas	MIPAS reference atmosphere (pptv)	2000 estimate (pptv)	2003 estimate
CO ₂ (ppmv)	369	367	375
CH ₄ (ppmv)	1.760	1.745	1.732
N ₂ O (ppmv)	3.17	3.14	3.18
CFC-11	265	268	256
CFC-12	545	533	538
HCFC-22	140	132	157
CFC-113	19	84	80
CFC-114	12	15	17
CFC-115	4	7	17
CH ₃ Cl	700	—	532 ^a , 510 ^b
CCl ₄	108	102	96

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Fig. 1. Equatorial profiles of all species covered by the IG2 climatologies, excluding CO_2 , N_2O , NO_2 , O_3 , HNO_3 and H_2O which are covered elsewhere.

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Fig. 2. Comparison of mean MIPAS CFC-12 offline retrievals (OPERA) for July 2003 with AT-MOS data updated using age of air and tropospheric trends. The corresponding tropical IG2 profile for July and the tropical standard atmosphere are also shown.

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Fig. 3. Left Panel: IG2 profiles for CO₂ for all 6 latitude bands for 2003. Right Panel: Detail of the variability with latitude and altitude below 40 km in the 2003 CO₂ data.

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Fig. 4. Comparison of MIPAS and reference atmosphere temperature data for latitudes from 20°N to 65°N for July 2003: Left panel - Operational MIPAS data from July 2007 for “cloudy” data (CI-A <2.5 – red dots) and “clear-sky” data (CI-A >2.5 – black dots). Means for the cloudy data (large red crosses) and clear data (large green crosses) at each scan altitude are also shown. Plotted in blue are the standard atmosphere profile and associated maximum and minimum values (dashed lines) for this latitude band and period. Orange lines represent the IG2 climatologies with 1 sigma values (dashed) for this latitude band and period. Right panel - The standard deviation of the MIPAS values at each scan altitude are shown in green diamonds (clear sky data only) and black asterisks (all data including cloudy data). One sigma values for the IG2 climatologies are shown in light orange dotted lines, while an estimate of one sigma values for the standard atmospheres is shown by the dark orange dashed lines.

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Fig. 5. Comparison for H₂O for latitudes from 0N to 20N for July 2003. Annotation as per Fig. 4.

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Fig. 11. CH₄ concentrations retrieved by MIPAS during July 2003 below 150 mb binned by associated CI-A value. Blue asterisks denote mean CH₄ concentrations for each CI-A bin. Red asterisks denote the standard deviation within each bin. The left panel shows data for latitudes from 50 S to 0 S, the right panel shows data for 0 N to 50 N.

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Fig. 13. Comparison for N_2O for latitudes from 65 S to 20 S for July 2003. Annotation as per Fig. 4.

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Fig. 14. N₂O concentrations retrieved by MIPAS during July 2003 below 150 mb binned by associated CI-A value. Blue asterisks denote mean N₂O concentrations for each CI-A bin. Red asterisks denote the standard deviation within each bin. The left panel shows data for latitudes from 50 S to 0 S, the right panel shows data for 0 N to 50 N.

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Fig. 16. Comparison for HNO_3 for latitudes from 65 S to 20 S for July 2003. Annotation as per Fig. 4.

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Fig. 18. Comparison for NO₂ for latitudes from 65 S to 20 S for July 2003. Annotation as per Fig. 4.

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