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Emissions of primary aerosol and precursor gases in the years 2000 and 1750, prescribed data-sets for AeroCom

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Abstract

Inventories for global aerosol and aerosol precursor emissions, and auxiliary information, have been collected, assessed and prepared for the year 2000 (present-day conditions) and for the year 1750 (pre-industrial conditions). These global datasets establish a reference for input in global modeling, when simulating the aerosol impact on climate with state-of-the-art aerosol component modules. These modules stratify aerosol by type, distinguishing among dust, seasalt, sulfate, organic matter and soot. The datasets are also intended to serve as systematic constraints in sensitivity studies of the AeroCom initiative, which aims to evaluate uncertainties in aerosol global modeling. The datasets comprise daily size-resolved emissions of sea-salt and dust and monthly-to-yearly emissions for all other currently known emissions of natural and anthropogenic aerosol (precursors). The emissions are a reference dataset for aerosol modeling in the coming years and benchmark the emissions according to our knowledge in the year 2004.

1 Introduction

Aerosol introduces large uncertainties in climate modeling (IPCC, 2001). These uncertainties are fueled by limitations to adequately represent aerosol amount, aerosol properties and aerosol interactions (e.g. the aerosol impact on cloud properties) in global modeling. In 2003 the Aerosol Inter Comparison project AeroCom (http://nansen.ipsl.jussieu.fr/AEROCOM/) has been initiated to identify the specific nature of these uncertainties. Detailed output (of aerosol modules) of more than 15 global models was systematically analyzed, mainly by comparing to available data from in-situ sampling and remote sensing. In initial comparisons of the AeroCom Experiment “A” modelers were allowed to use aerosol emissions (an essential model input) of their choice, without a clear connection to any particular year. This complicated the interpretation of model diversity (Kinne et al., 2005; Textor et al., 2005). It remained
unclear, if unusual model-tendencies for any particular aerosol component related to model input (e.g. emission-data) or to aerosol processing and transport. In an effort to harmonize aerosol input, two additional AeroCom simulations were requested with prescribed emission data for primary aerosol and for aerosol precursor gases. AeroCom Experiment “B” ties model input (e.g. aerosol emissions and meteorological data via nudging in GCMs) to the year 2000 in an effort to evaluate models with observational data available for that year. AeroCom Experiment “Pre” was designed to establish a pre-industrial (year 1750) modeling reference, to extract the anthropogenic aerosol impact in conjunction with Experiment “B” (Schulz et al., 2006\textsuperscript{1}). For both Experiments “B” and “Pre” global data sets for aerosol emissions needed to be defined. Background and details to these data-sets are here introduced.

2 Common features

For the years 1750 and 2000, global aerosol emission fields were developed at a spatial resolution of 1.0°[latitude]×1.0°[longitude] (in units of kg per grid-box). Temporal resolution ranges from daily to yearly depending on the species. Information was given for injection height and for sizes of particulate emissions. Emissions are categorized by their origin as either natural or anthropogenic.

Natural emissions from dust, sea-salt and volcanic sources are assumed to be identical in both Experiment “B” (year 2000) and Experiment “Pre” (year 1750). Details of common emissions are introduced in Sect. 3. All other emissions, which changed during

ing industrialization, are introduced for Experiment “B” in Sect. 4 and for Experiment “Pre” in Sect. 5. Temporal resolution of all data-sets is summarized in Sect. 6, injection heights are explained in Sect. 7 and size recommendations are addressed in Sect. 8. All emission data-sets are available (in netcdf data format) via a dedicated file transfer site at the Joint Research Center in Ispra, Italy: ftp://ftp.ei.jrc.it/pub/Aerocom/. To help in locating specific data, all ftp-site subdirectories and their content are listed in Appendix A.

3 Common emissions

Natural emissions (common to both emission scenarios) include wind-blown contributions of mineral dust (DU) and sea-salt (SS), sulfur (S) contributions from volcanoes and DiMethyl Sulphide (DMS, mainly oceanic) and Secondary Organic Aerosol (SOA) formed from natural Volatile Organic Compound (VOC) emissions. Most of S and all of SOA are formed within the atmosphere from precursor gases. Hereby, SOA is added to the flux of Particulate Organic Matter (POM). An overview of recommendations for injection height and particle size and a comparison of annual total fluxes to IPCC-TAR estimates are given in Table 1.

Table 1 also lists size-recommendations in terms of log-normal size-distribution parameters \( r_m \) (number mode radius) and \( \sigma \) (standard deviation), plus the resulting \( r_{\text{eff}} \). This radiatively most “effective radius” is defined by the ratio of sums by third and second radii moments, \( \Sigma r^3 / \Sigma r^2 \). A daily temporal resolution was adopted for emissions tied to near surface winds.

3.1 Dust

Emission estimates for dust (DU) are based on simulations with near surface winds of the year 2000 generated by the NASA Goddard Earth Observing System Data Assimilation System (GEOS DAS). The daily average DU flux output was dis-
distributed over four size-bins (radii ranges of 0.1–1.0, 1.0–1.8, 1.8–3.0, 3.0–6.0 µm) at 1.0°[latitude]×1.0°[longitude] horizontal resolution (Ginoux et al., 2001, 2003). To accommodate modal size schemes (frequently used in global modeling) DU flux data (after subdividing the original size-bins each into ten subsections of equal radius-range and number) were stratified according to size into two domains: accumulation mode (radii: 0.05–0.5 µm) and coarse mode (radii >0.5 µm). Then for each size domain the flux was distributed over a log-normal function, which is defined by the three parameters of mode-radius $r_m$ (radius at the peak concentration), standard deviation $\sigma$ (distribution width) and number $N$. Assuming a DU density of 2.5 g/cm$^3$ and prescribing the distribution width (a standard deviation of 1.59 for the accumulation mode and 2.0 for the coarse mode) the number mode-radius $r_N$ was determined for each mode at each grid-point and time-step (as data for number $N$ and mass-flux were provided by the original data). Daily global fields for mode-radius and number (along with prescribed values for density and standard deviation) established the recommended emission input for DU in modal size schemes. To accommodate aerosol modules with (size-) bin schemes, software is provided, which can extract the DU emission flux for any DU size-range. Based on the modal approach 98.6% of the DU flux mass is assigned to the coarse mode (and 1.4% to accumulation mode). The spatial distributions of DU emissions on an annual and a monthly basis are given in Fig. 1.

The characteristic DU size is about 4 µm in maximum dimension (a coarse mode radius of 0.63 µm combined with a standard deviation of 2.0 translates into an effective “radius” of 2.1 µm). DU emissions are prescribed to take place in the lowest model layer. Biases may have been introduced by limitations of the GEOS DAS meteorology and simplifying assumptions. The coarse (daily) temporal resolution and injections into the lowest model layer only are expected to contribute to DU flux underestimates. Inaccuracies are introduced by the simple modal size representation and by model specific implementations (e.g. inconsistencies with boundary layer mixing (e.g. convection) and/or adaptations to different model resolutions).
3.2 Sea-salt

Sea-salt (SS) daily emission data are based on year 2000 ECMWF near surface winds. SS mass fluxes are provided over 24 size bins (covering radii from 0.005 to 20.48 µm) at 1.175°[latitude]×1.175°[longitude] horizontal resolution (Gong et al., 2003). Contributions of SS emissions associated with radii larger 10 µm were ignored and SS contributions over sea-ice were removed according to monthly ECMWF sea-ice-free-fractions for the year 2000. The data were regridded to a horizontal resolution of 1.0°×1.0° and redistributed (after subdividing the original size-bins each into ten subsections of equal radius-range and number) over three size domains: Aitken mode (radii <0.05 µm), accumulation mode (radii: 0.05–0.5 µm) and coarse mode (radii >0.5 µm). Then for each size domain SS fluxes were distributed over a log-normal function, which is defined by the three parameters of mode-radius $r_m$, standard deviation $\sigma$ and number $N$. Assuming a SS density of 2.2 g/cm$^3$ and prescribing the distribution width (a standard deviation of 1.59 for Aitken and accumulation modes and 2.0 for the coarse mode) the number mode-radius $r_m$ was determined for each mode at each grid-point and time-step (since particle number $N$ and mass-flux were provided by the original data). Daily global fields for mode-radius and number (along with prescribed values for density and standard deviation) establish to recommended emission input for SS in modal size schemes. To accommodate aerosol modules with (size-)bin schemes, software is provided, which can extract the SS emission flux within any size-range. Distributions of sea-salt emissions are displayed in Fig. 2.

The characteristic SS is about 5 µm in diameter (a mode radius of 0.74 µm in conjunction with a standard deviation of 2 for the coarse mode translates into an effective “radius” of 2.5 µm). SS emissions are prescribed to take place in the lowest model layer. Uncertainty issues are similar to those for DU, including the numerical spread by the modal representation. For example, even though original SS fluxes of radii larger than 10 µm were rejected, the coarse size mode flux fit with a large distribution width (standard deviation of 2.0) will lead to non-negligible SS flux contributions for radii up
3.3 DMS

Daily DMS emission data are based on six hourly data of simulations with the LMDZ general circulation model LOA (Boucher et al., 2003) at 2.5°[latitude]×3.75°[longitude] resolution. Oceanic DMS emissions are derived by applying a parameterization for air-sea transfer velocities (Nightingale, 2000) to simulated climatology tied to measurement samples (Kettle and Andreae, 2000). Continental DMS emissions of biogenic origin (Pham et al., 1995) are generally much lower. To exclude unrealistic high contributions over coastal land regions (gridboxes with more than 5% land fraction), values of the nearest completely continental pixel were adopted. Then daily average data were interpolated onto a 1°×1° grid. Distributions of DMS emissions are given in Fig. 3.

3.4 Volcanic emissions

Annual volcanic (sulfur) emissions data, displayed in Fig. 4, consider both continuous degassing and explosive volcanos. Volcanic sulfur is emitted as 97.5% SO$_2$ and 2.5% SO$_4$. Data are based on the GEIA inventory (http://www.igac.noaa.gov/newsletter/22/sulfur.php; http://www.geiacenter.org/) (Andres and Kasgnoc, 1998). However, since there are a number of ambiguities which may lead to implementation differences, an interpreted and updated dataset is provided, which is summarized next.

The data on explosive emissions are based on observational evidence including the Aerosol Index (AI) of the TOMS satellite sensors. The multi-annual total emission (at 2 TgS/year) of explosive emissions is equally distributed over all grid boxes with volcanoes that had been active over the last 100 years (Halmer et al., 2002). The emissions are assumed to be continuous released, because only about 1/3 of explosive emissions are linked to violent explosive phases and placed between 500 and 1500 m above each volcano peak.

Continuous degassing is equally distributed over all grid points with GEIA volcano
locations. The annual continuous degassing emissions recommend by GEIA are considered to be an underestimate (Graf et al., 1998; Textor et al., 2004). Thus, continuous sulfur-containing emissions including the non-SO$_2$-species given in the GEIA inventory are multiplied by a factor of 1.5. The new estimate (at 12.6 TgS/year) is still conservative, since many volcanoes are either not monitored or have emissions that cannot be accurately detected by satellite sensors presently at use. Continuous degassing emissions are placed in the upper third of the volcano (altitude) to account for degassing at the flanks of the volcanoes.

3.5 Secondary Organic Aerosol

Monthly secondary Organic Aerosol (SOA) emissions are provided. They are based on the assumption that 15% of natural terpene emissions form SOA, although SOA production is much more complicated (Kanakidou et al., 2005). It is assumed that SOA is formed on time-scales of a few hours and that SOA precursor emissions condense on pre-existing aerosol. In reality, substantial SOA formation can occur at higher altitudes (Kanakidou et al., 2005). Terpene emissions of 127 Tg/year were taken from GEIA (Guenther et al., 1995). This translates into an annual global average of 19.1 Tg POM/year, which is within bounds of other estimates of 10 to 60 Tg POM/year (Kanakidou et al., 2005). The spatial distribution of SOA emission, depending on the vegetation type is given in Fig. 5.

4 Anthropogenically modified emissions – year 2000

Emissions which have changed (usually sharply increased) with industrialization are: contributions by sulfur – mainly in form of sulfur-dioxide (SO$_2$) – and carbon. Carbonaceous emissions are commonly stratified into organic carbon in terms of Particulate Organic Matter (POM) and strongly absorbing black carbon (BC). BC is also often referred to as elemental carbon, although strictly speaking elemental carbon is only
the light absorbing fraction of BC. Primary sources for anthropogenic emissions are large scale (wildland) fires, bio fuel burning and fossil fuel burning. For the latter a distinction is made between sources from road traffic, international shipping, off-road (rail, inland shipping, non-specified transport, and transport by pipelines), industry and power-plants, to account for differences in injection height and particle size, as indicated in Table 2. Table 2 summarizes individual contributions recommended for AeroCom Experiment “B” year 2000 simulations. Details on other emissions are given below.

4.1 Large-scale (wildland) fire emissions of BC, POM and SO₂

Monthly data for large-scale (wildland) fire emissions of carbon and sulfur are based on the Global Fire Emission Database (GFED) inventory (van der Werf et al., 2004, available online at http://www.ess.uci.edu/~jranders/). In this data-set, satellite derived fire hot spots from TRMM-VIRS and ERS-ATSR are calibrated to burnt area from the MODIS sensor for selected regions, and combined with the CASA biogeochemical model that was previously adjusted to account for fires to estimate fuel loads (van der Werf et al., 2003). The peak in the seasonal cycle of biomass burning emissions derived from fire hot spots has a tendency to precede the peak as derived from atmospheric measurements of CO, especially in the southern hemisphere (Petron et al., 2004). Annual totals are in overall agreement to other estimates (e.g. Generoso et al., 2003; Bond et al., 2004; Hoelzemann et al., 2004), and are further evaluated in Appendix D. Due to strong year-to-year variations, two data-sets are offered: An estimate specifically for the year 2000 and a 5-year (1997–2001) average for climatological simulations. Year 2000 emissions were below average because of wet La Niña conditions in the tropics (van der Werf et al., 2004). The annual emission patterns of BC, POM and SO₂ from large scale biomass burning (wildland fires) representing the year 2000 are displayed in Fig. 6. Monthly emissions into the atmosphere are distributed over six ecosystem-dependent altitude regimes between the surface and 6 km (see Sect. 7).
4.2 Biofuel emissions of BC, POM, and SO$_2$

Biofuel emissions include the burning of charcoal and dung and charcoal making. Also included in this category is crop residue burning. Yearly average data (no annual cycle) for biofuel organic emissions are based on the Speciated Particulate Emissions Wizard (SPEW) inventory for 1996 (Bond et al., 2004). It was assumed that this inventory applies without changes to the year 2000. Sulfur-dioxide biofuel emissions for the year 2000 are based on energy statistics for the year 2000 (Cofala et al., 2005). Country and regional estimates (see Appendix B) were gridded following EDGAR (Olivier et al., 2002) distribution patterns (Dentener et al., 2005) and projected from EDGAR estimates for the year 1995. Emission patterns of BC and POM from SPEW and EDGAR based sulfur-dioxide emissions from domestic and off-road sources are displayed in Fig. 7.

4.3 Fossil-fuel emissions of BC, POM, and SO$_2$

Yearly average data (no annual cycle) for fossil-fuel organic emissions are based on the Speciated Particulate Emissions Wizard (SPEW) inventory for 1996 (Bond et al., 2004). It was assumed that this inventory applies without changes to the year 2000. Sulfur-dioxide fossil-fuel emissions are based on energy statistics for the year 2000 using technology controlled emission factors from IIASA/RAINS (Cofala et al., 2005). The country and region estimates (see Appendix B) were gridded following EDGAR (Olivier et al., 2002) distribution patterns (Dentener, 2005). Ship traffic for the year 2000 is assumed to have increased by 1.5% per year since 1995 over the EDGAR3.2 values. Fossil fuel emission pattern for BC and POM and SO$_2$ emission patterns from road traffic, power-plants, ship-traffic and industry for the year 2000 are displayed in Fig. 8.
4.4 More emissions

For additional emission data (e.g. for full chemistry simulations) it is recommended to use the EDGAR 3.2, 1995 data-base (Olivier et al., 2002; http://www.mnp.nl/edgar). No specific recommendations are given for oxidant fields.

5 Anthropogenically modified emissions – year 1750

With the exception of fossil fuel emissions, which can be neglected, anthropogenic emissions in pre-industrial times were small but not zero. In particular contributions from wildland fires (open burning) and biofuel emissions in inhabited regions must be considered. In the absence of observational data, emission estimates for AeroCom Experiment “Pre” (year 1750) are derived from educated-guess assumptions. Recommendations for AeroCom Experiment “Pre” (year 1750) simulations are summarized in Table 3. More details on these emissions are provided below.

5.1 Large-scale (wildland) fire emissions of BC, POM and SO$_2$

Pre-industrial wildland fire emissions are based on scaled 5 year averages (1998–2002) of monthly data of the Global Fire Emission Database (GFED) inventory (van der Werf et al., 2003, 2004; Randerson et al., 2005). Central to a rescaling is the “year1750-to-year2000” population ratio from HYDE data-set (http://www.rivm.nl/hyde, see also Table C3 in Appendix C).

Actual scaling corrections are then performed according to present day land cover (Olsen et al., 1985). Emissions from deforestation fires are scale by population whereas emissions over all other land-surfaces (e.g. grassland, shrub/bush, agricultural activity) scale only to 60% by population (as it is assumed that 40% burns anyhow). Forest emissions in high latitudes of the northern hemisphere (Europe, N. America, Russia) are doubled from current estimates, to account for less fire suppression in
the past (Brenkert et al., 1997).

5.2 Biofuel emissions of BC, POM, and SO₂

Pre-industrial biofuel estimates are derived separately for carbonaceous aerosol and sulfur emissions. The BC and POM contributions are scaled back to the year 1750 based on statistics for population and crop production (where differences between developing and developed countries are considered). Pre-industrial carbon emissions are tied to the wood consumption, where the switch from electricity or natural gas as predominant cooking fuel back in time to wood was considered (Ito and Penner, 2005). For sulfur-dioxide a CO biofuel inventory for the year 1890 (van Aardenne et al., 2001) is multiplied by 0.00346, based on the ratio of emission factor estimates (Andreae and Merlet, 2001) for SO₂ and CO (0.27 g and 78 g per kg of burned dry biomass, respectively). A “year1750-to-year1890” population ratio from the HYDE data-set (see Appendix C) establishes the emissions for the year 1750. Emissions at high latitudes in the northern hemisphere (Europe, N. America and Russia) are doubled to account for a higher per person use (Brenkert et al., 1997).

6 Temporal resolution

The temporal resolution of the individual datasets is given in Tables 1 to 3. For simplification all anthropogenic enhanced emissions have no inter-annual variation annual – except for large scale wildland fires, where a monthly resolution is necessary to represent their (dry-) seasonal character. For all natural emissions that are tied to near surface winds (e.g. DU, SS, DMS) a daily temporal resolution was selected as a compromise between detail and size of the data-sets. The simulations on which daily emissions are based were done on a much higher temporal resolution.
7 Injection height

Recommended injection heights above the earth surface of the individual emission datasets are given in Tables 1 to 3. Most emissions are assumed to be evenly distributed in the lowest model-layer (“surface”). Fossil fuel emissions from industry and power-plants should be injected between 100 and 300 m above the surface, because these emissions are usually released at the top of chimneys. Large-scale wildland fire emissions are released distributed over six altitude regimes: 0–100 m, 100–500 m, 500–1 km, 1–2 km, 2–3 km, 3–6 km according to wildland fire location and type based on detailed work by D. Lavoue (2003, personal communication). Table 4 lists the altitude distribution by type.

Emissions are distributed evenly within each altitude layer. Contributions assigned to heights below the actual surface altitude are moved into the lowest applicable height range while contributions assigned to the 0–100 m altitude are always emitted in the lowest model layer. For illustration purposes also the maximum emission height for wildland fire emissions is indicated in Fig. 9.

The most complex altitude assignment is for volcanic emission. It is based on a data-set for volcanic location and volcano top altitude \(V_{\text{TOP}}\) (Halmer et al., 2002). For each volcano, (continuous) explosive contributions should be evenly placed between 500 and 1500 m above \(V_{\text{TOP}}\) and continuous degassing should occur in the upper 1/3 altitude of each volcano.

8 Size choices for primary particulate emissions

Recommended aerosol sizes for particles of the individual emission datasets are given in Tables 1 to 3. Size recommendations are given in terms of log-normal distributions parameters, where the mode radius \(r_m\) describes the peak concentration and the standard deviation \(\sigma\) describes the distribution width. From both log-normal values a (radiatively) characteristic size has been determined \(r_{\text{eff}}\), as the ratio between the
sums of the third and the second moment of the radius: $\Sigma r^3/\Sigma r^2$.

For wildland fire (open burning) and biofuel aerosol the recommended characteristic size of $r_{\text{eff}} \sim 0.1 \, \mu m$ is based on an analysis of numerous field-measurements (see Appendix C). For fossil-fuel two different size-distributions are defined. A large size mode with $r_{\text{eff}} \sim 1.6 \, \mu m$ is recommended for power-plant and industrial emissions (representing fly-ash, and components formed on it). A relatively small size mode with $r_{\text{eff}} \sim 0.04 \, \mu m$ is recommended for other fossil fuel emissions (e.g. traffic) based on kerbside measurements in several EU-cities (Putaud et al., 2004; van Dingenen et al., 2004). For particles from volcanic emissions, half of the mass is assigned each to the small fossil-fuel size ($r_{\text{eff}} \sim 0.04 \, \mu m$) and to the biofuel size ($r_{\text{eff}} \sim 0.1 \, \mu m$). For dust and sea-salt, size recommendations are more complex, because they are defined by two and three size-modes, respectively, with variations permitted between consecutive days. However, since the mass flux is dominated by contributions of the coarse size domain, the average characteristic size is well represented by the coarse mode with $r_{\text{eff}} \sim 2.1 \, \mu m$ for dust and at $r_{\text{eff}} \sim 2.5 \, \mu m$ for sea-salt.

9 Discussion and conclusion

The above emissions, recommended for AeroCom, represent the state-of-the-art for global aerosol emissions inventories in the year 2004. In cases where several alternative datasets were available, such as three large-scale burning wildland fire inventories (van der Werf et al., 2004; Generoso et al., 2003; Hoelzemann et al., 2004; see Appendix D), selection criteria included global annual coverage, lack of biases and emission-fluxes within expected bounds. The overall goal was to provide global emission fields and recommendations for seasonality, emission height, and size distribution of all aerosol sources that are simple enough to be easily implemented in global models.

The data-sets explained in this paper are considered an initial first step to establish a global and temporal sufficiently comprehensive aerosol emission database. This
work did not attempt to assess uncertainties; nevertheless it should be recognized
that there is more confidence in some sources (like SO$_2$ from anthropogenic sources)
than in other sources (e.g. emissions of EC and POM) where there are deviations of
a factor of 3 between estimates found in the literature. There is certainly potential for
improvement, such as (1) using identical meteorological fields, when deriving emission
for sea-salt, dust and DMS, (2) improved temporal resolution for all aerosol emissions
(e.g. introducing seasonality to all emissions) or (3) consistencies to trace-gas emis-
sions that are relevant to atmospheric chemistry. In addition, an extension to include
emission-estimates associated with future scenarios is planned.

Appendix A

Names and location of data-files on IES, JRC ftp-web site

All AeroCom emission data-sets are available (in netcdf data format) via a dedicated file
transfer site at the Joint Research Center in Ispra, Italy: ftp://ftp.ei.jrc.it/pub/Aerocom/.
To help locate specific data ftp-site subdirectories and their content is outlined below:

/dust_ncf dust data “dust200001.nc” (January 2000), . . . , “dust200012.nc” (De-
cember 2000)

/seasalt_ncf salt data “salt200001.nc” (January 2000), . . . , “salt200012.nc” (De-
cember 2000)

/DMS_ncf DMS data “dms200001.nc” (January 2000), . . . , “dms200012.nc” (De-
cember 2000)

/volcanic_ncf volcanic data
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– degassing “contineous_volc.nc”
– explosive “explosive_volc.nc”

/other_ncf_2000 all other emissions for the year 2000
– BC-biofuel “BC1bfuel.nc”
– BC-fossil fuel “BC1ff.nc”
– BC-wildland fire “GFED_2000_BC.nc”
“GFED_average_1997–2002_BC.nc”
– POM-wildland fire “GFED_2000_POM.nc”
“GFED_average_1997–2002_POM.nc”
– SO2-wildland fire “GFED_2000_SO2.nc”
“GFED_average_1997–2002_SO2.nc”
– POM-biofuel “POMbfuel.nc”
– POM-fossil fuel “POMff.nc”
– SO2-domestic “SO2_Domestic_2000bau.nc”
– SO2-industry “SO2_Industry_2000bau.nc”
– SO2-ships “SO2_International_Shipping_2000bau.nc”
– SO2-off road “SO2_Off-road_2000bau.nc”
– SO2-powerplant “SO2_Powerplants_2000bau.nc”
– SO2-traffic “SO2_RoadTransport_2000bau.nc”

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Appendix B

Regional aspects of AeroCom emissions

For a quick reference on regional contributions of BC, POM and SO$_2$, annual AeroCom emissions are stratified into 18 regions according to the IMAGE integrated assessment model, as illustrated in Fig. B1 (courtesy of B. Eickhout at RIVM, Netherlands).

Regional AeroCom emissions by species and source are summarized for current conditions (year 2000) in Table B1. Figures B2, B3 and B4 display the global distribution of annual emission totals for BC, POM and SO$_2$ for the year 2000, combining all sources.

Corresponding to Table B1 regional distributions to for pre-industrial conditions (year 1750) are given in Table B2. Figures B5, B6 and B7 display the global distribution of annual emission totals for BC, POM and SO$_2$ for the year 2000, combining all sources.
Differences of emissions between years 2000 and 1750 provide estimates on anthropogenic contributions.

A central role, when back-scaling current emission to obtain estimates on pre-industrial emissions has been the population ratio of the HYDE data-set (http://www.rivm.nl/hyde). For the 18 IMAGE regions (see Fig. B1) these population ratios are listed in Table B3.

Appendix C

Background to the sizing of primary aerosol from biomass burning

Biomass burning is one of the main sources for carbonaceous aerosol in the atmosphere. Globally it contributes to about 40% of CO$_2$, 32% of CO, 38% of tropospheric ozone, 7% of total particulate matter and 39% of particulate organic carbon. The majority of biomass aerosol (ca. 80%) occurs in the tropics as seasonal event (e.g. August–October: S. Africa and S. America).

To provide background on choices for the sizing of freshly emitted (young) biomass aerosol. We compiled a summary of measurements (Allen and Miguel, 1995; Anderson et al., 1996; Andreae and Merlot, 2001; Cachnier et al., 1996, LeCanut et al., 1996, Radke et al., 1991; Scholes et al., 1996, Suscott et al., 1991). Measured size distributions have been fitted to a multi-model lognormal distribution, which is defined by

$$\frac{dN}{d \log r} = \frac{N \ln 10}{\sqrt{2\pi} \ln \sigma} \cdot \exp \left[ - \frac{(\ln r - \ln r_m)^2}{2(\ln \sigma)^2} \right]$$

where $N$ is total particle number, $r_m$ is the mode radius and $\sigma$ is the standard deviation. A comparison of fits is presented in Fig. C1.

The accumulation size-mode (radii smaller than 0.5 $\mu$m) usually contains more then 90% of the biomass burning aerosol mass. Figure C2 presents the log-normal param-
eters data-pairs for $r_m$ (actually the diameter is shown) and $\sigma$ of only the accumulation size-mode associated with the size distributions of Fig. C1.

Data points in the lower right of Fig. C2 indicate log-normal parameters from fits to young biomass aerosol. As the biomass aerosol ages, the mode radius increases and the distribution width narrows, resulting in an increase to effective radius (from 0.10 to 0.15 $\mu$m), which is small in the context of orders of magnitudes among different aerosol sizes. As biomass burning emissions are associated with young aerosol, recommendations are 0.04 $\mu$m for (number-) mode radius $r_m$ and 1.8 for standard deviation $\sigma$ (as listed in Tables 1 to 3). These values are consistent with estimates of volatility analysis on young aerosol emitted by North American fires (Reid et al., 2003), where sampled aerosol ($r_m=0.55 \mu m$, $\sigma=1.7$) has been pre-heated to 310°C, to bake off volatile compounds (organics, sulfate, nitrate, water etc.), resulting in core aerosol (mainly BC and low volatility organics) of $r_m=0.32 \mu m$ and $\sigma=1.9$.

Appendix D

Comparison of available wildland fire inventories

By 2003, three different global large-scale wildland fire inventories had been developed (van der Werf et al., 2004; Generoso et al., 2003; Hoelzemann et al., 2004). When comparing the inventories on a regional basis, as for POM in Fig. D1, there is good agreement in a relative sense, although estimates by Hoelzemann (GWEM 1.2) are smaller, especially in South America. This lower estimate is erroneous and associated with a failure in the underlying area burnt satellite product, so that the impact of small but intense fires is underestimated in that region. A more recent version of GWEM (version 1.4) is now more consistent with the other inventories. For more discussion on the uncertainty of the POM/EC inventories, the reader is referred to the description of SPEW (Bond et al., 2004). A particular problem is the uncertainty of emission factors, which for BC can exceed an order of magnitude (C. Liouesse, private communication,
Thus, the certainty for carbonaceous emissions can not be expected to be better than a factor of 2.

Acknowledgements. AeroCom was sponsored by the European Union FP5 project “Phoenics” EVK2-CT-2001-00098.

References


Halmer, M., Schmincke, H., and Graf, H.: The annual volcanic gas input into the atmosphere, in


Table 1. AeroCom common (natural) emissions.

<table>
<thead>
<tr>
<th>Type</th>
<th>Time</th>
<th>Injection Type</th>
<th>Altitude</th>
<th>ln $r_m$ $[\mu m]$</th>
<th>ln $\sigma$ $[\mu m]$</th>
<th>$r_{eff}$ $[\mu m]$</th>
<th>Flux $[Tg/yr]$</th>
<th>Flux $[Tg/yr]$</th>
</tr>
</thead>
<tbody>
<tr>
<td>Dust</td>
<td>Daily</td>
<td>DU</td>
<td>Surface</td>
<td>.650$^C$</td>
<td>2.0$^C$</td>
<td>2.10</td>
<td>1678</td>
<td>2150 $\pm/50%$</td>
</tr>
<tr>
<td>Sea-salt</td>
<td>Daily</td>
<td>SS</td>
<td>Surface</td>
<td>.740$^C$</td>
<td>2.0$^C$</td>
<td>2.50</td>
<td>7925</td>
<td>3340 $\pm/80%$</td>
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<tr>
<td>DMS</td>
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<td>S</td>
<td>Surface</td>
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<td>18.2</td>
<td>25 $\pm/60%$</td>
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<tr>
<td>Volcanic</td>
<td>Yearly</td>
<td>S</td>
<td>(V$_T$+500 m)–(V$_T$+1500 m)</td>
<td>.040 $^{50%}$</td>
<td>1.8</td>
<td>.059</td>
<td>2.0</td>
<td></td>
</tr>
<tr>
<td>Explosive</td>
<td>Yearly</td>
<td>S</td>
<td>(0.67 *V$_T$)–(1.0 *V$_T$)</td>
<td>.040 $^{50%}$</td>
<td>1.8</td>
<td>.059</td>
<td>12.6</td>
<td>9.3 $[4–20]$</td>
</tr>
<tr>
<td>SOA</td>
<td>Monthly</td>
<td>POM</td>
<td>Surface</td>
<td>.015 $^{50%}$</td>
<td></td>
<td></td>
<td>19.1</td>
<td>[12–70]$^K$</td>
</tr>
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</table>

* DU-dust, SS-sea-salt, S-sulfur (=0.50×SO$_2$, or 0.33×SO$_4$), POM-particulate org matter (=1.40×organic carbon)

+ only 2.5% of sulfur (S) should be emitted as particulate SO$_4$, most sulfur (S) is emitted as gaseous SO$_4$

$^C$ log-normal size-distribution parameters of the coarse size mode

$^K$ based on a recent review by Kanakidou et al. (2005)
Table 2. AeroCom anthropogenically (full molecular mass) emissions for the year 2000.

<table>
<thead>
<tr>
<th>type</th>
<th>data source</th>
<th>time resolution</th>
<th>aero type</th>
<th>injection altitude</th>
<th>size (ln)</th>
<th>size $\sigma$</th>
<th>$r_{eff}$ [µm]</th>
<th>flux [Tg/yr]</th>
<th>AeroCom</th>
<th>IPCC-TAR</th>
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<td>wild-fire</td>
<td>GFED</td>
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<td>BC</td>
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<td>.095</td>
<td>3.1</td>
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<tr>
<td></td>
<td>GFED</td>
<td>monthly</td>
<td>POM</td>
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<td>.095</td>
<td>34.7</td>
<td>54</td>
<td>[45–80]</td>
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<td>GFED</td>
<td>monthly</td>
<td>S $^+$</td>
<td>6 layers $^H$</td>
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<td>1.8</td>
<td>.095</td>
<td>2.1</td>
<td>2.2</td>
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<td>BC</td>
<td>surface</td>
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<td>1.8</td>
<td>.095</td>
<td>1.6</td>
<td>in wild fire</td>
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<tr>
<td></td>
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<td>yearly</td>
<td>POM</td>
<td>surface</td>
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<td>1.8</td>
<td>.095</td>
<td>9.1</td>
<td></td>
<td></td>
</tr>
<tr>
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<td>IIASA</td>
<td>yearly</td>
<td>S $^+$</td>
<td>surface</td>
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<td>.036</td>
<td>9.6</td>
<td>11.4$^B$</td>
<td></td>
</tr>
<tr>
<td>fossilfuel</td>
<td>SPEW</td>
<td>yearly</td>
<td>BC</td>
<td>surface</td>
<td>.015</td>
<td>1.8</td>
<td>.036</td>
<td>3.0</td>
<td>6.6</td>
<td>[6–8]</td>
</tr>
<tr>
<td>roads</td>
<td>IIASA</td>
<td>yearly</td>
<td>S $^+$</td>
<td>surface</td>
<td>.015</td>
<td>1.8</td>
<td>.036</td>
<td>3.2 (+19.1)</td>
<td>28</td>
<td>[10–30]$^A$</td>
</tr>
<tr>
<td>shipping</td>
<td>EDGAR</td>
<td>yearly</td>
<td>S $^+$</td>
<td>surface</td>
<td>.500</td>
<td>2.0</td>
<td>1.66</td>
<td>7.8</td>
<td>7.3$^B$</td>
<td></td>
</tr>
<tr>
<td>off-road</td>
<td>IIASA</td>
<td>yearly</td>
<td>S $^+$</td>
<td>surface</td>
<td>.015</td>
<td>1.8</td>
<td>.036</td>
<td>1.6</td>
<td>1.9$^B$</td>
<td></td>
</tr>
<tr>
<td>industry</td>
<td>IIASA</td>
<td>yearly</td>
<td>S $^+$</td>
<td>100–300 m</td>
<td>.500</td>
<td>2.0</td>
<td>1.66</td>
<td>39.2</td>
<td>67.5$^B$</td>
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</tr>
<tr>
<td>power-pl.</td>
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<td>yearly</td>
<td>S $^+$</td>
<td>100–300 m</td>
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<td>2.0</td>
<td>1.66</td>
<td>48.4</td>
<td>53.6$^B$</td>
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</tbody>
</table>

* S-sulfur ($=0.50 \times$SO$_4$, or $0.33 \times$SO$_2$), POM-particulate org matter ($=1.40 \times$organic carbon), BC-black carbon
+ 2.5% of sulfur should be emitted as particulate SO$_2$, most sulfur (S) is emitted as gaseous SO$_2$
$^H$ 0–100 m, 100–500 m, 500–1000 m, 1–2 km, 2–3 km, 3–6 km, assignment according to Table 4
$^A$ Cooke et al. (1999) report a more moderate amount of 10.1 Tg OC-C/yr.
$^B$ based on EDGAR3.2 FT2000 (http://www.rivm.nl/edgar) and Olivier et al. (2005)
Table 3. AeroCom anthropogenically modified (full molecular mass) emissions for the year 1750.

<table>
<thead>
<tr>
<th>Time Resolution</th>
<th>Aerosol Type</th>
<th>Injection Altitude</th>
<th>Size ln ( r_m ) [( \mu m )]</th>
<th>Size ln std.dev</th>
<th>( r_{eff} ) [( \mu m )]</th>
<th>Amount [Tg/yr]</th>
</tr>
</thead>
<tbody>
<tr>
<td>Wildland Fire</td>
<td>Monthly BC</td>
<td>6 layers (^H)</td>
<td>.040</td>
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<td>.095</td>
<td>1.02</td>
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<td></td>
<td>Monthly POM</td>
<td>6 layers (^H)</td>
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<td>1.8</td>
<td>.095</td>
<td>12.8</td>
</tr>
<tr>
<td></td>
<td>Monthly S(^+)</td>
<td>6 layers (^H)</td>
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<td>1.8</td>
<td>.095</td>
<td>0.7</td>
</tr>
<tr>
<td>Biofuel</td>
<td>Yearly BC</td>
<td>Surface</td>
<td>.040</td>
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<td>.095</td>
<td>0.26</td>
</tr>
<tr>
<td></td>
<td>Yearly POM</td>
<td>Surface</td>
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<td>1.8</td>
<td>.095</td>
<td>2.5</td>
</tr>
<tr>
<td></td>
<td>Yearly S(^+)</td>
<td>Surface</td>
<td>.040</td>
<td>1.8</td>
<td>0.95</td>
<td>0.06</td>
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</table>

* S-sulfur (=0.50\( \times \)SO\(_4\), or 0.33\( \times \)SO\(_2\)), POM-part.org matter (=1.40\( \times \)organic carbon), BC-black carbon

\(^+\) 2.5% of sulfur should be emitted as particulate SO\(_2\), most sulfur (S) is emitted as gaseous SO\(_2\)

\(^H\) 0–100 m, 100–500 m, 500–1000 m, 1–2 km, 2–3 km, 3–6 km, assignment according to Table 4
Table 4. Fractional distribution (in %) of emission heights for wild-land fires.

<table>
<thead>
<tr>
<th></th>
<th>0–100 m *</th>
<th>100–500 m</th>
<th>500–1000 m</th>
<th>1000–2000 m</th>
<th>2000–3000 m</th>
<th>3000–6000 m</th>
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<tr>
<td>agricultural waste</td>
<td>100</td>
<td>–</td>
<td>–</td>
<td>–</td>
<td>–</td>
<td>–</td>
</tr>
<tr>
<td>tropical (30 S–30 N)</td>
<td>20</td>
<td>40</td>
<td>40</td>
<td>–</td>
<td>–</td>
<td>–</td>
</tr>
<tr>
<td>Temperate (30 N–60 N, 30 S–60 S)</td>
<td>20</td>
<td>20</td>
<td>20</td>
<td>40</td>
<td>–</td>
<td>–</td>
</tr>
<tr>
<td>Boreal (Eurasia)</td>
<td>10</td>
<td>10</td>
<td>20</td>
<td>20</td>
<td>40</td>
<td>–</td>
</tr>
<tr>
<td>Boreal (Canada)</td>
<td>10</td>
<td>10</td>
<td>10</td>
<td>10</td>
<td>20</td>
<td>40</td>
</tr>
</tbody>
</table>

* contributions assigned to heights below the actual altitude are moved into the lowest applicable height range and contribution of the 0–100 m altitude are always emitted in the lowest modeling layer.
Table B1. Regional distributions of AeroCom emissions for the year 2000.

<table>
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<tr>
<th>category</th>
<th>BC</th>
<th>BC</th>
<th>BC</th>
<th>OM</th>
<th>OM</th>
<th>OM</th>
<th>OM</th>
<th>OM</th>
<th>OM</th>
<th>SO₂</th>
<th>SO₂</th>
<th>SO₂</th>
<th>SO₂</th>
<th>SO₂</th>
<th>SO₂</th>
<th>SO₂</th>
<th>SO₂</th>
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</thead>
<tbody>
<tr>
<td></td>
<td>wildf</td>
<td>biof</td>
<td>fos.f</td>
<td>all</td>
<td>wildf</td>
<td>biof</td>
<td>fos.f</td>
<td>all</td>
<td>wildf</td>
<td>v.ex</td>
<td>v.co</td>
<td>dom</td>
<td>road</td>
<td>pow</td>
<td>off-r</td>
<td>ship</td>
<td>indu</td>
<td></td>
</tr>
<tr>
<td>CANADA</td>
<td>0.01</td>
<td>0.01</td>
<td>0.03</td>
<td>0.05</td>
<td>0.18</td>
<td>0.05</td>
<td>0.02</td>
<td>0.69</td>
<td>0.95</td>
<td>0.02</td>
<td>0.00</td>
<td>0.00</td>
<td>0.01</td>
<td>0.54</td>
<td>0.05</td>
<td>0.00</td>
<td>1.19</td>
<td>1.88</td>
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<td>0.07</td>
<td>0.06</td>
<td>0.27</td>
<td>0.40</td>
<td>1.11</td>
<td>0.45</td>
<td>0.2</td>
<td>1.23</td>
<td>2.99</td>
<td>0.11</td>
<td>0.22</td>
<td>1.01</td>
<td>0.31</td>
<td>0.17</td>
<td>12.4</td>
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<td>0.11</td>
<td>3.12</td>
</tr>
<tr>
<td>C AMERICA</td>
<td>0.13</td>
<td>0.03</td>
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<td>0.22</td>
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<td>0.62</td>
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<td>0.23</td>
<td>2.24</td>
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<tr>
<td>S AMERICA</td>
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<td>0.22</td>
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<td>0.24</td>
<td>6.58</td>
<td>15.6</td>
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<td>0.47</td>
<td>4.76</td>
<td>0.16</td>
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<td>0.54</td>
<td>0.12</td>
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<td>1.61</td>
</tr>
<tr>
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<td>0</td>
<td>0.05</td>
<td>0.05</td>
<td>0.00</td>
<td>0.03</td>
<td>0.06</td>
<td>0.03</td>
<td>0.12</td>
<td>0.00</td>
<td>0.00</td>
<td>0.00</td>
<td>0.00</td>
<td>0.05</td>
<td>0.07</td>
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<td>0.68</td>
<td>0.01</td>
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<td>0.97</td>
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<td>0.00</td>
<td>0.44</td>
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<td>3.47</td>
<td>0.19</td>
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</tr>
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<td>0.14</td>
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<td>0.01</td>
<td>0.00</td>
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<td>0.04</td>
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<td>0.17</td>
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<td>0.06</td>
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Table B2. Regional distributions of AeroCom emissions for the year 1750.

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### Table B3. Population by region for the years 1750 and 2000 (by HYDE) – and ratios.

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Fig. 1. Global fields of annual dust emission fluxes and dust emissions for four selected months (January, April, July and October). Dust emissions are given in kg/m² according to the linear color scale and its minimum and maximum values.
Fig. 2. Global fields of annual sea-salt emission fluxes and sea-salt emissions for four selected months (January, April, July and October). Sea-salt emissions are given in kg/m² according to the linear color scale and its minimum and maximum values.
Fig. 3. Global fields of annual DMS flux emissions and DMS emissions for four selected months (January, April, July and October) DMS emissions are given in kg/m² according to the linear color scale and its minimum and maximum values.
Fig. 4. Annual distribution of volcanic explosive (upper panel) and volcanic continuous SO$_2$ emissions (lower panel).
Fig. 4. Continued.
Fig. 5. Annual SOA emissions (tied to natural terpene emissions) for the year 2000.
Fig. 6. Annual wildland fire emissions of BC (upper panel), POM (center panel) and SO$_2$ (lower panel) and for year 2000.
Fig. 6. Continued.
Fig. 6. Continued.
Fig. 7. Annual biofuel emissions of BC (first panel), POM (second panel) and SO$_2$ from domestic (third panel) and off-road activity (fourth panel) for the year 2000.
Fig. 7. Continued.
Fig. 7. Continued.
Fig. 7. Continued.
Fig. 8. Annual fossil-fuel emissions of BC (first panel), POM (second panel) and SO$_2$ from road transport (third panel), powerplants (fourth panel), shipping (fifth panel) and industrial activity (sixth panel) for the year 2000.
Fig. 8. Continued.
Fig. 8. Continued.
Fig. 8. Continued.
Fig. 8. Continued.
Fig. 8. Continued.
**Fig. 9.** Maximum emission height (in meter) for (large-scale) wild-fire aerosol.
Fig. B1. Regional choices for continental regions by the IMAGE project (by B. Eickhout).
Fig. B2. Global distribution of all BC emissions for the year 2000.
Fig. B3. Global distribution of all POM emissions for the year 2000.
Fig. B4. Global distribution of all SO$_2$ emissions for the year 2000.
Fig. B5. Global distribution of all BC emissions for the year 1750.
Fig. B6. Global distribution of all POM emissions for the year 1750.
Fig. B7. Global distribution of all SO₂ emissions for the year 1750.
Background to the sizing of primary aerosol from biomass burning

Biomass burning is one of the main sources for carbonaceous aerosol in the atmosphere. Globally it contributes to about 40% of CO$_2$, 32% of CO, 38% of tropospheric ozone, 7% of total particulate matter and 39% of particulate organic carbon. The majority of biomass aerosol (ca 80%) occurs in the tropics as seasonal event (e.g. Aug-Oct: S. Africa and S. America).

To provide background on choices for the sizing of freshly emitted (young) biomass aerosol, Luisa Marelli (JRC) has compiled a summary of measurements (Allen and Miguel, 1995; Anderson et al., 1996; Andreae and Merlot, 2001; Cachnier et al., 1996, LeCanut et al., 1996, Radke et al., 1991; Scholes et al., 1996, Suscott et al., 1991).

Measured size distributions have been fitted to a multi-model lognormal distribution, which is defined by

$$N \propto \exp \left[ -\frac{1}{2} \left( \frac{r - \mu}{\sigma} \right)^2 \right]$$

where $N$ is total particle number, $r_m$ is the mode radius and $\sigma$ is the standard deviation. A comparison of fits is presented in Figure C1.

**Fig. C1.** Multi-modal log-normal fits to measured biomass size samples.
The accumulation size-mode (radii smaller than 0.5 µm) usually contains more than 90% of the biomass burning aerosol mass. Figure C2 presents the log-normal parameters data-pairs for \( r_m \) (actually the diameter is shown) and \( \sigma \) of only the accumulation size-mode associated with the size distributions of Figure A1.

**Fig. C2.** Relationship between (number-) mode diameter (in nm) and standard deviation (describing the distribution width) from log-normal fits to the size-distributions of Fig. A1. Data points in the lower right refer to young biomass aerosol. Data points towards the upper left refer to aged biomass aerosol indicating an aging trend to larger mode-radii and narrower distribution width.

The equation for the line of best fit is:

\[
y = -426.51x + 837.27
\]

\[ R^2 = 0.882 \]
By 2003, three different global large-scale wildland fire inventories had been developed (van der Werf et al, 2004; Generoso et al, 2003; Hoelzemann et al, 2004). When comparing the inventories on a regional basis, as for POM in Figure D1, there is good agreement in a relative sense, although estimates by Hoelzemann (GWEM 1.2) are smaller, especially in South America. This lower estimate is erroneous and associated with a failure in the underlying area burnt satellite product, so that the impact of small but intense fires is underestimated in that region. A more recent version of GWEM (version 1.4) is now more consistent with the other inventories. For more discussion on the uncertainty of the POM/EC inventories, the reader is referred to the description of SPEW (Bond et al 2004). A particular problem is the uncertainty of emission factors, which for BC can exceed an order of magnitude (C. Liousse, private communication). Thus, the certainty for carbonaceous emissions can not be expected to be better than a factor of 2.

Figure D1. Comparison of three different global large-scale burning (wildland fire) inventories for regional POM emission estimates by Generoso et al. (2003), Hoelzemann et al. (2004) and van der Werf et al. (2004).