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**Global terrestrial  
COS sink strength**

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# Global uptake of carbonyl sulfide (COS) by terrestrial vegetation: Estimates corrected by deposition velocities normalized to the uptake of carbon dioxide (CO<sub>2</sub>)

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## Abstract

COS uptake by trees, as observed under dark/light changes and under application of the plant hormone abscisic acid, exhibited a strong correlation with the CO<sub>2</sub> assimilation rate and the stomatal conductance. As the uptake of COS occurred exclusively through the stomata we compared experimentally derived and re-evaluated deposition velocities ( $V_d$ ) for COS and CO<sub>2</sub>. We show that  $V_d$  of COS is generally significantly larger than that of CO<sub>2</sub>. We therefore introduced this attribute into a new global estimate of COS fluxes into vegetation. The global COS uptake by vegetation as estimated by the new model ranges between 0.69–1.40 Tg a<sup>-1</sup>, based on the Net Primary Productivity (NPP). Taking into account Gross Primary Productivity (GPP) the deposition estimate ranges between 1.37–2.81 Tg a<sup>-1</sup> (0.73–1.50 Tg S a<sup>-1</sup>). We believe that in order to obtain accurate and reliable global NPP-based estimates for the COS flux into vegetation, the different deposition velocities of COS and CO<sub>2</sub> must be taken into account.

## 1. Introduction

Carbonyl sulfide (COS) is a substantial source for stratospheric sulfate aerosol and plays an important role in stratospheric ozone chemistry (Crutzen, 1976; Andreae and Crutzen, 1997). According to Watts (2000) and Kettle et al. (2002) total global sources and sinks are balanced within the uncertainties of the estimates. Deposition to vegetation and soils represents the main sink for this trace gas (Logan et al., 1979; Brown and Bell, 1986; Chin and Davis, 1993, 1995). Soils have been recognized as a global sink for COS only recently and the uncertainty is rather large as parameterization of the uptake has been performed with only one soil type to date (Kesselmeier et al., 1999), which clearly warrants further studies. The role of vegetation as a major global tropospheric sink for COS has been studied for 20 years and is undisputed, but the uncertainty in the quantitative estimates of this sink is still large. Two common methods for

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the estimate of the global COS sink strength were reported. The estimate by Brown and Bell (1986) is based on the deposition velocity of COS and its atmospheric concentration. Another approach is described by Chin and Davis (1993) who used the correlation between the COS deposition and the CO<sub>2</sub> assimilation, assuming the same deposition velocities for CO<sub>2</sub> and COS. Recent estimates refer to this method (Kesselmeier et al., 1993; Andreae and Crutzen, 1997; Watts, 2000; Kettle et al., 2002). However, a simple 1:1 relation for the uptake ratio of COS/CO<sub>2</sub> appears insufficient as a preferential uptake of COS on a leaf as well enzymatic basis has been reported (Kesselmeier and Merk, 1993; Protoschill-Krebs et al., 1996). Therefore, we studied the stomatal uptake of COS separately. Moreover we investigated the close correlation between the rate of photosynthesis and the COS uptake for several European tree species and considered the differences in deposition velocities for CO<sub>2</sub> and COS. The observed ratios of the deposition velocities were used to improve global estimates of the COS vegetation sink based on Net Primary Production.

## 2. Materials and methods

### 2.1. Plant material and growth

For all experiments young trees (3–4 years old) from German and English tree nurseries were used. Tree species investigated were Holm oak (*Quercus ilex* L.), European beech (*Fagus sylvatica* L.), Norwegian spruce (*Picea abies*) and Scots pine (*Pinus sylvestris*). The trees were kept in 20 l pots with gardener's compost for the years 1997–2000 and fertilized with commercially available fertilizer (Baumfit, Spiess-Urania, Germany). For the Mediterranean tree species *Quercus ilex* sand was mixed with the soil in a 1 to 2 ratio. The trees were grown in a greenhouse at 25°C under a 12/12 h light-dark regime with a light intensity of 600  $\mu\text{mol m}^{-2} \text{s}^{-1}$  of photons (PAR) and a relative humidity of 70% under 350 ppm CO<sub>2</sub>.

## 2.2. Purification of ambient air

Compressed air was purified by passing it ( $6 \text{ l min}^{-1}$ ) through multistage gas purification system consisting of (1) silica gel (Merck, Germany), (2) molecular sieve (0.5 nm, Merck, Darmstadt), (3) charcoal (Merck, Germany), and soda lime (Merck, Germany), 3 l each. COS and  $\text{CO}_2$  mixing ratios were adjusted to desired values by mixing the purified compressed air with known gas mixtures produced from a permeation device (Haunold, Germany) with COS permeation tubes (VICI Metronics, Santa Clara, California) and  $\text{CO}_2$  from a pressurized bottle (Messer-Griesheim, Germany). All flows were regulated by mass flow controllers (MKS, Massachusetts, USA).

## 2.3. Enclosure system (cuvettes) and exchange measurements

Gas exchange of enclosed tree branches was investigated using a dynamic (flow-through) Teflon-film-cuvette system consisting of a plant measuring and an empty reference cuvette with all inner surfaces made of Teflon to avoid interference with the investigated trace gases. All experiments were performed in a climate chamber with identical conditions as compared to the growth chamber. Trace gas sampling was accompanied by measurements of ambient  $\text{CO}_2$ ,  $\text{CO}_2$  exchange and transpiration by an infra-red gas analyzer. For details see Kuhn et al. (1999, 2000) and Kuhn and Kesselmeier (2000). The enclosures were constantly flushed with  $1 \text{ l min}^{-1}$  of purified and conditioned ambient air (see above) which was artificially moistened (r.H. >70%) before entering the cuvettes. COS was quantified in the ppt range by an automated analytical system according to Von Hobe et al. (2000) by consecutive sampling at both cuvettes. All trace gas exchange rates were calculated by concentration differences between the branch enclosing cuvette and the empty reference. Accuracy and precision of the analytical system were better than 2% plus any uncertainties introduced by the cuvette sampling, mainly by the accuracy of mass flow controllers. The stomatal conductance for water vapor was determined according to Pearcy et al. (1989).

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## 2.4. Induction of stomatal closure

Stomatal closure was induced by infiltration of abscisic acid (ABA) to an oak branch cut from the tree two days before this application. The ABA treatment was performed by cutting a small branch of *Quercus ilex* under water (to prevent air from penetrating into the water-conducting elements) and dipping it into a nutrient solution of 1.0 mM KCl, 0.1 mM NaCl, 0.1 mM CaCl<sub>2</sub>. Measurements of CO<sub>2</sub>-exchange and water vapor transpiration proved the unrestricted viability of this sample branch. ABA was applied by exchanging the nutrient solution against a fresh one containing additional 0.1 mM ABA, which was transported into the leaves by the transpiration stream. For details see Gabriel et al. (1999).

## 3. Results

### 3.1. Stomatal uptake

The automated analytical COS analyzer allowed a dense protocol to follow the exchange behavior with high time resolution. Figure 1 shows the COS uptake by an enclosed branch of *Fagus sylvatica* (European beech) in relation to assimilation and stomatal conductance over nearly three days of measurements. The COS-uptake closely followed the light/dark cycle. This behavior can be related to the calculated stomatal aperture and consequently directly compared to the exchange of CO<sub>2</sub>, i.e. net photosynthetic assimilation rate. Low uptake rates were found under dark conditions. As the stomata did not completely close in the dark and a respiration activity was detectable, the low COS exchange may be understood as a physiological consumption at a lower rate. On the other hand we do not exclude fluctuations caused by the non-simultaneous COS sampling at the sample and reference cuvette due to switching of the automated system from one cuvette to the other. Nevertheless, a clear relation to light and stomatal aperture is obvious, though the enzymatic pathway of COS con-

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sumption by carbonic anhydrase itself is generally light independent (Protoschill-Krebs et al., 1996).

The close relation between light and COS uptake in contrast to the light independent consumption by the enzyme carbonic anhydrase supports the assumption of an exclusively stomatal uptake pathway, as light represents a considerable trigger of stomatal movement. The final proof for a stomatal related exchange was demonstrated by the application of abscisic acid (ABA), a plant hormone which causes stomatal closure. Treatment with ABA was performed by cutting a small branch of *Quercus ilex* and incubating the stem into a vial filled with a buffer solution. Under these conditions, the branch was performing a normal behavior of gas exchange and COS uptake (Fig. 2) as followed for two light/dark episodes. In the course of the third light phase we infiltrated ABA and observed a fast decline of CO<sub>2</sub> exchange down to zero (no respiration measurable) closely accompanied by the decrease of COS uptake (Fig. 2). Stomatal conductance also showed a decrease to night values under the influence of ABA under light conditions, though with some delay which may be understood as inhomogeneous stomatal apertures or slight water condensation on the cuvette walls interfering with the water vapor measurements. However, the prompt decline of assimilation to a zero-exchange of CO<sub>2</sub> under light is a most convincing argument for the strict regulation of this trace gas exchange by stomatal aperture.

### 3.2. Deposition velocities of COS and CO<sub>2</sub>

As described above, an exclusive uptake of COS via the stomatal pathway in close relation to the CO<sub>2</sub> exchange could be demonstrated. The simultaneous measurements of the CO<sub>2</sub>-exchange during all enclosure studies has the potential for a better quantification of the COS uptake. We could directly link the uptake rates of COS and CO<sub>2</sub> by comparing the deposition to the leaves after normalization by the ambient atmospheric concentration of each trace gas, i.e. by comparing deposition velocities ( $V_d$ ). The obtained COS and CO<sub>2</sub> deposition velocities for *F. sylvatica*, *Q. ilex*, *P. sylvestris* and *P. abies* are summarized in Table 1. A clear preference for COS deposition is indicated

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by the uptake ratios of  $V_d$  COS versus  $V_d$  CO<sub>2</sub> and was found for all tree species investigated. The results reflect a significantly pronounced uptake of COS over CO<sub>2</sub> by a factor between 1.4 and 3.4.

In order to widen our basis for further calculations, deposition velocities as reported in or calculated from published data sets were additionally taken into account. In some cases we were able to re-estimate data on the basis of the published figures and tables. For our own data sets published recently, we calculated on the basis of the original data. The results are shown in Table 2. The  $V_d$ -ratios are sensitive to the rate of CO<sub>2</sub> uptake taken into account. In case of enclosure measurements we used the net exchange as measured with the enclosed branch or leaf (Net Primary Productivity; NPP<sub>cuvette</sub>) in order to compare with other published data. However, in case of one data set (spruce forest, Xu et al. 2002) from flux studies above the forest, a correction might be necessary, as this flux value is representing the net exchange as a result of gross photosynthesis minus autotrophic (=NPP) and heterotrophic (soils) respiration. For the enclosure related data we found a range of  $V_{d_{\text{COS}}}/V_{d_{\text{CO}_2}}$  between 1.3 and 5.5 with the exception of 0.4 for non fertilized pea plants (Kesselmeier and Merk, 1993), 7.0 for young corn plants (Hoffmann, 1993) and 8.7–10.3 for a spruce enclosure study (Huber, 1994). These extraordinary numbers may be subject of further speculation.

### 3.3. Corrected estimate of a global sink strength for COS

The data presented above on the uptake of COS normalized to the net assimilation or gross assimilation, respectively, allows a new global sink strength to be estimated for the vegetation (Table 3). Our calculations were based on deposition velocity ( $V_d$ ) ratios of COS versus CO<sub>2</sub> fluxes according to Eq. (1).

$$J_{\text{COS}} = J_{\text{CO}_2} \times [\text{COS}]/[\text{CO}_2] \times V_{d_{\text{COS}}}/V_{d_{\text{CO}_2}} \quad (1)$$

with

$J_{\text{COS}}$       global COS uptake (mol m<sup>-2</sup> s<sup>-1</sup>)  
 $J_{\text{CO}_2}$       global CO<sub>2</sub> uptake (mol m<sup>-2</sup> s<sup>-1</sup>)

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[COS] atmospheric COS mixing ratio ( $\text{mol m}^{-3}$ )  
 [CO<sub>2</sub>] atmospheric CO<sub>2</sub> mixing ratio ( $\text{mol m}^{-3}$ )  
 $V_{d\text{COS}}$  COS deposition velocity ( $\text{m s}^{-1}$ )  
 $V_{d\text{CO}_2}$  CO<sub>2</sub> deposition velocity ( $\text{m s}^{-1}$ ).

5 For global CO<sub>2</sub> fluxes we considered net primary productivity (NPP). NPP dry matter data according to Whittaker and Likens (1975) and Lieth (1975) were recalculated to NPP carbon according to Larcher (1994). Based on our observations of different deposition velocities we assigned ecotype related  $V_d$ -ratios and recalculated the COS sinks adequately. For each ecotype we tried to take into account the exchange behavior of  
 10 typical plant species. Based on our compilation in Table 2 we were able to assign a  $V_d$ -ratio for each ecotype except for “Extreme desert” and “Swamp and marsh”. Instead a best guess was used. The results show that tropical and boreal forests, as well as savannas, are of highest significance for a global estimate of the vegetation sink strength. A ranking of ecotype-significance for the global budget shows that tropical  
 15 rainforest contributes most, followed by tropical seasonal forest, savannah, boreal forest, cultivated land, temperate deciduous and evergreen forest, to an overall NPP based total COS sink strength of  $0.686\text{--}1.404 \text{ Tg a}^{-1}$ .

#### 4. Discussion

20 The close relation of COS uptake to photosynthesis and the clear consumption pathway via stomatal uptake allowed a recalculation of the COS uptake by terrestrial vegetation. We regard this  $V_d$ -ratio-corrected estimate of the COS sink strength to be necessary when estimating a COS sink strength from NPP and regard earlier estimates not taking such a correction into account as to be too low. Furthermore, it has to be kept  
 25 in mind, that this estimate is based on a NPP related CO<sub>2</sub> deposition according to Whittaker and Likens (1975). This NPP value represents only around 50% of the total gross CO<sub>2</sub> uptake by vegetation. A 50% loss by autotrophic respiration, i.e. respira-

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tion by the photoautotrophic biological organism itself, has been subtracted from the Gross Primary Productivity (GPP). In contrast, the data resulting from the enclosure measurements represent a net exchange based on the gross uptake of CO<sub>2</sub> minus the respiration of the leaves and some branches only. This loss is small as compared to the overall respiration of a whole tree with stem and roots. Hence, in a first attempt we doubled the number for the COS sink in order to relate to the real gross CO<sub>2</sub> flux which was the basis for the NPP calculated by Whittaker and Likens (1975). This correction leads to an estimate of a COS sink strength between 1.372–2.808 Tg a<sup>-1</sup>, equivalent to 0.732–1.498 Tg a<sup>-1</sup> on a sulfur basis.

Table 4 gives an overview, comparing the different estimates as reported within the last years. Different procedures have been used based on several parameters such as atmospheric COS concentration, deposition velocity, leaf and ground area as well as the relation of COS deposition to the uptake of CO<sub>2</sub>. The highest estimates between 2 and 5.6 Tg a<sup>-1</sup> were calculated not taking into account the close relation to CO<sub>2</sub> assimilation, thus excluding any diurnal and seasonal effect. Much lower values, between 0.2 and 1.0 Tg a<sup>-1</sup>, supported by a modeling study with similar low estimates (Kettle et al., 2002), were obtained by relating the deposition of COS to the CO<sub>2</sub> assimilation data bases. The latter group, however, did not take into account the preferred uptake and enzymatic consumption of COS as related to the CO<sub>2</sub> assimilation. Including such a preference by correction with the deposition velocity ratios, the new data presented here show that this new procedure results in a significant increase of the COS sink strength calculation again, as productivity, seasonality as well as the preferred uptake is taken into account.

Nevertheless, uncertainties remain large and depend on a still limited data set. Based on a set of enclosure studies we report the net uptake of CO<sub>2</sub> in relation to the net uptake of COS. However, under normal atmospheric concentrations of both gases this means to compare the one-directional COS uptake and consumption with the bi-directional exchange of CO<sub>2</sub>, i.e. COS gross uptake with CO<sub>2</sub> net uptake. This causes an unknown uncertainty as this number is highly dependent on the actual respiration

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rates. A better way would be to relate to gross primary production, i.e. gross photosynthesis. To achieve this goal we need exchange studies covering longer episodes and day/night studies in order to estimate autotrophic respiration for enclosure studies and autotrophic plus heterotrophic respiration for flux studies. Furthermore, we need a better experimental data set especially for tropical and boreal forest trees, as both ecotypes may represent a dominant contribution to the global terrestrial sink strength for COS.

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**Table 1.** Leaf area based exchange and deposition velocities ( $V_d$ ) of COS and  $CO_2$  for *Fagus sylvatica* and *Quercus ilex*, growing under 350 ppm  $CO_2$  over 2–3 years. Three tree individuals ( $T_1$ – $T_3$ ) were grown and investigated. COS-exchange data were obtained under an atmospheric COS mixing ratio of 600 ppt. *Pinus sylvestris* and *Picea abies* were only measured in one year.

Plant Species		COS Uptake ( $pmol\ m^{-2}\ s^{-1}$ )	$\pm$ SD	$V_d$ COS ( $mm\ s^{-1}$ )	$\pm$ SD	(n)	$CO_2$ Uptake ( $\mu mol\ m^{-2}\ min^{-1}$ )	$\pm$ SD	$V_d\ CO_2$ ( $mm\ s^{-1}$ )	$\pm$ SD	(n)	$V_d\ COS / V_d\ CO_2$
<b><i>F. sylvatica</i> Summer 1998</b>	$T_1$	13.8	4.8	0.551	0.211	19	171	12	0.199	0.022	>57	2.77 (1.54 – 4.31)
	$T_2$	12.5	3.0	0.441	0.168	21	193	8	0.225	0.021	>51	1.96 (1.11 – 2.99)
	$T_3$	12.3	4.7	0.429	0.226	36	238	7	0.277	0.025	>69	1.55 (0.67 – 2.60)
<b><i>F. sylvatica</i> Summer 1999</b>	$T_1$	21.7	4.8	0.873	0.217	28	317	29	0.370	0.046	>84	2.36 (1.58 – 3.36)
	$T_2$	26.1	13.7	1.105	0.486	33	332	27	0.387	0.046	>99	2.86 (1.43 – 4.67)
	$T_3$	19.1	8.5	0.994	0.268	43	354	25	0.290	0.078	>129	3.43 (1.97 – 5.95)
<b><i>F. sylvatica</i> Fall 1999</b>	$T_1$	11.3	5.1	0.460	0.221	54	204	6	0.238	0.022	>162	1.93 (0.92 – 3.15)
	$T_2$	11.5	6.4	0.470	0.271	71	241	17	0.281	0.031	>213	1.67 (0.64 – 2.96)
	$T_3$	14.1	0.5	0.575	0.094	99	201	5	0.234	0.021	>297	2.46 (1.89 – 3.14)
<b><i>Q. ilex</i> Summer 1998</b>	$T_1$	15.1	6.9	0.612	0.279	27	283	44	0.330	0.059	>81	1.86 (0.86 – 3.29)
	$T_2$	12.9	4.9	0.542	0.210	50	197	14	0.229	0.026	>150	2.37 (1.30 – 3.70)
	$T_3$	14.8	6.4	0.679	0.189	19	180	9	0.210	0.021	>57	3.23 (2.12 – 4.59)
<b><i>Q. ilex</i> Summer 1999</b>	$T_1$	16.4	4.6	0.751	0.166	33	347	25	0.404	0.045	>99	1.86 (1.30 – 2.55)
	$T_2$	14.4	2.7	0.514	0.164	44	308	22	0.359	0.040	>152	1.43 (0.88 – 2.13)
	$T_3$	15.4	3.4	0.700	0.113	29	318	17	0.371	0.037	>87	1.89 (1.44 – 2.43)
<b><i>Q. ilex</i> Winter 1999/2000</b>	$T_1$	25.8	2.1	0.673	0.372	120	179	16	0.208	0.026	>360	3.24 (1.29 – 5.74)
	$T_2$	25.1	1.5	0.575	0.263	81	318	17	0.371	0.037	>243	1.55 (0.76 – 2.51)
	$T_3$	23.1	2.5	0.733	0.219	94	289	23	0.336	0.039	>282	2.18 (1.37 – 3.21)
<b><i>P. sylvestris</i> Fall 2002</b>	$T_1$	21.1	1.7	0.743	0.035	23	298	22	0.427	0.021	>69	1.74 (1.58 – 1.92)
<b><i>P. abies</i> Fall 2002</b>	$T_1$	12.6	1.7	0.435	0.060	43	275	20	0.459	0.023	>129	0.95 (0.78 – 1.14)

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**Table 2.** Carbonyl sulfide (COS) deposition velocities ( $V_d$ ) and normalized net uptake as expressed in the  $V_d -$  ratio of COS and  $\text{CO}_2$ . Data as obtained in the course of this study by measurements and recalculations from published data compared to available literature values.

Plant species	COS deposition velocity ( $\text{mm s}^{-1}$ )	Normalized relation COS/ $\text{CO}_2$ net uptake ( $V_{d\text{COS}}/V_{d\text{CO}_2}$ )	Atmospheric concentration (ppt)	Remarks	Reference
<b>CROPS</b>					
<i>Allium cepa</i>	0.29 – 0.35		4000	Lab., light. encl.	1
<i>Brassica campestris</i> ssp.	0.47 – 0.56		4000	Lab., light. encl.	1
<i>Brassica napus</i>	1.24	1.25	50 – 300	Lab., light. encl.	7
<i>Brassica oleracea</i> ssp.	0.46 – 0.52		4000	Lab., light. encl.	1
<i>Glycine max</i>	3.1		2000	Lab., light. encl.	2
<i>Glycine max</i>	0.71		500	Lab., light. encl.	3
<i>Lactuca sativa</i>	0.25 – 0.35		4000	Lab., light. encl.	1
<i>Lolium perenne</i>	0.78		4000	Lab., light. encl.	1
<i>Lycopersicon esculentum</i>	0.4		2000	Lab., light. encl.	2
<i>Medicago sativa</i>	1.6		500	Lab., light. encl.	3
<i>Phaseolus vulgaris</i>	1.4		2000	Lab., light. encl.	2
<i>Pisum sativum</i> , non fert.	0.2	0.4	300 – 900	Lab., light. encl.	7
<i>Pisum sativum</i> , fert.	1.1	3.8	300 – 900	Lab., light. encl.	7
<i>Pisum sativum</i> , fert.	1.03 <sup>a</sup>	3.1 <sup>a</sup>	820	Lab., light. encl.	8
<i>Raphanus sativus</i>	0.57		4000	Lab., light. encl.	1
<i>Spinacia oleracea</i>	0.4		4000	Lab., light. encl.	1
<i>Triticum aestivum</i>	1.54		500	Lab., light. encl.	3
<i>Triticum aestivum</i> (10 d)	1.08 <sup>a</sup>	3.2 <sup>a</sup>	360	Lab., light. encl.	8
<i>Zea mays</i>	1.14		500	Lab., light. encl.	3
<i>Zea mays</i>	0.69	2.85	100 – 900	Lab., light. encl.	7
<i>Zea mays</i> (9–12 d)	2.36 <sup>a</sup>	7.0 <sup>a</sup>	740	Lab., light. encl.	8
<b>TREES</b>					
<i>Fagus sylvatica</i>	0.66 ± 0.26	2.33 ± 0.62	600	Lab., light. encl.	This work
<i>Picea abies</i>	0.4 – 1.8	<7 <sup>b</sup>	atmospheric	Field, light. REA	4
<i>Picea abies</i>	0.26 (0.5 max)	8.7 – 10.3 <sup>a</sup>	300 – 650	Field, light. encl.	6
<i>Picea abies</i>	0.46 ± 0.023	0.95	700	Lab., light. encl.	This work
<i>Pinus sylvestris</i>	0.43 ± 0.021	1.74	700	Lab., light. encl.	This work
<i>Porterandia cladantha</i>	0.23 (max)	2.37 <sup>§§</sup>	400 – 600	Field, light. encl.	9
<i>Quercus agrifolia</i>	0.44 <sup>§</sup>	1.8 <sup>§</sup>	300 – 500	Field, light. encl.	5
<i>Quercus ilex</i>	0.64 ± 0.09	2.18 ± 0.66	600	Lab., light. encl.	This work
<i>Sacoglottis gabonensis</i>	0.04 <sup>§§</sup>	1.74 <sup>§§</sup>	400 – 600	Field, light. encl.	9
<i>Sacoglottis gabonensis</i>	0.09 <sup>§§</sup>	5.49 <sup>§§</sup>	400 – 600	Field, light. encl.	9

**Note:** No correction was applied to take into account the decrease of the measured net  $\text{CO}_2$  exchange by respiration (Net uptake). Hence, considering a gross photosynthetic uptake by increasing the  $\text{CO}_2$  uptake can lead to significantly reduced values for the normalized COS/ $\text{CO}_2$ -uptake in case of data obtained by flux studies above the forest with the influence of stem and soil respiration.

<sup>§</sup>recalculated based on published fluxes and atmospheric mixing ratios of 500 ppt for COS and 350 ppm for  $\text{CO}_2$

<sup>§§</sup>recalculated considering maximal  $V_d$ -values only

<sup>a</sup>calculated on actual data as given in the paper

Literature cited: <sup>1</sup> Kluczewski et al. (1985); <sup>2</sup> Taylor et al. (1983); <sup>3</sup> Goldan et al. (1988); <sup>4</sup> Xu et al. (2002);

<sup>5</sup> Kuhn et al. (1999); <sup>6</sup> Huber (1994); <sup>7</sup> Kesselmeier and Merk (1993); <sup>8</sup> Hofmann (1993); <sup>9</sup> Kesselmeier et al. (1993).

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**Table 3.** Estimate of global sink strength for carbonyl sulfide (COS) based on deposition velocity ( $V_d$ ) ratios of COS versus  $\text{CO}_2$  of typical plant species and net primary productivity (NPP). NPP dry matter data according to Whittaker and Likens (1975) and Lieth (1975) recalculated to NPP carbon according to Larcher (1994).

Ecosystem type	Area ( $10^6 \text{ km}^2$ )	NPP dry matter ( $\text{g m}^{-2} \text{ a}^{-1}$ )	NPP dry matter ( $10^9 \text{ t a}^{-1}$ )	NPP C ( $10^{15} \text{ g a}^{-1}$ )	NPP $\text{CO}_2$ ( $10^{15} \text{ g/a}$ )	$V_d$ ratios COS / $\text{CO}_2$ min – max	$F_{\text{COS}}$ Tg /a min - max	Plant species related $V_d$ assigned (see Table 2)
Tropical rain forest	17	2200	37.40	16.46	60.34	1.7 – 3.6	0.246 – 0.508	<i>S. gabonensis</i> , <i>P. cladantha</i>
Tropical seasonal forest	7.5	1600	12.00	5.28	19.36	1.7 – 3.6	0.079 – 0.163	<i>S. gabonensis</i> , <i>P. cladantha</i>
Temperate evergreen forest	5	1300	6.50	2.86	10.49	1.5 – 2.9	0.037 – 0.071	<i>Q. ilex</i> , <i>Q. agrifolia</i>
Temperate deciduous forest	7	1200	8.40	3.70	13.55	1.7 – 3.0	0.054 – 0.095	<i>F. sylvatica</i>
Boreal forest	12	800	9.60	4.22	15.49	1 – 1.7	0.036 – 0.063	<i>P. abies</i> <i>P. sylvestris</i>
Woodland and scrubland	8.5	700	5.95	2.62	9.60	1.5 – 2.9	0.034 – 0.065	<i>Q. ilex</i> , <i>Q. agrifolia</i>
Savannah	15	900	13.50	5.94	21.78	1.5 – 2.9	0.076 – 0.148	<i>Q. ilex</i> , <i>Q. agrifolia</i>
Temperate grassland	9	600	5.40	2.38	8.71	2.0 – 3.0	0.041 – 0.061	<i>Z. mays</i> , <i>T. aestivum</i>
Tundra and alpine	8	140	1.12	0.49	1.81	2.0 – 3.0	0.008 – 0.013	best guess
Desert and semi desert scrub	18	90	1.62	0.71	2.61	1.5 – 2.9	0.009 – 0.018	<i>Q. ilex</i> , <i>Q. agrifolia</i>
Extreme desert, rock, sand, ice	24	3	0.07	0.03	0.12	1.0 – 3.0	0.000 – 0.001	best guess
Cultivated land	14	650	9.10	4.00	14.68	1.3 – 3.8	0.043 – 0.130	<i>Z. mays</i> , <i>T. aestivum</i> , <i>B. Napus</i> , <i>P. sativum</i>
Swamp and marsh	2	3000	6.00	2.64	9.68	1.0 – 3.0	0.023 – 0.068	best guess
<b>TOTAL</b>							<b>0.686 – 1.404</b>	

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**Table 4.** Estimates of the global COS sink strength for terrestrial vegetation.

$Tg\ a^{-1}$	Source	Parameters considered
2 – 5	Brown and Bell (1986)	$V_d$ , LAI, Area, COS atmospheric conc.
5.6	Servant (1989)	$V_d$ , LAI, Area, COS atmospheric conc.
0.2 – 0.6	Goldan et al. (1988)	$R_{COS} = R_{CO_2}$ atmospheric conc. CO <sub>2</sub> uptake
$0.93 \pm 0.07$	Kesselmeier and Merk. (1993)	$V_d\ COS = V_d\ CO_2$ atmospheric ratios CO <sub>2</sub> uptake
0.16 – 0.91	Chin and Davis (1993)	$V_d\ COS = V_d\ CO_2$ atmospheric ratios CO <sub>2</sub> uptake
$0.56 \pm 0.1$	Watts (2000)	$V_d\ COS = V_d\ CO_2$ atmospheric ratios CO <sub>2</sub> uptake
0.21 – 0.27	Kettle et al. (2002)	Surface flux model $V_d\ COS = V_d\ CO_2$ atmospheric ratios CO <sub>2</sub> uptake
1.37 – 2.81	This work	$V_d\ COS > V_d\ CO_2$ atmospheric ratios CO <sub>2</sub> uptake

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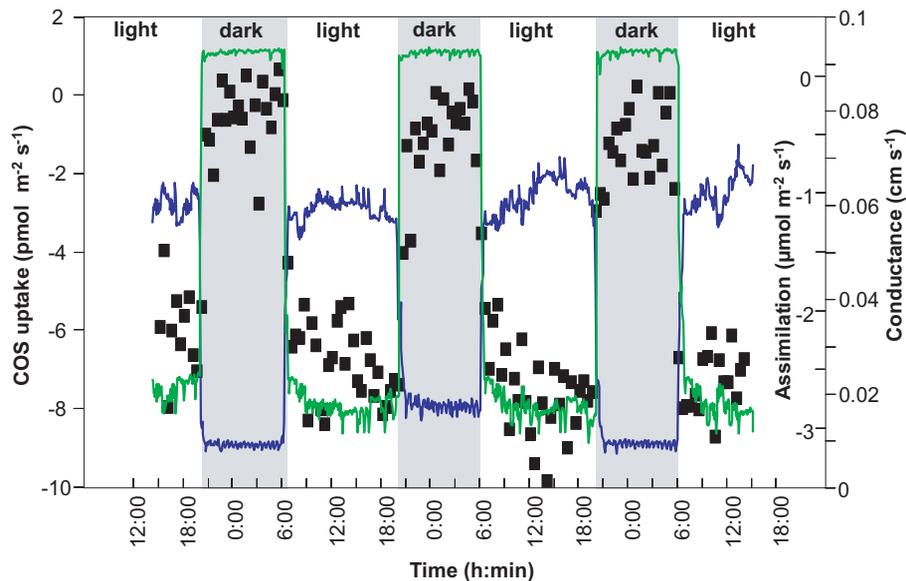
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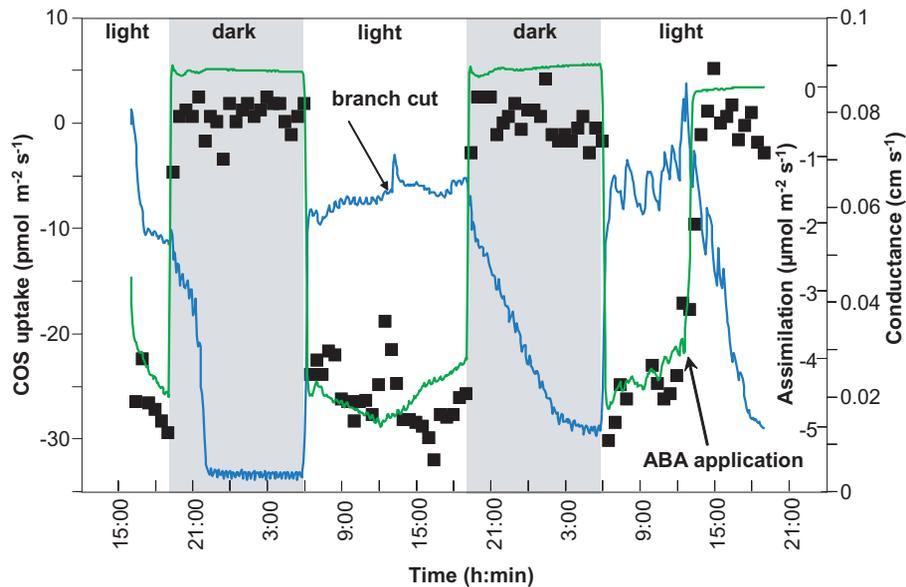
**Fig. 1.** Exchange (negative=uptake) of carbonyl sulfide (COS, filled squares) in relation to branch conductance (blue line) and CO<sub>2</sub> exchange (green line) for European beech (*Fagus sylvatica* L.).

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**Fig. 2.** Exchange (negative=uptake) of carbonyl sulfide (COS, filled squares) in relation to branch conductance (blue line) and  $\text{CO}_2$  exchange (green line) for Holm oak (*Quercus ilex* L.). Note: Conductance calculated from climate chamber conditions ( $25^\circ\text{C}$ , 70 % r.H.).

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