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Past and future simulations of NO₂ from a coupled chemistry-climate model in comparison with observations

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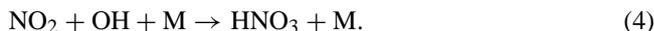
Abstract. Trends in NO₂ derived from a 45 year integration of a chemistry-climate model (CCM) run have been compared with ground-based NO₂ measurements at Lauder (45° S) and Arrival Heights (78° S). Observed trends in NO₂ at both sites exceed the modelled trends in N₂O, the primary source gas for stratospheric NO₂. This suggests that the processes driving the NO₂ trend are not solely dictated by changes in N₂O but are coupled to global atmospheric change, either chemically or dynamically or both. If CCMs are to accurately estimate future changes in ozone, it is important that they comprehensively include all processes affecting NO_x (NO+NO₂) because NO_x concentrations are an important factor affecting ozone concentrations. Comparison of measured and modelled NO₂ trends is a sensitive test of the degree to which these processes are incorporated in the CCM used here. At Lauder the 1980–2000 CCM NO₂ trends (4.2% per decade at sunrise, 3.8% per decade at sunset) are lower than the observed trends (6.5% per decade at sunrise, 6.0% per decade at sunset) but not significantly different at the 2σ level. Large variability in both the model and measurement data from Arrival Heights makes trend analysis of the data difficult. CCM predictions (2001–2019) of NO₂ at Lauder and Arrival Heights show significant reductions in the rate of increase of NO₂ compared with the previous 20 years (1980–2000). The model results indicate that the partitioning of oxides of nitrogen changes with time and is influenced by both chemical forcing and circulation changes.

1 Introduction

It has been recognised for some time that the reactive species NO_x are important in the altitude range from approximately 20 km to 35 km in determining the concentration of stratospheric ozone (Crutzen, 1970). NO_x destroys ozone through the catalytic cycle shown in Reactions (1) and (2).



In the lower stratosphere (approximately 10 km to 20 km) the most significant influence of NO_x is its interaction with the ClO_x and BrO_x ozone loss cycles via the formation of reservoir species ClONO₂ and BrONO₂. NO₂ also reacts with OH (Eq. 4), reducing the HO_x concentration and thus inhibiting the HO_x catalysed destruction of ozone.



Oxidation of N₂O (Reaction (5)) is the major source of stratospheric NO_y (NO+NO₂+NO₃+HNO₃+2N₂O₅+HNO₄+ClONO₂+BrONO₂) and hence NO_x (Minschwener et al., 1993).



N₂O concentrations are predicted to continue to increase over the coming century due to anthropogenic surface emissions, mostly attributed to agricultural nitrogen fixation (IPCC, 2001; WMO, 1999). Trends in the concentration of atmospheric N₂O for the period 1980 to 1988 have been estimated to be +0.25±0.05% per year (IPCC (2001), p253). Zander et al. (1994) quote a trend in N₂O of +0.33±0.04%

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per year based on remote measurements made at Jungfraujoch from 1984 to 1996. In contrast to N₂O, stratospheric halogen concentrations are expected to decline over the coming 50 years as anthropogenic emissions of halogenated ozone-depleting compounds decrease (WMO, 2003). It is therefore important to understand the combined effect of changing halogen and NO_x concentrations on the future evolution of the ozone layer.

Randeniya et al. (2002) studied the effect of increasing N₂O and methane on northern mid-latitude ozone columns for the period 2000–2100 using the CSIRO two dimensional chemical model (Randeniya et al., 1997). Their model results show partial recovery of ozone columns through to the middle of the century due to reductions in halogen concentrations. This is followed by reduction in ozone columns from 2050 to 2100 which Randeniya et al. (2002) attribute to increased destruction of ozone by NO_x, modulated by the amount of methane present in the model. Methane concentrations influence the ozone recovery through the photo-oxidation of methane by halogen radicals in the lower stratosphere and above 20 km through changes in OH concentrations leading to changes in the rate of conversion of NO₂ to HNO₃. Comparing model ozone profiles for 2100 with profiles from 2000, Randeniya et al. (2002) find increases in lower stratospheric ozone concentrations over the 2000–2100 integration. This is explained by a reduction in the amount of ozone destroyed by halogen catalytic cycles. However, the recovery in lower stratospheric ozone is offset by reductions of up to 7% in ozone concentrations in the middle stratosphere due to increased NO_x catalytic destruction of ozone.

Modelling future ozone change requires realistic representations of the anticipated changes in NO_x and halogens, in addition to the many other interacting chemical and dynamical components of the stratospheric system. Three-dimensional coupled chemistry-climate models (CCMs) are designed to capture the interaction between climate change and changes in atmospheric chemistry and are now becoming important tools in the prediction of future stratospheric chemistry and dynamics (WMO, 2003). To have confidence in CCM predictions of future stratospheric change, it is necessary to assess their reliability. An important methodology for validating components of atmospheric chemical models is comparison of long time-series of modelled and measured amounts of trace gases.

Since climate models underpinning all CCMs are chaotic and exhibit unforced variability, comparisons of CCM output with observations are necessarily statistical in nature. Although comparisons between CCM output and observations on individual days are not meaningful in anything other than a climatological sense, the response of the model to long term forcings should agree with the response observed in the real atmosphere. For this reason, long time-series of observations are required to compare the modelled response to atmospheric forcings (WMO, 2003).

High quality measurements of NO₂ slant column densities have been routinely made at Lauder, New Zealand (45° S) since 1980 (Johnston and McKenzie, 1989) and Arrival Heights, Antarctica (77.8° S) since 1982 (McKenzie and Johnston, 1984; Keys and Johnston, 1986, 1988). The measurements are taken at twilight (sunrise and sunset). Measuring at a solar zenith angle (SZA) of 90° with zenith sky viewing geometry increases the sensitivity of the observed columns to stratospheric NO₂ amounts (Solomon et al., 1987).

Liley et al. (2000) studied the Lauder NO₂ data set using least squares regression. Indices of QBO, ENSO, solar cycle and the El Chichon and Pinatubo volcanic events in addition to the NO₂ annual cycle and secular trend were fitted to the measured NO₂ time-series. Liley et al. (2000) concluded that the mean 90° SZA linear trend in NO₂ over Lauder from 1980 to 1999 was 5±1% per decade (sunrise trend 5.9±2.4, sunset trend 4.6±1.5 % per decade). These trends are significantly greater than the trend in N₂O estimated to be between 2.5% and 3.3% per decade (WMO, 1999). As N₂O is the dominant source of stratospheric NO₂, this result implies that there has been a dynamical or chemical change at southern mid-latitudes which has increased NO₂ disproportionately.

Using 1995–2002 Fourier Transform Infra-Red (FTIR) measurements at Kitt Peak (31.9° N), Rinsland et al. (2003) derived an NO₂ trend of 10.3±5.5% (2σ) per decade. Liley et al. (2000) show however, that using short time periods in regression analyses results in large uncertainties in the trends and therefore the results of Rinsland et al. (2003) cannot be meaningfully compared with those of Liley et al. (2000) or those presented here. NO₂ data derived from a combination of FTIR and differential optical absorption spectroscopy measurements taken at the Network of the Detection of Stratospheric Change (NDSC) station at Jungfraujoch (46.5° N) have been analysed and suggest a linear increase for the period 1985–2001 of 6±2% per decade (WMO, 2003). These results suggest that the NO₂ trend being greater than the N₂O trend is unlikely to be a local feature at Lauder but is likely to be of global extent.

Fish et al. (2000) used a column model to determine the sensitivity of southern mid-latitude NO₂ to changes in stratospheric temperature, ozone and water vapour. A number of possible mechanisms that could change the amount of NO₂ relative to the amount of N₂O were identified:

- direct emission of NO_x
- changes in stratospheric circulation
- changes in the shape of the mean NO₂ profile
- changes in the partitioning of NO_y .

Changes in the partitioning of NO_y could arise due to changes in ozone, temperature, stratospheric water vapour and sulfate aerosol. The focus of Fish et al. (2000) was to determine whether changes in the partitioning of NO_y

could explain the observed trend in NO₂. The model was forced with observed changes in ozone, temperature and water vapour and a 2.5% per decade increase in N₂O. The model failed to reproduce the observed NO₂ trend (model trends +4.0±0.6% per decade at sunrise, +2.0±0.4% per decade at sunset). Including a 20% decrease in stratospheric aerosol in the model forcing gave NO₂ trends in agreement with observations (model trends +5.9±0.6% per decade at sunrise, +4.3±0.4% per decade at sunset).

McLinden et al. (2001) used a combination of a three-dimensional chemical transport model, a static column chemistry model and a radiative transfer model to generate NO₂ slant column densities and compared the change in the columns over a 20 year time period with the Lauder NO₂ observations. Their results show a trend in NO₂ of 4.3% per decade for a prescribed increase in N₂O of 3% per decade. Differences in the trends of NO₂ and N₂O were attributed to the less than equivalent conversion of N₂O to NO_y and repartitioning of NO_y due to ozone and halogen changes. The sensitivity of their system to temperature and aerosol changes was also investigated. McLinden et al. (2001) show that the trend in NO₂ vertical column density varies diurnally with large changes in the slant column density trends at sunrise and sunset.

Both the studies of Fish et al. (2000) and McLinden et al. (2001) do not explicitly model the influence of circulation changes on NO₂ amounts. They use rather different models and forcings to estimate NO₂ slant column density trends. Both conclude that their models reproduce the observed NO₂ trend but differ in their conclusion as to why there is a distinction between the trends in NO₂ and N₂O. More work is required to fully understand the mechanisms responsible for the observed difference in trends.

In this paper we calculate NO₂ slant column densities for Lauder and Arrival Heights using results from the Unified Model with Eulerian Transport and Chemistry (UMETRAC), a three-dimensional CCM. NO₂ values are derived from a 40 year simulation of the model, (1980 to 2019). Model NO₂ slant columns for Lauder and Arrival Heights are compared with the measurements for the period 1980 to 2000 to test the model's ability to reproduce the greater than than proportional increase in NO₂ relative to N₂O that has been observed in the measurements. The UMETRAC predictions of NO₂ trends at Lauder and Arrival Heights for the period 2001 to 2019 are introduced and discussed in light of the 1980 to 2000 model/observation comparison.

2 Measurements

The NO₂ measurement technique and retrieval algorithm are discussed by Johnston and McKenzie (1989) and Liley et al. (2000). The automated scanning spectrometers measure wavelengths from 435 nm to 450 nm with a spectral resolution of 1.2 nm. Twilight spectra are ratioed with mid-

day reference spectra to remove Fraunhofer absorption lines present in sunlight. Each twilight measurement is also corrected for the Ring effect (Grainger and Ring, 1963) using the “offset Ring” approach (Johnston and McKenzie, 1989).

The instruments at Lauder and Arrival Heights have been calibrated and intercompared to the standard required of the Network for Detection of Stratospheric Change (NDSC).

Model ozone and 20 hPa temperatures are also compared with measurements as changes in these quantities have been identified by Fish et al. (2000) and McLinden et al. (2001) as being important in determining changes in NO₂. The 20 hPa level approximately coincides with the peak in the NO₂ mixing ratio and was therefore used as the level for the temperature comparison. It has been shown (Keys and Johnston, 1986) that NO₂ slant columns measured over Arrival Heights correlate strongly with stratospheric temperatures.

Modelled ozone data are compared with the NIWA assimilated ozone data-set (Bodeker et al., 2001). This data-set contains total column ozone values derived from a combination of TOMS and GOME satellite measurements which are corrected against the global ground-based network of Dobson spectrophotometers. NCEP/NCAR (Kalnay et al., 1996) 20 hPa temperature data are used for the temperature comparison.

3 Least squares regression analysis

Linear trends in the NO₂ slant column density time-series are calculated using a least squares regression model. The regression model has been applied to ozone time-series in previous studies (Bodeker et al., 1998, 2001).

A number of forcings are known to influence the southern mid-latitude stratosphere:

- El Niño southern oscillation (ENSO)
- Quasi-biennial oscillation (QBO)
- 11 year solar cycle
- Volcanic injection of sulfate aerosol and water vapour. For the time period 1980–2000, two volcanic events occurred which were significant for the southern mid-latitude stratosphere, El Chichon (1982) and Pinatubo (1991).

The regression model is designed to allow fitting of indices for the above forcings, in addition to fitting the seasonal cycle and linear trend in the data. Inclusion of additional, higher order Fourier components to the basis function describing the trend allows the model to fit seasonally varying trends in the data. A description of the indices for the externally forced basis functions (ENSO, QBO, solar cycle and volcanoes) within the regression model is given in Bodeker et al. (1998).

The residual terms from the fitting may be autocorrelated leading to an underestimation of the uncertainty in the resulting trends. An autocorrelation model is applied to the residuals to correct for any autocorrelation.

For the observational record, all basis functions (ENSO, QBO, solar cycle volcanoes) are included in the regression analysis. For the model time-series, only the ENSO and QBO terms are applied (in addition to the annual cycle (offset) and linear trend). The model does not include an 11 year solar cycle and uses an invariant, background sulfate aerosol field. The non-orographic gravity wave forcing scheme within the model produces a QBO in the tropical zonal winds (Scaife et al., 2000). Mean, equatorial zonal winds produced by the model at 50 hPa were used to generate the model QBO basis function.

4 Model description

4.1 The coupled chemistry-climate model

UMETRAC uses the Met Office's Unified Model (UM) (Cullen and Davies, 1991) as the underlying climate model. The model has a resolution of 3.75° (longitude), 2.5° (latitude) and 64 vertical levels from the surface to 0.01 hPa. A non-orographic gravity wave forcing scheme (Warner and McIntyre, 1999) is used to parameterise gravity wave breaking.

The UMETRAC chemistry scheme is based on a families approach (Austin, 1991). 15 chemical tracers and one dynamical tracer are advected by the model. The dynamical tracer is used to parameterise the long lived species H₂O, CH₄, Cl_y, Br_y, H₂SO₄ and NO_y. The chemistry scheme includes 65 gas phase chemical reactions, 9 heterogeneous chemical reactions and 27 photolysis reactions. A PSC scheme based on liquid ternary solutions and water ice and a simple sedimentation scheme are also part of the model. Chemical reaction rates are taken from DeMore et al. (1997) and Sander et al. (2000).

There are some differences between the UMETRAC configuration used in this paper and the configuration used in previous work (Austin and Butchart, 2003). These include a new tracer advection scheme, the chemistry scheme being applied over the whole vertical domain of the model rather than being restricted to the stratosphere and lower mesosphere, two additional chemical tracers are included (CO and CH₃OOH) and an extension of the chemical reaction set.

Water vapour concentrations in the chemistry module were taken from the UARS reference atmosphere project (URAP) reference atmosphere (<http://code916.gsfc.nasa.gov/Public/Analysis/UARS/urap/home.html>) and held fixed over the integration. Stratospheric aerosol was derived from the aerosol surface area densities given in Table 8.8 of WMO (1991). Again, the aerosol amounts were held constant over the model integration.

Importantly for this study, the rate of increase of N₂O was fixed at +2.6% per decade for the entire length of the integration (1980 to 2019). The model uses a parameterisation to determine the amount of NO_y present at each time step, based on the transport of a conserved tracer and the compact relationships of Plumb and Ko. (1992). In a subsequent step, the model partitions the NO_y according to the chemical scheme within the model. The global rate of increase of NO_y is fixed to the rate of increase in N₂O (2.6% per decade) but local rates of increase in NO_y may differ from this in response to circulation changes in the model via changes in the conserved tracer.

Results used in this paper come from a 45 year integration of UMETRAC (1975 to 2019), with the first five years used as model spinup. The IS92a IPCC scenario was used to determine greenhouse gas concentrations (CO₂, CH₄ and N₂O). Halogen concentrations were taken from the WMO (1999) assessment. Observed sea surface temperatures (SST) were used in the model from 1975 to 1999. For years after 1999, sea surface temperature data were taken from simulations of a coupled ocean-atmosphere version of the UM.

Global fields of the 15 chemical tracers, 6 long-lived species and chemical families and temperature are output from UMETRAC every five days at 00:00 UT. These data are used as input for the UMETRAC column model which produced NO₂ profiles.

4.2 UMETRAC column model

Since the output from the 45 year UMETRAC integration provided only concentrations of chemical families, an off-line chemical model was required to determine the partitioning amongst the chemical families output from the full three-dimensional model. Vertical profiles of chemical families and daily mean temperature for the grid boxes covering Lauder and Arrival Heights were extracted from the UMETRAC output and used as initial conditions for the column model. The chemistry scheme within the column model is identical to the chemical scheme used in the full three-dimensional model.

The column model was run for one day to establish the NO₂ diurnal cycle. The chemical time step was reduced from 15 min in the full three-dimensional model to 3 min for the column model. This allows NO₂ profiles to be calculated with a SZA resolution of 1°, ensuring the diurnal cycle of the model is adequately represented in the radiative transfer model for the calculation of slant column densities.

Over the one day integration of the column model, the NO₂ concentrations will differ from the corresponding NO₂ concentrations calculated in the full UMETRAC model due to the use of the diurnal mean temperature profile. Using simple sensitivity tests, it can be shown that these inconsistencies between the full model and the off-line column model have little impact on the slant column densities due to the

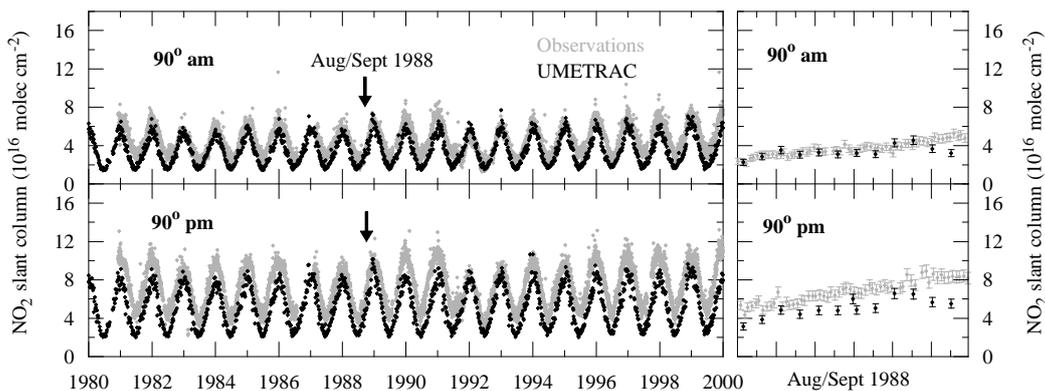


Fig. 1. Comparison of UMETRAC modelled (black) and observed (grey) 90° NO₂ slant column density time-series for Lauder (1980–2000). Upper left panel sunrise, lower left panel sunset. A subset of the data is shown in the right hand panels to illustrate the estimated errors in the model and observed values and the data frequency.

small diurnal temperature variations in the lower and middle stratosphere where the slant column weighting is greatest.

Finally, the NO₂ profiles calculated by the column model were interpolated from terrain following model levels to altitude surfaces from the ground to 72 km in 500 m steps.

4.3 Radiative transfer model

In transforming the measured NO₂ slant column densities to vertical column densities, some prior knowledge of the vertical profile shape is required (McKenzie et al., 1991). Assuming a vertical profile shape in an air mass factor calculation can lead to errors due to the seasonal and annual variation of the true vertical profile, which the assumed profile shape may not capture. In this work the slant column densities were calculated using the radiative transfer model (RTM) developed by Schofield et al. (2003) from the profiles determined by UMETRAC. This allows direct comparison of slant column densities derived from UMETRAC output with observations.

The RTM constructs a model atmosphere using temperature, pressure and ozone profiles taken from the UMETRAC column model. The RTM uses a spherically curved atmospheric geometry, divided into discrete atmospheric shells. For this study, ninety 1 km thick shells were used to describe the model atmosphere. The effects of refraction, Rayleigh scattering, Mie scattering and molecular absorption at a wavelength of 450 nm (see Sect. 2) were included in the RTM path description of the NO₂ zenith-sky measurement. Refractive indices were taken from Bucholtz (1995). A single scattering approximation was used for the zenith-sky viewing geometry.

The diurnal variation of the NO₂ profile adds complexity to the slant column density calculation. The variation of the NO₂ abundance along the slant path with solar zenith angle is explicitly taken into account. NO₂ profiles at 1° solar zenith angle intervals between 30° and 97° calculated by

the UMETRAC column model were used to construct a two-dimensional profile grid. This was then interpolated to the relevant altitude and solar zenith angle along each scattered light path. The slant column for the solar zenith angle of 90° was calculated by integrating the amount of NO₂ over all light paths scattered from the zenith.

5 Discussion of Lauder results

5.1 Lauder NO₂ time-series

The time-series of modelled and measured NO₂ slant column densities for Lauder are shown in Fig. 1. The model results capture the absolute value of the NO₂ slant columns and the amplitudes of the seasonal cycle in NO₂ for both sunrise and sunset very well. The model in general underestimates the observed slant columns for both the sunrise and sunset time-series. The agreement for the sunset case is somewhat worse than the sunrise which implies that the model has a suppressed NO₂ diurnal cycle.

The right hand panels of Fig. 1 show a subset of the data to illustrate more clearly the data frequency and estimated errors. The time period (August and September 1988) was chosen as representative of the whole data record, excluding the time periods affected by the El Chichon and Pinatubo eruptions. Observation errors are estimated at 0.2×10^{16} molec cm⁻² + 5% of the observed value. This prescription comes from an analysis of the error characteristics of the measured spectra and the retrieval algorithm used to derive the NO₂ slant column measurements.

The regression model used to fit the time-series requires an error estimate associated with each input data point. In the absence of an ab initio estimate of the error in the model NO₂ slant column densities, errors are prescribed using the same error estimate as that used for the observations.

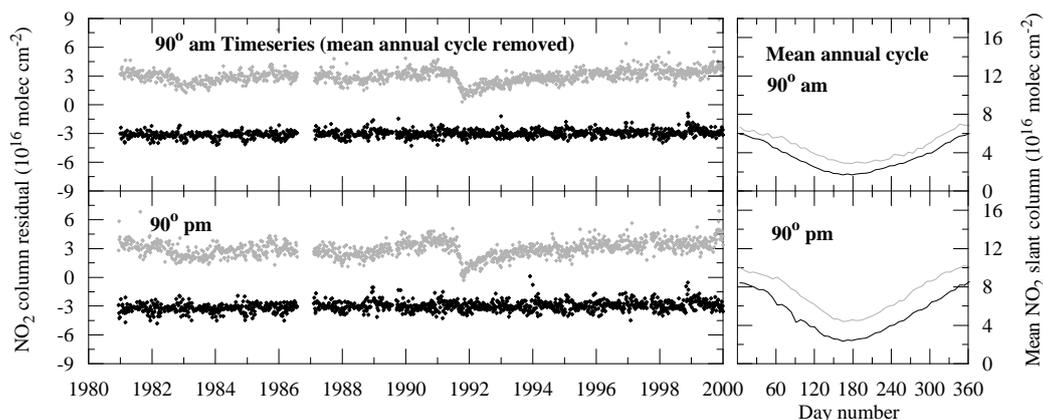


Fig. 2. Lauder NO₂ slant column density time-series with the mean annual cycle removed. Observed residuals (grey) have been offset by +3 units and the model residuals (black) offset by −3 units for clarity. The time-series have been thinned by retaining data only when both observation and model values are present on a given day. Right hand panels show the mean annual cycles (1980–2000) for sunrise (top) and sunset (bottom).

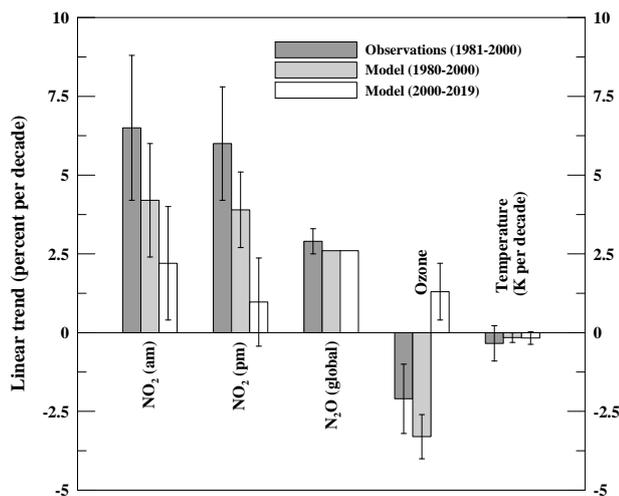


Fig. 3. Seasonally independent trends in NO₂ for Lauder in % per decade, derived from observations and UMETRAC model results. Also given are the estimated trends in global N₂O over the same time periods and the observed and modelled trends in ozone and temperature (temperature trends are shown in K/decade). The errors are $\pm 2\sigma$. Note, model N₂O trends are the prescribed global value and thus has no error associated with them.

The observed data and model results from Fig. 1 were re-sampled to produce new time-series in which values were only retained for a given day, if both modelled and observed slant columns were present on that day. This processing removes sampling biases when comparing modelled and observed time-series. Figure 2 shows these time-series, with the mean annual cycles removed to show more clearly factors influencing the variability in the data. For clarity, the observation and model time-series are offset by +3 and −3 ($\times 10^{16}$ molec cm^{−2}), respectively.

The observed time-series in Fig. 2 clearly show the reduction in NO₂ amounts following the Pinatubo eruption. This effect is also seen to a lesser extent for the El Chichon eruption (1982). This reduction in NO₂ occurs via the hydrolysis of N₂O₅ to HNO₃ on the volcanic sulphate aerosol (Rodríguez et al., 1991). N₂O₅ is an important night-time NO_x reservoir and therefore removal of stratospheric N₂O₅ causes denoxification of the stratosphere. Other interannual variability is evident in the observational record. In contrast the model results show little variation other than the short term daily variability and a linear trend.

The right hand panels of Fig. 2 show the mean annual cycles for the four Lauder data-sets. As indicated above, the model slant columns underestimate the observed data with the sunrise comparison being in better agreement than the sunset case. For both cases the amplitude of the modelled mean annual cycle is consistent with observations with the whole cycle offset lower (by approximately 1×10^{16} molec cm^{−2} for sunrise and 2×10^{16} molec cm^{−2} for sunset). A constant shift of the modelled mean seasonal cycle NO₂ relative to observations suggests that too little NO_y is present in the model. This does not explain the suppression of the diurnal cycle. More work is required to fully explain the differences between the modelled and observed mean annual cycles.

5.2 Lauder past trends and factors affecting them

Seasonally independent linear trends in the observation and model time-series, calculated using the least squares analysis algorithm are given in Fig. 3.

The NO₂ trends derived from observations differ from the results of Liley et al. (2000) (5% per decade) even though the same observational data-set was used in both analyses. This is because different time periods were used in the two

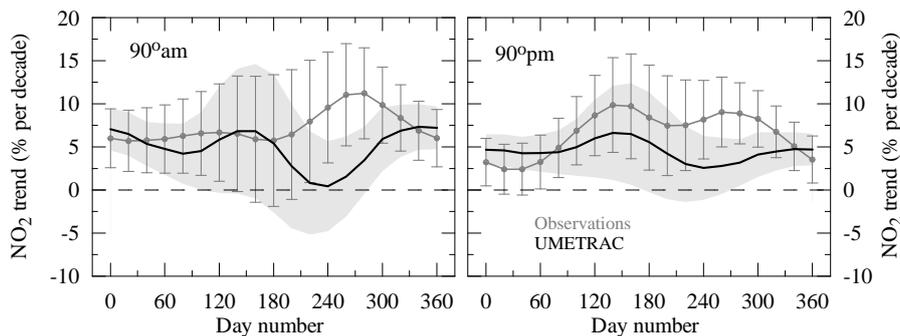


Fig. 4. Annually varying NO₂ trends for Lauder for the period 1980–2000. The grey shaded region represents the 2 σ uncertainty range in the trends derived from model output. Error bars on the observed trends indicate $\pm 2\sigma$.

analyses (Liley et al. (2000) studied the period 1981–1999). The 26 parameter model with autocorrelation (AC) correction used by Liley et al. (2000) for the period 1981–1999 most closely resembles the regression model used in this study. Applying our regression model to the Lauder measurements for the period 1981–1999 gives linear NO₂ trends that agree with the results of Liley et al. (2000).

The model NO₂ results underestimate the NO₂ trends derived from the measurements for both sunrise and sunset but the model and observed trends do agree within the 2 σ uncertainty range. The trends derived from observations are significantly higher than the observed trends in N₂O (assumed to be between 2.5 and 3.3% per decade). This result is in agreement with the conclusions of Liley et al. (2000). The model results show a significant difference (at the 2 σ level) in the NO₂ and N₂O trends for the sunset case only. The NO₂ results from Fig. 3 indicate that UMETRAC is not fully capturing the magnitude of the change in relative amounts of NO₂ and N₂O seen in the observations over the period 1980 to 2000. This will be discussed more fully in Sect. 5.2.1 where the seasonally dependent NO₂ trends are examined.

Fish et al. (2000) and McLinden et al. (2001) both show that modelled NO₂ trends are sensitive to changes in ozone and temperature. Fish et al. (2000) suggest this is primarily due to the reaction,



which has a strong temperature dependence. To have confidence in the modelled NO₂ trends it is therefore important for the model to reproduce observed ozone and temperature trends. The model slightly overestimates the seasonally independent negative ozone trend (see Fig. 3) although the difference from observations is within the 2 σ uncertainty range. Temperature trends agree well with both the observations and model showing a small negative 20 hPa temperature trend. Both the observed and modelled temperature trends are not significantly different from zero at the 2 σ level.

The NO_y chemical family and HNO₃ are relatively long lived tracers in the lower stratosphere (Brasseur and

Table 1. Lauder – seasonally independent UMETRAC model trends in NO₂, N₂O, ozone, temperature, NO_y and HNO₃ for the periods 1981–2000 and 2001–2019. Trends are given in percent per decade (temperature trends are shown in K/decade). The quoted errors are $\pm 2\sigma$.

	Model (1981–2000)		Model (2001–2019)	
	am	pm	am	pm
NO ₂	4.2 \pm 1.8	3.8 \pm 1.2	2.2 \pm 1.8	0.96 \pm 1.3
N ₂ O	2.6		2.6	
Ozone	−3.3 \pm 0.7		1.3 \pm 0.8	
Temperature	−0.16 \pm 0.14		−0.18 \pm 0.18	
NO _y	2.5 \pm 1.1		2.7 \pm 1.2	
HNO ₃	1.8 \pm 1.3		3.4 \pm 1.5	

Solomon, 1986). Therefore their concentrations can be influenced by both circulation changes and in-situ chemical processing. The concentrations of these species in turn control the concentration of NO₂. To further investigate the role these species have in determining trends in NO₂ in the model, NO_y and HNO₃ vertical column densities (vcds) were calculated from the model output. Every five days NO_y and HNO₃ vertical column densities were taken from the model output and the regression model applied to these time-series to determine linear trends for the period 1980–2000.

Modelled NO_y and HNO₃ seasonally independent trends at Lauder for the period 1980–2000 were calculated to be +2.5 \pm 1.1% per decade and +1.8 \pm 1.3% per decade, respectively (Table 1). The 2.5% per decade trend in NO_y matches the global rate of increase in N₂O and is significantly less than the observed trends in NO₂.

The fact that HNO₃ is increasing at a lower rate than NO_y suggests that the higher trend in NO₂ over the period 1980–2000 is gained at the expense of HNO₃, and therefore the partitioning in NO_y over this period is shifting towards the more chemically active NO_x species and away from the HNO₃ reservoir. Additional analysis is required to determine what

fraction of the difference between the NO_y and HNO₃ trends can be attributed to circulation changes and what fraction of the difference is due to chemical forcing.

5.2.1 Annually varying NO₂ trends

Figure 4 shows the seasonally dependent NO₂ trends calculated by adding two Fourier components (annual and semi-annual) to the basis function describing the trend in the regression model (Bodeker et al., 1998). The trends are taken as a percentage of the daily mean. That is, the trends are calculated in units of molec cm⁻² per decade and then taken as a percentage of the annual mean over the whole period (given in the right hand panels of Fig. 2).

The results shown in Fig. 4 indicate that for much of the year, the model NO₂ trends are in good agreement with the trends derived from observations. The greatest differences occur in spring time for both sunrise and sunset cases, where the model is underestimating the trends seen in the observations. The model underestimation of the springtime increase in NO₂ is the cause of the model failing to fully reproduce the difference between the seasonally independent NO₂ and N₂O trends for the period 1980 to 2000 (see Fig. 3).

The same regression model (with seasonally dependent basis functions for the linear trend term) was used to fit the modelled and measured ozone and 20 hPa temperature time-series. The seasonal cycle in ozone and 20 hPa temperature trends was reproduced well by the model (not shown). Therefore ozone and temperature were ruled out as the cause of the model underprediction of springtime NO₂ trends.

Differences in the timing of the vortex breakup and subsequent mixing of vortex and mid-latitude air can also be ruled out as the cause of the differences in the modelled and observed Lauder springtime NO₂ trends. Generally, the mixing of vortex air to mid-latitudes occurs in early summer (Ajtic, 2004), from approximately day 288 (15 October) which is after the maximum in the trend differences (approximately day 240–28 August). The model, if anything, tends to delay the breakup of the Antarctic polar vortex further reinforcing this conclusion.

Stratospheric NO₂ amounts have been shown to be sensitive to changes stratospheric water vapour (Fish et al., 2000; McLinden et al., 2001). Measured trends in stratospheric water vapour are uncertain (SPARC, 2000; Rosenlof et al., 2001; Randel et al., 2004). In this study, UMETRAC used a fixed climatology of stratospheric water vapour which may contribute to the discrepancy between modelled and observed NO₂ trends.

5.3 Model predictions: Lauder 2001–2019

Table 1 (and Fig. 3) compares the seasonally independent, predicted NO₂ model trends for the period 1 January 2001 to 31 December 2019 with the model trends for the period 1 January 1980 to 31 December 2000.

Predicted NO₂ trends (2001–2019) are lower than the trends for the period 1980–2000 even though the N₂O trend stays the same. This difference is statistically significant at the 2σ level for the sunset case. Both the absolute and relative differences between trends for the two periods are larger for the sunset case.

NO_y and HNO₃ vertical column density trends for the period 2001–2019 are compared with the 1980–2000 trends in Table 1. It is clear from Table 1 that the shift in NO₂ trends from 1980–2000 to 2001–2019 is associated with a change in HNO₃ trends between the two periods. As was the case for the 1980–2000 period, it is a change in partitioning of the NO_y that is resulting in the less than proportional rate of increase in NO₂ relative to NO_y (and N₂O) rather than local changes in the amount of NO_y relative to N₂O.

Quantification of how chemical changes (for example halogen loading and ozone) and circulation changes affect NO_y partitioning requires additional analysis of model results which is outside the scope of this paper.

It should be reiterated that the results shown here are for a solar zenith angle of 90°. McLinden et al. (2001) show that the trends in NO₂ vary significantly with solar zenith angle. It is therefore to be expected that different conclusions may be drawn from analysis of data at different solar zenith angles.

6 Discussion of Arrival Heights results

6.1 Arrival Heights NO₂ time-series

Arrival Heights slant column density comparisons are shown in Fig. 5. Twilight (SZA=90°) data at Arrival Heights are only available during the autumn (17 February to 21 April) and spring (22 August to 25 October). Because of the large annual cycle, daily and seasonal variability and the discontinuous nature of the data, it is difficult to assess the level of agreement in the data from Fig. 5.

Figure 6 shows the processed Arrival Heights time-series comparison with the annual mean removed (see Sect. 5.1 and Fig. 2). Significantly less data are available for Arrival Heights than for Lauder because of the limited period of twilight (spring and autumn) at Arrival Heights and gaps in the observation record, particularly in the years 1982–1989. Both the observed and modelled data for Arrival Heights show greater variability about the mean than the Lauder data.

There is good qualitative agreement between observations and model for the mean seasonal cycle at both sunrise and sunset. However, during autumn, the model tends to systematically underpredict the observations, possibly for reasons discussed in Sect. 7.

6.2 Arrival Heights NO₂ trends

Regression analysis of the Arrival Heights data is complicated by the discontinuous nature of the data. We analyse the

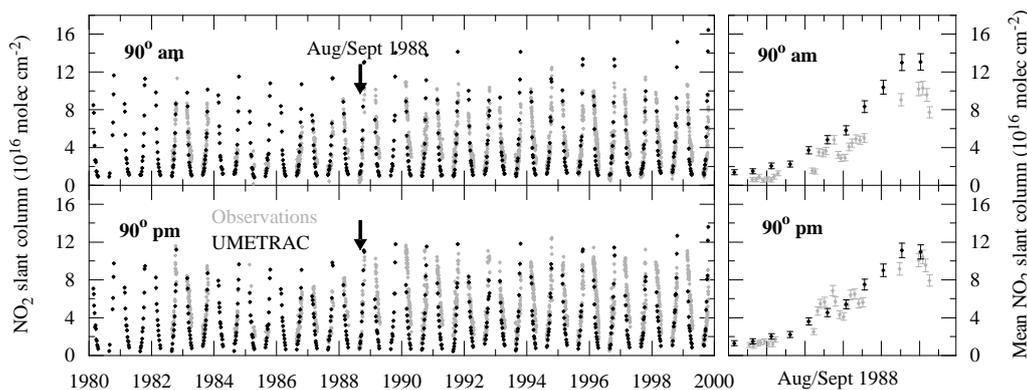


Fig. 5. Comparison of UMETRAC modelled (black) and observed (grey) 90° NO₂ slant column density time-series for Arrival Heights (1980–2000). Upper left panel sunrise, lower left panel sunset. A subset of the data is shown in the right hand panels to illustrate the estimated errors in the model and observed values and the data frequency.

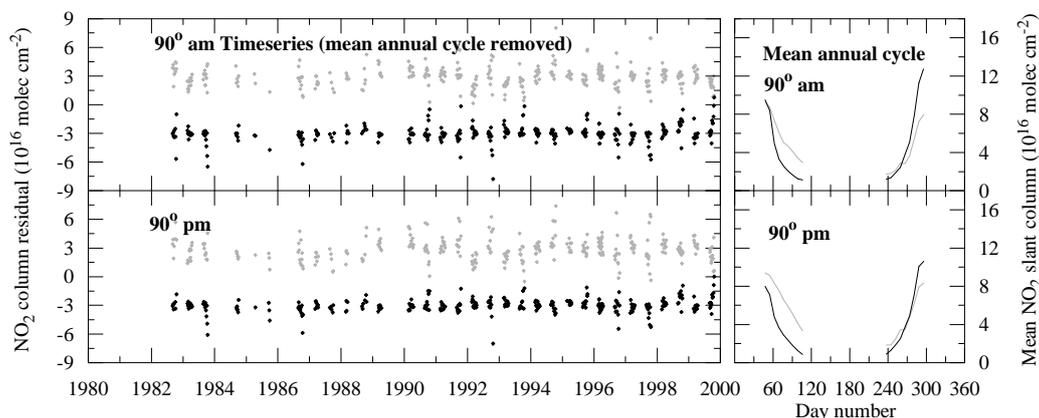


Fig. 6. Arrival Heights NO₂ slant column density time-series with the mean annual cycle removed. Observed residuals (grey) have been offset by +3 units and the model residuals (black) offset by −3 units for clarity. The time-series have been thinned by retaining data only when both an observation and model value is present on a given day. Right hand panels show the mean annual cycles (1980–2000) for sunrise and sunset.

autumn and spring data separately as the chemical and dynamical regimes during the two seasons are markedly different and therefore the NO₂ response is expected to be different for the two seasons. No seasonal components were included as part of any of the basis functions used in the regression model.

The basis function describing the effect of the El Chichon eruption was not used in the regression analysis of the Arrival Heights data because the first observation available is 30 August 1982, approximately five months after the El Chichon eruption. The El Chichon basis function would include only a decay term and can potentially induce spurious trend results if included.

Because the NO₂ slant columns are correlated with 20 hPa temperature, particularly in the spring (Keys and Johnston, 1986), an additional basis function was added to the regression model which describes 20 hPa temperature anomalies

over Arrival Heights. Episodic events of anomalously cold stratospheric air over Arrival Heights are associated with transport of air from the cold vortex core to Arrival Heights, which generally sits below the relatively poorly mixed edge of the polar vortex.

The temperature basis function was generated by taking daily Arrival Heights 20 hPa temperatures from the NCEP/NCAR reanalysis data-set and fitting them using the same regression model as is used for the NO₂ data. The residual temperatures from this fitting were then used as the temperature basis function for the regression analysis of the NO₂ data. The same process was applied to the model 20 hPa temperatures to generate the model temperature basis function.

Figure 7 gives the linear, (autumn–upper panel and spring–lower panel) trends in NO₂, N₂O, ozone and temperature for Arrival Heights from the model and observations for the period September 1982 to October 2000. It is evident from this

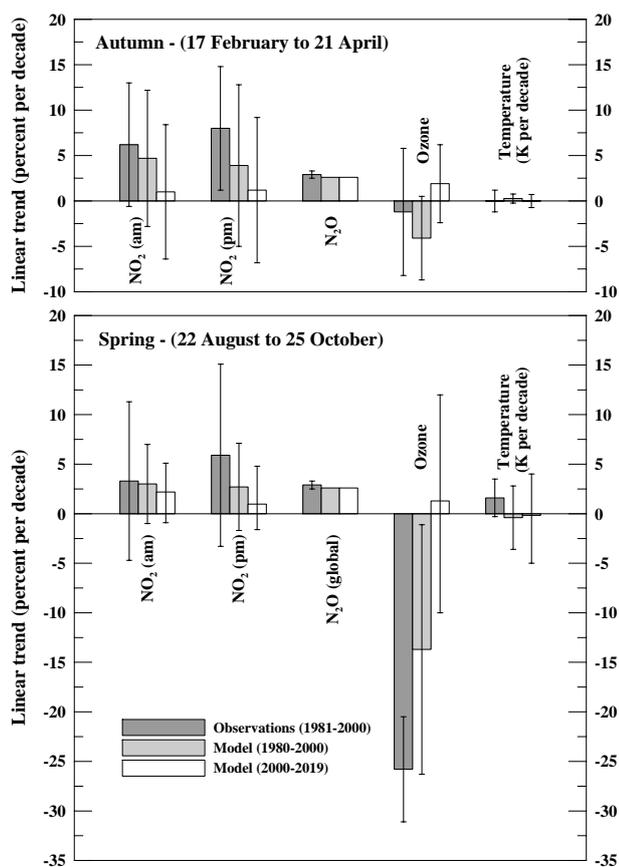


Fig. 7. Autumn (upper panel) and spring (lower panel) trends in NO₂, N₂O, ozone and temperature derived from observations and model results for Arrival Heights. Trends are given in % per decade (temperature trends are shown in K/decade). The errors are $\pm 2\sigma$.

figure that the 2σ uncertainties in the derived trends for both the observations and model are large compared to the trends themselves. Although there is agreement between modelled and observed trends for all cases it is difficult to attribute any significance to the comparison given the large uncertainties.

Modelled and observed autumn NO₂ trends (Fig. 7 upper panel) are greater than the spring NO₂ trends. It is expected that increases in NO_y and NO₂ in the springtime polar regions will be offset to some extent by denitrification (and denoxification) of the lower stratosphere which is implied by these results but the amounts have not been quantified. As is the case with the Lauder comparison, the model underpredicts the rate of increase in NO₂ relative to the observations in all cases.

The autumn NO₂ trends shown in Fig. 7 are similar to the Lauder results from Fig. 3. The ozone and 20 hPa temperature results are also similar. Thus, at high latitudes in autumn, the stratosphere is showing a similar response in NO₂, ozone and temperature to mid-latitudes. This is expected given the quiescent and well mixed nature of the summer and autumn extratropical stratosphere.

The chemically and dynamically disturbed high latitude spring on the other hand, shows some deviations from previous results (Fig. 7 lower panel). As expected, the negative ozone trends are significantly larger than the autumn case and as mentioned previously, the NO₂ trends are closer to the N₂O trends than for autumn.

6.3 Model predictions: Arrival Heights 2000–2019

The predicted (2001–2019) NO₂, ozone and temperature trends for Arrival Heights are shown in Table 2 and Fig. 7. As with the autumn model/observation comparison discussed in the previous section, the autumn model predictions are similar to the Lauder results (see Table 1). NO₂ increases at a greater than equivalent rate with respect to N₂O for the period 1982–2000. This is followed by a shift to a less than equivalent increase in NO₂ for 2001–2019. The change in NO₂ trends is associated with a change in ozone trends from a negative trend of -4.1% per decade to a small positive trend ($+1.9\pm 4.3\%$ per decade).

Spring NO₂ trends for both the 1982–2000 period and the 2001–2019 period are close to the N₂O trend applied to the model. There is a small decrease in the 2001–2019 predicted trends relative to the 1982–2000 results but the difference is less than the autumn case. The large change in model ozone trends ($-13.7\pm 16.6\%$ per decade to $+1.0\pm 11.0\%$ per decade) in this case is not associated with a significant change in NO₂ trends.

Table 2 gives the autumn and spring NO_y and HNO₃ vertical column trends derived from Arrival Heights model output. HNO₃ trends follow a similar pattern to the Lauder data (see Table 1), with a change from low values for 1982–2000 to higher values for the 2001–2019 period, the change in HNO₃ trends being mirrored by the changes in NO₂ trends.

A problem with this picture arises in the autumn data, where the predicted trend in NO_y is 5.4% per decade, approximately equal to the HNO₃ trend of 5.2% per decade. If NO_y and HNO₃ trends only were controlling the NO₂ trends, then the predicted NO₂ trend would also be expected to be close to 5.4% per decade, rather than approximately 1% per decade as predicted by the model.

It is clearly important to try and reduce the uncertainty associated with the regression model's estimates of linear trends from the Arrival Heights data. At this time, the large level of uncertainty means no significant conclusions can be drawn from the Arrival Heights data. A more thorough investigation of the sources of variability in the Arrival Heights data is required.

7 Conclusions

7.1 Lauder

Model slant column densities compare well with the measured 90° slant columns. Both the annual cycle and diurnal

Table 2. Autumn and spring trends in NO₂, N₂O, ozone, temperature, NO_y and HNO₃ derived from model results for Arrival Heights for the periods 1982–2000 and 2001–2019. Trends are given in % per decade (temperature trends are shown in K/decade). The quoted errors are $\pm 2\sigma$.

	Autumn				Spring			
	Model (1982–2000)		Model (2001–2019)		Model (1982–2000)		Model (2001–2019)	
	am	pm	am	pm	am	pm	am	pm
NO ₂	4.7±7.5	3.9±8.9	1.0±7.3	1.2±8.0	3.0±4.0	2.7±4.4	2.1±3.0	1.6±3.2
N ₂ O	2.6		2.6		2.6		2.6	
Ozone	−4.3±4.6		1.9±4.3		−13.7±12.4		1.0±11.0	
Temperature	0.3±0.5		0.0±0.72		−0.4±3.2		−0.5±4.5	
NO _y	2.3±5.0		5.4±7.6		4.6±4.0		3.0±7.6	
HNO ₃	1.2±3.1		5.2±4.1		−0.6±4.6		6.0±5.1	

cycle are reproduced by the model. Systematic underprediction of the measurements by the model is most likely due to an underprediction of the amount of NO_y. Although this affects the absolute values of the NO₂ slant columns, the percentage trends from the model are not directly affected by any mis-specification of the NO_y amounts.

Modeled and measured linear trends for the period 1980–2000 agree within the 2σ uncertainty range, for both the sunrise and sunset cases. Both the model and measurements show a greater rate of increase of NO₂ compared with N₂O suggesting a change in the partitioning of the oxides of nitrogen over Lauder. Model NO_y and HNO₃ trends over the same period confirm that, in the model, there is a change in partitioning of the NO_y family from the reservoir HNO₃ to the more chemically active NO_x species.

The largest discrepancy between modeled and measured NO₂ trends occurs in springtime when model NO₂ trends are lower than the trends derived from measurements for both sunrise and sunset cases. The annually varying trends still agree at the 2σ level. It is not clear at this time what is causing the model to underpredict the NO₂ trends during this time of the year. Polar vortex air mixing into mid-latitudes does not occur until early summer which rules out vortex air being the cause. Failure of the model to capture changes in chemistry or dynamics or both could explain the discrepancy in NO₂ trends.

Significant changes in the rate of increase of NO₂ (at a SZA of 90°) are predicted by the model for the period 2001–2019 compared with the period 1980–2000. Trends in NO₂ are greater than the trends in N₂O (and NO_y) for the period 1980–2000 with a reverse for the period 2001–2019. Changes in the NO₂ trends are associated with changes in HNO₃ trends, the ratios NO_x/NO_y and HNO₃/NO_y being anticorrelated.

Concurrent changes in ozone and stratospheric halogen trends are also found in the model results when comparing the 1980–2000 period with 2001–2019 but without additional sensitivity experiments it is not possible to fully deter-

mine the nature of the coupling between the NO_y partitioning and other variables within the model.

7.2 Arrival Heights

As was the case with the Lauder comparison, the model results at Arrival Heights reproduce well the measured seasonal and diurnal cycles of NO₂ slant column densities at sunrise and sunset. Autumn model results consistently underestimate the measured values. The most likely reason for this is an underestimation of the amount of NO_y in the model during this season.

Large variability in both the measured and modeled slant column densities result in large estimates in the errors of the trends in NO₂, ozone and temperature relative to the trends themselves. This makes interpretation of the trend results difficult and care must be taken in assigning significance to the conclusions drawn.

Autumn trend results for Arrival Heights are similar to the annually averaged results from Lauder. NO₂ increases at a faster rate than N₂O (and NO_y in the model) for the period 1980–2000, followed by a decrease in the NO₂ trends for the 2001–2019 period. Modeled HNO₃ indicates that the partitioning of NO_y follows a similar course to the Lauder results with an increase in the relative amount of NO_x at the expense of HNO₃ up to approximately the year 2000. This change is reversed in the model for the period 2001–2019.

For the chemically and dynamically active Antarctic spring, secular changes in NO_y partitioning are less pronounced. NO₂, HNO₃ and N₂O (NO_y) trends are all similar over the whole period of the integration. Denitrification of the polar stratosphere is an important process which occurs during the winter/spring period and is expected to significantly influence the trends in the oxides of nitrogen over Arrival Heights during spring. Currently the model uses a simple NAT/ice sedimentation scheme using a fixed sedimentation velocity as thus are not expected to capture the full detail of this process.

Results from both Lauder and Arrival Heights demonstrate that changes in NO₂ slant column densities derived from UMETRAC output follow the observed trends for the period 1981–2000. This suggests processes influencing the nitrogen chemistry within the model are reasonably well represented. Interestingly, the model prediction of the trends in NO₂ and HNO₃ for the period 2001–2019 show a reversal compared with the period 1981–2000. NO₂ trends are predicted to fall below the N₂O trends, associated with an increase in the HNO₃ trends indicating a shift in the partitioning of NO_y away from its active form at Lauder and Arrival Heights (autumn) over the coming two decades. Additional analysis of the model data is required to examine to what fraction of the change in NO₂ and HNO₃ trends can be attributed to circulation changes and what fraction is associated with chemical forcing.

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