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The role of wetlands in the hydrological cycle

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Abstract

It is widely accepted that wetlands have a significant influence on the hydrological cycle. Wetlands have therefore become important elements in water management policy at national, regional and international level. There are many examples where wetlands reduce floods, recharge groundwater or augment low flows. Less recognised are the many examples where wetlands increase floods, act as a barrier to recharge, or reduce low flows. This paper presents a database of 439 published statements on the water quantity functions of wetlands from 169 studies worldwide. This establishes a benchmark of the aggregated knowledge of wetland influences upon downstream river flows and groundwater aquifers. Emphasis is placed on hydrological functions relating to gross water balance, groundwater recharge, base flow and low flows, flood response and river flow variability. The functional statements are structured according to wetland hydrological type and the manner in which functional conclusions have been drawn. A synthesis of functional statements establishes the balance of scientific evidence for particular hydrological measures. The evidence reveals strong concurrence for some hydrological measures for certain wetland types. For other hydrological measures, there is diversity of functions for apparently similar wetlands. The balance of scientific evidence that emerges gives only limited support to the generalised model of flood control, recharge promotion and flow maintenance by wetlands portrayed throughout the 1990s as one component of the basis for wetland policy formulation. That support is confined largely to floodplain wetlands, while many other wetland types perform alternate functions — partly or fully. This paper provides the first step towards a more scientifically defensible functional assessment system.

Keywords: wetlands, hydrological functions, flood reduction, groundwater recharge, low flows, evaporation

Introduction

Open any book on wetland conservation and it will encourage the maintenance of wetlands partly because of their role in the water cycle. Wetlands are said to perform “hydrological functions”; to “act like a sponge”, soaking-up water during wet periods and releasing it during dry periods (eg. Bucher *et al.*, 1993). As Maltby (1991) reports “... the case for wetland conservation is made in terms of ecosystem functioning, which result in a wide range of values including groundwater recharge and discharge, flood flow alteration, sediment stabilization, water quality”. Since wetlands cover around 6% of the land surface of the earth (OECD, 1996) and many exist in the upstream parts of river catchments, the total downstream area over which a hydrological influence may be exerted is substantial. Yet the hydrological processes and behaviour of wetland ecosystems has certainly lacked the scientific integration received by other land surface systems, such as forests.

Kusler and Riexinger (1985) reported that “the science base and efforts to assimilate existing studies are still inadequate with regard to some functions, particularly hydrology”.

The basic references on the hydrological functions of wetlands are summaries of studies collated in the USA in the 1980s (Adamus and Stockwell, 1983; Bardecki, 1984; Carter, 1986). These summaries have been used by organisations, such as IUCN-The World Conservation Union (Dugan, 1990), Wetlands International (Davis and Claridge, 1993) and the Ramsar Convention on Wetlands of International Importance (Davis, 1993). They have influenced international wetland policy (OECD, 1996) and its uptake at the national (eg. Zimbabwe and Uganda), and continental levels e.g. Europe (CEC, 1995) and Asia (Howe *et al.*, 1992).

Recent emphasis at the Second World Water Forum in The Hague 2000 and the World Summit on Sustainable Development 2002 in Johannesburg was placed on the need

to ensure the integrity of ecosystems as part of integrated water resources management. Also receiving high prominence was the use of water to meet basic human needs and economic development. Thus, it is essential to re-examine, periodically, the conclusions of scientific studies on wetland functions. This ensures that policy at all levels is underpinned by a consensus of sound scientific opinion.

The scientific literature contains a range of studies that describe the water quantity functions of individual or groups of wetlands. They represent a substantial accumulation of hydrological knowledge. The majority of these papers supports the notion that wetlands have a significant influence on the hydrological cycle. However, many recognise that “it is difficult to make definitive statements regarding the role of various types of wetlands in runoff production or storm water detention” (Carter, 1986). Furthermore, some studies have produced evidence that contradicts previous widely accepted knowledge. For example, the classic hydrological studies of Hewlett and Hibbert (1967) identified headwater wetlands along river margins as flood generating areas. Burt (1995) concluded that “... most wetlands make very poor aquifers; ... accordingly, they yield little base flow, but in contrast, generate large quantities of flood runoff. Far from regulating river flow, wetlands usually provide a very flashy runoff regime”.

This paper has three objectives: first, to present an ordered database of published papers on hydrological functions of wetlands; second, to provide a collation of scientific evidence among hydrological measures and wetland type; third, to stimulate debate and further research. The focus of this paper is limited to water quantity functions, including impacts on water resource availability, groundwater replenishment and flood control. It does not consider other aspects of wetlands, such as water quality or biodiversity, which are part of a wider case for wetland conservation.

Creating a literature-based review of water quantity functions

With the objective of creating a comprehensive and consistent database of past studies, a literature review of water quantity functions was undertaken by keyword searches on the major databases of abstracts, and by tracking citations to earlier and related studies. Consequently, the database is drawn from 169 publications that report the results of scientific study that quantify hydrological functions of wetlands. Papers that report other authors' findings or give only qualitative descriptions of wetland process are not included.

Certain guidelines were followed, namely that:

- the review is restricted to freshwater wetlands, excluding lakes;
- conclusions of wetland function must be supported by hydrological data and not based on the original author's opinion alone;
- double-accounting is avoided, whereby repetition of conclusions for an individual wetland in successive publications is not duplicated;
- unsubstantiated generic statements, such as wetlands reduce flooding, are not included.

Consistency is ensured by extracting common elements from the diverse sources. Important information is maintained in the detail of the particular hydrological function, wetland type and the manner of conclusion. The approach adopted was to complete the following general statement (where bracketed and underlined phrases relate to elements in Annex 1) for each study:

“(Author(s)) undertook a study in a given location (country, or US State/Canadian Province or Territories) of a particular hydrological type of wetland (wetland type), also referred to by a more general or locally-specific wetland term (local term). Based on results from a particular type of study (categorisation of wetland study) and drawing inferences in a particular manner (basis of inference), the authors conclude (page number) that the wetland performs a particular function with respect to a specific hydrological measure (hydrological measure), as can be summarised by a functional statement (summary functional statement) and a summary function (summary of wetland water quantity function)”.

There are, therefore, ten elements extracted from each publication, each entered into the database. Explanation of each of these elements is expanded upon below. Because the format of the review is tabular, abbreviated codes are adopted for some elements for purposes of brevity.

Author(s): Citation to original source.

Country, or US State/Canadian Province or Territories: Location of wetland study.

Wetland type: For the purpose of this study, wetlands are categorised into five types according to three broad hydrological features (Table 1), based, with modification, upon the scheme proposed by Novitski (1978) for Wisconsin wetlands and subsequently applied by Adamus and Stockwell (1983). The three hydrological features are general catchment location, connectivity with the groundwater system and connectivity with the downstream channel network. General catchment location distinguishes

Table 1. Categorisation of wetland type by hydrological features

Type	Wetland type	Code	Features
HEADWATER	Surface water depression	SW/D	No hydraulic connectivity with groundwater. Outlet has no direct connectivity with river system
	Surface water slope	SW/S	No hydraulic connectivity with groundwater. Outlet has direct connectivity with river system
	Groundwater depression	GW/D	Hydraulic connectivity (permanent or periodic) with groundwater. Outlet has no direct connectivity with river system
	Groundwater slope	GW/S	Hydraulic connectivity (permanent or periodic) with groundwater. Outlet has direct connectivity with river system
FLOODPLAIN	Floodplain	FP	Inputs are dominantly upstream river flows
GENERAL			Wetland type, or one element of the type, cannot be specified

between headwater and floodplain; the distinction is that headwater wetlands are not fed by significant stream sources. Further subdivision applies only to headwater types. The connectivity with the groundwater system distinguishes ‘groundwater’ types that are in hydraulic connectivity with the groundwater system for all, or part of, the time, from ‘surface water’ types, which are not. Connectivity with the downstream channel network distinguishes ‘slope’ types, which are characterised by an outlet to the downstream river system, from ‘depression’ types, which are not. This categorisation deviates from that of Novitski and Adamus and Stockwell by including a floodplain type and in the ‘surface–slope’ type which, in that scheme, categorises lakeshore wetlands. Therefore, the two schemes are similar but are not directly comparable. An unspecified category is

added, and applied where the hydrological context of the wetland cannot be discerned.

Local term: Many local terms are applied to wetlands, including such general anglicised terms as ‘marsh’, ‘swamp’, ‘bog’ etc, and regionally specific terms such as dambo, pakihi, pocosin. There is no known means of providing a direct association between local terms and hydrological type in a fully inclusive manner.

Categorisation of wetland study: Wetland studies have adopted a number of experimental frameworks, ranging from intensive long-term monitoring of the water balance of wetland and non-wetland at the most complex extreme, to analyses based on single flood event hydrographs. Table 2

Table 2. Categorisation of methodological approach to wetland studies

Category of wetland study	Code	
Conceptual catchment model	CCM	Calibration and application of a conceptual catchment model
Water balance	WB	Quantification of the terms of the catchment and/or wetland water balance
Long-term hydrograph	LTH	Analysis of the characteristics of long time series of river flows
Single-event hydrograph	SEH	Analysis of the characteristics of a single river flow event
Trend analysis in time series	TS	Analysis of trends in hydrological time series (associated with detecting the impacts of drainage)
Component process	COMP _a	Investigation of an individual water balance component or hydrological process. (See Table 4 for definitions of ‘a’)
Chemical balance	CHEM	Quantification of a chemical process or chemical balance

presents categories of methodological approach with abbreviated codes for brevity in the tabular review.

Basis of inference: Many studies draw conclusions of the kind that, for example, wetlands reduce floods or augment

dry season river flows. This kind of conclusion, when taken out of context, leaves unanswered the basis for that conclusion, notably that the wetland reduces (or increases) river flows — compared with what? Table 3 presents a set of the comparative scenarios used amongst the various

Table 3. Basis for inference of wetland function

<i>Code</i>	<i>Baseline for inference</i>	<i>Methodology</i>	<i>Limitations</i>
With/out	Comparison of the same basin, with or without a wetland	This approach is restricted to catchment model simulations, in which the model is calibrated in either the “with” or “without” wetland case, as occurs in the modelled catchment. Model runs with the wetland case reversed generate simulated hydrological outputs. Differences between observed and simulated outputs are attributed to the presence of the wetland.	One case is simulated only. Stability of model parameters. Response of wetland replacement zone.
Drained	Comparison of the same wetland basin before and after drainage; or neighbouring drained and undrained wetlands	Hydrological outputs from a wetland are observed prior to drainage. The wetland is drained, and outputs are observed after drainage. Differences in the pre- and post- drainage outputs are attributed to the wetland. alternately, outputs from two adjacent catchments, each with wetlands, are observed. One of the catchments is drained, and differences between the outputs between the two catchments are attributed to the presence of the wetland.	Response of replacement land-use in drained wetland. Initial short-term responses to drainage may differ from long-term responses.
Pair	Comparison of paired basins, one with wetland the other without	Hydrological outputs are observed from two catchments, similar in all respects except that one that one catchment contains a wetland while the second does not. Differences between the outputs are attributed to the presence of the wetland.	If two catchments are otherwise identical, why does only one have a wetland?
Multiple	Comparison of several basins with varying proportions of wetland	Hydrological outputs are observed from several catchments, each containing different proportions of wetland. Differences between the outputs are attributed to greater or lesser presence of wetland.	Variability in non-wetland characteristics amongst several catchments
In-out	Comparison of inflows and outflows of a wetland system	Hydrological inputs and outputs associated with a single wetland are measured. Differences between inputs and outputs are attributed to the presence of the wetland.	
Same	Comparison of wetland hydrological response with response elsewhere in the same basin	Hydrological outputs from a single wetland are measured as well as outputs from other non-wetland portions of the same catchment. Differences between the outputs are attributed to the wetland.	
Comp _a	Conclusions relating to individual components of the wetland through the development of an understanding of a component process	Individual component processes are observed and understood within a single wetland. The understanding of the processes is the basis for inferring the influence of that process on the hydrological output. (Substituting for _a : T = topography, V = vegetation, S = Soil, WC = water content, GW = groundwater, ET = evaporation)	Extrapolation of a single process in isolation. Processes may not be homogeneous across the entire wetland.

publications as the basis for inferring wetland function. Each basis has some limitations, and these are summarised. Again, abbreviated codes are set out.

Page number: Page number in the original publication on which the conclusion is drawn.

Hydrological measure: There are many different measures in hydrology to describe and define aspects of the flow regime. While non-hydrologists might refer generically to floods, the hydrologist would be concerned with measures such as the magnitude of the peak flow during the flood event, the volume of runoff contained in the event, and the time-to-peak. Published studies in wetland hydrology are not consistent in their attention to different measures; it is possible to find one study analysing the return period of flood peaks extracted from a 20 or 30 year flow record, and another analysing the flood volume of a single event, with both drawing conclusions on wetland influences on floods.

Table 4 presents and defines different hydrological measures within five broad groupings of hydrological response, namely; gross water balance, groundwater recharge, base flow and low flows, flood response and river flow variability, including some seasonal variations.

Summary of wetland water quantity function: Conclusions regarding water quantity functions extracted directly, or in paraphrased form, from the original text are presented.

Summary functional statement for hydrological measure: Functional statements of the form 'wetlands increase low flows' are expressed as the sign of the wetland influence upon the hydrological measure; thus '+' indicates an increasing influence upon the hydrological measure, '-' indicates a reducing influence and '.' represents a neutral influence (i.e. no significant influence exists or can be detected). In the case of groundwater recharge and groundwater discharge sites, there is interest in the conservation-based literature whether either of these functions is, or is not, present in a wetland. Therefore, '=' indicates that this function is present and 'x' indicates that it is not.

Global data base of wetland water quantity functions

The first objective of this paper is to redress the deficiency in availability of hydrological information on wetland functions by providing an accessible and consistent database

of past studies. Annex 1 presents the product of the application of the global review. The database is composed of 169 different published studies with 440 functional statements, representing the fullest sample of studies that could be traced, conforming to the principles adopted. It would not be claimed that the sample is exhaustive, but it is considered to be comprehensive.

BALANCE OF SCIENTIFIC EVIDENCE FOR PARTICULAR HYDROLOGICAL MEASURES

Table 5(a to e) collates the number of functional statements for each wetland type for the five principal groups of water quantity measures. For example, interpreting Table 5a for floodplain-type wetlands, two studies conclude that an example of this wetland type increases mean annual flow, two studies concludes that no significant influence can be detected and eight studies conclude that examples of this wetland type reduce mean annual flow. Total numbers of functional statements are presented across all hydrological measures and all wetland types.

Analysis of the "balance of scientific evidence" draws on comparison of the number of papers that conclude a particular function. This is seen to be an important step and a precursor to more detailed exploration of the evidence for particular measures for specific wetland types. The results cannot yet be considered to reflect a "balance of scientific opinion", because there has been no inter-comparison amongst the different studies.

There are some cautionary perspectives and some limitations on the comparison that must be stated.

- (1) The number of papers reporting a particular influence on the water cycle does not necessarily indicate the total picture. Some hydrological functions have been studied more than others.
- (2) The number of functional statements cannot be interpreted as the number of wetlands performing a function; for example, a functional statement based on multiple catchments commonly involves a large number of individual wetlands.
- (3) There is no certainty that the 169 publications represent all past studies of wetlands. Although not exhaustive, the sampling method has been applied independently of any initial bias associated with policy interests. However, it cannot be discounted that there is potential bias in the wetlands that were selected for study by the original authors.
- (4) The distribution of wetland types within the sample of reviewed studies does not represent the distribution of wetland types worldwide; this is particularly true given

Table 4. Hydrological measures and their definition

<i>CODE</i>	<i>Hydrological measure</i>	<i>Definition</i>
GROSS WATER BALANCE		
MAF	Mean annual flow	Volume (or rate) of river flow during a year (on average)
MAAE	Mean annual actual evaporation	Volume (or rate) of evaporation during a year (on average)
WPF	Wet period flow	Volume (or rate) of river flow during, or in response to, periods of rainfall
WPAE	Wet period actual evaporation	Volume (or rate) of evaporation during, or in response to, periods of rainfall
DPAE	Dry period actual evaporation	Volume (or rate) of evaporation during periods without rainfall
	Dry period flow	See 'DPFV' - Dry period flow volume
GROUNDWATER RECHARGE		
AGR	Annual groundwater recharge system during a year	Volume of water moving vertically from the wetland into the underlying groundwater
WPGR	Wet period groundwater recharge system during, or in response to, periods of rainfall	Volume of water moving vertically from the wetland into the underlying groundwater
DPGR	Dry period groundwater recharge system during periods without rainfall	Volume of water moving vertically from the wetland into the underlying groundwater
BASE FLOW AND LOW FLOWS		
GDS	Groundwater discharge site	Movement of groundwater into surface water through the wetland
DPFV	Dry period flow volume	Volume of flow during dry periods
DPFD	Dry period flow duration	Duration for which flow is sustained during dry periods
DPRR	Dry period recession rate in flow, a low rate indicates sustained low flows	Rate of flow recession during periods without rain; a high rate indicates a rapid decrease
FLOOD RESPONSE		
FPLM (T < 2 years)	Flood peak of low return period every two years or less	Peak flow during a flood event, where the flow is exceeded on average
FPHM (T > 2 years)	Flood peak of high return period every two years or more	Peak flow during a flood event, where the flow is exceeded on average
FEV	Flood event volume	Total volume of flow during an individual flood event
FTTP	Flood time-to-peak	Time between the onset and peak of a flood event
FRR	Flood recession rate	The recession rate of a flood event
RIVER FLOW VARIABILITY		
FVa	Flow variability	Flow variability within the full flow regime
WPFVa	Wet period flow variability	Flow variability during, or in response to, periods with rainfall
DPFVa	Dry period flow variability	Flow variability during periods without rainfall

the focus of studies on North America. Consequently, one cannot necessarily transfer the results of this study to the general grouping of 'worldwide wetlands'.

- (5) Conclusions are presented as stated in the original paper; no attempt is made to evaluate or uphold the conclusions that are drawn. Further work in the critical evaluation of past studies would represent a valuable contribution

to the science of wetland hydrology.

- (6) No attempt has been made at cross-correlation of hydrological measures within single studies. For example, if a study has concluded that a particular wetland increases evaporation, no conclusion is drawn that flow is reduced, unless that explicit statement is made by the author.

Table 5

a Gross water balance

			FP	SW/D	SW/S	GW/D	GW/S	General	Total
Mean annual flow	MAF	+		2				2	4
		-		8	1	5	1	7	23
		.		2	1	4		8	16
Mean annual actual evaporation	MAAE	+	12	1	4	1	6	1	25
		-				1	2		3
		.		1	1	1			3
Wet period flow	WPF	+		1		3		5	6
		-		1			2	2	5
		.		1	1		1	1	4
Wet period actual evaporation	WPAE	+							1
		-				1	1		2
		.							0
Dry period actual evaporation	DPAE	+	4		3		15		22
		-					1		1
		.							0
			3	0	4	1	11	6	25
			25	2	12	2	30	5	76
			4	3	5	1	9	1	23

b. Groundwater recharge

			FP	SW/D	SW/S	GW/D	GW/S	General	Total	
Annual groundwater recharge	AGR	+		1			1	2	1	5
		-		1		2			4	7
		.				2	1	1		4
		X		1		5	1	7	2	16
Wet period groundwater recharge	WPGR	=	8	4	2	4	6	2	26	
		+		1					1	
		-							0	
		.							0	
Dry period groundwater recharge	DPGR	X		1					1	
		=	1	1		1	1		4	
		+							0	
		-		1	1				2	
		.							0	
		X		1					1	
		=		1	1	1			2	
		increased recharge	1	1	0	1	2	1	6	
decreased recharge	1	1	3	0	0	4	9			
not increased or decreased	0	0	2	1	1	0	4			
recharge does not occur	1	2	5	1	7	2	18			
recharge occurs	9	6	3	5	7	2	32			

c. Base flow and low flow

			FP	SW/D	SW/S	GW/D	GW/S	General	Total
Groundwater discharge site	GDS	=		2	4		3	18	27
		X			2	2	1		5
Dry period flow volume	DPFV	+	3	1	1		6	3	14
		-	5		11	1	22	8	47
		.	1		4	1	2	2	10
Dry period flow duration	DPFD	+							0
		-			1		1		2
		.							0
Dry period recession rate	DPRR	+					1	1	2
		-							0
		.							0
			3	1	1	0	7	4	16
			5	0	12	1	23	8	49
			1	0	4	1	2	2	10
			0	2	2	0	1	0	5
			2	4	0	3	18	0	27

d. Flood response

			FP	SW/D	SW/S	GW/D	GW/S	General	Total
Floodpeak low magnitude (T>5 yrs)	FPLM	+	1		7		2	1	11
		-	12	1	7	1	10	7	38
		.	2		3		2	2	9
Floodpeak high magnitude (T<5 y)	FPHM	+					1	1	1
		-	4	1		1	1	1	8
		.			1		2	4	7
Flood event volume	FEV	+	1		3		6	1	11
		-	4		1		3	1	9
		.							0
Flood time to peak	FTTP	+	3		2		1		6
		-	1						1
		.			1				1
Flood recession rate	FRR	+			1				1
		-			5		3		8
		.							0
floods increased or advanced or recession reduced			3	0	15	0	12	2	32
floods reduced or delayed or recession increased			23	2	11	2	15	9	62
not increase, reduced, delayed or advanced			2	0	5	0	4	6	17

e. River flow variability

			FP	SW/D	SW/S	GW/D	GW/S	General	Total
Flow variability	FVa	+			5		4	1	10
		-	6	1	2		0	1	10
		.	1		3		2		6
Wet period flow variability	WPFVa	+			1		1		2
		-							0
		.							0
Dry period flow variability	DPFVa	+							0
		-							0
		.							0
flow variability increased			0	0	6	0	5	1	12
flow variability decreased			6	1	2	0	0	1	10
not increased or decreased			1	0	3	0	2	0	6

The association between hydrological types and local terms is presented in Table 6. It is immediately evident that there is no strong linkage between hydrological categorisation as applied in this paper and the use of local or ecological wetland terms; the terms peat, bog, marsh (and several others) recur in different hydrological types. Thus, grouping by hydrological type is seen as more meaningful than grouping by local terms. For example, the term ‘bog’ can be found in all five hydrological types.

From a hydrological perspective, the content of the database may be perceived as limited due to its emphasis on functions rather than hydrological processes — given that the concept of functions is not well-established within the hydrological community. However, while more process information can be extracted from the set of publications, the target of this paper is the use of functional generalisations to represent wetland hydrology to the wetland management and policy arena. Clearly, there is a strong case for bringing hydrological processes and function closer together.

Geographically, the dataset is dominated by 92 studies

from North America (including 23 different U.S. States and six Canadian Provinces/Territories), with additionally 33 studies from 14 countries in Europe, 27 studies from 10 countries in Africa and 17 from elsewhere (including New Zealand (2), Australia, Brazil (3), India, Indonesia and Malaysia). This distribution reflects the substantial investment in scientific enquiry in North America compared with other regions of the world and a relative dearth of accessible information relating to Asian and South American wetland hydrology.

The term ‘wetland’ embraces a wide variety of land types, from springs to large inland deltas. As a result, a lack of consistency in the impact of wetlands on the water cycle was anticipated. Unique conclusions concerning any specific hydrological function cannot be drawn for all wetlands. Taking flow variability as a single hydrological measure, for example, there are 28 statements with good geographical coverage, of which 10 show that variability is increased by wetlands, 11 that variability is reduced, and 6 that wetland influence is neutral. When wetlands are sub-divided into

Table 6. Wetland hydrological types and local terms

Type	Code	Local terms
Floodplain	FP	Sudd, river valley, swamp, floodplain, valley bog, marsh, cypress swamp, delta (inland), dambo, marsh, inland valley swamps and bolis, pantanal
Surface water depression	SW/D	Bog, prairie pothole, slough, kettle-hole, shallow depressions
Surface water slope	SW/S	Perched bog, peat, peat bog, pond, raised bog, blanket peat, muck, pocosin, saturated heath, palsa, aapa-mire, infilled lakes, constructed wetland
Groundwater depression	GW/D	Marsh, prairie pothole, fen, pond, bog
Groundwater slope	GW/S	Swamp, marsh, river valley, bog, dambo, fen bogs, peat bog, peat, wetland, pocosin, macrophytes, cypress domes, peat fen, pakahi, fen, mire, blanket bog, quaking fen, pond, domed bog and fen
General		Swamp, bog, marsh, 'basin storage', ponds, hydromorphic gleys, forest wetland

hydrologically similar types, greater consistency of conclusion emerges — six of seven conclusions from studies of floodplains conclude that flow variability is reduced by wetlands (including the Sudd and Okavango in Africa and Barito in Indonesia). But ambiguity still exists: amongst 19 studies of headwater wetlands (all from USA and Europe), 10 studies conclude that flow variability is increased; 5 that the wetland influence is neutral; and 4 that variability is reduced. Even apparently similar wetlands perform functions that are seemingly in opposition (e.g. peat bogs occur in each of the three categories; increasing, decreasing and not-affecting flow variability). But unanimity of function is not anticipated — there is no prior assumption that all wetlands of a particular type perform the same function. This study has not yet investigated the detailed climatology, catchment conditions and internal wetland structure, any of which can mean a particular wetland will perform differently from other wetlands that are otherwise similar.

Whether hydrological functions of wetlands are considered to be beneficial or not depends upon one's point of view. For example, ecologists may see evaporation from wetlands as an essential process supporting plant growth, whilst water resource managers may see it as a loss of a vital downstream resource. Those living in flood-prone areas downstream of wetlands that generate floods may view them negatively while those living downstream of wetlands that reduce floods may not view them in the same light. Ecologists see floods as essential elements of the river flow regime maintaining channel structure through sediment transport, and interactions between the river and its floodplain that drives nutrient exchange and breeding cycles (Junk *et al.*, 1989, Poff *et al.*, 1997).

The main conclusions of the analysis are as follows.

1. Wetlands are significant in altering the water cycle. The 169 scientific studies published during the period 1930–2002 (as traced by this paper) provide 439 statements on the hydrological significance of wetlands. Of the 439 statements, only 83 (19%) conclude the wetland influence on the water cycle to be neutral or insignificant. The vast majority conclude that wetlands either increase or decrease a particular component of the water cycle. It is this evidence that has led to the notion that wetlands perform hydrological functions. Since wetlands cover approximately 6% of the world's land area, with many linked directly to rivers and aquifers, this is an issue of importance to water management.
2. There are some significant generalisations that emerge from the published hydrological evidence. These are different from the long-standing generalisation that wetlands always reduce floods, promote groundwater recharge and regulate river flows. Most, but not all, studies (23 of 28) show that floodplain wetlands reduce or delay floods, with examples from all regions of the world. This same influence on floods is also seen, but less conclusively (30 of 66) for wetlands in the headwaters of river systems (e.g. bogs and river margins). A substantial number (27 of 66) of headwater wetlands increases flood peaks. These studies were mostly from Europe, but included work from West Africa and Southern Africa. Around half of the statements (11 of 20) for flood event volumes and 8 of 13 for wet period flows) show that headwater wetlands increase the immediate response of rivers to rainfall,

generating higher volumes of flood flow, even if the flood peak is not increased. The coverage of these studies is world-wide including Africa and South America. This function occurs because headwater wetlands tend to be saturated and convey rainfall rapidly to the river; thus they are a principal mechanism for generation of flood flows.

3. There is strong evidence that wetlands evaporate more water than other land types, such as forests, savannah grassland or arable land. Two thirds of studies (48 of 74) conclude that wetlands increase average annual evaporation or reduce average annual river flow. About 10% of studies (7) conclude the opposite; for example some woodlands in Zambia had greater evaporation than the adjacent wetlands. The remaining 25% are neutral. There is no obvious distinction amongst different wetland sub-types or geographical regions of the world.
4. Two-thirds of studies (47 of 71) conclude that wetlands reduce the flow of water in downstream rivers during dry periods. Evidence is mainly from North America and Europe, but includes floodplains in Sierra Leone and wetlands in Southern Africa. This is backed by overwhelming evidence (22 of 23 studies) that shows evaporation from wetlands to be higher than from non-wetland portions of the catchment during dry periods. There is no discernible difference for different wetland sub-types. In 20% of cases, wetlands increase river flows during the dry season.
5. Many wetlands exist because they overlie impermeable soils or rocks and there is little interaction with groundwater. The database contains 69 statements referring to groundwater recharge; 32 conclude merely that recharge takes place, and 18 conclude there is no recharge. There are similar numbers of studies that report wetlands either to recharge more (6) or less than (9) other land types. Some wetlands, such as floodplains in India and West Africa on sandy soils, recharge groundwater when flooded. Many wetlands have formed at springs and are fed by groundwater. The direction of water movement between the wetland and the ground may change in the same wetland, such as in some peatlands in Madagascar, according to hydrological conditions.
6. Conclusions have been drawn above on flow variability. The over-riding picture appears to be a reduction in flood peaks by floodplains. In some cases, such as many headwater wetlands, increasing flood flows combines with decreasing dry season flows to widen the overall range of flows.

Implications of the results for wetland research and policy formulation

This paper confirms that wetlands exert a strong influence on the hydrological cycle. It strengthens the view that management of wetlands must be an important part of integrated water resources and flood management of all river basins. Where wetlands reduce floods, recharge groundwater and increase dry season flows, wetland hydrology is working in sympathy with water resources managers and flood defence engineers. Where wetlands have high evaporation demands or generate flood-runoff, they may create or exacerbate water management problems. Whatever the hydrological functions they perform, decisions on wetland conservation will inevitably be taken in a wider context and will also depend on water scarcity and on other functions, such as human health, fisheries, navigation, recreation, cultural heritage and biodiversity.

Successful water management requires knowledge of the extent to which wetlands are performing different hydrological functions. Since it is not feasible to study every wetland in detail, rapid assessment methods are required to identify likely functions. Furthermore, a major objective of this work is to stimulate discussion on hydrological functions. It is relatively easy to add to the database, either within additional previously published work or through new research. It is harder to account for variations in functions. For it is clear that there is no simple relationship between wetland types and the hydrological functions they perform. Part of the problem stems from the lack of a simple classification of wetlands that consistently relates hydrology, vegetation, substrate type and geomorphology. It is unlikely that any sophisticated classification scheme would be able to explain the variation of function in evidence.

Various methods have been developed in a number of countries around the world to assess hydrological functions of wetlands. Some are merely classification systems that group wetlands according to botanical, geomorphological and/or water regime characteristics (e.g. Cowardin *et al.*, 1979; Brinson, 1993; Gilvear and McInnes, 1994; Wheeler, 1984). Other methods give each wetland a grade for a function, such as high medium or low (e.g. Adamus *et al.*, 1987) or produce a quantitative estimate of performance of a given function (e.g. Amman and Stone, 1991; Hruby *et al.*, 1995). Maltby *et al.* (1996) have developed a framework of functional analysis through characterisation of distinct ecosystem/landscape units (termed hydrogeomorphic units). The objective is to provide a simple and rapid procedure, but the system still needs to be operationalised. In addition, guides have been produced for extending the functional assessment to produce an economic value for the functions

(Lipton *et al.*, 1995, Barbier *et al.*, 1995). Recent work in the UK (Wheeler and Shaw, 2000) used data from over 80 wetlands in Eastern England to develop a classification and assessment system called WETMECS that combines landscape situation, water supply mechanism, hydrotopographical elements, acidity (base-richness) and fertility. The outcome of this study is that apparently similar wetlands are driven by very different hydrological processes; almost invariably, some data need to be collected at a site to identify its functional role.

Consequently, generalised and simplified statements of wetland function are discouraged because they demonstrably have little practical value. As a minimum, this paper encourages the future representation of diversity and complexity amongst wetlands and the portrayal of diverse hydrological functions.

Hydrologists must be more imaginative and proactive in contributing scientific knowledge to underpin policy formulation and management decisions. For the hydrological community, a new challenge is set for wetland policy.

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References

- Ackroyd, E.A., Walton W.C. and Hills D.L., 1967. *Groundwater contribution to streamflow and its relation to basin characteristics in Minnesota*. Report of Investigations No.6, University of Minnesota, USA.
- Adamus, P.R. and Stockwell, L.T., 1983. *A method for wetland functional assessment: Volume I, Critical review and evaluation concepts*. Federal Highway Agency FHWA-IP-82-23, U.S. Dept. of Transportation, Washington, USA. 176pp.
- Adamus, P.R. *et al.*, 1987. *Wetland evaluation technique (WET) volume II: Methodology*. US Army Corp. Engineers, Vicksburg, Mississippi State.
- Amman, A. and Stone, A.L., 1991. *Method for comparative evaluation of non-tidal wetlands in New Hampshire*. New Hampshire Department of Environmental Services. NHDES-WRD-1991-3
- Baden, W. and Eggelsmann, R., 1986. *The hydrologic budget of the Highbogs in the Atlantic region*. Proc. Third International Peat Congress, Dept. Energ. Mines & Resourc.; Nat. Research Council, Ottawa, Ontario, Canada, 206–211.
- Balek, J. and Perry J.E., 1973. Hydrology of seasonally inundated African headwater swamps. *J. Hydrol.*, **19**, 227–249.
- Balek, J., 1977. Chapter 7: Swamps. In: *Hydrology and water resources in tropical Africa*, A. Millington (Ed) Elsevier, Amsterdam, The Netherlands. 154–168.
- Barbier, E.B., Acreman, M.C. and Knowler, D., 1997. *Economic valuation of wetlands: a guide for policy makers and planners*. Ramsar Convention, Gland, Switzerland.
- Bardecki, M., 1984. What value wetlands? *J. Soil Water Conserv.*, 166–169.
- Bardecki, M., 1987. Cumulative impacts of agricultural land drainage on watershed hydrology. Proc. Nat. Wetland Symp. on 'Wetland Hydrology', Association of Wetland Managers Inc, Chicago, USA. 124–129.
- Bavina, L.G. 1975. Water balance of swamps in the forest zone of the European part of the USSR. In: *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences. UNESCO/IAHS, Paris. 297–304.
- Bay, R.R., 1967. Groundwater and vegetation in two peat bogs in Northern Minnesota. *Ecology*, **48**, 308–310.
- Bay, R.R., 1969. Runoff from small peat watersheds. *J. Hydrol.*, **9**, 90–102.
- Bedinger, M.S., 1981. Hydrology of bottomland hardwood forests of the Mississippi Embayment. In: *Wetlands of bottomland hardwood forests*, J.R.Clark and J. Benforado (Eds.), Elsevier, The Netherlands. 161–176.
- Boelter, D.H. and Verry, E. S., 1977. *Peatland and water in the Northern Lake States*. General Tech. Rep. NC-31, US Dept Agricultural Forestry Service, North Central Experimental Station, St. Paul Minnesota, USA. 22pp.
- Boeye, D. and Verheyen, R.F., 1992. The hydrological balance of a groundwater discharge fen. *J. Hydrol.*, **137**, 149–163.
- Brandesten, C.O., 1988. *Seasonal variation in streamflow recessions in the mire complex Komosse, Southern Central Sweden*. Symposium on the Hydrology of wetlands in temperate and cold regions: Vol 1. Academy of Finland, Helsinki. 84–91.
- Brinson, M.M., 1993. *A hydro-geomorphic classification for wetlands*. Technical Report WRP-DF-4. US Army Engineers Waterways Experimental Station, Vicksburg, Mississippi, USA.
- Brooks, K.N. and Kreft, D.R., 1989. A hydrologic model for Minnesota Peatlands. *J. Minnesota Acad. Sci.*, **55**, 113–119.
- Brown, S.L., 1990. Structure and dynamics of basin forested wetlands in North America. In: *Forested wetland ecosystems of the world*. A.E. Lugo, M.M. Brinso and S.L. Brown (Eds.) Elsevier, Amsterdam, The Netherlands. 171–199.
- Brun, L.J., Richardson, J.L., Enz, J.W. and Larsen, J.K., 1981. Streamflow changes in the Southern Red River Valley of North Dakota. *North Dakota Farm Research*, **38**, 11–14.
- Bucher, E.H., Bonetto, A., Boyle T., Canevari P., Castro G, Huszar P. and Stone, T., 1993. *Hidrovia- an initial environmental examination of the Paraguay - Parana waterway*. Wetlands for the Americas Publication No. 10, Manomet, MA, USA, 72pp.
- Bulavko A.G. and Drozd V.V., 1975. Bog reclamation and its effect on the water balance of river basins. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 461–468.
- Bullock, A., 1992a. The role of dambos in determining river flow regimes in Zimbabwe. *J. Hydrol.*, **134**, 349–372.
- Bullock, A., 1992b. Dambo hydrology in Southern Africa: review and reassessment. *J. Hydrol.*, **134**, 373–396.
- Burke, W., 1968. *Drainage of blanket peat at Glenamoy*. Proc. 2nd International Peat Congress. HMSO, Edinburgh, UK. 809–817.

- Burke, W., 1975. Aspects of the hydrology of blanket peat in Ireland. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 171–182.
- Burt, T.P., 1995. The role of wetlands in runoff generation from headwater catchments. In: *Hydrology and hydrochemistry of British wetlands*, J. Hughes and L. Heathwaite (Eds.) Wiley, Chichester, UK. 21–38
- Campbell, R.E. and Drecher, F.C., 1970. *A proposed streamflow data program for Wisconsin*: Madison, Wisconsin, U.S. Geol. Surv. Open File Rep., 55pp.
- Carter, V., 1986. An overview of the hydrologic concerns related to wetlands in the United States. *Can. J. Bot.*, **64**, 364–374.
- CEC, 1995. *Wise use and conservation of wetlands*. Communication from the Commission to the Council and the European Parliament, Commission of the European Communities COM(95) 189 (final), 54pp.
- Choi, J. and Harvey, J.W., 2000. Quantifying time-varying groundwater discharge and recharge in wetlands of the Northern Florida Everglades. *Wetlands*, **20**, 500–511.
- Conger, S., 1971. *Estimating magnitude and frequency of floods in Wisconsin*: Madison, Wisconsin, U.S. Geological Survey open-file report, 200pp.
- Cowardine, L.M. et al., 1979. *Classification of wetlands and deepwater habitats of the United States*. US Fish and Wildlife Service FWS/OBS-79/31. 103 pp.
- Daniel, C., 1981. Hydrology, geology and soils of pocosins: a comparison of natural and altered systems. In: *Pocosin Wetlands*, C. J. Richardson (Ed.). Hutchinson Ross Publ. Co., Straudsbury, 109–124.
- Darmer, K.I., 1970. *A proposed streamflow data program for New York*: Albany, New York, U.S. Geological Survey open-file report. 29pp.
- Davies, J., Claridge C.F., 1993. *Wetland benefits: the potential for wetlands to support and maintain development*. Asian Wetland Bureau Publication No. 87: IWRB Spec. Publ. 27: Wetlands for the America Publication No. 11. Asian Wetland Bureau, Kuala Lumpur, Malaysia, 45pp.
- Davis, T.J., 1993. *Towards the wise use of wetlands*. Ramsar Convention Bureau, Gland Switzerland, 180pp.
- Demissie, M., Khan A. and al-Mubarak, R., 1991. *Influence of wetlands on streamflow in Illinois*. Proc. ASCE Nat. Conf. on Hydraulic Engineering, 949–954.
- Demissie, M. and Khan A., 1991. *Wetland drainage and streamflow trends in Illinois*. Proc. ASCE Nat. Conf. on Hydraulic Engineering, 1050–1054.
- Doss, P.K., 1995. Physical-hydrogeologic processes in wetlands. *Nat. Area. J.*, **15**, 216–226.
- Doyle, A.F., 1987. *The Charles River Watershed: a dual approach to floodplain management*. Proc. Nat. Wetland Symp on Wetland Hydrology. Assoc. of State Wetland Managers Inc., Chicago 110–112.
- Drayton, R.S., Kidd, C.H.R., Mandeville, A.N. and Miller, J.B., 1980. A regional analysis of river floods and low flows in Malawi. *Institute of Hydrology Report No. 72*, Wallingford, UK
- Dubreuil, P.L., 1985. Review of field observations of runoff generation in the tropics. *J. Hydrol.*, **80**, 237–264.
- Dugan, P. J., 1990. *Wetland conservation - a review of current issues and required action*. IUCN - The World Conservation Union, Gland, Switzerland, 96pp.
- Eisenlohr, W.S., 1972. Hydrologic investigations of prairie potholes in North Dakota, 1959–1968. *Geological Survey Professional Paper 585-A*, US Govt. Printing Office, Washington, 102 pp.
- Eisenlohr, W.S., 1975. Hydrology of marshy ponds on the coteau du Missouri. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 305–312.
- Egglesmann, R., 1975. The water balance of lowland areas in north-western regions of the FRG. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 355–368.
- Eggelsmann, R., Heathwaite, A.L., Grosse-Brauckmann, G., Kuster, E., Naucke, W., Schuch, M. and Schweickle, 1993. Physical processes and properties of mires. In: *Mires: process, exploitation and conservation*, A.L. Heathwaite and Kh. Gottlich (Eds.). Wiley, Chichester, UK. 171–262.
- Faulkner, R.D. and Lambert, R.A., 1991. The effect of dambo irrigation on dambo hydrology: a case study. *J. Hydrol.*, **123**, 147–161.
- Ferrari, M.R., Miller, J.R. and Russell, G.L., 1999. Modeling the effect of wetlands, flooding and irrigation on river flow: application to the Aral Sea. *Water Resour. Res.*, **35**, 1869–1876.
- Flippo, H.N. Jr., 1977. Floods in Pennsylvania: Pennsylvania Department of Environmental Resources, *Water Resour. Bull.*, **13**, 59.
- Ford, J. and Bedford, B.L., 1987. The hydrology of Alaskan wetlands, U.S.A.; a review. *Arctic Alp. Res.*, **19**, 209–229.
- Forest, W.E. and Walker, P.N., 1970. *A proposed streamflow data program for Maryland and Delaware*, Towson, Maryland, U.S. Geol. Surv. Open File Rep. 41pp.
- Freeze, R.A., 1969. *Hydrology of the Good Spirit Lake Drainage Basin, Saskatchewan - a preliminary analysis*. Inland Waters Branch Tech. Bull. No.14, Dept. of Energy, Mines and Resources, Canada.
- Gehrels J. and Mulamootil, G., 1990. Hydrologic process in a southern Ontario wetland. *Hydrobiologia*, **208**, 221–234.
- Genereux, D. and Slater, E., 1999. Water exchange between canals and surrounding aquifer and wetlands in Southern Everglades, USA. *J. Hydrol.*, **219**, 153–168.
- Gerla, P.J., 1999. Estimating the groundwater contribution in wetlands using modelling and digital terrain analysis. *Wetlands*, **19**, 394–402.
- Gerla, P.J. and Matheny, R.K., 1996. Seasonal variability and simulation of groundwater flow in a prairie wetland. *Hydrol. Process.*, **10**, 903–920.
- Gibson, J.J., Edwards, T.W.D. and Prowse, T.D., 1993. Runoff generation in a high boreal wetland in northern Canada. *Nord. Hydrol.*, **24**, 213–224.
- Gilvear, D.J., Andrews, R., Tellam, J.H., Lloyd, J.W. and Lerner, D.N., 1993. Quantification of the water balance and hydrogeological processes in the vicinity of a small groundwater-fed wetland, East Anglia, U.K.. *J. Hydrol.*, **144**, 311–334.
- Gilvear, D. and McInnes, R.J., 1994. Wetland hydrological vulnerability and the use of classification procedures: a Scottish case study. *J. Environ. Manage.*, **42**, 403–414.
- Gilliam, J.W. and Skaggs, R.W., 1981. Drainage and agricultural development: effects on drainage waters. In: *Pocosin Wetlands*, C. J. Richardson (Ed.), Hutchinson Ross Publ. Co., Straudsbury, 109–124.
- Glazacheva, L.I., 1975. The effect of reclamation on rivers and lakes in Latvia. In *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 513–518.
- Gonthier, G.J., 1996. Groundwater flow conditions within a bottomland hardwood wetland, Eastern Arkansas. *Wetlands*, **16**, 334–346.

- Grillot, J.C. and Dussarrat, B., 1992. Hydraulique des unites d'interfluves et de bas-fond tourbeux: un exemple en zone de socle altere (Madagascar). *J. Hydrol.*, **135**, 321–340.
- Gurnell, A.M. and Gregory, K.J., 1986. Water table level and contributing area: the generation of runoff in a heathland catchment. In: *Conjunctive Water Use*, S.M. Gorelick (Ed.), IAHS Publ. no. 156, 87–95.
- Hall, F.R., Rutherford, R.J. and Byers, G.E., 1972. *The influence of a New England wetland on water quantity and quality*. Res. Rep. 4, Wat. Resour. Res. Centre, New Hampshire University, Durham, NH, USA.
- Hamilton, S.K., Sippel, S.J., Calheiros, D.F. and Melack, J.M., 1997. An anoxic event and other biogeochemical effects of the Pantanal wetland on the Paraguay River. *Limnol. Oceanogr.*, **42**, 257–272.
- Hardy, R.J., Bates, P.D. and Anderson, M.G., 2000. Modelling suspended sediment deposition on a fluvial floodplain using a two-dimensional dynamic finite element model. *J. Hydrol.*, **229**, 202–218.
- Hayashi, M., van der Kamp, G. and Rudolph, D.L., 1998a. Water and solute transfer between a prairie wetland and adjacent uplands. 1. Water balance. *J. Hydrol.*, **207**, 42–55.
- Hayashi, M., van der Kamp, G. and Rudolph, D.L., 1998b. Water and solute transfer between a prairie wetland and adjacent uplands. 2. Chloride balance. *J. Hydrol.*, **207**, 56–67.
- Heikurainen, L., 1976. *Comparison between runoff conditions on a virgin peatland and a forest drainage area*. Proc. Fifth International Peat Congress. 76–86.
- Heimburg, K., 1984. Hydrology of North-Central Florida cypress domes. In: *Cypress Swamps*, K.C. Ewel. and H.T. Odum (Eds.), Univ. of Florida, Gainesville, USA. 72–82.
- Hemond, H.F., 1980. Biogeochemistry of Thoreau's Bog, Concord, Massachusetts. *Ecol. Monogr.*, **50**, 507–526.
- Hensel, B.R. and Miller, M.V., 1991. Effects of wetland creation on groundwater flow. *J. Hydrol.*, **126**, 293–314.
- Hewlett, J.D. and Hibbert, A.R., 1967. Factors affecting the response of small watersheds to precipitation in humid regions. In: *Forest Hydrology*, W.E. Sopper and H.W. Lull (Eds.) Pergamon Press, Oxford., UK, 275–290.
- Hey, D.L., Barrett, K.R. and Biegen, C., 1994. The hydrology of four experimental constructed wetlands. *Ecol. Eng.*, **3**, 319–343.
- Hickok, E.A., Hannaman, M.C. and Wenck, N.C., 1977. *Urban runoff treatment methods. Volume 1. Non-structural wetlands treatment*. U.S. EPA-600/2-77-217. NTIS Report No. PB-278-172/2ST.
- Hill, J.L. and Kidd, C.H.R., 1980. *Rainfall-runoff relationships for 47 Malawi catchments*. Water Resources Branch, Malawi, No. TR7. 22pp.
- Hillman, G.R., 1998. Floodwave attenuation by a wetland following a beaver dam failure on a second order boreal stream. *Wetlands*, **18**, 21–34.
- Hodnett, M.G., Vendrame, I., de O. Marques, F., Oyama, M.D. and Tomasella, J., 1997. Soil water storage and groundwater behaviour in a catenary sequence beneath forest in central Amazonia. *Hydrol. Earth Syst. Sci.*, **1**, 265–277.
- Hollis, G.E., 1990. Environmental impacts of development on wetlands in arid and semi-arid lands. *Hydrolog. Sci. J.*, **35**, 411–428.
- Hollis, G.E., 1996. Hydrological inputs to management policy for the Senegal river and its floodplain. In: *Water management and wetlands in Sub-Saharan Africa*, M.C. Acreman and G.E. Hollis (Eds.), IUCN. 155–184.
- Homik, K.T. and Madissoon, G.I., 1975. The influence of fen bog drainage on annual runoff in the Estonian SSR. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 487–492.
- Hooijer, A., 1996. *Floodplain hydrology: an ecologically oriented study of the Shannon Callows, Ireland*. Amsterdam University, The Netherlands.
- Howard-Williams, C., 1983. Wetland and watershed management: the role of aquatic vegetation. *J. Limnol. Soc. S. Afr.*, **9**, 54–62.
- Howe, C.P., Claridge, G.F., Hughes, R. and Zuwendra, 1992. *Manual of guidelines for scoping EIA in tropical wetlands. 2nd Edition*, Asian Wetland Bureau-Indonesia, Bogor, Indonesia, 261pp.
- Hruby, T. et al., 1995. Estimating relative wetland values for regional planning. *Wetlands*, **15**, 93–107.
- Hunt, R.J., Krabbenhoft, D.P. and Anderson, M.P., 1996. Groundwater inflow measurement in wetland systems. *Water Resour. Res.*, **32**, 495–507.
- Hurst, H.E., 1933. The Sudd region of the Nile. *J. Roy. Soc. Arts.*, **81**, 721–736.
- Iritz, L., Johanson, B. and Lundin, L., 1994. Impacts of forest drainage on floods. *Hydrolog. Sci. J.*, **39**, 637–661.
- Ingram, H.A.P., 1983. Hydrology. In: *Ecosystems of the World. 4A: Mires: swamp, bog, fen and moor*, A.J.P. Gore (Ed.), Elsevier, Amsterdam, The Netherlands. 67–158.
- Jackson, R.J., 1987. Hydrology of an acid wetland before and after draining for afforestation, western New Zealand. In: *Forest Hydrology and Watershed Management*, R.H Swanson (Ed), Proc. of the Vancouver Symposium, IAHS Publ. no. 167, 465–474.
- Johansson, B. and Seuna, P., 1994. Modelling the effects of wetland drainage on high flows. *Aqua Fennica* **24**, 59–68.
- John, D.M., Leveque, C. and Newton, L.E., 1993. Chapter on West Africa. In: *Wetlands of the World I: Inventory, ecology and management*. D.F. Whigham, D. Dykyjova and S. Hejny (Eds.), Kluwer, Dordrecht, The Netherlands.
- Junk, W.J., Bayley, P.B. and Sparks, R.E., 1989. The flood pulse concept in river-floodplain systems. *Can. J. Fisheries Aquat. Sci.*, **106**, 110–127.
- Keough, J.R. and Pippen, R.W., 1984. The movement of water from peatland into surrounding groundwater. *Can. J. Bot.*, **62**, 835–839.
- Khan, N., 1995. Protection of North Selangor Peat Swamp Forest, Malaysia. *Parks*, **5**, 24–31.
- Kiselev, P.A., 1975. Study of the water balance of stratified aquifers based on the analysis of their regime during hydrogeological investigations on marshy lands and adjacent areas. In: *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 35–46.
- Klepper, O., 1992. Model study of the Negara River Basin to assess the regulating role of its wetlands. *Regul. River*, **7**, 311–325.
- Kloet, L., 1971. *Effects of drainage and runoff and flooding within the Pembina River Basins: North Dakota - Manitoba*. Bureau of Sport, Fisheries and Wildlife, Bismarck, North Dakota
- Klueva, K.A., 1975. The effect of land reclamation by drainage on the regime of rivers in Byelorussia. In: *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 419–438.
- Koerselman, W., 1989. Groundwater and surface water hydrology of a small groundwater-fed fen. *Wetland Ecol. Manage.*, **1**, 33–43.

- Konyha, K.D., Robbins, K.D. and Skaggs, R.W., 1988. Evaluating peat-mining drainage using DRAINMOD. *J. Irrig. Drain. Eng.*, **114**, 490–504.
- Kovrigo, P.A. and Yatshukno, V.M., 1988. Optimization of the microclimate of bog geosystems. *Symposium on the Hydrology of wetlands in temperate and cold regions: Vol 2*. Academy of Finland, Helsinki, 22–29.
- Kowalik, P.J., Banach, K. and Toczyłowska, I., 1988. Seepage component in water balance of peat soils. *Symposium on the Hydrology of wetlands in temperate and cold regions: Vol 1*. Academy of Finland, Helsinki, 178–183.
- Kusler, J.A. and Riexinger, P., 1986. *National wetland assessment symposium*. Association of State Wetland Managers, New York, USA.
- Lipton, D.W. et al., 1995. *Economic valuation of natural resources – a handbook for coastal resource policy-makers*. US Environmental Protection Agency. EPA/600/R-92/167.
- Littlejohn, C., 1977. An analysis of the role of natural wetlands in regional water management. In: *Ecosystem modelling in theory and practice* C.A.S. Hall and J.W. Day (Eds.), Wiley, Chichester, UK. 451–476.
- Logan, W.S. and Rudolph, D.L., 1997. Micro depression-focused recharge in a coastal wetland, La Plata, Argentina. *J. Hydrol.*, **194**, 221–238.
- Ludden, A.P., Frank, D.L. and Johnson, D.H., 1983. Water storage capacity of natural wetland depressions in the Devils Lake Basin of North Dakota. *J. Soil Water Conserv.*, **38**, 45–48.
- Lundin, L., 1988. Impacts of drainage for forestry on runoff and water chemistry. *Symposium on the Hydrology of wetlands in temperate and cold regions: Vol 1*. Academy of Finland, Helsinki, 197–205.
- Maltby, E., 1991. Wetland management goals: wise use and conservation. *Landscape Urban Plan.*, **20**, 9–18.
- Maltby, E., Hogan, D.V. and McInnes, R.J., 1996. *Functional analysis of European wetland ecosystems*. Report to the European Commission. EC DGXII STEP-CT90-0084.
- Matheney, R.K. and Gerla, P.J., 1996. Environmental isotopic evidence for the origins of ground and surface water in a prairie discharge wetland. *Wetlands*, **16**, 109–120.
- McCartney, M.P., 2000. The water budget of a headwater catchment containing a dambo. *Phys. Chem. Earth B*, **25**, 611–616.
- McComas, M.R., Kempton, J.P. and Hinkley, K.C., 1972. *Geology, soils and hydrogeology of Volo Bog and vicinity, Lake County, Illinois*. Environmental Geology Note 57, Ill. State Geol. Surv. 27pp.
- McKay, E.D., Elzaftawy, A. and Cartwright, K., 1979. *Groundwater geology of selected wetlands in Union and Alexander Counties, Illinois*. Environmental Geology Note No. 85, Illinois Inst. of Natural Resources, Urbana, 41pp.
- Meigh, J., 1995. The impact of small farm reservoirs on urban water supplies in Botswana. *Natur. Resour. Forum*, **19**, 71–83.
- Meyboom, P., 1964. Three observations on streamflow depletion by phreatophytes. *J. Hydrol.*, **2**, 248–261.
- Mikulski, Z. and Lesniak, E., 1975. Hydrological research on a peatbog in the upper Suprasl basin. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris, 55–68.
- Miller, E.G., 1965. Effect of Great Swamp, New Jersey, on streamflow during base-flow periods. *U.S. Geol. Surv. Prof. Pap.* 525-B, 177–179.
- Millington, A.C., 1985. Inland Valley Swamps and bolls in Sierra Leone: hydrological and pedological considerations for agricultural development. *Z. Geomorph. N.F. Suppl.* **36**, 201–222.
- Mitsch, W.J., Dorge, C.L. and Wiemhof, J.R., 1977. *Forested wetlands for water resources management in southern Illinois*. Report 132, Water Resources Centre, Univ. of Illinois, Urbana.
- Mokylak, V.I., Kubyshkin, G.P. and Karkutsiev, G.N., 1975. The effect of drainage works on streamflow. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 439–446.
- Moskvin, Y.P., 1988. Water and energy regime of palsa bogs in West Siberia. *Symposium on the Hydrology of wetlands in temperate and cold regions: Vol 2*. Academy of Finland, Helsinki, 18–21.
- Mustonen, S.E. and Seuna, P., 1975. Influence of forest drainage on the hydrology of an open bog in Finland. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris 519–530.
- NERC, 1975. Flood routing studies: Vol III. *Flood Studies Report*. Natural Environment Research Council, London, UK.
- Newson, M.D., 1981. Mountain streams, In: *British rivers*, J. Lewin (Ed.), George Allen & Unwin. 59–89.
- Nielsen, S.A., Refsgaard, J.C. and Mathur, V.K., 1991. Conceptual modelling of water loss on floodplains and its application to River Yamuna upstream of Delhi. *Nord. Hydrol.*, **22**, 265–274.
- Nisula, T. and Kuittinen, R., 1988. Snowmelt and runoff in a mire in Finnish Lapland. *Symposium on the Hydrology of wetlands in temperate and cold regions: Vol 1*. Academy of Finland, Helsinki, 76–83.
- Nortcliff, S. and Thornes, J.B., 1981. Seasonal variations in the hydrology of a small forested catchment near Manaus, Amazonas, and the implications for its management. In: *Tropical Agricultural Hydrology*, R. Lal and E.W. Russell, (Eds.) Wiley, Chichester, UK. 37–58.
- Novitski, R.P., 1978. Hydrologic characteristics of Wisconsin's wetlands and their influence on floods, streamflow and sediment. In: *Wetland functions and values: the state of our understanding*, Amer. Water. Resour. Assoc., Minneapolis, MN, USA, 377–388.
- Novitski, R.P., 1982. *Hydrology of Wisconsin wetlands*. Univ. of Wisconsin, Geol. and Nat. Hist. Surv. Info. Circular No. 40, 22pp.
- Novitski, R.P., 1985. The effects of lakes and wetlands on flood flows and base flows in selected northern and eastern States. Proc. Conf. on *Wetlands of the Chesapeake, Easton, Maryland*, Environmental Law Institute, 143–154.
- Nuckles, D., 1970. *Virginia streamflow data program analysis*. Richmond, Virginia, U.S. Geol. Surv. Open File Rep. 54pp.
- O'Brien, A.L., 1977. Hydrology of two small wetland basins in Eastern Massachusetts. *Water. Resour. Bull.*, **13**, 325–340.
- O'Brien, A.L., 1980. The role of groundwater in stream discharges from two small wetland controlled basins in eastern Massachusetts. *Ground Water*, **18**, 359–365.
- O'Brien, A.L., 1988. Evaluating the cumulative effects of alteration on New England wetlands. *Environ. Manage.*, **12**, 627–636.
- OECD, 1996. *Guidelines for aid agencies for improved conservation and sustainable use of tropical and sub-tropical wetland*. Organisation for Economic Cooperation and Development Development Assistance Committee: Guidelines on Aid and Environment No.9, 69pp.
- Ogawa, H. and Male, J.W., 1986. Simulating the flood mitigation role of wetlands. *J. Water Resour. Plan. Man. ASCE*, **112**, 114–127.

- Owen, C.R., 1995. Water budget and flow patterns in an urban wetland. *J. Hydrol.* **169**, 171–187.
- Panu, U.S., 1988. Environmental assessment of peatland development operations in Newfoundland. *Symposium on the Hydrology of wetlands in temperate and cold regions*: Vol 1. Academy of Finland, Helsinki, 291–300.
- Phillips, P.J. and Shedlock, R.J., 1993. Hydrology and chemistry of groundwater and seasonal ponds in the Atlantic Coastal Plain in Delaware, USA. *J. Hydrol.*, **141**, 157–178.
- Poff, N.L., Allan, J.D., Bain, M.B., Karr, J.R., Prestegard, K.L., Richter, B.D., Sparks, R.E. and Stromberg, J.C., 1997. The natural flow regime. *Bioscience*, **47**, 769–784.
- Price, J.S., 1992. Blanket bog in Newfoundland. Part 2. Hydrological processes. *J. Hydrol.*, **135**, 103–119.
- Price, J.S. and Maloney, D.A., 1994. Hydrology of a patterned bog-fen complex in Southeastern Labrador, Canada. *Nord. Hydrol.*, **25**, 313–330.
- Raisin, G., Bartley, J. and Croome, R., 1999. Groundwater influence on the water balance and nutrient budget of a small natural wetland in North Eastern Victoria, Australia. *Ecol. Eng.*, **12**, 133–147.
- Riekirk, H. and Korhnaak, L.V., 2000. The hydrology of Cypress wetlands in Florida pine flatwoods. *Wetlands*, **20**, 448–460.
- Riggs, H.C., 1964. The base-flow recession as an indicator of groundwater. Symposium on Surface Waters, Proceedings of the General Assembly at Berkeley of I.U.G.G., *IAHS Publ. no. 63*, 352–363.
- Robertson, S.B., 1987. Hydrology of Arctic Wetlands. *Proc. Nat. Wetland Symp: Wetland Hydrology*. Assoc. of State Wetland Managers Inc., Chicago. 262–269.
- Robinson, M., Gannon, B. and Schuch, M., 1991. A comparison of the hydrology of moorland under natural conditions, agricultural use and forestry. *Hydrolog. Sci. J.*, **36**, 565–577.
- Romanov, V.V., 1968. *Hydrophysics of bogs*. Israel Program for Scientific Translations, Jerusalem.
- Roulet, N.T. and Woo, M.K., 1986. Hydrology of a wetland in the continuous permafrost region. *J. Hydrol.*, **89**, 73–91.
- Sander, J.E., 1976. An electric analogue approach to bog hydrology. *Ground Water*, **14**, 30–35.
- Sellers, C. D., 1981. A floodplain storage model used to determine evaporation losses in the Upper Yobe River, Northern Nigeria. *J. Hydrol.*, **52**, 257–268.
- Serban, P., Stanescu, V.A. and Simota, M., 1988. Contributions to the study of peatlands influence on maximum flow. *Symposium on the Hydrology of wetlands in temperate and cold regions*: Vol 1. Academy of Finland, Helsinki, 92–99.
- Seyhan, E., Hope, A.S. and Schulze, R.E., 1983. Estimation of streamflow loss by evapotranspiration from a riparian zone. *S. Afr. J. Sci.*, **79**, 88–90.
- Sharma, T.C., 1984. Some hydrologic characteristics of the Zambian headwaters. *Zambian J. Sci. Technol.*, **7**, 12–21.
- Sharma, T.C., 1988. An evaluation of evapotranspiration in tropical central Africa. *Hydrolog. Sci. J.*, **33**, 31–40.
- Shedlock, R.J., Wilcox, D.A., Thompson, T.A. and Cohen, D. A., 1993. Interactions between groundwater and wetlands, southern shore of Lake Michigan, USA. *J. Hydrol.*, **141**, 127–156.
- Shjeflo, J.B., 1968. *Evapotranspiration and the water budget of prairie potholes in North Dakota*. Geol. Surv. Prof. Pap. 585-B, United States Government Printing Office, Washington.
- Shiklmanov, I.A. and Novikov, S.M., 1988. Drainage effect on the environment. *Symposium on the Hydrology of wetlands in temperate and cold regions*: Vol 2. Academy of Finland, Helsinki, 66–71.
- Siegel, D.I., 1983. Groundwater and the evolution of patterned mires, glacial Lake Agassiz peatlands, Northern Minnesota. *J. Ecol.*, **71**, 913–921.
- Siegel, D.I., 1988. The recharge-discharge influence of tectonic setting on the hydrological characteristics of deeply weathered terrains: evidence from Uganda. *J. Hydrol.*, **218**, 44–71.
- Thompson, J.R. and Hollis, G.E., 1995. Hydrological modelling and the sustainable development of the Hadejia-Nguru wetlands, Nigeria. *Hydrolog. Sci. J.*, **40**, 97–108.
- Vecchioli, J., Gill, H.E. and Lang, S. M., 1962. Hydrologic role of the Great Swamp and other marshland in Upper Passaic River Basin. *J. Amer. Water Works Assoc.*, **54**, 695–701.
- Verry, E.S., 1988. The hydrology of wetlands and man's influence on it. *Symposium on the Hydrology of wetlands in temperate and cold regions*: Vol 2. Academy of Finland, Helsinki, 41–61.
- Verry, E.S. and Boelter, D.H., 1975. The influence of bogs on the distribution of streamflow from small bog upland catchments. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp. (June, 1972), International Association of Hydrological Sciences UNESCO/IAHS, Paris. 469–478.
- Verry, E.S. and Boelter, D.H., 1978. Peatland hydrology. In: Amer. Wat. Resour. Assoc. *Wetland functions and values: the state of our understanding*. Amer. Water Resour. Assoc, New York. 389–402.
- Verry, E.S. and Timmons, D.R., 1982. Waterborne nutrient flow through an upland-peatland watershed in Minnesota. *Ecology*, **63**, 1456–1467.
- Waddington, J.M., Roulet, N.T. and Hill, A.R., 1993. Runoff mechanisms in a forested groundwater discharge wetland. *J. Hydrol.*, **147**, 37–60.
- Walton, R., Chapman, R.S. and Davis, J.E., 1996. Development and application of the wetlands dynamic water budget model. *Wetlands*, **16**, 347–357.
- Wharton, C.H., 1970. *The Southern River Swamp - a Multiple-use environment*. Bureau of Business and Economic Research, School of Business Administration, Georgia State University.
- Wheeler, B.D., 1984. British Fens: a review. In: *European Mires*, P.D. Moore, (Ed.). Academic Press, London, UK. 237–281
- Wheeler, B.D. and Shaw, S., 2000. *A wetland framework for impact assessment at statutory sites in Eastern England*. Technical Report W6-068/TR1. Environment Agency, Bristol, UK.
- Wilcox, D.A., Shedlock, R.J. and Hendrickson, W.H., 1986. Hydrology, water chemistry and ecological conditions in the raised mound of Cowles Bog. *J. Ecol.*, **74**, 1103–1117.
- Williams, R.E., 1968. Flow of groundwater adjacent to small, closed basins in glacial till. *Water Resour. Res.*, **4**, 777–783.
- Wilson, B.H. and Dincer, T., 1976. An introduction to the hydrology and hydrography of the Okavango Delta. *Proc. Symp. on the Okavango Delta and its future utilisation*, Botswana Society, Gaborone, 33–48.
- Wilson, T.V. and Wisler, E.H., 1974. Groundwater-baseflow relationships and baseflow predictions from groundwater levels on piedmont watersheds. *Trans. Amer. Soc. Civ. Engrs*, 269–279.
- Winner, M.D. and Simmons, C.E., 1977. *Hydrology of the Creeping Swamp watershed, North Carolina, with reference to potential effects of stream channelization*. U.S.G.S./WRI-77/26.
- Wolski, A., 2002. Hydrological processes in the Okavango wetland, Botswana. *Proc. EGS*. Nice, April 2002.
- Woo, M.K. and Heron, R., 1987. Effects of forests on wetland runoff during spring. In: *Forest Hydrology and Watershed Management*. R.H. Swanson (Ed.), Proc. of the Vancouver Symposium, *IAHS Publ. no. 167*, 297–307.
- Woo, M.K. and Rowsell, R.D., 1993. Hydrology of a prairie slough. *J. Hydrol.*, **146** 175–207.
- Woo, M.K. and Winter, T.C., 1993. The role of permafrost and seasonal frost in the hydrology of northern wetlands in North America. *J. Hydrol.*, **141**, 5–32.

Zivert, A.A., Rieksts, I.A. and Skinkis, C.M., 1975. Investigation of runoff in drainage outlets and subsurface drained soils in the Latvian SSR. *Hydrology of marsh-ridden areas*, Proc. of the Minsk Symp (June, 1972), International Association of Hydrological Sciences, UNESCO/IAHS, Paris. 117–128.

Annex 1

GLOBAL REVIEW OF WETLAND WATER QUANTITY FUNCTIONS

Author	Country (state, province)	Wetland type	Local term	Basis of inference	Category of study	Summary of functional statement	Function summary
Hurst (1933)	Sudan	FP	Sudd	In-Out	LTH	FVa: swamp discharge varies very little; fluctuations are rapidly damped out MAAE: 14 milliards of flow is lost in the swamps (p.731-2)	FVa: - MAAE: +
Vecchioli <i>et al.</i> (1962)	New Jersey, USA	GW/S	Swamp, marsh	In-out	LTH	DPFV: baseflow is reduced to 75% of the input ... DPAE: ... by high summer evapotranspiration MAF: swamplands have little effect on the total flow FPLM: the swamp decreases the downstream floods FEV: seasonal runoff is greater than that of the upland (p.699-700)	DPFV: - DPAE: + MAF: . FPLM: - FEV: +
Riggs (1964)	N. Carolina, Tennessee, USA	General	Swamp	Paired	LTH	DPRR: comparison of a flat recession for the swampy Haw, with a steep recession of the non-wetland New River (p.359).	DPRR: +
Meyboom (1964)	Saskatchewan, Canada	GW/S	River valley	CompV	SEH	DPFV: 70% of flow depletion can be accounted for by phreatophytic vegetation (p.254)	DPFV: -
Meyboom (1966)	Saskatchewan, Canada	FP	River valley	CompV	SEH	DPFV: phreatophytes in the valley diminished flow by at least 20% (p.257).	DPFV: -
					SEH	DPFV: phreatophytes depleted 100% of river flow (p.259)	DPFV: -
					Comp _{GW}	AGR: "Groundwater was recharged during the period of study" (p.60)	AGR+
Miller (1965)	New Jersey, USA	GW/S	Swamp, marsh	In-out	WB	DPFV/DPAE: summer evapotranspiration causes a significant reduction in baseflow GDS: the swamp is an area of discharge for the regional groundwater body (p. B179)	DPFV: - DPAE: + GDS: =
Ackroyd <i>et al.</i> (1967)	Minnesota, USA	General	Swamp, bog, marsh	Multiple	LTH	DPFV: basins having lake or wetland areas in excess of 5% have more than twice as much annual groundwater runoff. (p.27)	DPFV: +
Bay (1967)	Minnesota, USA	SW/S	Perched bog	Comp _{GW}	Comp _{GW}	GDS: the water table reacts quite independently from the regional groundwater system (p. 309)	GDS: X
		GW/S	Bog	Comp _{GW}	Comp _{GW}	GDS: groundwater from surrounding mineral soil recharges the bog (p.309)	GDS: =
Burke (1968)	Ireland	SW/S	Peat	Drained	WB	MAF: outflow is the same from drained and undrained area DPFV: flow eventually becomes zero from both areas FPLM: much higher peaks occur for the undrained area FTIP: the undrained area flows sooner than the drained area FRR: after flood peaks, the drained area is still discharging water at a faster rate WPFVa: sustained flows are more uniform in the drained area (p.814-6)	MAF: . DPFV: . FPLM: + FTIP: + FRR: + WPFVa: +

Romanov (1968)	Former USSR	SW/D	Bogs	Same	WB	MAAE/MAF: bog evaporation and runoff approaches that of unbogged areas FVa: drainage of highmoor bogs causes the redistribution to become more marked. (p.232-3) WPGR: it was assumed that no seepage took place during the winter months when the potholes were frozen (p.B48)	MAAE: . MAF: . FVa: - WPGR: X
Shjeflo (1968)	N. Dakota, USA	SW/D	Prairie pothole	Comp _{GW}	WB	GDS/AGR: some marshes behave as groundwater sinks, others as a groundwater mound (p.782)	GDS: = AGR: =
Williams (1968)	Illinois, USA	GW/D	Marsh	Comp _{GW}	Comp _{GW}	DPFD: storage was not available to sustain summer flow FVa: bogs were not effective as streamflow regulators	DPFD: - FVa: .
Bay (1969)	Minnesota, USA	SW/S	Peat bog	Same	LTH	FPLM/ FRR: low peak flows and long-drawn out recessions suggest that the bogs do store short-term runoff (p.101)	FPLM: - FRR: -
Freeze (1969)	Saskatchewan, Canada	SW/D	Slough	Comp _{GW}	Chem	DPGR: among 76 sloughs 27 are classified as Fast Recharge, 10 as Slow Recharge and 39 do not recharge GDS: 14 are classified as Fast discharging sloughs, 10 as Slow Discharging sloughs; 50 do not discharge. (p. 12-14)	DPGR: = DPGR: X GDS: = GDS: X
Campbell & Drecher (1970) in Novitski (1985)	Wisconsin, USA	General	Basin storage	Multiple	LTH	WPF/AGR/DPFV: in basins with large lake and wetland area, more water runs off in spring and only a small amount recharges the aquifer; thus base flow is reduced in summer, fall and winter (Novitski, p. 147)	WPF: + AGR: - DPFV: -
Darner (1970) in Novitski (1985)	New York, USA	General	Lakes and ponds	Multiple	LTH	FPHM: basin storage is statistically insignificant in explaining the variability of flood peaks (T = 20 yrs) AGR/DPFV: large percentage of storage in basins results in reduced recharge and consequently reduced baseflow (Novitski, p. 145-8)	FPHM: . AGR: - DPFV: -
Forest & Walker (1970) in Novitski (1985)	Delaware, Maryland, USA	General	Basin storage	Multiple	LTH	FPHM: basin storage is statistically insignificant in explaining the variability of flood peaks (T = 5-100 yrs) AGR/WPF/DPFV: spring runoff is greater and recharge to groundwater (and baseflows) lower in basins with a larger percentage of lake and wetland area (Novitski, p. 145-9)	FPHM: . AGR: - WPF: + DPFV: -
Nuckles (1970) in Novitski (1985)	Virginia, USA	General	Basin storage	Multiple	LTH	FPHM: basin storage is statistically insignificant in explaining the variability of flood peaks (T = 5-100 yrs) WPF/DPFV: annual minimum series are lower for basins with large percentages of lakes and wetlands (Novitski, p. 149)	FPHM: . WPF: + DPFV: -
Wharton (1970)	Georgia, USA	FP	Swamp	Paired	LTH	FPLM: a comparison of Alcovy (swamp) and Yellow hydrographs suggests a damping influence FTTP: peaks of the Alcovy lag 24 hours behind the Yellow FVa: Alcovy flow duration curves are smooth and many minor fluctuations shown for the Yellow are missing. It is likely that the Alcovy (alluvial) aquifer beneath the floodplain does strongly influence the variability of base flow DPFV: although the Yellow drains an area 36% larger, its low flows are close to those of the Alcovy, suggesting some possible influence of the swamp on base flow (p. 15-18).	FPLM: - FTTP: - FVa: . DPFV: +

Conger (1971) in Novitski (1985)	Wisconsin, USA	General	Basin storage	Multiple	LTH	FPHM: basin storage is significant in explaining variability in flood peaks (T = 2-100 yrs); the higher the storage term, the lower the flood peaks. (p.144-5)	FPHM: -
Kloet (1971) in Bardecki (1987)	N. Dakota, USA & Manitoba, Canada	General	General wetland	Drained	LTH	FPLM: peak flow is increased due to drainage in a comparative study of a drained and undrained basin (Bardecki, p.126);	FPLM: -
Millar (1971)	Canada	GW/S	Prairie pothole	Comp _{GW} Comp _E	WB	AGR: "Shoreline related water loss accounts, on average, for 60% or more of total water loss in sloughs less than 1.0 acres in size and not more than 30-35% in sloughs larger than 1 acre" (p.259) DPAE: "Reduction in evaporation loss (is due to) the sheltering effect of topography and marginal and emergent vegetation" (p.279)	AGR+ DPAE: -
Eisenlohr (1972)	N. Dakota, USA	GW/D	Prairie pothole	Comp _{GW}	Comp _{GW}	GDS: ponds in prairie potholes, in effect, are "outcrops" of the water table (p.A82)	GDS: =
Hall et al. (1972)	New Hampshire, USA	SW/S	Pond	CompV	CompV	DPFV: there is no evidence of the wetland sustaining baseflow (p.41)	DPFV: .
McComas et al. (1972)	Illinois, USA	GW/S	Bog	In-out	WB	DPFV: no groundwater left the basin as baseflow FEV: materials are saturated in the spring and little further retention takes place. (p.17)	DPFV: - FEV: +
Stewart & Kantrud (1972)	N. Dakota, USA	SW/D	Prairie pothole	Comp _{GW}	Chem	AGR: seepage accounted for some water loss in ponds at higher elevations on glacial till. (p.D4-D5)	AGR: =
Balek & Perry (1973)	Zambia	GW/S	Prairie pothole	CompGW	Chem	GDS: Ponds at lower elevations, on outwash, were subject to greater seepage inflow. (p.D4-D5)	GDS: =
		GW/S	Dambo	Same	CCM	MAF: total runoff is independent of the size of the dambos WPF: from a higher percentage of dambo area, a higher surface runoff volume occurs FRR: duration of surface runoff as compared with a non-swampy area is delayed by dambo resistance MAAE: evapotranspiration by woodland is three times higher than that from dambos DPFD: base flow from the dambo ceases earlier than from the transitive region DPRR: recessions are steeper from a catchment containing 10% dambo compared with 5% dambo (239-248)	MAF: . WPF: + FRR: - MAAE: - DPFD: - DPRR: +
Wilson & Wiser (1974)	S. Carolina, USA	FP	Floodplain	In-out	LTH	DPFV/DPAE: not only does the lower (floodplain) zone fail to contribute water but through it some of the potential baseflow is lost to evapotranspiration (p.274)	DPFV: - DPAE: +
Bavina (1975)	European USSR	SW/S	Raised bogs	Paired	WB	MAF: runoff from the swamp and from the (non-swamp) river basins was the same (p.301)	MAF: .
		FP	Valley bogs	Paired	WB	MAF: runoff from the swamp is in general agreement with runoff from the swampy catchment (p.302)	MAF: .

Bulavko & Drozd (1975)	USSR	General	Peat	Drained	LTH	MAF/MAAE: initially after drainage annual runoff is increased due to a decrease in evapotranspiration WPF: spring flows either increase or decrease DPFV: minimum and summer low flows increase considerably FV _a : the proportion of the groundwater contribution to flow increases, improving the distribution of river runoff (p.466)	MAF: - MAAE: + WPF: + WPF: - DPFV: - FV _a : +
Burke (1975)	Ireland	SW/S	Blanket peat	Drained	WB	MAF/FV _a /FEV: the drained area runoff was 60% greater than the undrained area. Flow was much more uniform and did not show the sharp peaks evident in the undrained area AGR: some seepage may occur, albeit small, but both peat and the subsoil have extremely low permeability FPLM/DPFV: after drainage, floods will be reduced in frequency and amount and summer flow of streams will be increased in the short-term (p.176)	MAF: - FV _a : - FEV: - AGR: . FPLM: + DPFV: -
Eggelsmann (1975)	Germany	SW/S	Peat	Same	WB	MAAE/MAF: there is higher evaporation from peat and reduction of runoff with respect to mineral soils FV _a : there is no causal relations between the steady runoff and the water storage of peat (p.359)	MAAE: + MAF: - FV _a : .
Eisenlohr (1975)	N. Dakota, USA	GW/D	Prairie Pothole	CompGW	CompGW	AGR: although there is evidence that a very small portion of seepage moves vertically downward, most of the seepage outflow moves laterally (p.309)	AGR: .
Glazacheva (1975)	Latvia	General	Marsh	Drained	LTH	FV _a /FPLM/DPFV: river flow variability increased following drainage as the maximum discharges increased and the minimum decreased (p.514)	FV _a : - FPLM: - DPFV: +
Hommik & Madissoo (1975)	Estonia	GW/S	Fen bogs	Drained	WB	MAF: annual runoff from drained fen bogs increases by 92mm (p.489)	MAF: -
Kiselev (1975)	Byelarus	GW/S	Swamp	CompGW	CompGW	AGR/GDS: swamps may be supplied by groundwater and may also serve as sources of groundwater recharge (p.38-40)	AGR: = GDS: =
Clueva (1975)	Byelorussia	SW/S	Marsh	Drained	LTH	DPFV: the minimum discharge increased by 30-150% FPLM: maximum discharges decreased by 17-30% on 7 basins; in other basins no significant change occurred FTTP: no significant changes were observed in the timing of the snowmelt flood, or in the date of the peak FV _a : 6 basins were characterised by decrease of spring flow of 10 to 30%; summer and autumn flow increased 20 to 60% MAF: MAF of 9 basins increased by 10 to 20%. On the other 7 basins there was no significant change (p.424-6)	DPFV: - FPLM: + FPLM: . FTTP: . FV _a : + MAF: - MAF: .
Moklyak <i>et al.</i> (1975)	Ukraine	FP	Marsh	Drained	LTH	MAF: runoff tends to decrease after drainage WPF: decrease in spring flows is generally found but cases are encountered where the flow is unchanged or tends to increase FPLM: drainage does not always affect the maximum discharge, although it may either decrease or increase (p.442)	MAF: + WPF: + WPF: . WPF: - FPLM: + FPLM: .
Mikulski & Lesniak (1975)	N.E. Poland	GW/S	Peat bog	Drained	WB	MAAE: evaporation decreased by about 15% as compared with the pre-reclamation period MAF/WPF/DPFV: results show a 20% increase in runoff, as compared to the pre-reclamation period... there were increases in the runoff, in summer and autumn (p.59)	MAAE: + MAF: - WPF: - DPFV: -

Author (Year)	Country	SW/S	Peat	Drained	LTH	MAAE/MAF: decrease in evapotranspiration led to an increase in total runoff FPLM: an acceleration of flow caused by the drainage network led to an increase in maximum runoff DPFV: the minimum runoff for both winter and summer increased markedly (p.523)	MAAE: + MAF: - FPLM: - DPFV: -
NERC (1975)	UK	FP	Floodplain	In-out	SEH	FPLM/FPHM: the most important change induced by a large floodplain on the shape of a flood hydrograph is the attenuation of the peak discharge (p.9)	FPLM: - FPHM: -
Smith (1975)	Florida, USA	FP	Cypress swamp	Same	CompGW	AGR: clays between the water table and the Floridan aquifer allow virtually no vertical movement MAAE: vegetation is acting as a pump, removing water from the water table via transpiration (p.128-135)	AGR: X MAAE: +
Verry & Boelter (1975)	Minnesota, USA	GW/S	Lake-filled bog	Same	LTH, SEH	FVa: the groundwater bog has no regulating effect FPLM: the bog does reduce storm flow peaks FPHM: maximum peaks are independent of the bog FRR: the bog delays the release of storm flow (p.472)	FVa: . FPLM: - FPHM: . FRR: -
		SW/S	Lake-filled bog	Same	LTH, SEH	FVa: the perched bog has no regulating effect FPLM: the bog does reduce individual storm peaks FPHM: maximum peaks are independent of the bog FRR: the bog delays the release of storm flow (p.472)	FVa: . FPLM: - FPHM: . FRR: -
Zivert <i>et al.</i> (1975)	USSR	GW/S	Fen bog	Same	WB	MAF/WPVa: in fen bogs drainage flow is 20 to 40% greater than mineral soils. The hydrograph of flood discharge is more uniform in peat than in mineral soils (p.121)	MAF: + WPVat: -
Heikuranen (1976)	Finland	SW/S	Peat	Drained	WB	WPF: peak flow caused was considerably lower on the drained peatland FTTP: the flood from the drained peatland began earlier and lasted longer FPLM: peak flow of the former remained lower DPFV: runoff from the drained peatland during the dry summer was greater FRR: flood from the drained peat began earlier and lasted longer (p.84-5)	WPF: + FTTP: + FPLM: - DPFV: - FRR: -
		GW/S	Bog	In-out	WB /CCM	FVa: the bog's chief role is to superimpose a significant seasonal fluctuation on discharge WPF: the bog releases excess water during wet periods DPFV/DPAE: the bog depletes available supplies during dry periods through evapotranspiration GDS: the wetland receives water from groundwater (p.35)	FVa: + WPF: + DPFV: - DPAE: + GDS: =
Wilson & Dincer (1976)	Botswana	FP	Delta	In-out	WB	MAAE/MAF: outflow in the Boteti amounts to only 2% of inflow due to the evapotranspiration losses (p.36).	MAAE: + MAF: -
Balék (1977)	Zambia	GW/S	Dambo	Same	WB/CCM	FRR: the duration of the surface runoff is prolonged until early June (p.159)	FRR: -
Boelter & Verry (1977)	Northern Lake States, USA	SW/S	Perched peat bog	CompGW	WB	DPFV/DPAE/AGR: late spring, summer and early fall evapotranspiration is at the expense of flow and deep seepage; peatland does not sustain streamflow during dry summer months by slowly releasing stored water FRR: stormflows are modified by peatland. Storm hydrographs have long-drawn out recession curves FPLM: peatland does reduce the peak rates of flow FVa: neither bogs nor fens maintain an even distribution of streamflow (p.14-18)	DPFV: - DPAE: + AGR: - FRR: - FPLM: - FVa: +

Author (Year)	Location	GW/S	Groundwater peat fen	CompGW	WB	FV/a/WPF/DPFV/DPAE: instead of regulating flow, the fen may do the opposite by releasing excess water more quickly than mineral aquifers during periods of high precipitation and losing more water by evapotranspiration during dry periods (p. 15-16)	FV/a: + WPF: + DPFV: - DPAE: +
Flippo (1977) in Novitski (1985)	Pennsylvania, USA	General	Basin storage	Multiple	LTH	FPHM: basin storage is statistically insignificant in explaining the variability of flood peaks (T = 5-100 yrs) (p.145)	FPHM: .
Hickok <i>et al.</i> (1977)	Minnesota, USA	GW/S	Wetland	Same	WB	AGR/GDS: the wetland is a point of discharge for the local glacial till. Groundwater losses are considered zero DPFV: the wetland reduced minimum flow to less than the estimated base flow (p.46)	AGR: X GDS: = DPFV: -
Littlejohn (1977)	Florida, USA	FP	Cypress swamp	Drained	CCM	FV/a: retention of cypress swamps contribute to greater stability of water regimes (p.472)	FV/a: -
Mitsch <i>et al.</i> (1977)	Illinois, USA	FP	Forest swamp	Same	WB	FEV: water retained by the swamp is 7.8% of an individual event. No effect was seen on any other storm occurrences DPFV: flow maintenance can be very significant (p.77-8,90)	FEV: - FEV: . DPFV: +
O'Brien (1977)	Massachusetts, USA	GW/S	Peat	Paired, Same	WB, LTH	WPAE: spring evapotranspiration is depressed relative to non-wetland WPF: the wetland was responsible for high spring flows DPAE: fall rates are high relative to non-wetland areas DPFV: baseflow during the low flow period was greatly depressed GDS: the wetland receives water from the regional groundwater body (p.336-338)	WPAE: - WPF: + DPAE: + DPFV: - GDS: =
Winner & Simmons (1977)	N. Carolina, USA	SW/S	Muck wetland	Paired, same	WB, LTH	WPAE: spring evapotranspiration is depressed relative to non-wetland WPF: the wetland was responsible for high spring flows DPAE: fall rates are high relative to non-wetland areas DPFV: baseflow was greatly depressed DPGR: during summer the swamp recharges groundwater (p.336-8)	WPAE: - WPF: + DPAE: + DPFV: - DPGR: =
Novitski (1978)	Wisconsin, USA	General	Floodplain swamp	Drained	LTH	FPLM/DPFV/MAF: higher flows (> 5 exceedance-percentile) increase; low flows increase. Total runoff would not change (p.55)	FPLM: - DPFV: - MAF: .
Verry & Boelter (1978)	Minnesota, USA	GW/S	Wetland and lake	Multiple	LTH	FPLM: flows may be 80% lower in basins with much wetland and lake WPF: more spring runoff occurs in basins with much wetland and lake AGR/DPFV: less groundwater recharge (and baseflow) occurs in basins with much lake and wetland (p.384-6) AGR: the GW/S fen does not discharge through the peat (p.398)	FPLM: - WPF: + AGR: - DPFV: - AGR: x
McKay <i>et al.</i> (1979)	Illinois, USA	SW/S	Peat	Same	WB	FV/a: peats do not regulate flow from one season to another FPLM: peat does reduce the peak flow DPFV/DPGR: peat summer evapotranspiration is at the expense of streamflow or recharge (p.398)	FV/a: - FPLM: - DPFV: - DPGR: -
		GW/S	Swamp	CompGW	WB	GDS: the volume of (ground-) water annually entering the swamp by discharge from the bedrock is quite large (p. 31).	GDS: =

Author (Year)	Location	FP	Dambo	Multiple	LTH	Summary	Key Findings
Drayton <i>et al.</i> (1980)	Malawi	FP	Dambo	Multiple	LTH	FPHM: dambo provides a lot of floodplain storage DPFV: dambo catchments are not significantly distinguishable from unaffected neighbours for Q75 MAAE/MAF: presence of dambo increases evaporation with a corresponding decrease in average annual yield (p.58)	FPHM: - DPFV: . MAAE: + MAF: -
Hemond (1980)	Massachusetts, USA	SW/D	Kettle hole	In-out	WB	GDS: the bog is characterised by an absence of recharge (p.522)	GDS: X
Hill & Kidd (1980)	Malawi	GW/S	Dambo	Multiple	LTH	MAAE/MAF: average annual runoff volume is reduced by 6.4mm for every 1% of dambo (p.16)	MAAE: + MAF: -
O'Brien (1980)	Massachusetts, USA	GW/S	Wetland	CompGW	LTH	WPF: groundwater was the major component of all flood peaks (rather than originating from the wetlands) (p.359)	WPF: .
Bedinger (1981)	Mississippi, USA	FP	Floodplain	CompGW	LTH	FPLM: floodplains ameliorate downstream flooding AGR: significant recharge occurs on floodplains (p.168,173)	FPLM: - AGR: =
Brun <i>et al.</i> (1981)	N. Dakota, USA	GW/D	Sloughs	Drained	TS	MAF/FPHM/DPFV: approximately 50% of the increase in flow, 36% of the increase in maximum flow, and 70% of the increase in spring flow is due to increased drainage (p.13)	MAF: - FPHM: - DPFV: -
Daniel (1981)	N. Carolina, USA	GW/S	Pocosin	Drained	WB	FPLM: hydrographs from the Albemarle Canal show five floods while the wetland Van Swamp show none DPFV/DPAE: from raised wetlands there is usually little discharge after June because of evapotranspiration MAF: runoff from the two catchments is nearly equal (p.90-3)	FPLM: - DPFV: - DPAE: + MAF: .
Gilliam & Skaggs (1981)	N. Carolina, USA	GW/S	Pocosin	Drained	WB	FPLM/FTTP: peaks are higher and earlier on developed sites MAF: there was little difference in total flow (p.115)	FPLM: - FTTP: + MAF: .
Newson (1981)	Wales, UK	GW/S	Peat	Same	SEH	DPFV: very low yields prevailed, excepting two headwater areas characterised by periglacial deposits and peat (p.69).	DPFV: +
Nortcliff & Thornes (1981)	Brazil	FP	Floodplain	CompWC	WB	FEV: the dominant and rapid hydrograph response comes essentially from saturation overland flow on the floodplain (p.54)	FEV: +
Sellars (1981)	Nigeria	FP	Floodplain	CompET	CCM	MAF/DPFV/AGR: losses in the Upper Yobe are 70% of MAF. Storage causes flooding to continue into the dry season. 10% of losses contribute to regional groundwater (p.267)	MAF: - DPFV: + AGR: =
Novitski (1982)	Wisconsin, USA	SW/D	SW/D wetland	Same	WB	FPLM: ... an effect in reducing flood peaks DPFV: ... some effect on increasing base flows WPF: ... increases spring time runoff MAAE: ... no effect on ET (p.19-20)	FPLM: - DPFV: + WPF: + MAAE: .
		SW/S	SW/S wetland	Same	WB	A surface-water depression wetland has ... FPLM: ... an effect in reducing flood peaks DPFV: ... some effect on increasing base flows WPF: ... increases spring time runoff MAAE: ... no effect on ET (p.19-20)	FPLM: - DPFV: + WPF: + MAAE: .
		GW/D	GW/D wetland	Same	WB	A ground-water depression wetland has ... FPLM: ... an effect in reducing flood peaks DPFV: ... no effect on increasing base flows WPF: ... no effect on increasing spring time runoff MAAE: ... an effect in increasing ET AGR: some recharge occurs (p.19-20)	FPLM: - DPFV: . WPF: . MAAE: + AGR: =

	GW/S	GW/S wetland	Same	WB		FPLM: - DPFV: - WPF: + MAAE: +
Verry & Timmons (1982)	General	Wetland and lake	Multiple	LTH	A groundwater slope wetland has ... FPLM: an effect in reducing flood peaks DPFV: no effect on increasing base flows WPF: an effect in increasing spring time runoff MAAE: some effect in increasing ET (p.19-20)	FPLM: - DPFV: - WPF: + MAAE: +
	SW/S	Peat	Same	WB	FPLM: flood peaks are 80% lower in a basin with 40% lake and wetland area than in a basin with no lake or area (p.16)	FPLM: -
Howard-Williams (1983)	GW/S	Macrophytes	Comp _{ET}	LTH	MAF: peat converts 36% of its precipitation to water yield, or nearly the same as the upland FEV: during high water table periods, the bog can transmit flow much quicker than the mineral soil AGR: because hydraulic conductivities of well-decomposed peat and glacial till are similar, seepage rates over long periods probably are similar (p.1461)	MAF: . FEV: + AGR: .
Ludden <i>et al.</i> (1983)	SW/D	Shallow depressions	Multiple	LTH	DPAE: distinct diurnal rhythms in hydrographs were caused by evaporative losses during the day, which could not be detected when the macrophyte growth was minimal (p.57)	DPAE: +
Seyhan <i>et al.</i> (1983)	FP	Marsh	Same	LTH	MAF/FPHM: the depressions store 72% of 2-yr return period flow and 41% of 100-year return period flow (p.45)	MAF: - FPHM: -
Siegel (1983)	GW/S	Raised bogs Fens	CompGW	CCM	DPAE: recession hydrographs exhibit diurnal fluctuations with greater evapotranspirational losses in the riparian zone during the day (p. 88)	DPAE: +
Smith-Carrington (1983)	GW/S	Dambo	Same	WB	AGR: recharge zones in the peats are the raised bogs, and the discharge zones are the adjacent fens. Precipitation on the fens does not enter the groundwater system (p. 918)	AGR: = GDS: = AGR: X
Heimburg (1984)	GW/S	Cypress domes	CompGW	WB	MAAE: actual annual evaporation losses from dambo is 640mm compared with 692mm from shallow interfluvial and 760mm from wooded interfluvial FEV: hydrographs show more flashy responses because there is less dambo area for temporary retention of water. DPFV/DPAE: high dambo evapotranspiration results in a decrease in flow rate and cessation of river flow (p.36).	MAAE: - FEV: - DPFV: - DPAE: +
Keough & Phippen (1984)	GW/D	Peat bog	In-out	Chem	AGR: usually Sewage Dome 2 recharges the groundwater. At Austin Cary there is little deep percolation as the underlying aquifer is artesian (p. 80).	AGR: = AGR: X
Sharma (1984)	FP	Floodplain	Same	CompJET	WPGR: bog waters are moving locally into groundwater DPFV: wetlands retain rain and incoming runoff and groundwater until it evaporates or percolates (p.839)	WPGR: = AGR: =
Dubreuil (1985)	General	Hydromorphic/ gley soil	Same	Comp	MAAE: evaporation from flooded areas on the Flats is considerably higher than from non-flooded areas. (p.12-21)	MAAE: +
					FPLM: minimum rainfall required to produce runoff from hydromorphic/gley soils is lower than from ferrallitic soils FEV: on freely drained forest soils, no runoff occurs under rainfall of 120mm hr ⁻¹ , while runoff attained 80% of input under rainfall of 30mm h ⁻¹ on hydromorphic soils WPAE/WPF: losses are far greater during severe floods where there are floodplains. In Mauritania, this represents almost 30% of the total flood volume AGR: when a stream disappears into its floodplain, the alluvial water table may be replenished (p.244-257)	FPLM: + FEV: + WPAE: + WPF: - AGR: +

Millington (1985)	Sierra Leone	FP	Inland valley swamps and bolls	Same	LTH	DPFV: during the dry season, swamps retain surface water which is lost by evaporation and seepage (p.19)	DPFV: - AGR: = DPAE: +
Novitski (1985)	Northern and Eastern States, USA	General	Wetland and lake	Multiple	LTH	In basins with large percentages of lakes and wetlands... FPLM: Flood peaks are less WPF: spring runoff is greater DPFV: baseflow is less (p.151)	FPLM: - WPF: + DPFV: -
Baden & Eggelsmann (1986)	Germany	SW/S	Raised bog	Drained	WB	FV'a: run-off of the predrain bog was more extreme FPLM/DPFV: run-off in downstream areas are such that the risk of highwater and dryness becomes smaller (p.206-8)	FV'a: + FPLM: + DPFV: -
Gurnell & Gregory (1986)	England, UK	SW/S	Saturated heath	Same	LTH	FPLM: storm runoff volume can be related to the area of the catchment that is saturated or has a near-surface water table prior to a storm (p.94)	FPLM: +
Ogawa & Mate (1986)	Massachusetts, USA	FP	Wetland	With-out	CCM	FPLM: the worth of an upstream wetland for flood mitigation is negligible. Downstream main-stem wetlands were more effective in reducing downstream flooding (p.114)	FPLM: - FPLM: -
Roulet & Woo (1986)	NW Territories, Canada	GW/S	Peat	Comp _{wc}	WB	DPFV/DPAE/FV'a: post-spring water loss is mainly due to evaporation and not lateral runoff. Wetlands do not play an important role in streamflow regulation (p.89)	DPFV: - DPAE: + FV'a: .
Wilcox <i>et al.</i> (1986)	Indiana, USA	GW/S	Peat fen	Comp _{ew}	Comp _{ew}	AGR: a water table in the peat mound causes a pattern of shallow groundwater flow away from the peat mound GDS: seepage through marl into peat is deduced (p.1111-3)	AGR: = GDS: =
Bardecki (1987)	S.W. Ontario, Canada	General	General	Drained	LTH	MAF/FPLM/DPFV: neither strong evidence nor clear suggestions of any change in flow attributable to drainage was found (p.127)	MAF: - FPLM: . DPFV: .
Doyle (1987)	Massachusetts, USA	FP	Floodplain	Paired	LTH	FPHM: maximum floods in the Charles are extremely low compared to the adjacent low wetland Blackstone river (p.111)	FPHM: -
Ford & Bedford (1987)	Alaska, USA	General	General wetland	In-out	WB	AGR: the contribution of Alaskan wetlands to groundwater is probably negligible FPLM: during snowmelt, wetland soils do not contribute significantly to flood storage (p.209).	AGR: X FPLM: .
Jackson (1987)	South Island, New Zealand	GW/S	Pakhi wetland	Drained	WB	FEV: natural undrained wetlands are highly responsive to rainfall. Over 70% of total annual runoff is quickflow DPFV: low flows increased after vegetation was removed, but decreased after draining. FPMH: increased frequency of large peak flows is probably the most important impact of drainage works (p.471-3)	FEV: + DPFV: - DPFV: + FPHM: -
Robertson (1987)	Alaska, USA	General	General	Same	LTH	DPFV: wetlands contribute little to the mid-summer budget of tundra streams (p.267)	DPFV: .
Woo & Heron (1987)	N. Ontario, Canada	GW/S	Bogs and fens	CompWC	WB	GDS: meltwater from open bogs and fens is supplied by the local snow cover and from adjacent forested areas. Most of the runoff from the bog occurred as groundwater flow (p.303-4)	GDS: =

Brandesten (1988)	S.C. Sweden	GW/S	Mire	Multiple	LTH	DPFV: the principal difference between raised bogs and till is groundwater levels in the bogs are always sufficiently high to provide the streams from these areas with runoff (p.90)	DPFV: +
Konyha <i>et al.</i> (1988)	North Carolina, USA	GW/S	Peat	Drained	CCM	MAF/FPLM: peat mining alone increases the annual runoff and peak outflow rates (p.490)	MAF: - FPLM: -
Kovrigo & Yatsukhno (1988)	Byelorussia	SW/S	Bog	Drained	WB	MAAE: evaporation decreases 7-10% following drainage reclamation (p.24)	MAAE: +
Kowalik <i>et al.</i> (1988)	Poland	GW/S	Peat	Comp GW	Comp GW	GDS: upward seepage is about 30% of rainfall (p.178)	GDS: =
Lundin (1988)	Sweden	GW/S	Peat	Drained	WB	MAF: drainage changed MAF insignificantly DPFV: low discharges generally increased (p.204)	MAF: . DPFV: -
Moskvin (1988)	West Siberia	SW/S	Palsa bogs	Same	LTH	DPFV: flow ceases completely in some weeks or months of the warm period due to low storage capacity AGR: there is a lack of losses by percolation into deep aquifers through saturated frozen peat layers (p.20-1)	DPFV: - AGR: X
Nisula & Kuitinen (1988)	Finnish Lappland	SW/S	Aapa-mire	Same	WB	FEV: maximum runoff was caused by rapid runoff of the water from the mire (p.81)	FEV: +
Panu (1988)	Newfoundland, Canada	GW/S	Peat	Drained	LTH	FPLM: peak flows are increased by two to five folds DPFV: flow regime in the disturbed sub-basin experienced substantial changes such as decrease in low flows (p.295-6)	FPLM: - DPFV: +
Serban <i>et al.</i> (1988)	Romania	GW/S	Peat	Paired	LTH	FEV: runoff in the peat basin is 30-35% lower than in the control (p.93-94)	FEV: -
Sharma (1988)	Africa, Zambia	GW/S	Dambos	Same	WB	MAAE: a ratio of ET/PET (Penman) equal to 1.5 is not surprising for tropical wetlands (p.38-39)	MAAE: +
Shiklimanov & Novikov (1988)	Russia	FP	Swamps	Same	WB	MAF: because of the swamps, the mean annual runoff is reduced to 5600 x 10 ⁶ m ³ (p.38-9)	MAF: -
Siegel (1988)	Alaska, USA	GW/S	Blanket bog	Drained	LTH	MAF: drainage mainly increases MAF in the first years. There are some cases when MAF is reduced FPLM: high spring maxima tend to fall DPFV: minimum and summer runoff from swamps after drainage tends to increase FVa: drainage is manifested as more uniform and even flow distribution over seasons (p.69-71)	MAF: - MAF: + FPLM: + DPFV: - FVa: +
Verry (1988)	Minnesota, USA	GW/S	Mire	Same	SEH	AGR/GDS/DPFV: wetland recharge and discharge are very small. Groundwater discharge from wetlands is too small to measure (p.427)	AGR: X GDS: X DPFV: -
Brooks & Krefl (1989)	Minnesota, USA	GW/S	Fen peat and bog	Drained	CCM	FPHM: during large storms, peat looks like a reservoir - overland flow from the mineral soil was not observed (p.55) FEV: runoff is greater for mined peatlands than unmined DPFV: streamflow is generally reduced during summer months because of high evapotranspiration MAF: peat extraction in both bogs and fens appears to increase water yield over the short term (p.114-5)	FPHM: + FEV: - DPFV: - MAF: -

Author(s)	Location	GW/S	Quaking fen	Same	WB	GDS: the fen is a focus for groundwater discharge (p.31)	GDS: =
Koerselman (1989)	Netherlands	GW/S	Dambo	Same	CompPET	DPAE: the mean value of evaporation over all dambo regions over the 200 km ² was estimated to be 3.5mm d ⁻¹ and 3.2mm d ⁻¹ for the area outside the dambos (p.48-59)	DPAE: +
Stewart (1989)	Zimbabwe	GW/S	In-filled lakes	In-out	LTH	FPLM: wetlands do not have a substantial effect on the magnitude of flood peaks. The wetlands may have a lesser ability to attenuate flood peaks than would upland areas. AGR: wetlands have no significant role in groundwater recharge. (p.412)	FPLM: . FPLM: + AGR: X
Sundeen <i>et al.</i> (1989)	Colorado, USA	SW/S	Floodplain	Multiple	CCM	MAAE: losses in the Sudd and in the Niger are over half the annual inflow, in the lower Senegal, are insignificant, and in the Okavango are a very high proportion of the inflows FVa: Okavango and Sudd outflows are even less variable after damping (p.54-5).	MAAE: + MAAE: + MAAE: + MAAE: . FVa: - FVa: -
Sucliffe & Parks (1989)	Africa	FP	Forest wetland	CompGW	WB	AGR: the bog and pocomin wetlands have virtually no recharge, the cypress dome experiences losses through infiltration or seepage (p.172-173) AGR: recharge was found in the eastern portion of the wetland (p.225)	AGR: X AGR: =
Brown (1990)	North America, USA	General	Wetland	In-out	WB	AGR: the seasonally flooded floodplains provide 1400 x 10 ⁶ m ³ of recharge to the Chad formation (p.418)	AGR: =
Gehrels & Mulamoottil (1990)	S. Ontario, Canada	SW/S	Floodplain	Multiple	LTH	FPLM: peakflows decrease with increasing wetland FEV: flood volumes decrease with increasing wetland DPFV: wetlands consistently increased low flows (p.949)	FPLM: - FEV: - DPFV: +
Hollis (1990)	Nigeria	FP	Wetland	Drained	LTH	FPLM: in their pre-drained state, wetlands reduced flooding. DPFV: in their pre-drained state, wetlands reduced low flows (p.1050)	FPLM: - DPFV: -
Demissie <i>et al.</i> (1991)	Illinois, USA	General	Dambo	Same	WB	AGR: total amount of water lost from the dambo to groundwater is small DPAE: dry season evaporation was approximately 50% higher from the dambo than from dryland (p.153-5)	AGR: X DPAE: +
Faulkner & Lambert (1991)	Zimbabwe	GW/S	Pond	In-out	Comp _{GW}	AGR/DPFV: seepage from the two ponds overlying impermeable till is insignificant. Seepage from the two ponds overlying permeable sand and gravel is large enough to double groundwater discharge (p.313)	AGR: . AGR: + DPFV: +
Hensel & Miller (1991)	Illinois, USA	GW/S	Floodplain	With-out	CCM	FPLM/WPGR: flood volumes decrease due to large retention and infiltration losses on the wide sandy flood plains (p.274)	FPLM: - WPGR: =
Nielsen <i>et al.</i> (1991)	India	FP	Peat	Drained	WB	MAF/MAAE/FPLM/DPFV: drainage of wetlands increased flows and reduced evapotranspiration losses. Peak flows were greater and low flows were higher (p.275)	MAF: - MAAE: + FPLM: - DPFV: -
Robinson <i>et al.</i> (1991)	Germany	GW/S	Fen	CompPET	WB	DPAE/DPFV: summer evapotranspiration is a major output - the water table drops and direct runoff disappears (p.161)	DPAE: + DPFV: -
Boeye & Verheyen (1992)	Belgium	GW/D					

	Zimbabwe	GW/S	Dambo	Multiple	LTH	MAF: dambos are an indiscriminatory factor in determining the volume, persistence and variability of annual runoff DPFV: dambos do not maintain dry-season low flows and reduce low flows where they occur in association with regolith with significant baseflow components FPHM: the impact of dambo on flood magnitude, variability and frequency is insignificant (p.349)	MAF: . DPFV: . DPFV: - FPHM: .
Bullock (1992a)	Zimbabwe	GW/S	Dambo	Same	Comp _{ET}	DPAE: evaporation losses in August 1986 are 1 mm day ⁻¹ higher on the dambo margins than the central dambo and interfluvial vegetation. (p.389)	DPAE: +
Bullock (1992b)	Madagascar	GW/S	Peat	Same	WB	AGR/GDS: fluxes between aquifers vary between downward and upward directions during the year (p.321).	AGR: = GDS: =
Grillot & Dussarrat (1992)	Indonesia	FP	Floodplain	With-out	CCM	FV: the frequency distribution of river levels without swamps is considerably broader than the actual situation. MAF: without swamps, MAF decreases (p.322)	FV: - MAF: +
Klepper (1992)	Newfoundland, Canada	SW/S	Blanket bog	Same	WB	AGR:6% of losses was to groundwater seepage FV: pipe-flow and high near-stream gradients coupled with the high transmissivity produce a flashy gradient (p.103)	AGR: = FV: +
Price (1992)	Paraguay R., S. America	FP	Pantanal	In-out	LTH	FV/MAAE/MAF: modification of the Pantanal may lose its condition of natural sponge responsible for exceptional stability of flow. A faster passage of water may also decrease evapotranspiration and increase the amount of outflow (p.36).	FV: - MAAE: + MAF: -
Eggelsmann et al. (1993)	Germany	SW/S	Mires	Paired	WB	MAAE/MAF: mires are characterised by high evaporation and correspondingly low runoff relative to mineral soils AGR: mires are not important in the regeneration of groundwater. Seepage rates from mires to deep groundwater is 80 to 50% lower in mires than in sandy soils (p.234-6).	MAAE: + MAF: - AGR: -
Gibson et al. (1993)	NW Territories, Canada	SW/S	Wetland	CompWC	Chem	DPFV: wetland rivers commonly freeze by mid-October and negligible flows occur during winter (p.216)	DPFV: .
Gilvear et al. (1993)	England, U.K.	GW/S	Fen	In-out	WB	GDS: groundwater inflow accounted for about 90% of water inputs FV/DPAE/DPFV: seasonal pattern of wetland outflow mirrors that of groundwater inflow, though amplitude of change is greater because high evapotranspiration depletes surface outflow during the summer (p.325-6)	GDS: = FV: + DPAE: + DPFV: -
John et al. (1993)	West Africa	FP	Floodplain	In-out	LTH	FPLM: the peak is flattened by vast areas of swamp vegetation MAF/MAAE: the water budget of the Massenya floodplain has inflow of 1.7 10 ⁹ m ³ and outflow is 0.8 10 ⁹ m ³ ; the water budget of the Yaere floodplain is inflows 3.2 10 ⁹ m ³ and outflow is 1.1 10 ⁹ m ³ (p.52,67)	FPLM: - MAF: - MAAE: +
Phillips & Shedlock (1993)	Delaware, USA	GW/D	Ponds	Comp _{GW}	Comp _{GW}	GDS: the hydrology of shallow seasonal ponds is strongly influenced by the adjacent groundwater-flow system (p.176)	GDS: =
Shedlock et al. (1993)	Indiana, USA	GW/S	Fen	Comp _{GW}	CCM	AGR/GDS: the interior of Great Marsh are discharge zones, whereas the margins are recharge areas during wet periods. (p. 152)	AGR: = GDS: =

Waddington <i>et al.</i> (1993)	Toronto, Canada	GW/S	Swamp	CompGW	Chem	FEV: saturated overland flow from permanently saturated areas created by discharging groundwater was the major storm runoff mechanism (p.37)	FEV: +
Woo & Rowsell (1993)	Saskatchewan, Canada	GW/S	Prairie slough	In-out	WB	AGR: prairie depressions are likely more effective for groundwater recharge than the uplands (p.205)	AGR: +
Woo & Winter (1993)	Permafrost USA	SW/S	Wetland	Same	LTH	FEV/FRR: the magnitude and duration of overland flow in wetland areas are likely to be greater than in uplands because the higher degree of saturation... and gentler gradients (p.28)	FEV: + FRR: -
Hey <i>et al.</i> (1994)	Illinois, USA	SW/S	Constructed	Same	WB	AGR: seepage was very low (0-6%) as a component of outflow (p.340)	AGR: X
Iritz <i>et al.</i> (1994)	Sweden	GW/S	Peat	Drained	CCM	FPLM: Swedish rivers showed decreased peak flows after drainage because of lowered groundwater. Peat in the Huhtisuo catchment had different topography and special hydraulic characteristics which resulted in an increase in peak flows. The effects of peat drainage on floods were negligible in the Svarian basin. The effect of drainage depends on the groundwater level (p.657-9)	FPLM: + FPLM: . FPLM: -
Johansson & Seuna (1994)	Sweden and Finland	GW/S	Sweden: bogs Finland: fens	Drained	CCM	Sweden: FPLM/MAF: neither peaks nor runoff differed DPFV: summer low flows were slightly higher Finland: FPLM: there was a slight increase in peaks DPFV: there was a slight increase in low flows MAF: total runoff volume increased by 3.5% (p.62-66)	FPLM: . FPLM: - MAF: . MAF: - DPFV: -
Price & Maloney (1994)	S.E. Labrador, Canada	GW/S	Domed bog and fen	In-out	CompGW	FEV: storm flows flooded the fen and water was quickly discharged. DPAE: the large depression and detention storage of both systems enhanced evapotranspiration losses (p.328)	FEV: + DPAE: +
Burt (1995)	UK	SW/S	Peat	Multiple	LTH	Flow duration curves from peat-covered basins comparatively demonstrate DPFV: minimal baseflow in summer because there is virtually no drainage from the peat (or clay) FPLM: very high flood runoff given widespread production of surface flow FV/a: steep slope and therefore, greater flow variability (p.25-26) AGR: the volume of water that discharges through the basal peat into the mineral sediments may be low (p.224)	DPFV: - FPLM: + FV/a: +
Doss (1995)	Maine, USA	GW/D	Bog	CompGW	LTH	AGR: groundwater recharged from the swamp will be less than from surrounding areas (p.29)	AGR: X
Khan (1995)	Malaysia	FP	Floodplain	Same	WB		AGR: -
Meigh (1995)	Botswana	FP	Floodplain	Same	CCM	MAF/MAAE: the wetland was a major cause of water loss from the Bokaa catchment losing about 12 Mm ³ yr ⁻¹ DPAE: the wetland is recharged by river flows and loses water by evapotranspiration through the dry season (p.38)	MAF: - MAAE: + DPAE: +
Owen (1995)	Wisconsin, USA	SW/S	Peat	In-out	WB	AGR: wetland did not play an important role in attenuating flood peaks DPFV: the wetland did not make significant contributions to streamflow under low river flow conditions (p.185)	AGR: X FPLM: . DPFV: .
Thompson & Hollis (1995)	Nigeria	FP	Floodplain	Same	WB	AGR: the wetlands play a vital role in aquifer recharge. The key is the annual wet season flooding (p.97)	AGR: =

Hollis (1996)	Senegal	FP	Floodplain	Same	WB	AGR: the surface aquifer is recharged by the vertical percolation of floodwater (p. 172)	AGR: =
Gerla and Mathoney (1996)	North Dakota, USA	SW/D	Wetlands around lakes	Comp _{pw}	CCM	WPGR/DPGR/GDS: Results show that groundwater flow into the Lake ranged from -0.030m ³ /day during the late winter to +0.043m ³ /day several times during the summer (p. 914)	WPGR: + DPGR: - GDS: =
Gonthier (1996)	Eastern Arkansas, USA	FP	Floodplain	In-out	Comp _{pw}	AGR: During a significant percentage of the monthly water level measurements surface water inundating the wetland had sufficient hydraulic head to flow through the confining unit and the upper part of the alluvial aquifer to the lower part of the alluvial aquifer and away from the Black Swamp (intermediate recharge). The greater the distance from the river, the less likely the groundwater from the site will flow to the river and the more likely it will flow to the lower part of the aquifer (p. 338)	AGR: =
Hunt <i>et al.</i> (1996)	Southwestern Wisconsin	FP	Floodplain	In-out	WB	GDS: Groundwater inflow rates range from 0.2-0.8 cm/day on the site with areas of higher inflows located closer to the river (p. 505).	GDS: =
Hodnett <i>et al.</i> (1997)	Amazonia, Brazil	FP	Floodplain	Comp _{pw}	Comp _{pw}	GDS: floodplain water levels are controlled primarily by discharge of groundwater which maintains dry season streamflow	GDS: =
Mathoney and Gerla (1996)	North Dakota, USA	SW/D	Wetlands around lakes	Comp _{pw}	CCM	AGR/GDS: although hydraulic gradients are upward over most of the year, discharge of water from the Dakota aquifer contributes less than a few tenths of a percent to the wetland water budget... the recharge-discharge function of Lunby and Stewart wetland is in a state of dynamic equilibrium (p. 119) WPGR: deep penetration may occur during repeated, relatively brief periods of reversed hydraulic gradient corresponding to spring recharge and periods of above normal rainfall (p. 118).	AGR: = GDS: = WPGR: =
Walton <i>et al.</i> (1996)	Eastern Arkansas, USA	FP	Floodplain	In/out	LTH	FTTP: Hydrograph peaks downstream of the wetland occurred 4 to 8 days later than at the upstream gauge (p.283) FPLM: The mean reduction in peak discharge between the two gauges was about 20% (p. 283)	FTTP: + FPLM: -
Hooijer (1996)	Shannon, Ireland	FP	Callow	Comp _r	Comp _r	FEV: The floodplains "have a considerable flood control capacity; if the minimum 3500ha of callow land is flooded to an average depth of 1m, this represents a storage equivalent to one day of Shannon peak discharge (around 400m ³ s ⁻¹)" (p. 195)	FEV: -
Hamilton <i>et al.</i> (1997)	Pantanal, Brazil	FP	Floodplain	WB	In-out	MAAE: discharges of inflowing rivers is approximately equal; to the outflow from the Pantanal on an annual basis...water lost by evaporation is roughly balance by direct precipitation (p. 258)	MAAE: +
Logan and Rudolph (1997)	La Plata, Argentina	SW/D	Marshes	Comp _{pw}	CHEM	AGR: vertical gradients in the marsh are generally downward throughout the year, suggesting that the Marsh is a predominantly groundwater recharge area, especially the lower lying, wetter areas (p. 229).	AGR: =

Hayashi <i>et al.</i> (1998a and 1998b)	Saskatchewan, Canada	GW/D	Slough	In-out	CHEM/WB	AGR: The vertical hydraulic gradient was downward throughout the year... leading to a total of loss of the system to groundwater of 480mm although only 2mm of groundwater flux becomes net groundwater recharge to the aquifer; the rest is transferred to the surrounding vegetated areas where it is evaporated MAAE: evapotranspiration in the upland planted with wheat is greater than in the wetland (380mm versus 300mm) (p. 53).	AGR: = MAAE: -
Hillman (1998)	Alberta, Canada	FP	Floodplain	In/out	SEH	FPFM: Flow downstream of the wetland was 6% of the peak flow upstream (0.953 m ³ s ⁻¹ downstream compared with 15 m ³ s ⁻¹ upstream of the wetland area). (p.53). FEV/FTTP: the wetland greatly attenuated the flood wave in terms of volume and timing (p. 53).	FPLM: - FEV: - FTTP: +
Ferrari <i>et al.</i> (1999)	Aral Sea, Uzbekistan	FP	Deltaic floodplain	In-out	WB	MAF/FPLM: The addition of wetlands in model grid boxes reduced both mean and summer peak flows (p. 1874).	MAF: - FPLM: -
Genereux and Slater (1999)	Everglades, USA	General	Canal	In-Out	WB	AGR: between 50-90% of water entering the canal each month was seepage from the target wetland (p. 166)	AGR: =
Gerla (1999)	Central Minnesota, USA	GW/S	Headwater wetlands	Comp _{GW}	Comp _{GW}	GDS: On the Shingobe river, wetlands around the river provide 46.3l/s on average to a stream where flow is about 200 l/s. (p. 400).	GDS: =
Hardy <i>et al.</i> (2000)	Devon, England	FP	Floodplain	In-Out	SEH	FPLM: floodplain reduced flood peak by 7% (p.212) FPFM: floodplain reduced flood peak by 19% (p. 212) FTTP: For two events, floodplain increased lag by 35 and 4 hours respectively (p.212).	FPLM: - FPFM: - FTTP: +
Raisin <i>et al.</i> (1999)	North-east Victoria, Australia	GW/S	Reed swamp	In-Out	WB	DPFV: although flow into the wetland was negligible over extended periods... a baseline discharge from the wetland of around 0.65 ML day was usually measured (p. 139) GDS: During the relatively dry detailed study period the groundwater flow component comprised an estimated 97% of the surface flow leaving the wetland (p. 139)	DPFV: + GDS: =
Speksma (1999)	Lower Saxony, Germany	SW/S	Bog	Drained/Pair	LTH	DPFV: The previous data show that most water yield is produced during winter and that perennial storage is not available to sustain flow during dry periods... rewetted raised bogs are not very effective as long term storage areas and regulators of stream flow DPAE: Water yields during summer were usually zero or very low, since most of the summer rainfall was lost through evaporation and transpiration. GDS/AGR: groundwater flow is not a typical feature of raised peat bogs	DPFV: - DPAE: + GDS: x AGR: x
Taylor and Howard (1999)	Uganda	GW/S	Swamp-filled drainage channels	Pair	LTH	WPGR: Between 1988-1993, recharge occurred only in one year and was dependant upon years of exceptionally heavy rainfall (p. 46). MAF: Surface runoff in the Aroca catchment (with wetlands) is 3mm compared to 34mm in the Nyabesheki catchment (without) (p. 67). MAAE: Isotopic data show that river waters have been subjected to significantly less evaporation than wetland waters... surface water have a prolonged residency, which leads to greater exposure to evaporation (p. 66).	WPGR: = MAF: - MAAE: +

Choi and Harvey (2000)	Everglades, USA	SW/D	Drainage channel	In/out	CHEM, WB	AGR/GDS: Estimated net groundwater fluxes were almost entirely negative values indicating that ground-water recharge commonly exceeded ground-water discharge...approximately 31% of the water supplied provides recharge to the underlying aquifer system. Groundwater discharge was negligible in comparison (2.8%) (p. 510)	AGR: = GDS: =
McCartney (2000)	Zimbabwe	GW/S	Dambo	In/out	WB	FEV: "... saturation overland flow, arising within the area of the dambo, is the principal mechanism of storm runoff generation"	FEV: +
Riekirk and Korhnak (2000)	Florida, USA	FP	Pond cypress wetlands	In/out	WB	AGR: Annual deep seepage into the underlying aquifer was 17 cm yr^{-1} for two wetlands, but in wetland N, underlain by a thick layer of blue clay this was only one cm yr^{-1} (p. 452).	AGR: =
Wolski (2002)	Okavango, Botswana	FP	Swamp	Comp _{gw}	Comp _{gw}	AGR: "... shallow groundwater is recharged by infiltration and lateral flow from the floodplain. A difference in flood level of 0.6m was accompanied by a difference in groundwater storage, expressed in terms of groundwater level of at least 3m."	AGR: +