



The Léon Domain (French Massif Armoricain): a westward extension of the Mid-German Crystalline Rise? Structural and geochronological insights

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Abstract The Léon Domain in the NW part of the French Massif Armoricain is a stack of synmetamorphic nappes displaced from south to north in ductile conditions. From bottom to top, an orthogneissic basement is overthrust successively by (1) a Lower Nappe of gneiss including mafic eclogites, (2) an Intermediate Nappe of biotite–garnet–staurolite micaschists with mafic blocks, and (3) an upper nappe made up of Neoproterozoic phyllites covered by unmetamorphosed Paleozoic sedimentary series. This microstructural study documents a polyphase evolution with firstly a top-to-the-N shearing, secondly followed by upright folding of the stack of nappes coeval with migmatization, and lastly, a dextral wrenching along the North Armorican Shear Zone associated with emplacement of synkinematic plutons. New U–Th/Pb chemical dating of monazites from biotite–garnet–staurolite micaschists, migmatites, and granitoids argue for 340–335 Ma, 335–327 Ma, and about 320 Ma ages for synthrusting metamorphism, anatexis, and wrenching, respectively. A metagabbro from Le Conquet yields a zircon LA-ICP-MS age of 478 ± 4 Ma, which corresponds to magma emplacement time. The Léon Domain is interpreted as a microcontinent separated from Armorica by the Le Conquet–Penzé suture to the south and east, and from Laurussia by the Rheic suture to the north. A possible correlation with the Mid-German Crystalline Rise of Central Europe is discussed.

Keywords Variscan belt - Microcontinent collision - Crustal melting - U–Th/Pb chemical dating - Paleozoic - Massif Armoricain - Mid-German Crystalline Rise

Introduction

The Paleozoic Variscan Orogeny was responsible for the formation of the basement of Medio-Europa from Iberia to central Europe. It is widely accepted that the Variscan Belt resulted from a complex interplay between the three large continents Laurentia, Baltica, and Gondwana, and several intervening microcontinents such as Avalonia or Armorica (e.g., Matte 1986, 2001; Franke 2000; Paris and Robardet 1990; Pin 1990; Tait et al. 1997; Robardet 2003; Von Raumer et al. 2002). Although authors acknowledge that the Variscan Belt is a collisional orogen, the location and age of the sutures, the timing of the tectonic–metamorphic events, and the resulting structures remain controversial.

In France, a Silurian–Devonian east–west trending, Eo-Variscan suture, corresponding to the closure of the Medio-European Ocean is recognized in the southern part of the Massif Armoricain (Fig. 1; Matte 1986; Matte and Hirn 1988; Faure et al. 2005). This suture extends eastward, north of the Massif Central, below the southern part of the Mesozoic Paris basin. Another suture, called the Rheic suture, is inferred to exist north of the Massif Armoricain, in the British Channel, along the Bray fault and probably between the Vosges and the Ardennes. The domain located between these two main sutures corresponds to the Armorica

microcontinent (Matte 2001). Most of the Massif Armorica belongs to Armorica. However, its northwestern part, called Léon Domain, consists of metamorphic rocks different from the nearby Paleozoic rocks of Central Armorica. The boundary between the Léon Domain and Armorica, which forms the main part of the Massif Armorica, is tectonic, and marked by mafic magmatic rocks. The geodynamic evolution of the Léon Domain is still in debate. Several interpretations have been proposed (e.g., Cabanis et al. 1979; Balé and Brun 1986; Rolet et al. 1986, 1994; Le Corre et al. 1989, 1991; Faure et al. 2005, cf. below). This paper provides new structural and geochronological evidence suggesting that the Léon Domain can be considered as a distinct microcontinent drifted from Armorica and rewelded to it by north-directed thrusting.

Geological setting: the Léon Domain in the Massif Armorica

The Massif Armorica is one of the largest areas where the Variscan belt can be analyzed. This composite massif can be subdivided into several paleogeographic or tectonic domains (e.g., Cogné 1974; Le Corre et al. 1991; Rolet et al. 1994). Its structure is dominated by E–W trending dextral shear zones (Fig. 1). Both the North Armorican shear zone (NASZ) and the South Armorican shear zone (SASZ) split eastward into several branches. These hundred kilometer-long ductile faults are Middle to Late Carboniferous structures, active around 325–300 Ma, and developed within an already assembled continent. Thus, when dealing with early Variscan tectonics (i.e. before Viséan), three domains can be recognized there, namely (1) the South Armorican Domain, (2) the North-Central Armorican Domain, and (3) the Léon Domain (Fig. 1).

The South Armorican Domain is bounded to the North by the Nort-sur-Erdre fault and the SASZ. The former corresponds to the Eo-Variscan suture. The South Armorican Domain is lithologically and structurally similar to the Massif Central, both are formed by pre-Viséan metamorphic rocks, migmatites, granitoids, and Paleozoic sedimentary rocks. It is well acknowledged that the South Armorican Domain is a stack of nappes overthrust to the south (e.g., Marchand 1981; Ballèvre et al. 1994; Faure et al. 2005 and enclosed references). By analogy with the Massif Central (Ledru et al. 1989), the following litho-tectonic units are recognized from bottom to top. The Para-Autochthonous Unit, composed of micaschist, grauwacke, and quartzite, is recognized in the southern part of the Massif. The Lower Gneiss Unit (LGU) consists of orthogneiss, paragneiss, and micaschist. The Upper Gneiss Unit (UGU) formed by orthogneiss, paragneiss, and mafic rocks experienced a high-pressure metamorphism followed by crustal melting. In contrast to the Massif Central where the UGU represents the uppermost unit, in the Massif Armorica, the UGU was successively overthrust by (1) an ophiolitic nappe and (2) a basement nappe. The ophiolitic nappe consists of highly deformed metagabbro, amphibolite, serpentized ultramafics, and metasedimentary rocks. The basement nappe, also called the Mauges Nappe, is made up of Neoproterozoic (i.e. Brioverian) metamorphic rocks formed during the Cadomian orogeny, unconformably overlain by Cambrian-Ordovician terrigenous sedimentary rocks. The stack of nappes developed during the collision between the South and North-Central Armorican domains representing the North Gondwana margin and the Armorica microcontinent, respectively.

The North-Central Armorican Domain, the widest one, is formed by unmetamorphosed and weakly deformed Paleozoic sedimentary series underlain by a Neoproterozoic basement formed during the Cadomian orogeny. The structure of the North-Central Armorican Domain is rather simple, E–W trending upright folds associated with a vertical axial planar cleavage developed during the dextral shearing along the NASZ and the SASZ. Several Middle to Late

Carboniferous granitic plutons intrude the sedimentary succession; some of them are synkinematic bodies coeval with the dextral shearing.

The Léon Domain (Fig. 1) is the smallest one of the Massif Armoricain. In spite of numerous Carboniferous plutons that partly hide the structural relationships between the units, an antiformal stack of nappes is presently recognized (Le Corre et al. 1989; Rolet et al. 1994; Faure et al. 2005; Fig. 2). From bottom to top, the structure of the Léon Domain consists of the following succession. The Para-Autochthonous Unit corresponds to paragneiss intruded by augen orthogneiss (Tréglonou and Plounevez-Lochrist), the emplacement of which is dated at 385 ± 8 Ma by Rb–Sr method on whole rock, and 400 ± 40 Ma by U–Pb method on zircon fractions (Cabanis et al. 1979). Both the augen gneiss and country rocks experienced a pervasive migmatization, the age of which is provided for the first time in this paper. The Lower Nappe is formed by a biotite–garnet–sillimanite gneiss series that encloses eclogitic rocks mostly retrogressed into amphibolite and derived from tholeiitic gabbro or basalt. Pyroxenites and serpentinites crop out locally. Pressure and temperature (P–T) conditions for the eclogitic metamorphism are estimated around 14–15 kbar and 650–700°C, respectively (Cabanis and Godard 1987). U/Pb dates from zircons provide an age of 439 ± 12 Ma, which is considered as that of the high-pressure metamorphism (Paquette et al. 1987). The Intermediate Nappe or Le Conquet–Penzé Nappe consists of ortho- and para-derived amphibolite, metagabbro, biotite–garnet–staurolite micaschist, and metachert. The micaschists commonly include biotite–garnet–staurolite metapelites, and locally garnet porphyroblasts reach 1 cm in size. The close association of blue kyanite with quartz and muscovite suggests that it crystallized during a postmetamorphic hydrothermal event rather than during the regional metamorphism (Chauris and Hallégouet 1989). The Upper Nappe consists of Late Proterozoic micaschists intruded by Ordovician granodiorites, presently deformed into orthogneiss, called the Brest orthogneiss. The Brest orthogneiss yields zircon U–Pb and whole rock Rb–Sr ages of 460 ± 70 and 466 ± 25 Ma, respectively. In spite of their poor quality, these dates argue for an Early Ordovician age for the pluton emplacement (Deutsch and Chauris 1965; Cabanis et al. 1979). Weakly deformed Ordovician white quartzite and pelite form the Paleozoic cover of the Proterozoic basement of the Central-North Armorican Domain that crops out immediately south of the Brest orthogneiss. Locally, NE of Brest and east of Penzé, a decollement separates the Paleozoic sedimentary cover from its Proterozoic basement (Fig. 2). To the south, the stack of nappes of the Léon Domain is cut by the dextral strike-slip Elorn fault, which is a branch of the North Armorican Shear Zone. To the east, around the Penzé River, the general trend of the bedding of the sedimentary rocks, the metamorphic foliation, and the tectonic boundaries turn N–S. The following section provides microtectonic analyses of the Le Conquet and Penzé areas that document the polyphase deformation and north-directed shearing. Several granitic plutons intrude the nappe pile (Fig. 2). The Aber-Ildut, Kernilis, Brignogan, and Plouescat plutons are Late Carboniferous meta or peraluminous granites or monzogranites (Le Corre et al. 1989, 1991). The St-Renan-Kersaint pluton is an E–W elongated body whose sigmoidal shape, internal structure, and S–C structures indicate an emplacement contemporaneous with the dextral shearing of the NASZ (Goré and Le Corre 1987; Le Corre et al. 1989; Schultz et al. 2007).

The tectonic evolution of the Léon Domain can be summarized as a sequence of four main events:

1. The HP metamorphism, dated at 439 ± 12 – 13 Ma (Paquette et al. 1987), developed in oceanic mafic rocks. Because of the retrogression under amphibolite facies conditions, the early syn-HP deformation of the eclogites is poorly understood.

2. The main event observed in all units is the top-to-the-NW synmetamorphic shearing developed under amphibolite facies conditions. In the Para-Autochthonous Unit, the Penzé-Le Conquet Intermediate Nappe, and the Paleozoic series of the Upper Nappe, this synmetamorphic deformation is the earliest one recorded there, and responsible for the stacking of nappes (Balé and Brun 1986, Le Corre et al. 1989; Rolet et al. 1994; Faure et al. 2005).
3. As it is the case in many orogens, crustal thickening was responsible for anatexis of the most pelitic and hydrated rocks. In the Léon Domain, crustal melting developed mainly in the Lower Nappe in relation to the formation of the antiformal structure that refolded the nappe stack.
4. Lastly, several generations of granitic plutons cut across the earlier structures. The St-Renan-Kersaint pluton, which is the largest one in the Léon Domain, is a synkinematic body emplaced along the dextral North Armorican Shear Zone (Goré and Le Corre 1987).

Structural analyses of the Le Conquet and Penzé areas

Le Conquet area

Because of excellent exposure conditions and its mean N–S orientation, the coast line between the Pointe St-Mathieu and Kerhornou is the most convenient place to analyze the structure of the southern boundary of the Léon Domain (Fig. 3; Chauris and Hallégouet 1989). The tectonic units and the geometry described in the previous section are well exposed. From south to north, they are (1) the Brest orthogneiss and its Neoproterozoic Elorn micaschists host rocks, (2) the Conquet micaschists with mafic rocks, and (3) the Kerhornou migmatitic gneiss with numerous pods and dykes of anatectic granitoids. To the north, near Kerhornou, migmatites are well developed and the protoliths of the migmatite cannot be identified any more. The North Armorican shear zone separates the Le Conquet metamorphic rocks from the St-Renan-Kersaint pluton. Along the NASZ, the granite is pervasively foliated with a subvertical attitude and contains a subhorizontal mineral lineation. S–C shear criteria confirm the dextral sense of shear along the NASZ.

The Brest orthogneiss exhibits a steeply southward dipping NE–SW trending foliation in which two sets of stretching lineation can be distinguished (Fig. 3). The dominant one is well observed close to the northern boundary of the orthogneiss with Le Conquet metamorphic rocks. The vertical foliation (XY plane) contains a subhorizontal or weakly plunging to the southwest stretching lineation (X axis) represented by biotite streaks, quartz rods, elongated or boudinaged feldspar augens. In agreement with Balé and Brun (1986), kinematic indicators such as sigma-type porphyroclast systems, meter-scale shear zones, or shear bands provide a dextral sense of shear. Moreover, an early N155E to N170E steeply plunging lineation is also observed on the orthogneiss foliation. In sections perpendicular to the foliation and parallel to the N–S trending lineation, sigma-type porphyroclast systems and shear bands indicate top-to-the-N ductile shearing. This lineation is less developed than the NW–SE one as strike-slip shearing partly erased it. The boundary between the Brest orthogneiss and the Le Conquet micaschists is a strike-slip shear zone. However, an older, north-directed thrust contact, coeval with the N–S stretching lineation, can be documented on the basis of the microtectonic analysis of Le Conquet micaschists.

North of the wrench fault, metagabbro, diabase, and amphibolite boudins form a remarkable lithology within the Le Conquet micaschists series. Quartzite and conglomeratic lenses also develop in the metapelites (Fig. 4). Because of the poor outcrop conditions inland, the

geometry and significance of these lenses remain uncertain. They might be continuous layers, interstratified with the metapelites, or olistoliths. Quartz lenses are also pervasively distributed throughout the Conquet micaschists (Fig. 4). The locally high concentration of veins suggests that they formed along ductile shear zones with abundant fluid circulation. In this unit, a pervasive NNE–SSW lineation is marked by elongated amphiboles in the mafic rocks, stretched clasts in the paragneiss, quartz pressure shadows around garnet in micaschists, or elongated quartz rods. Kinematic indicators such as asymmetric pressure shadows, sigma-type porphyroclasts systems, sigmoidal biotite, rotating inclusions in garnet or asymmetric quartz veins show a top-to-the-N sense of shear (Figs. 4, 5, 6).

The northernmost unit of the Le Conquet area is the Kerhornou biotite–garnet–sillimanite paragneiss. These rocks contain abundant centimeter- to decimeter-sized pockets of quartz–feldspar–biotite leucosomes related to an incipient crustal melting. The transition from high-grade gneiss to sillimanite metatexites and diatexites is progressive. In places, meter- to several meter-sized granitoid dykes with numerous xenoliths of biotite–sillimanite gneiss correspond to anatectic granites (Fig. 4). As a whole, the Kerhornou gneiss, migmatites and anatectic granites suite is the lowest unit observed in the Le Conquet area. It belongs to the Lower Nappe of the Léon Domain (Fig. 2) in which a high-temperature metamorphic and plutonic event overprinted the HP eclogitic metamorphism. However, in the Le Conquet section, HP rocks have not yet been identified. In this unit, the mineral lineation is less developed than in the ones described above, but NNE–SSW trending biotite streaks or quartz rods are also observed where migmatization is absent. In the Kerhornou gneiss, kinematic indicators are rare and the synmetamorphic sense of shear is difficult to assess. The stretching lineation of the Kerhornou gneiss is parallel to that found in the Le Conquet micaschists. This observation suggests that in both units the lineation formed coevally, during north-directed nappe stacking. In gneiss where anatexis is not well developed, the older fabric of the protolith is still preserved. This shows that the above described north-directed ductile deformation is older than crustal melting. Moreover, the foliation is deformed by an E–W trending, heterogeneously developed crenulation cleavage. Millimeter- to centimeter-scale microfolds are overturned to the north. This structure, which is also observed in the migmatite, is likely related to the upright folding of the stack of nappes.

Penzé area

The eastern boundary of the Léon Domain is well exposed along the Penzé River (Fig. 7). There, the regional structural trend is north–northeasterly, but southeast of Penzé, the foliation turns into a NE–SW direction. From bottom to top, the following units are recognized. The Penzé micaschists consist of biotite–muscovite–oligoclase metapelites. Locally, staurolite and garnet are also present, but in contrast to the Le Conquet area, centimeter-sized porphyroblasts are rare. Mafic rocks such as metagabbro, amphibolite, mafic schists form elongated lenses scattered in the metapelites. Because of the intense ductile shearing and recrystallization, it is rather difficult to state if these rocks are olistoliths or tectonic boudins. The Penzé micaschists are overlain by sericite phyllite, metagrauwacke, and metasandstone. In the sericitic quartz phyllite, detrital plagioclase grains are common, while being quite rare in the underlying micaschists. To the east, this series is covered by unmetamorphosed conglomerate, quartzite, sandstone, and mudstone of Paleozoic age. Furthermore, Ordovician granites, changed into orthogneiss, intrude the sericitic phyllite (Fig. 7). Therefore, a Neoproterozoic age (Brioverian) is likely for the micaschists. At the scale of the Léon Domain, the Penzé micaschists, Neoproterozoic phyllites, and orthogneiss are correlated with the Le Conquet micaschists, Elorn micaschists, and Brest orthogneiss, respectively (Fig. 2).

Although never observed in the field, but in agreement with previous works (Rolet et al. 1994; Chauris et al. 1998), a tectonic contact between the Penzé micaschists and the overlying Neoproterozoic Cadomian micaschists is assumed (Figs. 7, 8). West of Morlaix, the sandstone and conglomerate that form the base of the Devonian sedimentary series show a well-developed bedding-parallel cleavage and a stretching lineation.

Except for the Paleozoic sedimentary rocks, the metamorphic rocks exhibit a N–S- to NE–SW-trending and E- or SE-dipping pervasive foliation and an E–W- or NW–SE-trending stretching lineation (Fig. 5). Shear criteria, such as sigmoidal micas in micaschists, sigma-type porphyroclast systems in orthogneiss, chlorite pressure shadows around hornblende clasts in metagabbro, indicate a top-to-the-west (or NW) sense of shear (Fig. 6). The foliation is also deformed by west-verging or upright NE–SW-trending folds often associated with a crenulation cleavage.

In addition to the Penzé and Le Conquet key areas, microstructural observations have been also carried out in several other places. For instance, the Tréglonou and Plounévez-Lochrist orthogneiss exhibit a NNW–SSW stretching lineation and top-to-the-NNW sense of shear (Fig. 5), supporting the interpretation that the nappes of the Léon Domain moved from south (or southeast) to north (or northwest).

Geochronological constraints

Previous radiometric dating

As radiometric constraints supporting the abovepresented evolution are rare and rather ancient (Deutsch and Chauris 1965; Leutwein et al. 1969; Cabanis et al. 1979), a geochronological investigation has been carried out. The Léon Domain has long been considered as a piece of the Proterozoic basement of northern Brittany and Normandy (e.g., Cogné 1974 and enclosed references). However, radiometric dating has clearly demonstrated that the main structure of the Léon Domain was acquired during the Paleozoic Variscan orogeny. The augen orthogneiss of Tréglonou and Plounévez-Lochrist, which form the lowermost unit, yield a whole rock (WR) Rb–Sr age of 385 ± 8 Ma and a zircon U/Pb age of 400 ± 40 Ma (Fig. 9; Cabanis et al. 1979). These Early Devonian ages are considered as the ages of the emplacement of the granitic protolith. The Brest orthogneiss yields a zircon U/Pb age of 460 ± 70 Ma (Deutsch and Chauris 1965). The zircon U/Pb measurements from an eclogite of the Lower Nappe gave a Concordia lower intercept at $439 + 12/- 13$ Ma that is interpreted as the age of the high-pressure metamorphism (Paquette et al. 1987).

The St-Renan-Kersaint pluton has been dated by the Rb/Sr method on whole rock at 320 Ma (Deutsch and Chauris 1965; Leutwein et al. 1969), and around 350–340 Ma by U/Pb method on zircon (Deutsch and Chauris 1965; Leutwein et al. 1969). A biotite concentrate of the same rock, yielded a K–Ar age of 340 ± 15 Ma (Leutwein et al. 1969). Thus, the ages of the metamorphic events are almost not documented. The age of the anatexis is sometimes assumed as that of the St-Renan pluton, since the migmatites are considered to have resulted from the thermal metamorphism related to pluton intrusion. However, for some authors, anatexis should have followed closely the high-pressure metamorphism and thus has been assumed as Silurian or Devonian (Le Corre et al. 1989; Ballèvre et al. 1994).

Monazite U–Th/Pb chemical ages

Over the past 15 years, chemical geochronology on monazite has become possible (e.g., Suzuki and Adachi 1991; Montel et al. 1996; Cocherie et al. 1998, 2005). Because of its high U–Th contents and negligible common Pb content (Parrish, 1990), monazite constitutes one of the phases frequently used in geochronology. Its use as radiochronometer is enhanced by the restricted lead diffusion in the monazite lattice (Montel et al. 1996; 2000; Cocherie et al. 1998) and by its temperature stability up to more than 900°C (Braun et al. 1998). To understand accurately the thermal and tectonic history experienced by a rock, geochronological data must be acquired in situ, that is to say, grains are directly analyzed with respect to their textural environment in thin section (e.g., Pyle et al. 2001; Williams and Jercinovic 2002; Be Mezème et al. 2005). Scanning electron microprobe (SEM) in back-scattered electron (bse) mode is one of the best means to distinguish constituent mineral species of the thin section. Detailed analytical procedures were given by Cocherie et al. (1998). The theoretical procedures to reduce the data and to calculate the average age from individual spot analyses are given by Cocherie and Albarède (2001), Cocherie et al. (2005), and Cocherie and Legendre (2007). The analyzed grains from thin section were prepared and analyzed by a Cameca SX 50 EPMA co-operated by BRGM-CNRS-Orléans University. The 2σ errors given on individual ages depend on U, Th, and Pb contents and are calculated by propagating the uncertainties of these elements (with 95% confidence level). However, because the mean age of an homogeneous age population can be calculated directly from the isochron plot Th/Pb versus U/Pb, only the uncertainty on these two parameters (Th/Pb and U/Pb) are required to calculate the uncertainty of the mean age. Age calculations were done using the Isoplot/Ex (3.1) program of Ludwig (2004) and a Microsoft Excel add-in program for determining U–Th–Pb ages from EPMA measurements (Pommier et al., 2002).

In the study area, one sample of biotite–garnet–staurolite micaschist (LE 5), three samples of migmatite (LE 21, LE 23, LE 27), one of anatectic granite (LE 17), and one of the St-Renan granite (LE 20) have been analyzed.

A biotite–garnet–staurolite micaschist (LE 05) has been sampled in Le Conquet (N 48°21'32"/W 4°47'00"). Five, non-zoned grains with sizes ranging from 50 to 100 μm have been analyzed. The intercept ages are similar within errors, since the U–Pb age (intercept with U–Pb axis) and Th–Pb age (intercept with Th–Pb axis) are 359 ± 31 – 36 and 322 ± 28 – 24 Ma, respectively. The MSWD (1.1 for 93 analyses) is compatible with a single age population. Monazite grains yield a mean age of 338 ± 5 Ma (at 2σ confidence level; Fig. 10).

Sample LE 21 is a metatexite that crops out north of the Kerhornou gneiss (Fig. 3; N 48°24'30"/W 4°46'22"). The five analyzed monazite grains are included in biotite; their size ranges from 40 to 150 μm . The intercept ages are similar within errors, since the U–Pb age (intercept with U–Pb axis) and the Th–Pb age (intercept with Th–Pb axis) are at 367 ± 59 – 75 and 302 ± 39 – 31 Ma, respectively. In addition, the MSWD of 1.4 indicates that the population of 157 analyses is statistically homogeneous within the analytical errors (Wendt and Carl 1991). As a consequence, it is possible to calculate the mean age at the centroid of the population, where the precision is the best within the error envelope. Thus, the monazite grains yield a mean age of 325 ± 5 Ma (at 2σ confidence level; Fig. 11).

Sample LE 23 is a metatexite from the right bank of Aber Benoît (N 48°33'50"/W 4°35'). Four monazite grains included in biotite with sizes ranging from 50 to 270 μm have been analyzed. The grains are not zoned and the intercept ages are similar within errors, since the

U–Pb age (intercept with U–Pb axis) and the Th–Pb age (intercept with Th–Pb axis) are at $323 + 50/-67$ and $343 + 56/-42$ Ma, respectively. On the other hand, an acceptable MSWD of 1.5 has been calculated from 115 analyses. Monazite grains yield a mean age of 334 ± 6 Ma (at 2σ confidence level; Fig. 12).

Sample LE 27 is a metatexite from an old quarry south of Plounévez-Lochrist (N $48^{\circ}36'27''$ /W $4^{\circ}12'20''$). Five nonzoned monazite grains included in biotite or in muscovite with sizes ranging from 60 to 170 μm have been analyzed. Because of rather constant Th/U ratio, the regression line is poorly defined and is responsible for a large error envelope. Nevertheless, the intercept ages are similar within errors, since the U–Pb age (intercept with U–Pb axis) and the Th–Pb age (intercept with Th–Pb axis) are at $306 + 127/-246$ and $347 + 148/-80$ Ma, respectively. The low MSWD parameter (1.2) also indicates that the 106 analyses can be associated to the same event. Monazite grains yield a mean age of 331 ± 6 Ma (at 2σ confidence level; Fig. 13).

The anatectic granite (LE 17) has been sampled south of the LE 21 migmatite (N $48^{\circ}22'21''$ /W $4^{\circ}45'31''$). Eight nonzoned grains with sizes ranging from 50 to 100 μm have been analyzed. The intercept ages are similar within errors, since the U–Pb age (intercept with U–Pb axis) and the Th–Pb age (intercept with Th–Pb axis) are at $230 + 135/-201$ and $368 + 84/-58$ Ma, respectively. The MSWD is also compatible with a single age population (1.2 for 54 analyses). The limited number of analyses and the low U + Th content lead to a higher error on the mean age. Monazite grains yield a mean age of 327 ± 15 Ma (at 2σ confidence level; Fig. 14).

The St-Renan-Kersaint granite (LE 20) has been sampled in a quarry west of Brest (N $48^{\circ}25'9''$ /W $4^{\circ}38'33.6''$). The size of nine monazite nonzoned grains included in biotite and muscovite ranges from 30 to 100 μm . The spread of the data is quite large, which lead to a narrow error envelope and subsequently rather precise age intercepts. They are similar within errors, since the U–Pb age (intercept with U–Pb axis) and the Th–Pb age (intercept with Th–Pb axis) are at $320 + 25/-26$ and 322 ± 9 Ma, respectively. In other words, the regression line (MSWD = 1.1 for 81 analyses) fits perfectly with the theoretical isochron at 321 Ma. Monazite grains yield a mean age of 321 ± 5 Ma (at 2σ confidence level; Fig. 15).

Zircon LA-ICP-MS age of Le Conquet gabbro sample LE 04

After a density separation of heavy minerals, hand-picked zircon grains were mounted in epoxy block, polished to obtain an even surface. Cathodoluminescence (CL) imaging was performed at the University of Sciences and Technologies of Lille (UMR 8110 PBDS) with a scanning electron microscope Quanta 200 (FEI, US). Analyses on single grains were made using the Neptune MC-ICP-MS (ThermoElectron, Bremen, Germany) at BRGM (Orléans, France) equipped with a multi-ion counting system, allowing a very high sensitivity (Cocherie and Robert 2007), and a laser ablation system (new wave frequency-quintupled Nd:YAG UV laser, distributed by VG, UK) operating at 213 nm. The crater was 20 μm in diameter and 15–20 μm deep. Argon gas was used as carrier gas. Zircon standard used is 91500. Standard bracketing was applied to correct both elemental fractionations during the ablation process and mass bias originating from the MC-ICP-MS itself. Detailed instrumentation and analytical accuracy description are given in Cocherie et al. (2007) and Cocherie and Robert (2008). Tera and Wasserburg's (1972) diagram was made using Isoplot/Ex (3.1) of Ludwig (2004).

Zircon grains are limpid, slightly pinky, and most of the time asymmetric, deformed, and with cracks. CL pictures showed two populations, one with a low CL (Fig. 16b), homogeneous unzoned or rarely slightly zoned, and the second with a higher CL (Fig. 16a) but often showing concentric or subconcentric growth zonations. Analyses data are given in Table 1.

On the Tera and Wasserburg's (1972) diagram (Fig. 16), two populations can be distinguished. The first one is characterized by a mean $^{238}\text{U}/^{206}\text{Pb}$ ratio of 8.5 ± 0.37 and a mean $^{207}\text{Pb}/^{206}\text{Pb}$ ratio of 0.0571 ± 0.0021 . These analyses show relatively low U, Th, and Pb concentrations and correspond mainly to the light CL zircons population, or cracked zircons. We interpret these analyses as a result of loss of U and a small amount of common Pb for the two points above the Concordia.

The analyses ($n = 8$) of the second population give concordant ellipses. However, three of them are significantly shifted indicating a slight radiogenic Pb loss. Thus, only five analyses were considered in the calculation of the mean $^{238}\text{U}/^{206}\text{Pb}$ age of 478 ± 4 Ma.

Discussion of the geochronological data

Our radiometric results confirm the relative chronology proposed above. In the biotite–garnet–staurolite micaschist of the Le Conquet metamorphic series (LE 05), monazite appears either as inclusion in biotite or along the biotite grain boundaries. These textural observations indicate that monazite crystallized during the intermediate pressure–intermediate temperature metamorphism. Thus, the approximate 340 Ma U–Th/Pb age can be considered as the age of the regional metamorphism. The close similarities between the monazite U–Th/Pb ages of the migmatites and anatectic granite indicate that crustal melting took place around 335–325 Ma. The 321 ± 5 Ma age of the St-Renan pluton appears slightly younger than the migmatization; however, the migmatite and granite yield the same age in the error limit. Thus, the genesis of the magma source of the St-Renan granite can be due to the same thermal event responsible for crustal melting.

The approximate 480 Ma age of the Le Conquet gabbro is close to that of the Brest orthogneiss. At the scale of the entire French Variscan Belt, Ordovician is acknowledged as the time when the northern margin of Gondwana experienced rifting responsible for the separation of several microcontinents such as Armorica and Avalonia (e.g., Matte 2001; Faure et al. 2005). Several alkaline or calcalkaline granitic plutons and mafic rocks (gabbro, diorite) were emplaced during this crustal thinning. These rocks were subsequently transformed into orthogneiss and metagabbro during Variscan tectonics. In the study area, the Brest orthogneiss probably belongs to the magmatic rocks formed in such an extensional setting. The mafic rocks of the Penzé-Le Conquet nappe might also be emplaced during the Ordovician rifting episode.

On the basis of field observation, the age of the synmetamorphic top-to-the-N (or NW) ductile shearing, coeval with nappe stacking in the Léon Block, is older than the anatexis dated here at 335–325 Ma and younger than the Devonian rocks involved in the deformation in the Penzé area (Fig. 7). Our monazite U–Th/Pb age of 338 ± 5 Ma (i.e. Viséan) is in agreement with this geological conclusion. Another ductile event is the dextral shearing along the NASZ and Elorn wrench faults (Figs. 3, 4). As observed in the Le Conquet area, the wrenching overprints the top-to-the-N shearing. In agreement with previous conclusions (e.g., Goré and Le Corre 1987; Le Corre et al. 1989; 1991; Rolet et al. 1994), the age of the dextral shearing can be inferred from that of the synkinematic St-Renan pluton. Our monazite U–Th/Pb dating

indicates that the NASZ was active at about 320 Ma (Fig. 9). Moreover, the HP metamorphism is dated at about 440 Ma (Paquette et al. 1987). Such a large time span of about 100 Ma between the HP metamorphism and crustal melting appears as quite unusual in orogens. Thus, even if the chronological data are correct, the Carboniferous metamorphism and crustal melting cannot be interpreted as a thermal and structural response of the exhumation of the HP rocks. However, the possibility of a younger age for the HP metamorphism remains an eventuality to be tested in forthcoming works.

The place of the Léon Domain within the Variscan framework

Possible geodynamic evolution of the Léon Domain

The Intermediate Nappe, composed of Penzé-Le Conquet micaschists and mafic rocks, and the Lower Nappe made of gneiss and eclogites, have been interpreted as relics of a suture zone separating the Léon and the North-Central Armorican domains (Fig. 2; Rolet et al. 1994; Faure et al. 2005). However, these rocks are not similar to typical ophiolitic suites, since mafic and ultramafic rocks are rare and true radiolarites are not recognized. Nevertheless, the geochemistry of the eclogites indicates that they are oceanic tholeiites formed in an oceanic basin (Cabanis and Godard 1987; Paquette et al. 1987). Thus, in spite of still poorly documented paleogeographic or geochemical constraints, it is possible that the Léon Domain and the North-Central Armorican Domain were separated by an oceanic basin or at least by a rift underlain by thinned continental crust. During the Variscan tectonics, the Léon Domain was overthrust to the North by the North-Central Armorican Domain. In this case, the Penzé-Le Conquet micaschists and mafic rocks could be interpreted as a suture separating the two domains. Nevertheless, the width of the oceanic domain remains unknown. A wide oceanic area has been speculated (Faure et al. 2005), but a smaller basin similar to the Brévenne basin that developed in Devonian in the eastern Massif Central (Fig. 1; Ledru et al. 1989; Leloix et al. 1999) is also possible.

The Léon Domain has also been considered as an “exotic terrane” initially belonging to the Upper Gneiss Unit of the South Armorican Domain and translated to the north by the Elorn strike-slip fault (Balé and Brun 1986; Le Corre et al. 1989). However, as already pointed out (Rolet et al. 1986, 1994; Faure et al. 2005), such an interpretation does not match the available structural data. In the Penzé area, at the northeastern termination of the Léon Domain, such a strike-slip fault is not recognized (Fig. 7). Conversely, the pervasive low-angle S- or SE-dipping foliation and NNW–SSE to E–W trending lineation show that the Central Armorican Domain has been overthrust onto the Léon Domain (Fig. 8). Therefore, the Penzé-Le Conquet Intermediate Nappe is rooted below the North-Central Armorican Domain. Obviously, in Le Conquet, the root zone was reworked along NE–SW-trending dextral wrench faults, which keep an E–W trend to the east but do not reach the Penzé area.

The northern limit of the Léon Domain is presently hidden below the sea. A gravity and magnetic anomaly that occupies the central part of the Channel is interpreted as the Rheic suture (Bois et al. 1994; Fig. 1). Several seismic lines through the Channel indicate dominantly south-dipping reflectors (Le Gall 1990). In particular, the ophiolitic nappe that crops out in the Lizard Point is dipping southward below crystalline rocks interpreted as the offshore extension of the Léon Domain. The Léon Domain appears as a continental block bounded by the Rheic and Le Conquet sutures on the north and south. Both sutures are distinct from the Eo-Variscan suture that separates the North-Central Armorican and the South Armorican domains. A simplified, crustal scale tectonic interpretation of the Massif

Armorica is depicted in Fig. 17. In this cross section, the late orogenic granitic plutons are not represented to emphasize the earlier structures.

Possible extension of the Mid-German Crystalline Rise

As the Léon Domain is not an “exotic terrane” displaced from the south, its equivalent must be located east of the Paris Basin (Fig. 1). In Central Europe, it is well acknowledged that the Ardenne Massif belongs to the Rheno-Hercynian Zone, which is the southern continental margin of Laurussia. The southern boundary of the Rheno-Hercynian Zone is represented by the Northern Phyllite Zone (NPZ), where a Carboniferous high-pressure metamorphism is associated with north-directed thrusting. South of the NPZ, the Paleozoic rocks that crop out in the Spessart and Odenwald massifs are parts of a microcontinent called the Mid-German Crystalline Rise (MGCR, e.g., Holder and Leveridge 1986; Hirschmann and Okrusch 1988; Oncken 1997; Altherr et al. 1999; Franke 2000; Will and Schmädicke 2001 and enclosed references).

In spite of a large lithological diversity and structural complexity, the MGCR contains Ordovician orthogneiss, Devonian-Carboniferous arc magmatic rocks, HT metamorphic and migmatitic rocks. Eclogites have also been recognized in the Odenwald Massif (Will and Schmädicke 2001). Farther south, the Saxothuringian basin separates the MGCR from the Tepla-Barrandian Zone that is correlated with the North-Central Armorican domain, both of them belonging to the Armorica microcontinent (Fig. 1; Franke 2000; Matte 2001). In this framework, the Léon Domain could be correlated with the MGCR (Fig. 1; Cabanis et al. 1979, Faure et al. 2005). Since the northern part of the Léon Domain is hidden below the sea, the limited surface exposure does not allow us to recognize subduction-related rocks as described in the Odenwald and Spessart massifs that would lie more to the north, and thus closer to the subduction zone (Oncken 1997; Franke 2000).

The boundary between the MGCR and Laurussia (i.e. Rheno-Hercynian Zone) corresponds to the Rheic or Rheno-Hercynian Suture, depending on authors (e.g., Franke 2000; Matte 2001). The continuation between the Rheic Suture and the NPZ is hidden below the Paris Basin and was possibly offset left-laterally along the Bray fault (Matte 1986; 2001). This correlation is supported by seismological, gravimetric, and magnetic data (Matte and Hirn 1988; Edel and Weber 2004). The boundary between the MGCR and Armorica does not crop out in Central Europe, except in the Bohemian Massif. A complete discussion of these elements is beyond the scope of this paper [see Franke (2001) for a comprehensive review].

Conclusion

The Léon Domain appears to be a microcontinent separated from Armorica in Early Ordovician time. This is suggested by the Brest orthogneiss and the mafic rocks observed in the Le Conquet and Penzé areas. The Léon microcontinent was rewelded to Armorica after a south-directed subduction coeval with high-pressure metamorphism, the age of which remains poorly constrained. The Léon Domain experienced a MP/MT metamorphism dated here at about 340 Ma by the EPMA method. The stack of the metamorphic nappes was subsequently folded into an antiform and experienced a HT metamorphism and crustal melting at about 335–327 Ma before the dextral strike-slip faulting along the North Armorican Shear Zone and the Elorn fault. Within the European Variscan framework, the Léon Domain is correlated with metamorphic and plutonic massifs of the MGCR. As a whole, the MGCR and its possible westward extension can be interpreted as a continental strip that separated from the Amoriga

microcontinent in the Early Ordovician. The oceanic or continental nature and the width of the area that separated Armorica and MGCR cannot be determined in the Léon area. Although less studied than Avalonia or Armorica, the MGCR probably played an important role in the complex plate tectonic evolution leading to the formation of the Variscan orogen.

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Figures

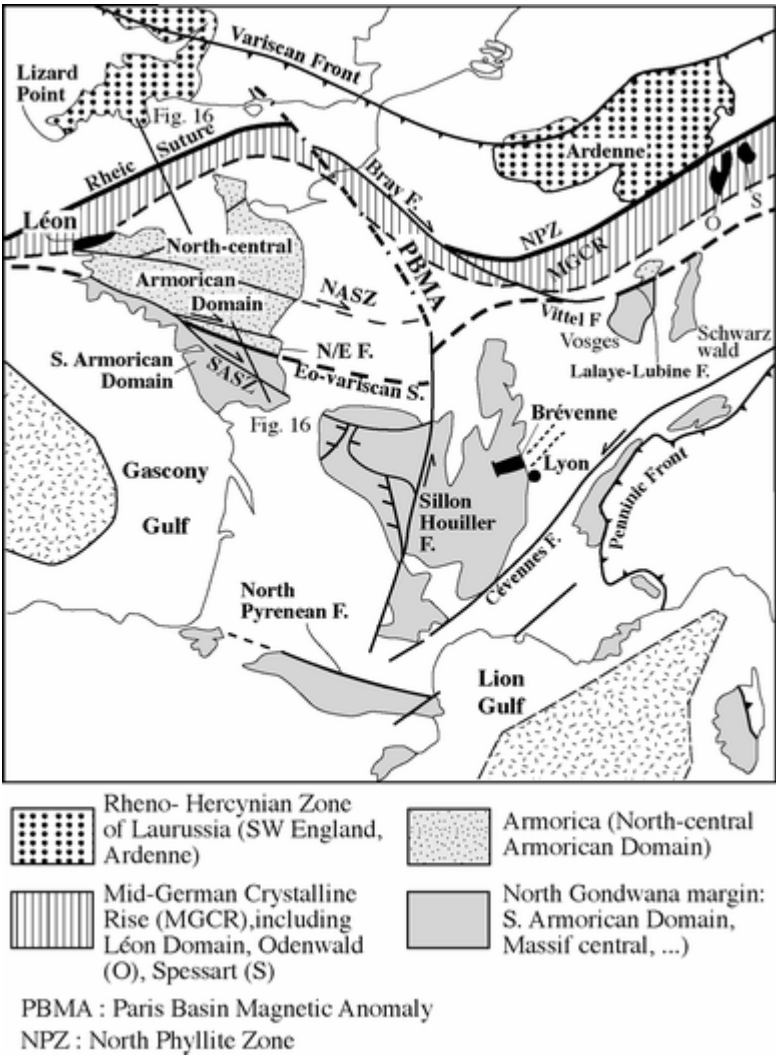


Fig. 1 Tectonic map of the Variscan Belt of Western Europe with emphasis on the main suture zones, litho-tectonic domains, and location of the Léon Domain. *NASZ* North Armorican Shear Zone, *SASZ* South Armorican Shear Zone, *N/E F* North-sur-Erdre Fault

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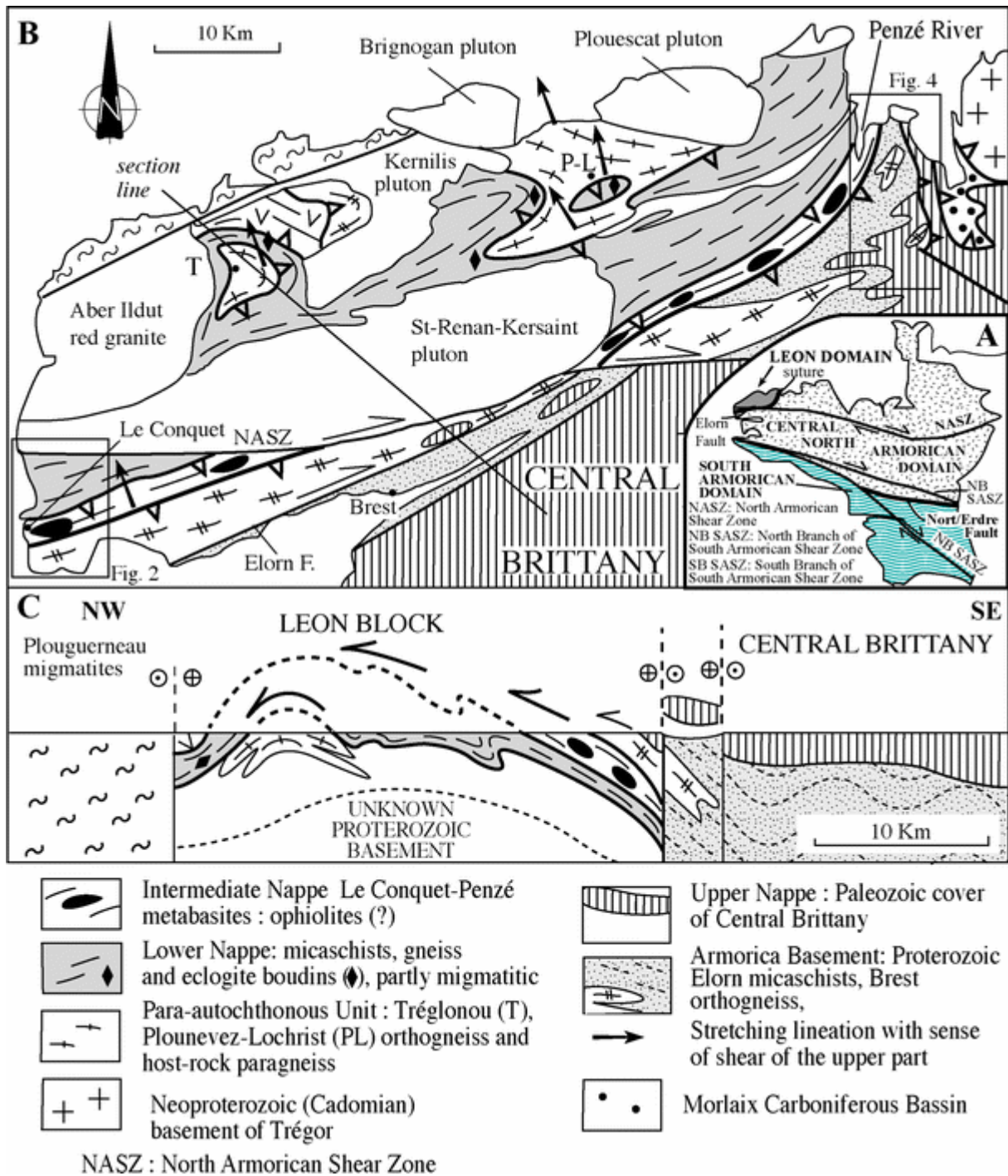


Fig. 2 Structural map and cross section of the Léon Domain showing the stack of nappes formed by top-to-the-N shearing: (a) subdivisions of the Massif Armoricain; (b) tectonic map; (c) section along the *line* located in (b)

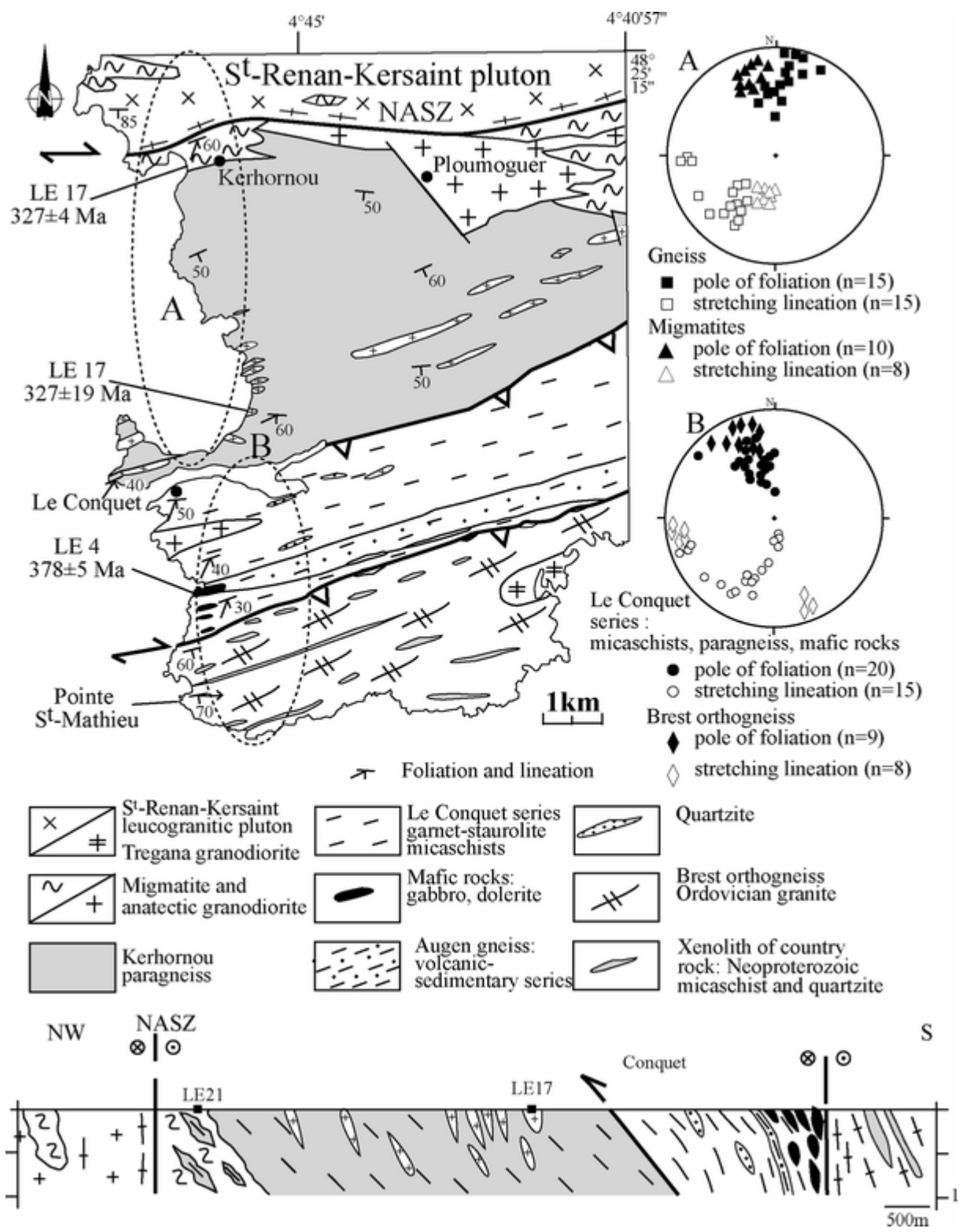


Fig. 3 Structural map of the Le Conquet area simplified from Chauris and Hallégouet (1989), with location of the dated samples. Simplified cross section showing the south-dipping foliation, and the gradual transition between the Kerhornou gneiss and migmatites. Stereograms (Schmidt projection, lower hemisphere) of the structural elements, (a) gneiss and migmatites, (b) Le Conquet micaschists and Brest orthogneiss

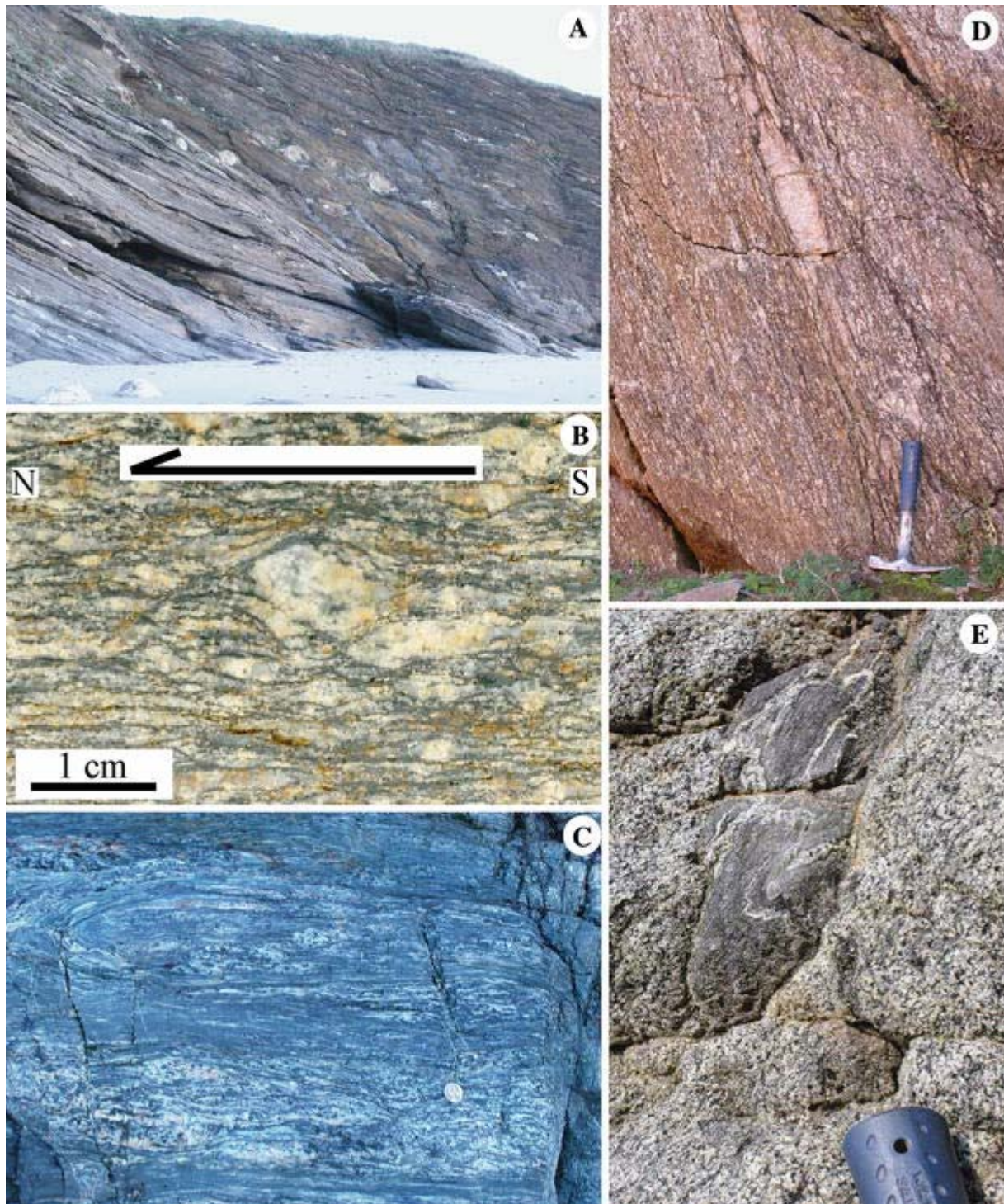


Fig. 4 Field photographs of Le Conquet area: (a) Le Conquet micaschists with abundant quartz veins; (b) asymmetric augen in the volcanic-sedimentary paragneiss of the Le Conquet series; (c) Mylonitic gabbro in Le Conquet; (d) Migmatite near Kerhornou; (e) Anatectic granite (sample LE 17) with xenolith of foliated and folded Kerhornou gneiss

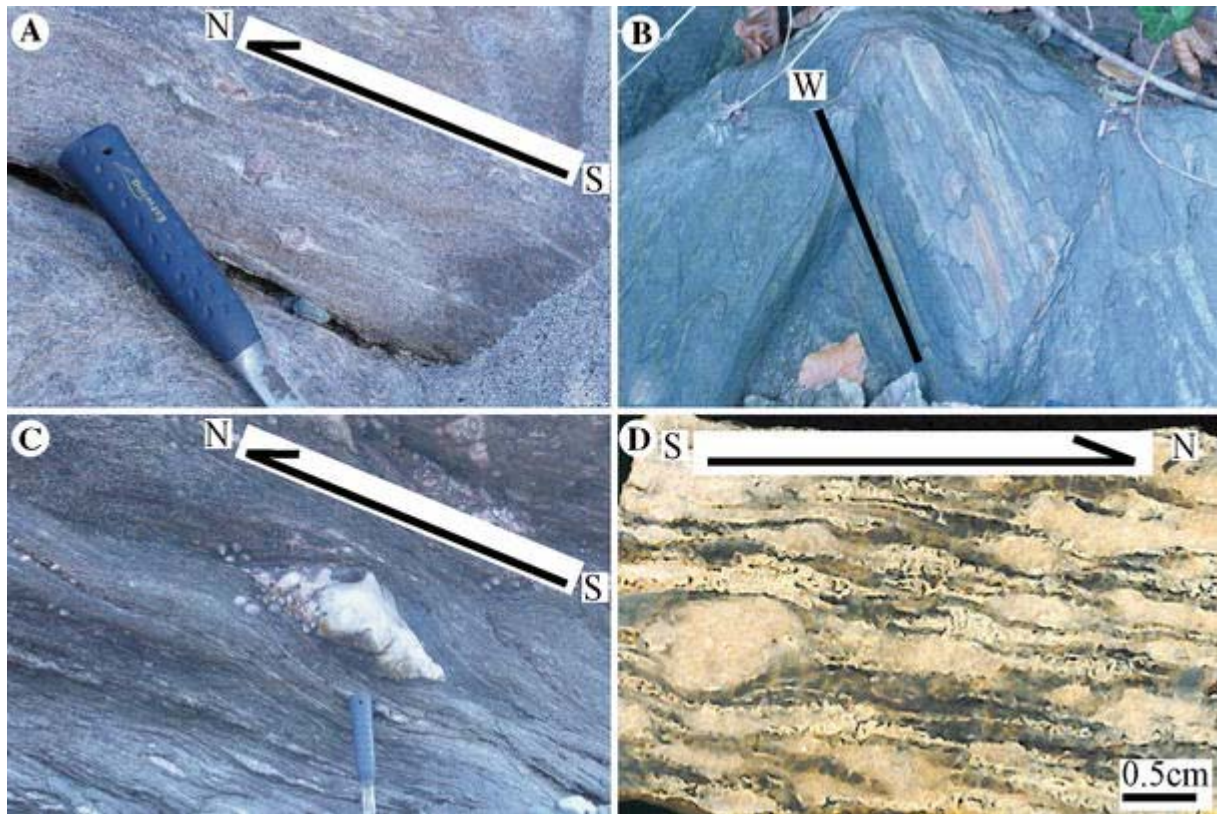


Fig. 5 Macroscopic structures: **(a)** centimeter-sized garnet porphyroblast with quartz pressure shadows in the Le Conquet micaschists, clast asymmetry indicates top-to-the-N shearing; **(b)** E–W trending stretching lineation in the Penzé micaschists; **(c)** top-to-the-N asymmetric quartz vein in Le Conquet micaschists; **(d)** Tréglonou augen orthogneiss with sigma-type porphyroclast systems indicating a top-to-the-N shearing

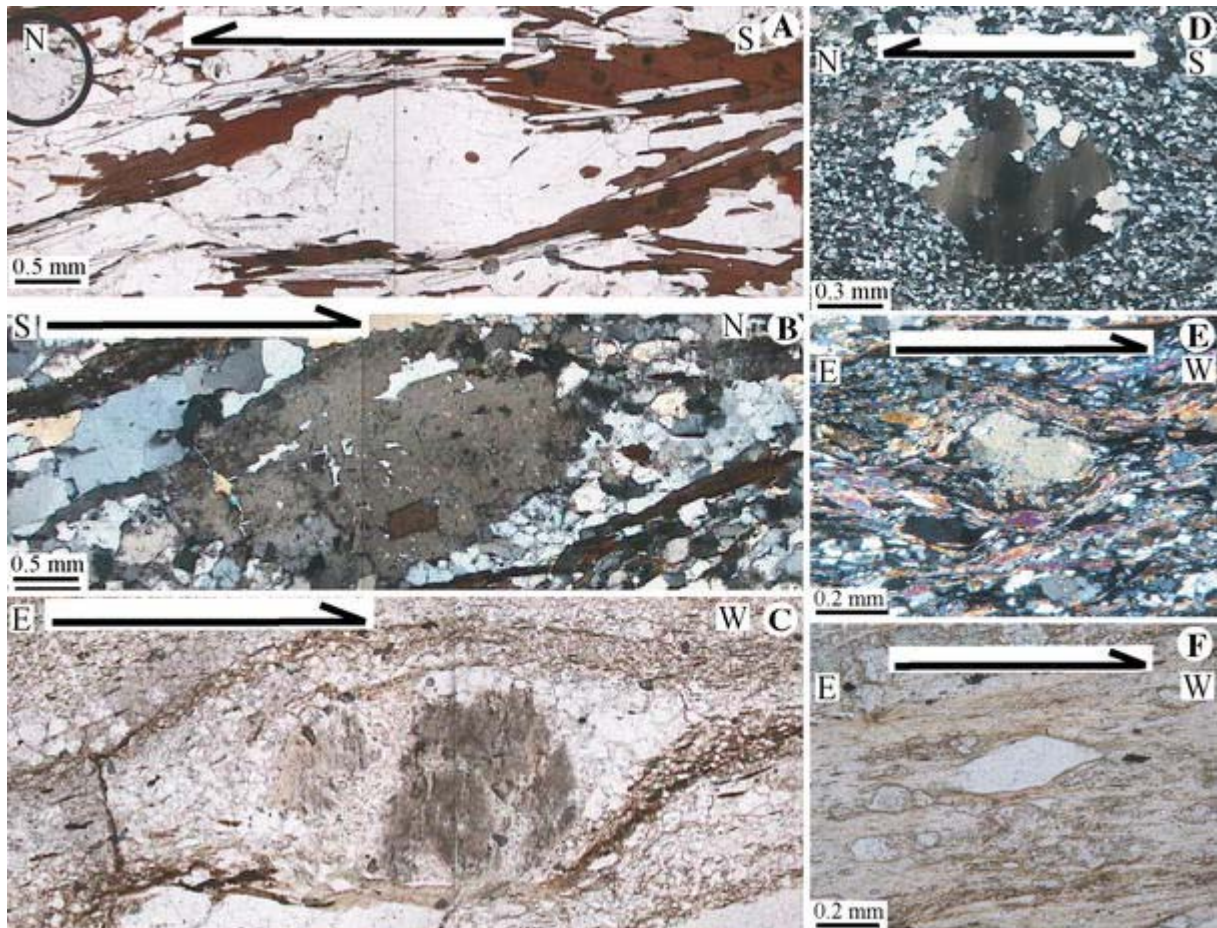


Fig. 6 Microscopic-scale kinematic indicators. **a** Sigma-type porphyroblast system in Kerhornou gneiss. **b** Asymmetric quartz pressure shadows around K-feldspar clast in Le Conquet paragneiss. **c** Sigma-type porphyroblast system with top-to-the-west shearing in Brest orthogneiss (Carantec, Penzé area). **d** Sigmoidal quartz clast in Le Conquet paragneiss. **e** Mylonitized amphibolite with shear bands and pressure shadows indicating a top-to-the-west shearing in Penzé metagabbro. **f** Sigmoidal clast in Devonian sandstone, East of Penzé

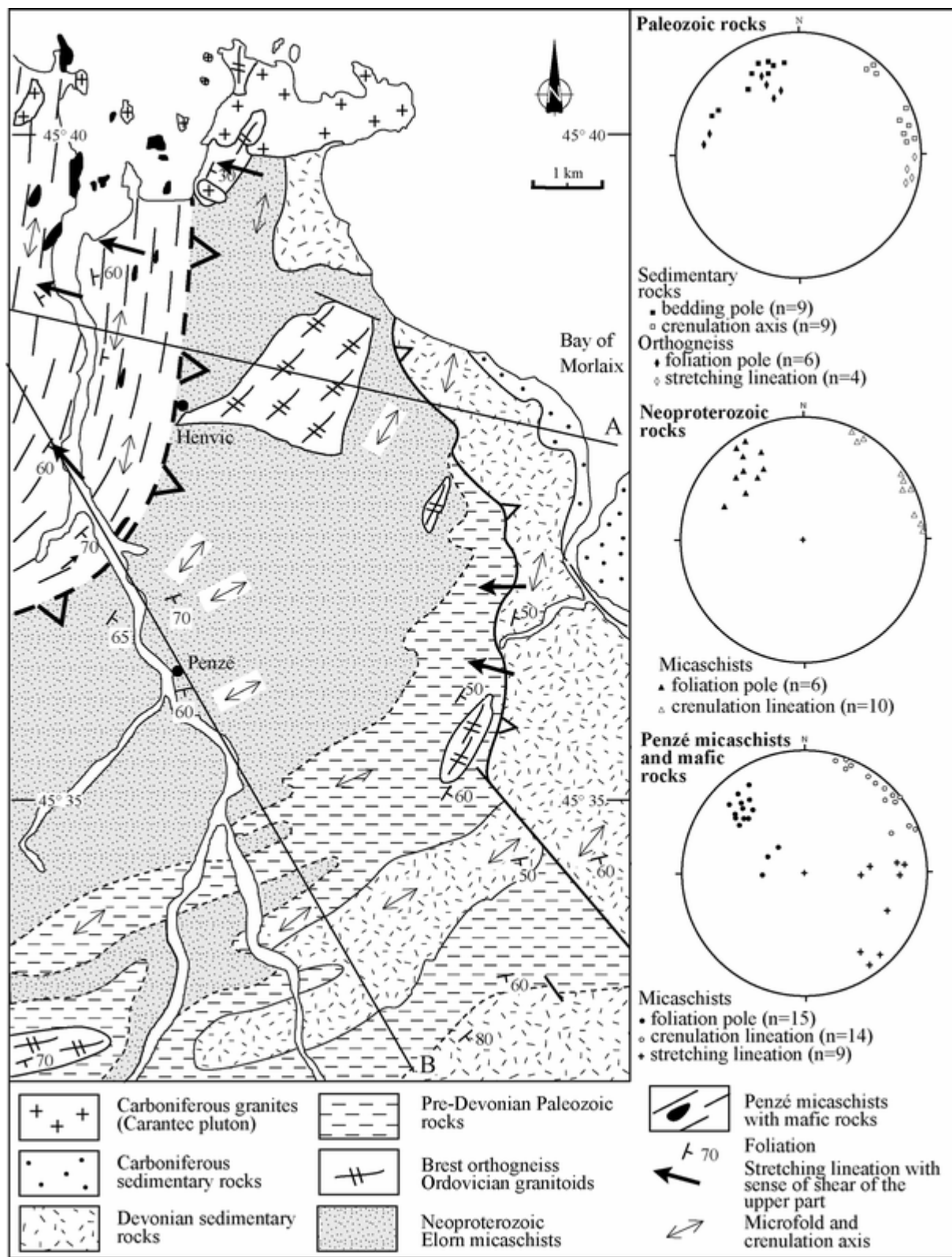


Fig. 7 Structural map of the Penzé area and stereograms (Schmidt projection, *lower hemisphere*) of the structural elements for Paleozoic sedimentary rocks and Brest orthogneiss, Neoproterozoic Elorn micaschists, and Penzé micaschists

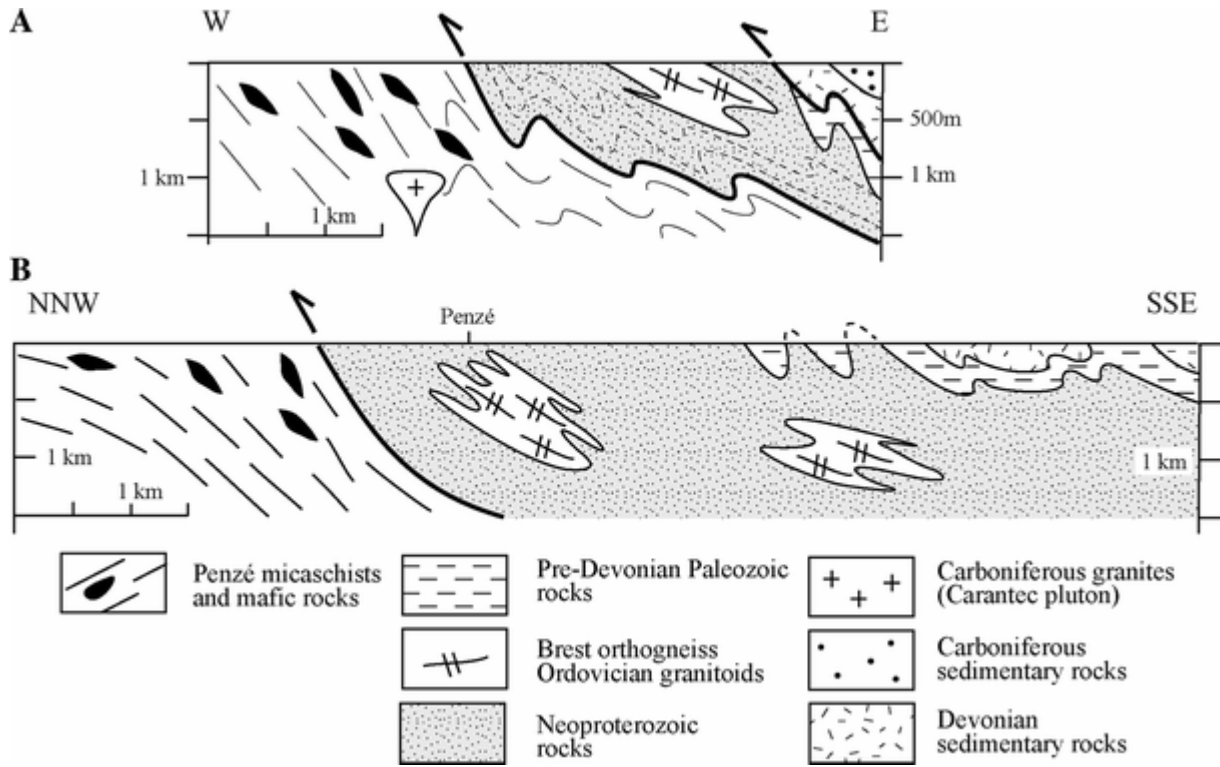


Fig. 8 Cross sections through the Penzé area (located in Fig. 7) showing the Westward displaced nappes

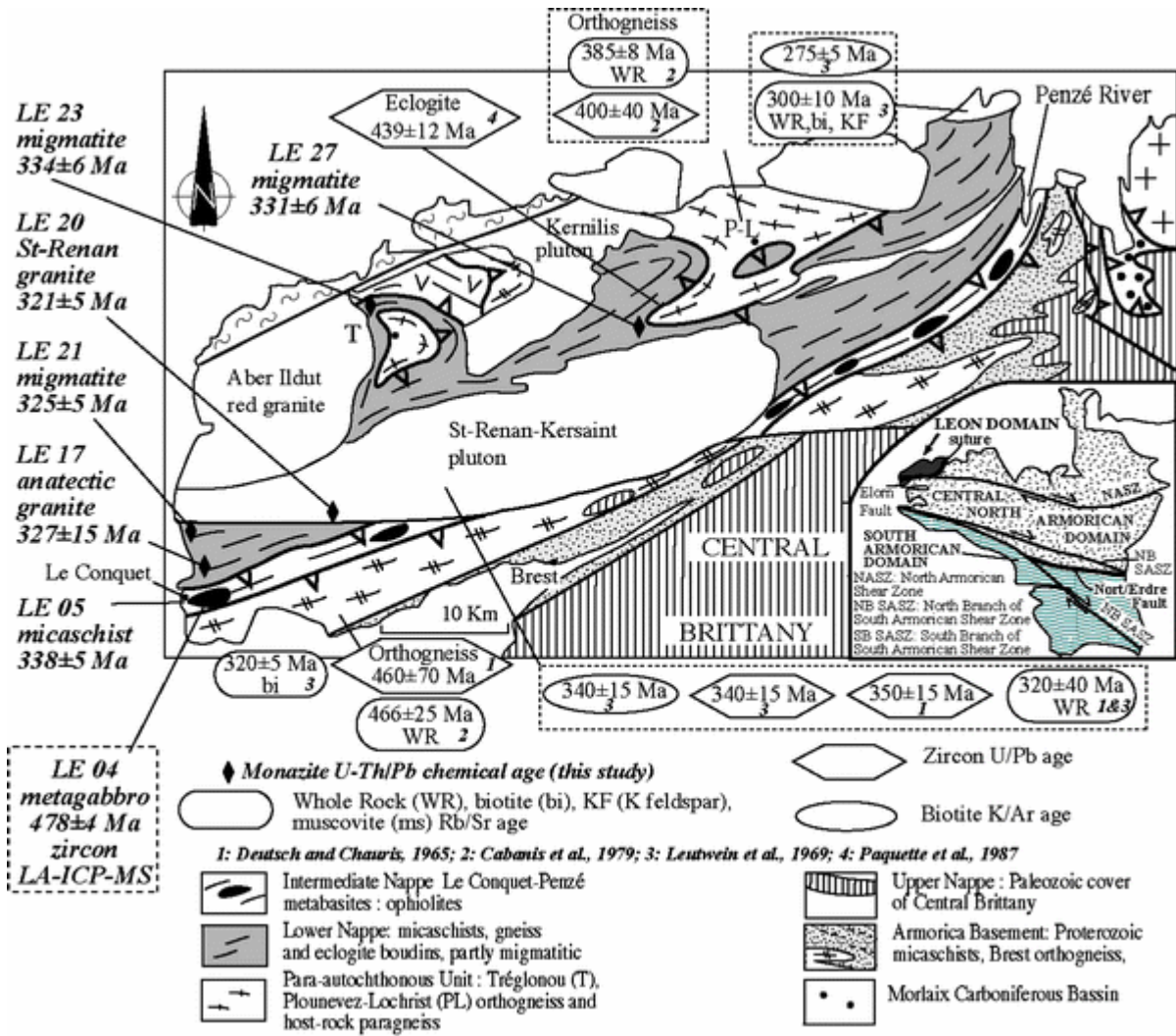


Fig. 9 Synthetic map of the Léon Domain with available radiometric ages (Deutsch and Chauris 1965; Leutwein et al. 1969; Paquette et al. 1987) and the new monazite U–Th/Pb chemical ages (this study)

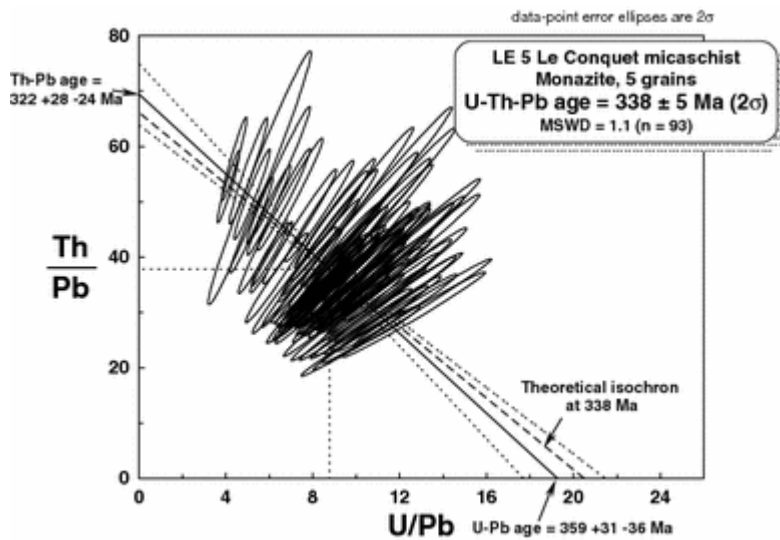


Fig. 10 Th/Pb versus U/Pb isochron diagram for biotite–garnet–staurolite micaschist at Le Conquet LE 05. The data led to the calculated regression line (*thick line*), while the theoretical isochron (*dashed line*) must be included between the two error hyperbolas (*thin lines*). The condition to allow the mean age to be calculated at the centroid of the population: 338 ± 5 Ma

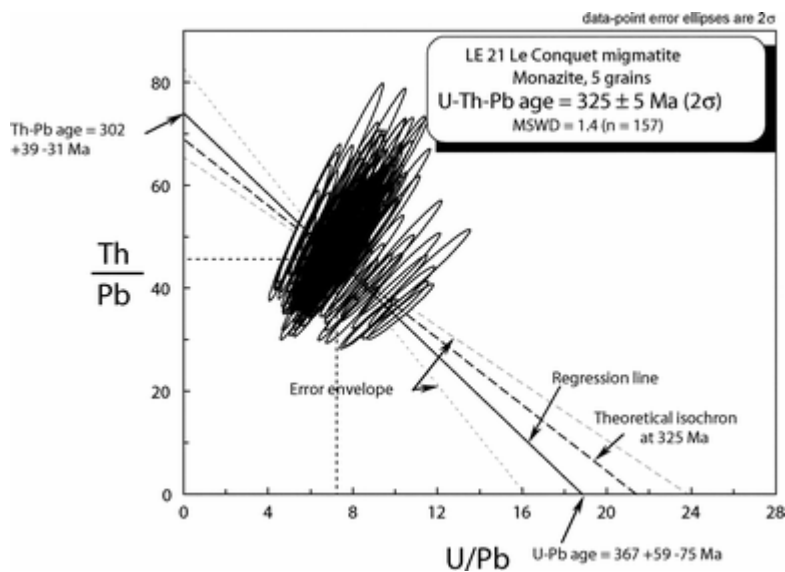


Fig. 11 Th/Pb versus U/Pb isochron diagram for migmatite sample LE 21. The data led to the calculated regression line (*thick line*), while the theoretical isochron (*dashed line*) must be included between the two error hyperbolas (*thin lines*). The condition to allow the mean age to be calculated at the centroid of the population: 325 ± 5 Ma

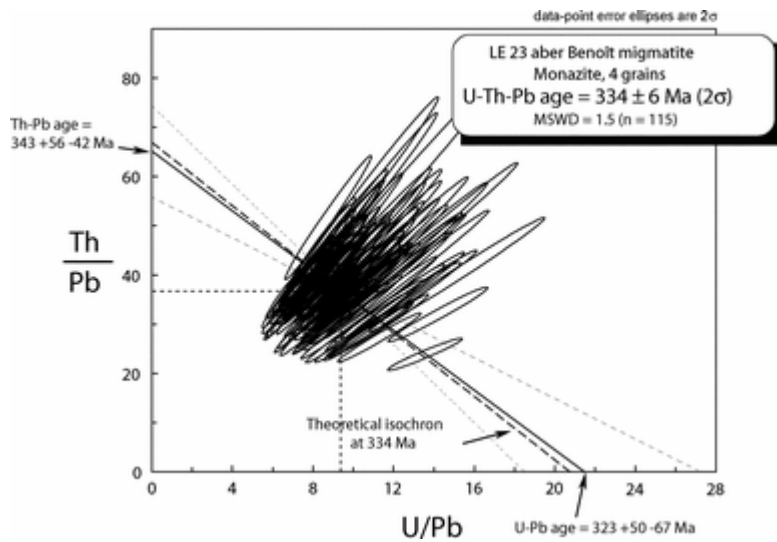


Fig. 12 Th/Pb versus U/Pb isochron diagram for migmatite sample LE 23. The data led to the calculated regression line (*thick line*), while the theoretical isochron (*dashed line*) must be included between the two error hyperbolas (*thin lines*). The condition to allow the mean age to be calculated at the centroid of the population: 334 ± 6 Ma

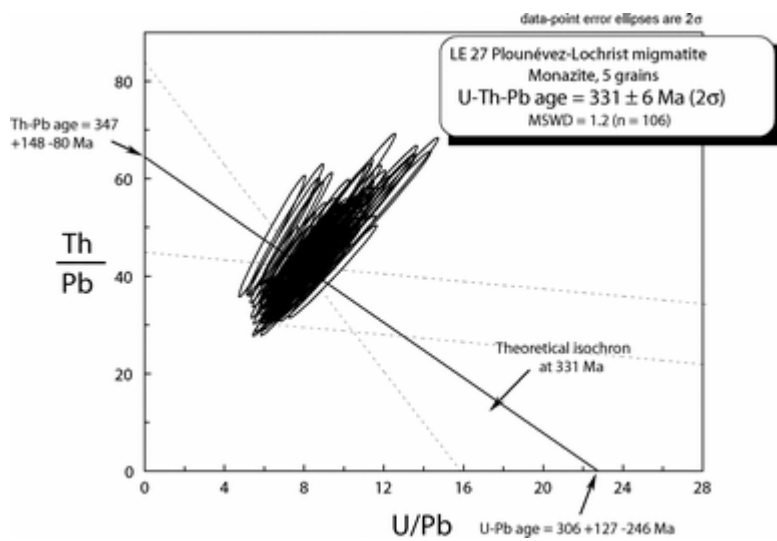


Fig. 13 Th/Pb versus U/Pb isochron diagram for migmatite sample LE 27. The data led to the calculated regression line (*thick line*), while the theoretical isochron (*dashed line*) must be included between the two error hyperbolas (*thin lines*). The condition to allow the mean age to be calculated at the centroid of the population: 331 ± 6 Ma

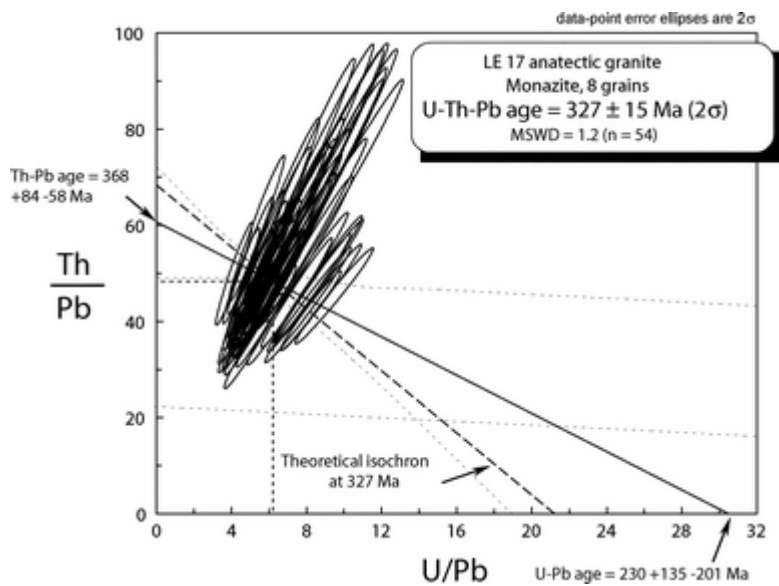


Fig. 14 Th/Pb versus U/Pb isochron diagram for anatectic granite sample LE 17. The data led to the calculated regression line (*thick line*), while the theoretical isochron (*dashed line*) must be included between the two error hyperbolae (*thin lines*). The condition to allow the mean age to be calculated at the centroid of the population: 327 ± 5 Ma

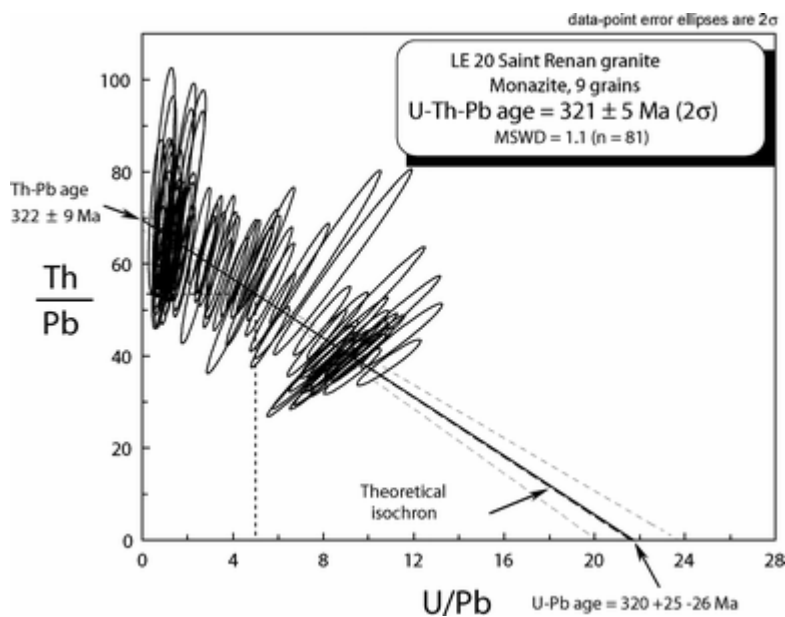


Fig. 15 Th/Pb versus U/Pb isochron diagram for St-Renan granite sample LE 20. Note that despite the wide variation of the composition for these monazite grains, all give the same age within the error as indicated by the slope of the regression line identical to that of the theoretical isochron. The mean age calculated at the centroid of the population is at 321 ± 5 Ma

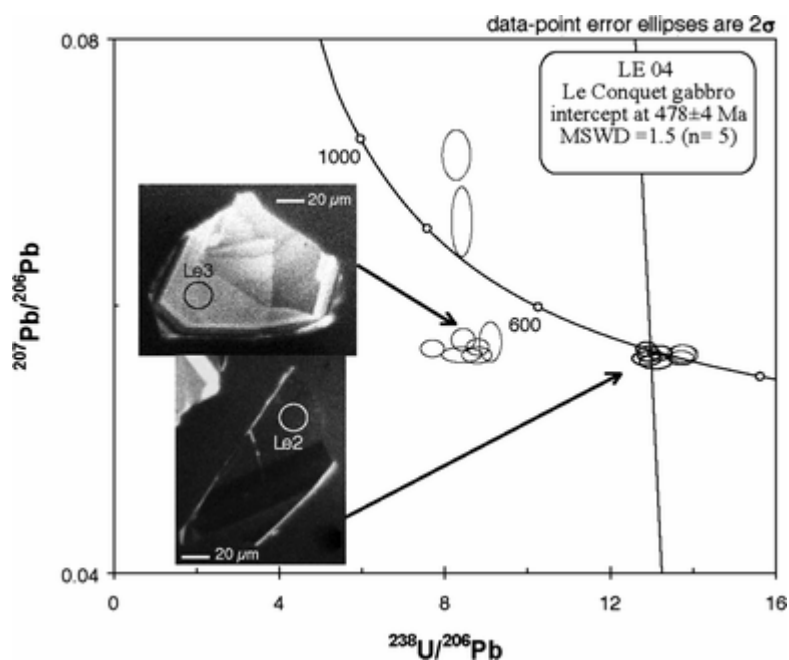


Fig. 16 Common-Pb uncorrected Tera-Wasserburg diagram for data obtained with LA-ICP-MS for Le Conquet gabbro (LE 04). *Bold ellipses* correspond to analyses integrated in age calculation. The gabbro yields a zircon U/Pb age of 478 ± 4 Ma

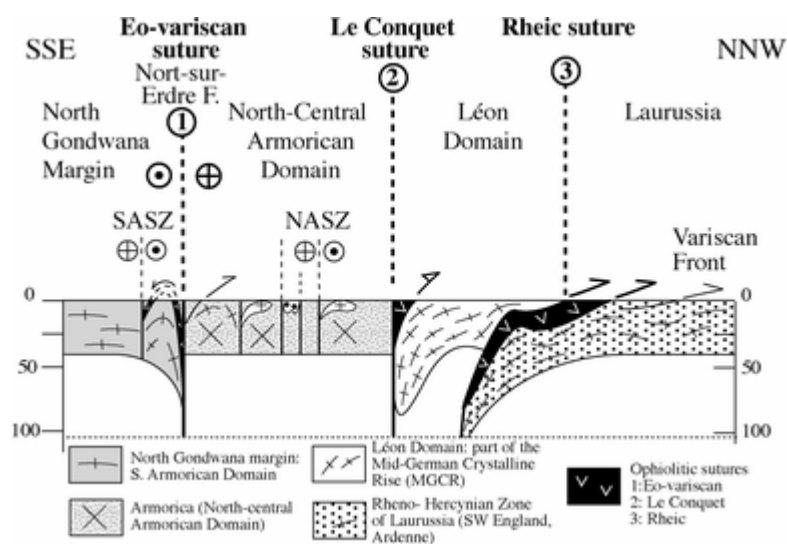


Fig. 17 Crustal scale cross section through the Massif Armoricain and North Britain showing the three suture zones and intervening microcontinents (located in Fig. 1) after the Variscan orogeny at the end of Middle Carboniferous. The Late Carboniferous plutons and the dextral wrench faults have not been represented to emphasize the oldest tectonic events related to microcontinent collision and north-directed thrusting

Table

Table 1 LA-MS-ICPMS U–Pb zircon data for Le Conquet gabbro (LE 04)

	U (ppm)	Th (ppm)	Th/U	Pb* (ppm)	²⁰⁶ Pb/ ²⁰⁴ Pb	f206%	²³⁸ U/ ²⁰⁶ Pb	±(1) ^a	²⁰⁷ Pb/ ²⁰⁶ Pb	±(1) ^a	^{207–206} Pb*/ ²³⁸ U	±(1) ^a	Age (Ma)	±(1) ^a
Concordant population analyses														
Le4-1	515.2	369.0	0.72	34.6	14,342	–	12.8795	0.1420	0.05605	0.0021	0.07764	0.00086	482.0	5.1
Le4-2	1,188.3	1,209.6	1.02	78.6	–	–	13.0926	0.1582	0.05585	0.00022	0.07638	0.00092	474.5	5.5
Le4-5	489.9	378.9	0.77	32.8	13,611	–	12.8727	0.1019	0.05675	0.00021	0.07768	0.00061	482.3	3.7
Le4-8	463.6	283.1	0.61	30.4	47,023	0.000	13.2028	0.1161	0.05647	0.00022	0.07574	0.00067	470.7	4.0
Le4-10	321.3	235.9	0.73	21.5	–	–	12.9262	0.1118	0.05638	0.00024	0.07736	0.00067	480.4	4.0
Discordant population analyses														
Le4-3	127.7	76.0	0.60	14.4	12,364	–	7.7017	0.1169	0.05679	0.00026	0.12984	0.00197	787.0	11.2
Le4-4	38.4	20.8	0.54	4.0	578	0.954	8.2793	0.1392	0.07129	0.00077	0.11963	0.00201	728.4	11.6
Le4-6	40.2	47.0	1.17	4.1	1,077	0.369	8.4092	0.1002	0.06628	0.00107	0.11848	0.00141	721.8	8.1
Le4-7	246.4	356.5	1.45	25.4	117,837	–	8.4038	0.1956	0.05632	0.00023	0.11899	0.00277	724.8	15.9
Le4-9	105.4	65.1	0.62	10.8	10,810	–	8.4452	0.1171	0.05744	0.00035	0.11841	0.00164	721.4	9.5
Le4-11	280.6	291.5	1.04	27.7	–	–	8.7735	0.1462	0.05630	0.00025	0.11398	0.00190	695.8	11.0
Le4-13	44.1	41.7	0.94	4.2	3 209	–	9.1012	0.1090	0.05728	0.00061	0.10988	0.00132	672.0	7.6
Le4-16	213.2	324.2	1.52	21.0	6,568	–	8.7924	0.1116	0.05689	0.00025	0.11373	0.00144	694.4	8.3

^aUncertainties are given at ±2σ level