

Interhemispheric comparison of average substorm onset locations: evidence for deviation from conjugacy

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Abstract. Based on 2760 well-defined substorm onsets in the Northern Hemisphere and 1432 in the Southern Hemisphere observed by the FUV Imager on board the IMAGE spacecraft, a detailed statistical study is performed including both auroral regions. This study focuses on the hemispheric comparisons. Southward pointing interplanetary magnetic field (IMF) is favorable for substorm to occur, but still 30% of the events are preceded by northward IMF. The magnetic latitude (MLat) of substorm onset depends mainly on the merging electric field (E_m) with a relationship of $|\text{dMLat}| = -5.2 E_m^{0.5}$, where dMLat is the deviation from onset MLat. In addition, seasonal effects on onset MLat are also detected, with about 2 degrees higher latitudes during solstices than equinoxes. Both IMF B_y and solar illumination have a significant influence on the magnetic local time (MLT) of onsets. An average relation, $\text{dMLT} = 0.25 B_y$ between IMF B_y and the deviation from onset MLT, was found. The B_y dependence varies slightly with the onset latitude. At lower latitudes (higher activity) it is reduced. After removal of the relationship with IMF B_y a linear relationship remains between the solar zenith angle and onset MLT with $\text{dMLT} = 1 \text{ min/deg}$. Therefore, both solar illumination and IMF B_y can contribute to hemispheric longitudinal displacements of substorm onset locations from conjugacy. No indications for systematic latitudinal displacements between the hemispheres have been found.

Keywords. Ionosphere (Auroral ionosphere) – Magnetospheric physics (Magnetospheric configuration and dynamics; Magnetotail)

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1 Introduction

Substorm onset, also called auroral breakup, is a sudden brightening of arcs taking place in the late evening sector. It indicates the impulsive unloading of energy previously stored in the magnetospheric tail, which leads to the precipitation of auroral particles (e.g. Rostoker et al., 1987; Baker et al., 1996; Lui, 1996).

It is well-known that auroral and geomagnetic activity undergoes important seasonal and diurnal variations. Equinox months are favorable for magnetic activity. Because of the tilt of the magnetic dipole relative to the rotational axis, the activity also undergoes diurnal (universal time) variation. The causes are still discussed controversially. Historically, there were mainly three methods proposed to explain the origin of the variations. These are the Russell-McPherron effect (Russell and McPherron, 1973), the axial model (Cortie, 1912), and the McIntosh equinoctial effect (McIntosh, 1959). Besides these explanations dominated by the solar wind interaction efficiency there are also some other attempts to explain the annual variation of magnetic activity, for example, the changing illumination of the polar ionosphere (e.g. Lyatsky et al., 2001; Nagatsuma, 2006). Recently, it was shown that the solar illumination could control the saturation of the polar cap potential (Nogatsuma, 2004) and the latitudinal position of the dayside field-aligned currents (Wang et al., 2005a).

The onset characteristics have been studied in the past in comparison with various solar wind parameters and solar illumination, which disclose to certain degree the origin of the seasonal and diurnal variations of the onset locations. Liou et al. (2001) found that the magnetic latitude (MLat) of the substorm is IMF B_z dependent, while the magnetic local time (MLT) of the onsets is IMF B_y dependent, based on 10 months of POLAR UVI observations. Gérard et al. (2004) found that the magnetic latitudes of the onsets were correlated best with the solar wind dynamic pressure.

This statement was based on a small subset (78) of substorm events observed in winter.

The average substorm onset location shows significant seasonal variations. Liou et al. (2001) found that the substorm onsets tended to start at earlier local times in northern hemispheric summer than in winter, and at intermediate local times in spring. Our previous study of northern hemispheric events disclosed that substorm onsets in sunlight tend to occur 1 h earlier in local time and 1.5° MLat more poleward than in darkness, suggesting that the ionospheric conductivity has an influence on the location of substorm onset (Wang et al., 2005b). Newell et al. (1996) found that an increased ionospheric conductivity decreases the likelihood of establishing energetic electron precipitation (which is needed to excite discrete auroras). Therefore, it is quite possible that ionospheric conductivity may play a role in the substorm process since it can control the auroral activity.

These general features of solar wind and solar illumination control of substorm onset location are derived primarily from Northern Hemisphere studies. One question we want to address here is whether these properties can be found also in the Southern Hemisphere. Recent studies have disclosed significant hemispheric asymmetry of substorm onset locations (e.g. Østgaard et al., 2006). They proposed that longitudinal asymmetry was primarily controlled by IMF clock angle defined in GSM coordinates, which is regarded as a follow on to the Northern Hemisphere study by Wang et al. (2005b). In this work we will investigate the statistical properties of substorm onset locations in both hemispheres, including their dependence on solar wind input and solar illumination. Furthermore, interhemispheric comparisons of substorm onset locations will be conducted in order to find out the cause for hemispheric discrepancies. Finally, seasonal and diurnal variations of substorm onset locations and their relation to the geomagnetic activity will be studied.

The systematic detection of substorm onsets by global imagers provides a more objective picture of the initial phase leading to the development of a substorm (Liou et al., 1999). By analyzing IMAGE satellite FUV-WIC measurements we have studied statistically the substorm onset characteristics in the auroral regions of both hemispheres. In the following section we describe the data set used in this work. In Sect. 3 we study the solar wind effects on the onset locations in terms of magnetic local time and latitude. Hemispheric comparison of the substorm onset locations is presented. A detailed discussion of the latitude and longitude variations of breakup locations in the two hemispheres is given in Sect. 4. Section 5 summarizes the main conclusions.

2 Data set

The IMAGE satellite is in a highly eccentric polar orbit of 1000 km by 45, 600 km altitude. The far ultraviolet (FUV) instrument on board the IMAGE spacecraft provides obser-

vations of the aurora in ultraviolet light on a global scale with a spatial resolution of 50 km and a temporal resolution of ~ 2 min. The prime data source used by Frey et al. (2004) to retrieve the information on substorm onset was the Wideband Imaging Camera (WIC), and if not available, the oxygen imaging Spectrographic Imager channel (SI-13). The WIC has a passband at wavelengths of 140–180 nm, which is sensitive to emissions from the N2-LBH band and atomic NI lines (149.3 nm). The SI-13 has a passband of 5 nm centered around the 135.6 nm OI doublet.

A detailed description of the instrument on board the IMAGE spacecraft and the approach how to determine auroral substorm onsets from the FUV instrument can be found in Frey et al. (2004). The following schemes are used for substorm onset identification. First, a clear local brightening of the aurora has to occur. Secondly, the aurora has to expand to the poleward boundary of the auroral oval and spread azimuthally in local time for at least 20 min. Finally, a substorm onset is only accepted if at least 30 min had passed after the previous onset (Frey et al., 2004). Figure 1 in Wang et al. (2005b) shows two auroral images from IMAGE FUV-WIC in the frame of magnetic latitude (MLat) and local time (MLT) coordinates. The coordinate frame used to order FUV observations is the Apex coordinate system described by Richmond (1995). A prominent auroral activity at local midnight occurs at 11:20 UT on 2 September 2001. In the next tens of minutes the aurora expands poleward and azimuthally, which are typical substorm features.

The list of substorm onset times and locations, observed by the IMAGE-FUV in the Northern Hemisphere covers the period May 2000 (start of the regular IMAGE-FUV operations) through December 2003 (Frey et al., 2004). In the southern auroral oval observations were made during January 2004 through December 2005 (Frey and Mende, 2007). The compiled list of substorm onset times and locations covers all seasons several times.

We use also 1 min interplanetary magnetic field (IMF) and solar wind data collected by the Advanced Composition Explorer (ACE) satellite. The satellite is located at the Lagrange Points (L1) approx. $220 R_E$ (~ 1.48 million km) in front of the Earth. The solar wind data have been propagated from the ACE satellite to the magnetopause. The solar wind parameters at the magnetopause are averaged over a one hour period before the onset (for justification see Wang et al., 2005b).

3 Observations

The large number of substorm events detected during the years considered (2760 events during 2000–2003 in the north and 1432 events during 2004–2005 in the south) is a good basis for studying the statistical properties of substorm onsets in the northern and southern polar region. Figure 1 shows the distribution of events versus the months of a year. All seasons are covered fairly evenly. Daylight does not seem

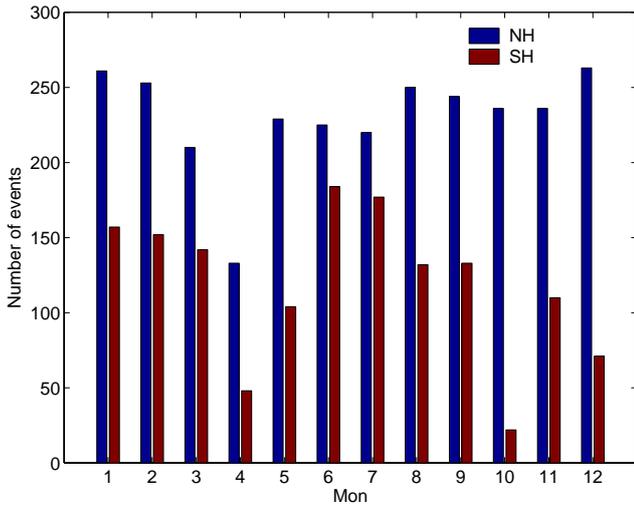


Fig. 1. Occurrence distribution of detected substorm onsets for the different months of the year, separately for the two hemispheres.

to cause a problem. The occurrence number varies around 250 for each month in the Northern Hemisphere and around 150 in the south. Markedly fewer events are detected during the months April and October in the Southern Hemisphere and during April in the Northern Hemisphere. During these months the camera was deactivated for most of the time in order to protect it against direct sunlight exposure.

3.1 Correspondence to solar wind parameters

At first let us see which interplanetary configurations are favorable for triggering a substorm onset. In a previous study Wang et al. (2005b) have performed a superposed epoch analysis of the substorm onset time with respect to various solar wind parameters (which had been time-shifted to the dayside magnetopause). That study was based on two years (2002–2003) of substorm onset events in the Northern Hemisphere. In this study, we will test the result with a much larger data set. Furthermore, we will perform the same analysis based on Southern Hemisphere events in order to find out whether the same features can be derived as earlier from Northern Hemisphere studies.

Figure 2 shows the average development of solar wind parameters, IMF B_z , E_m , merging electric field and N_p , proton density for the two hours around the auroral breakup time, as derived from WIC. The curves in the left frames are based on Northern Hemisphere events and in the right on southern events. No obvious differences could be found for the triggering conditions in the two hemispheres. The averaged IMF B_z attains a minimum about 20 min before the onset. At the time of onset an IMF northward turning is observed. The merging electric field, defined as $E_m = v_{sw}(B_y^2 + B_z^2)^{0.5} \sin^2(\theta/2)$ (e.g. Kan and Lee, 1979), where v_{sw} is the solar wind velocity and θ is the clock an-

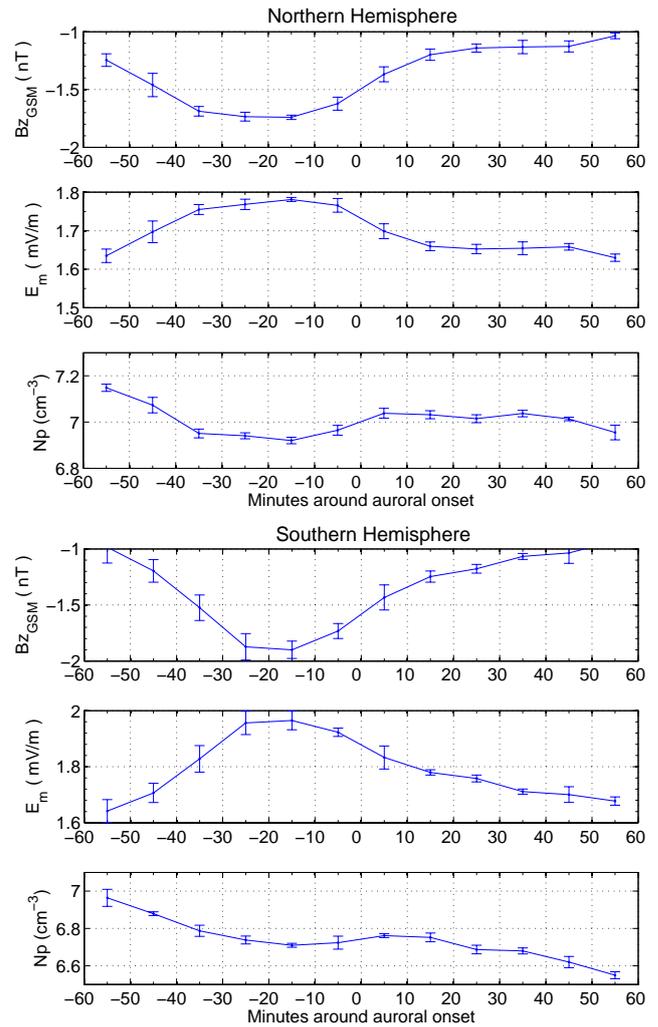


Fig. 2. Superposed epoch analysis of solar wind parameters around auroral breakup. The key time, “0”, corresponds to the onset time. Parameters from top to bottom are IMF B_z , merging electric field, E_m , and solar wind density, N_p , which have been propagated to the dayside magnetopause. The bars indicate the uncertainties of the 10 min averages.

gle of the IMF, performs a similar variation in anti-phase. The relative change of E_m is, however, significantly smaller than that of IMF B_z . The average number density, N_p , also attains a shallow minimum about 20 min before the onset. The density recovers at the time of breakup. Although this effect is small, it is significant and it shows up consistently in the independent datasets from both hemispheres. The same analysis was also performed with the other components of the IMF, the solar wind velocity, v_{sw} , the dynamic pressure, P_{dyn} , Akasofu coupling parameter, ε , and a more sophisticated coupling function, $\varepsilon^* = P_{dyn}^{1/6} v_{sw} B_T \sin^4 \vartheta / 2$ (Gérard et al., 2004), where B_T is the transverse component of the IMF and $\vartheta = \tan^{-1}(B_T/B_x)$ is the IMF cone angle. None of

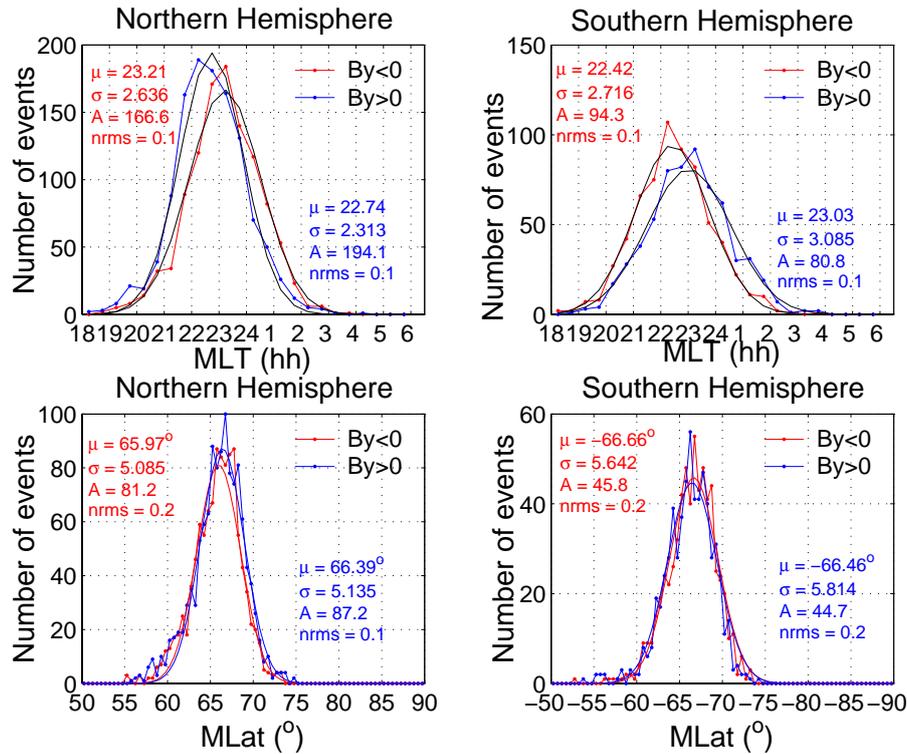


Fig. 3. Number of events of substorm onsets as a function of MLT and MLat for IMF $B_y > 0$ and IMF $B_y < 0$ in the Northern (left) and Southern (right) Hemisphere.

these parameters shows a persistent signature related to the substorm onset time.

In a next step we wanted to see, how IMF conditions influence the location of the breakup. First, let us have a look at the IMF B_y effect on the substorm onset location. For the purpose of this study, all substorm onset events are sorted by the sign of IMF B_y separately for the Northern and Southern Hemisphere. Then the distribution of events as a function of MLT and MLat are plotted and shown in Fig. 3. For a more quantitative description of the main characteristics of the distribution we have fitted and over-plotted a Gauss function to the points in this and also later figures. To estimate the goodness of Gauss fit, we have produced the normalized root-mean-square (nrms) as shown in this and also later figures. The nrms is defined as $\sqrt{\frac{\sum_{i=1}^n (y_i - y_i^*)^2}{\sum_{i=1}^n (y_i^*)^2}}$, where y_i is the measured and y_i^* the fitted. A nrms closer to 0 indicates a better fit. With regard to Fig. 3, we obtain all nrms no more than 0.2.

In both hemispheres there emerges a clear difference in local time of substorm onset for IMF $B_y > 0$ and $B_y < 0$. In the Northern Hemisphere substorm onsets shift toward earlier MLT by about half an hour for positive B_y and toward later local time during negative B_y . An opposite relationship with B_y , as mentioned above, is observed in the Southern Hemisphere. This effect of IMF B_y on onset location is

not new and has repeatedly been reported (e.g. Rodger et al., 1984; Liou et al., 2001). We have also investigated a possible B_y effect on the onset latitude. The average magnetic latitude is found at about 66° MLat. This value is independent of the B_y sign being valid for both hemispheres. The slight average poleward shift in the south we do not regard as significant.

The same analysis is performed with the other components of the IMF. The sign of IMF B_x also affects the MLT of onset location, but in opposite direction to that of the B_y component (not shown). This relation is not surprising since the IMF components B_x and B_y are correlated through the preferred Parker spiral geometry of the interplanetary magnetic field. It can be stated that neither B_x nor B_y signs have an influence on the onset latitude or on the occurrence rate of substorms. The situation is different for the IMF B_z component. As expected, the occurrence rate is strongly dependent on the B_z sign. About two-third of the events are preceded by a negative B_z . But it is still remarkable that 30% of the substorms are preceded by positive B_z . On the other hand, the mean local time of the onset is not affected by B_z , as can be seen in Fig. 4. We find furthermore that the average latitude of break-ups is about 1° more equatorward for negative B_z than for positive B_z .

We have also calculated the correlation of MLT and MLat with solar wind parameters like v_{sw} , P_{dyn} , E_m and ε . Generally there is no good correlation between MLT and these

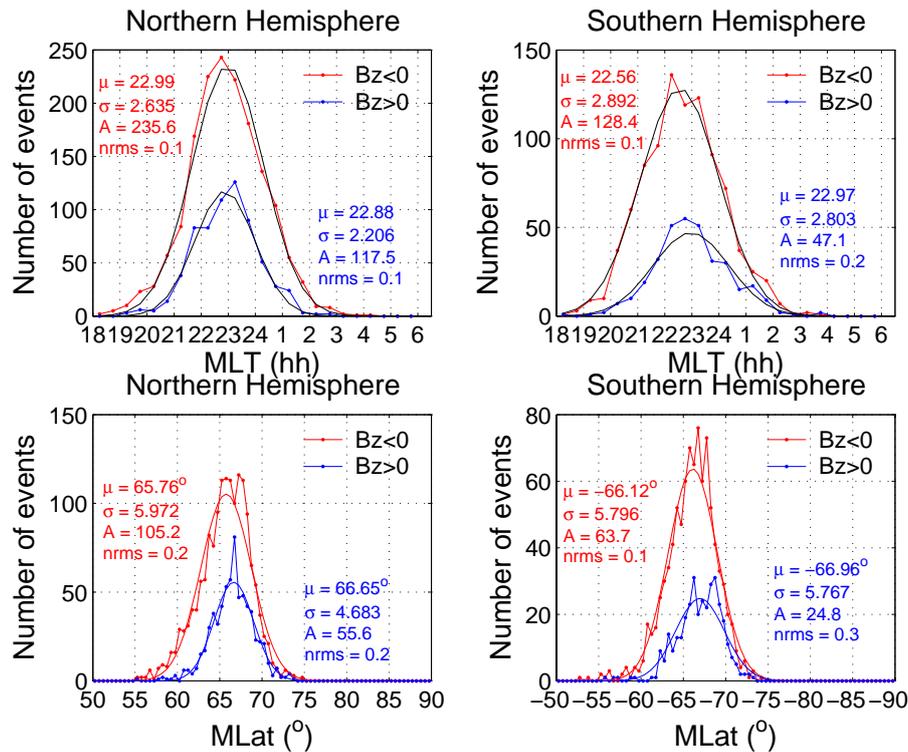


Fig. 4. Same as Fig. 3, but for IMF $B_z > 0$ and $B_z < 0$.

various solar wind parameters. However, E_m gives the best correlation with MLat yielding a correlation coefficient, $R=0.5$. This finding is consistent with the earlier study on substorm onsets (Wang et al., 2005b). In that case a linear relation between the polar cap area and the square root of E_m could be established. E_m has been identified as a suitable geoeffective solar wind coupling parameter for loading of the magnetosphere confirming the results of previous studies (e.g. Newell et al., 2006; Wang et al., 2005a).

3.2 Solar illumination effect

For the purpose of this study, all substorm onset events are sorted by solar illumination condition separately for the Northern and Southern Hemisphere. The substorm onsets are separated into two groups, depending on the solar zenith angle (SZA). Daylight is defined as $SZA < 100^\circ$ and darkness as $SZA > 100^\circ$ (e.g. Schlegel, 1988; Wang et al., 2005a). For each bin in magnetic latitude and local time the number of events was determined, as shown in Fig. 5. It is expected that a higher occurrence frequency of substorm onsets is observed in darkness because these events occur close to local midnight. This region is in sunlight only during local summer. In the Northern Hemisphere one-fourth of the events occur in sunlight and only one-fifth in the south.

In the Northern Hemisphere there is a clear difference in local time of the substorm onset between daylight and

darkness. In darkness substorm onsets occur on average at 23.1 MLT, while in sunlight they tend to occur earlier, at 22.7 MLT. The most likely onset locations are found at 66.6° MLat in sunlight and at 66.1° MLat in darkness, which suggests a slight dependence of the onset latitude on illumination conditions. This is consistent with previous studies based on relatively smaller numbers of events (e.g. Liou et al., 2001; Wang et al., 2005b).

In the Southern Hemisphere there is also a difference in local time of the substorm onset between daylight and darkness. Substorm breakups shift similar to the Northern Hemisphere to earlier time in daylight and to later time in darkness. However, there is an indication for an equatorward shift in daylight, which is opposite to the Northern Hemisphere. This indicates that solar illumination enhances the inter-hemispheric asymmetry of the substorm breakup locations in azimuthal direction, but has no effect in the meridional direction.

3.3 Annual variation of onset location

After having presented averaged characteristics of substorm onsets we will make use of the large number of events to look at the annual variation of some of the parameters. In Sect. 3.2 we have shown that the location of substorm onsets depends to a certain degree on the solar illumination conditions. The question arises, is this a seasonal or an illumination effect? In

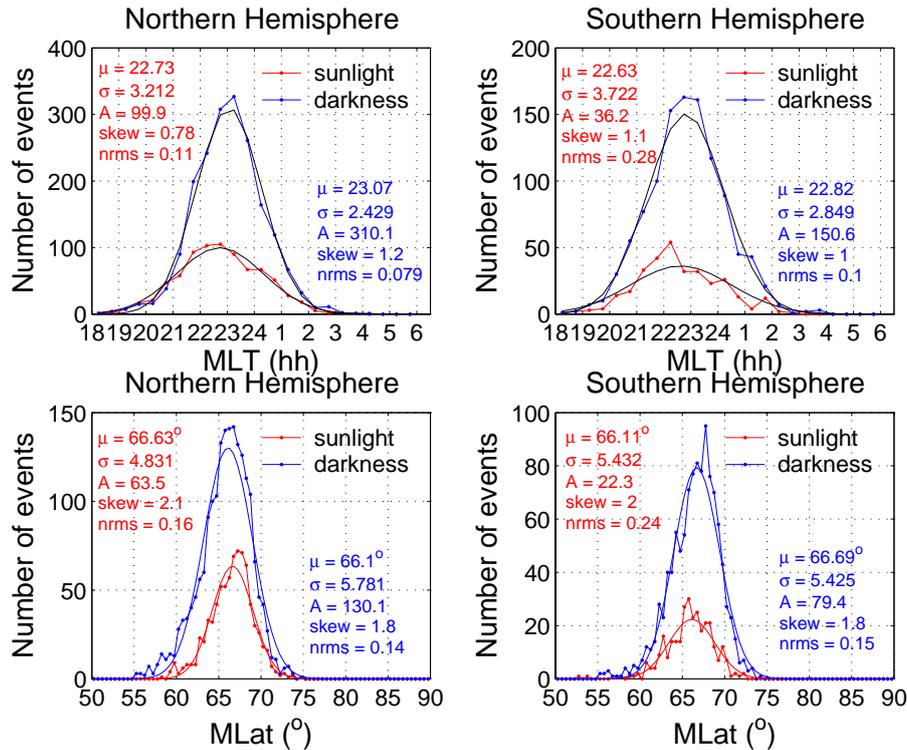


Fig. 5. Same as Fig. 3, but for solar zenith angle, $SZA > 100^\circ$ (darkness) and $SZA < 100^\circ$ (sunlight).

an attempt to clarify this open issue we plotted in Fig. 6 the distribution of magnetic latitudes of the onset locations in a month versus universal time coordinate frame. To visualize a possible relation of the distribution to the illumination conditions contour lines of the solar zenith angle are over-plotted.

In the Northern Hemisphere a clear horizontal four-band structures emerges reflecting alternating high and low latitudes. These bands are centered at the months January, April, July and October. Superimposed on that there is a UT modulation with higher onset latitudes around 05:00 UT and lower around 18:00 UT. In the Southern Hemisphere the distribution is not as clear, but still, a certain trend of anti-correlation with the north can be seen. Highest onset latitudes are detected during April and lowest around December solstice. Again, there is an UT dependence. In this hemisphere latitudes are higher around 14:00 UT and lower around 24:00 UT.

Similar to the latitude we also investigated the distribution of the onset local time. Figure 7 shows the variation of local times in a month versus universal time plot. The situation is not as clear as in the case of latitude. A reason for that is probably the additional influence of the IMF B_y component on the onset local time (cf. Sect. 3.1). In spite of that a certain trend can be deduced from the plots. There is a preference in the Northern Hemisphere for late local times in the first half of the year and for earlier onset times during the second half. Additionally, we find also here an UT depen-

dence shifting break-ups to earlier MLTs around 06:00 UT and to later MLTs around 18:00 UT. In the Southern Hemisphere the local time variation is even less evident, but some indications for opposite responses can be deduced, including both the seasonal and UT dependence. Some of these observations can be regarded as indications for a deviation of the break-up locations in the two hemispheres from their conjugate points. A discussion of possible causes will be given in the next section.

4 Discussion

In the previous section we have shown that the locations of substorm onsets depend both on IMF and illumination conditions. The large number of detected events allows us to look at both aspects simultaneously. To our knowledge this has never been done in a previous study on comparable scale.

4.1 Substorm latitude

It is known that for higher magnetic activity the polar cap area increases and with that substorm onsets occur at lower latitudes. A confirmation for that can be found in Fig. 4. For negative B_z the latitude is shifted on average by 1° in latitude equatorward. We have to correct for that effect before we can start a comparison between the hemispheres since the two data sets have been taken at different times therefore

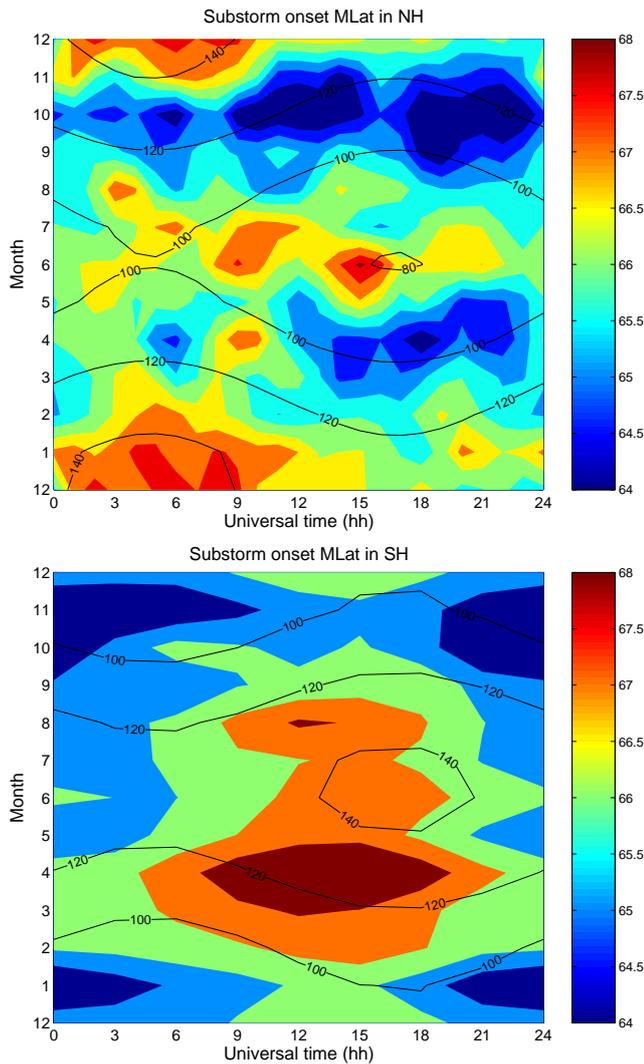


Fig. 6. Distribution of substorm onset magnetic latitude versus month and universal time for both hemispheres. Overplotted are SZA contours for the location 66.5°MLat and 23:00 MLT.

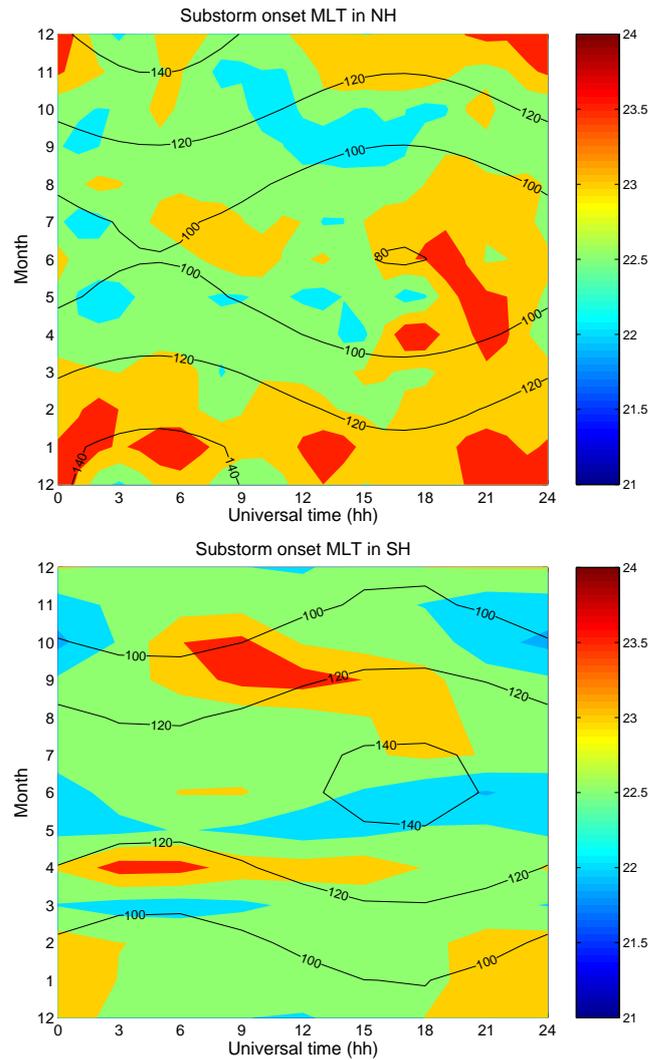


Fig. 7. Same as Fig. 6, but for onset local time.

experienced different interplanetary conditions. Dedicated investigations of Wang et al. (2005b) revealed that there is a good correlation in the Northern Hemisphere between the substorm latitude and the square root of the merging electric field, E_m . For events in darkness, which accounts for about 80% of the cases, they deduced an onset latitude, $\beta = 73^\circ - 5.2 E_m^{0.5}$. For events in sunlight the dependence was only half as strong.

In this study we make use of that derived latitude dependence and apply a latitude correction of the form, $dMLat = 5.2 E_m^{0.5}$ to the samples of both hemispheres. The new latitude distribution is shown in Fig. 8. Compared to Fig. 6 the average latitude has shifted by about 7° poleward, and the range of variation is reduced (some small-scale scatter has been introduced by the E_m correction). The result-

ing average latitude (valid for $E_m = 0$) is about 73° MLat for the Northern Hemisphere and about 72° for the south. This small difference cannot be considered as significant. We have compared the monthly mean latitudes, after applying the correction, with the monthly mean values of SZA and found no correlation in any of the hemispheres.

When looking at Fig. 8 some systematic latitude structure seem to remain, however, after the removal of the dependence on magnetic activity. Obviously, substorms occur about 2° more equatorward during equinoxes than during solstice. This implies a systematically larger polar cap area during equinoxes, which may be consistent with the Russell-McPherson and McIntosh effects. Liou et al. (2001) also found that the average onset latitude is lowest in spring, then followed by summer and winter. Wang et al. (2005b) have reported average onset latitudes 1.5° more poleward for events in sunlight compared to those in darkness. From the

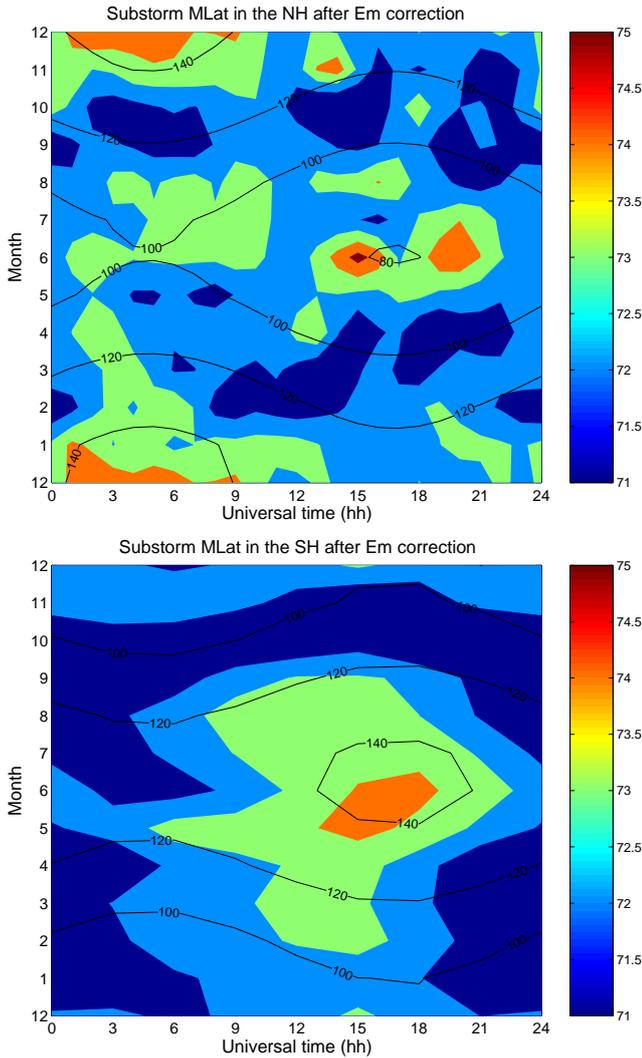


Fig. 8. Same as Fig. 6, but after removal of the E_m effect (see text).

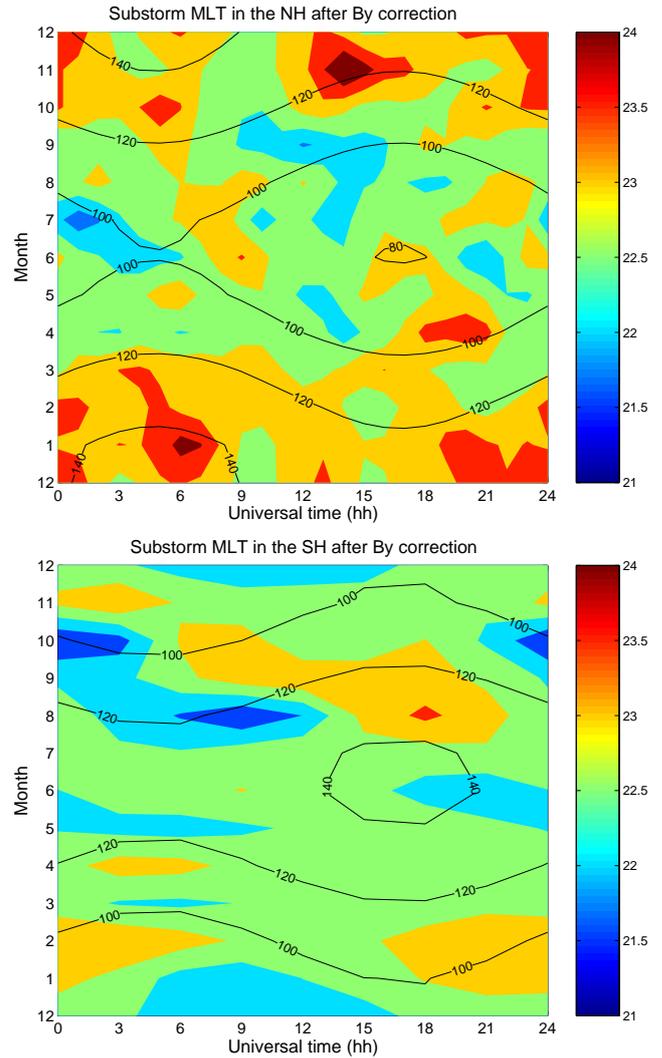


Fig. 9. Same as Fig. 7, but after removal of the IMF B_y effect (see text).

results obtained here we can say that this is a seasonal effect and not a response to solar illumination. Overall we may conclude that neither the IMF configuration nor the solar irradiation seem to cause deviations of onset locations from the conjugate latitudes in both hemispheres.

4.2 Onset local time

The situation is quite different for the local times at which the substorm onsets occur. Figures 3 and 5 seem to imply that both IMF B_y and SZA have an influence on the break-up time. To disentangle the two dependences we sorted the MLT value of each event into UT bins, as shown in Fig. 7. MLT averages for each UT hour were calculated, separately for both hemispheres. In the same way the corresponding IMF B_y components and SZA values were binned and averaged. As expected, the SZA is practically independent of UT. In this way the B_y influence on the MLT of the onsets can be

determined. The correlation analysis reveals the dependence, $dMLT = \pm 0.25 B_y$. In the Northern Hemisphere a positive B_y shifts the onset time to earlier hours and in the south to later times.

The obtained local time distribution after applying the IMF B_y correction is shown in Fig. 9. In the Northern Hemisphere MLT features are now rather well aligned with the SZA contour lines. Later onset times show up at large SZA and early times at small solar zenith angles. In the Southern Hemisphere no such clear picture emerges. For a more quantitative assessment we have calculated monthly averages of both the onset MLTs and SZAs. Figure 10 shows scatter plots for both hemispheres with regression lines. It is immediately evident that we have a linear dependence of these two quantities in the Northern Hemisphere while there is no correlation found in the south. The revealed dependence in the Northern

Hemisphere amounts to $dMLT=0.018\text{ h/degSZA}$, which is equivalent to an increase of about 1 min in local time for each degree in solar zenith angle. Over the whole range we find a local time shift of 45 min. This value is close to the one hour difference in MLT reported by Wang et al. (2005b) between onsets in sunlight and darkness. The question why solar irradiation has such a clear effect on the onset local time in the Northern Hemisphere and no effect in the south should be addressed in a dedicated study.

The influence of IMF B_y on various auroral phenomena is known already for quite a while. There are, however, only few quantitative statements in the literature giving numbers for the local time displacement of a substorm onset per nano Tesla of B_y . For example, Rodger et al. (1984) quote 0.2 h/nT and Liou et al. (2001) report 0.15 h/nT for $|B_z|$ and $|B_x| < 1nT$. From this study, based on a much larger data set, we determined a larger B_y effect (0.25 h/nT).

For an assessment of the apparent inconsistency between the different studies we follow the first order approach of Cowley (1981) on the magnetospheric asymmetry associated with the IMF B_y component. By assuming a dipolar field geometry he can give an analytical expression for the azimuthally displacement, $\Delta\phi$, of the ionospheric footprints in the substorm meridian:

$$\Delta\phi = \frac{b_y}{4B_0 \cos^6 \beta_0} [\cos^2 \beta_0 \tan \beta_0 (\cos^2 \beta_0 + 1.5) + 1.5\beta_0] \quad (1)$$

where b_y is the dawn-to-dusk magnetic field in the magnetosphere, B_0 the geomagnetic field strength at the equator ($\sim 30\,000\text{ nT}$), β_0 is the magnetic latitude of the substorm onset. It is evident from Eq. (1) that the amount of azimuthally displacement strongly depends on the latitude of the onset. Another question is, how much of the IMF B_y is penetrating into the magnetosphere? For the near-Earth space Maus and Luehr (2005) found the ratio, $b_y=0.23 B_y$. If we insert that number into Eq. (1) and search for the onset latitude that would give the obtained displacement (0.25 h/nT), we obtain $\beta_0=77.5^\circ$ MLat. From this resulting unrealistically high latitude we conclude that the dipole geometry is not a sufficient assumption for the fluxtube connecting to the substorm region. The dipole field is valid only in current free space and the influence of B_y on the displacement is strongest at the apex of the field line. Just in that region significant cross-tail currents are expected.

From Eq. (1) it can be deduced that the IMF B_y effect gets stronger on higher L-shells. In order to test whether there is a latitude-dependent sensitivity of the onset location on IMF B_y , we have divided the data set of Northern Hemisphere events into two equal groups, one for onsets at latitudes above 66.2° MLat and the other for those with lower latitudes. The same procedure was applied, as described above, for the determination of the mean MLT dependence on IMF B_y . We found a dependence $dMLT=0.3 B_y$ for the higher latitude events and $dMLT=0.2 B_y$ for those at lower latitudes. This shows that onset latitudes have an effect on

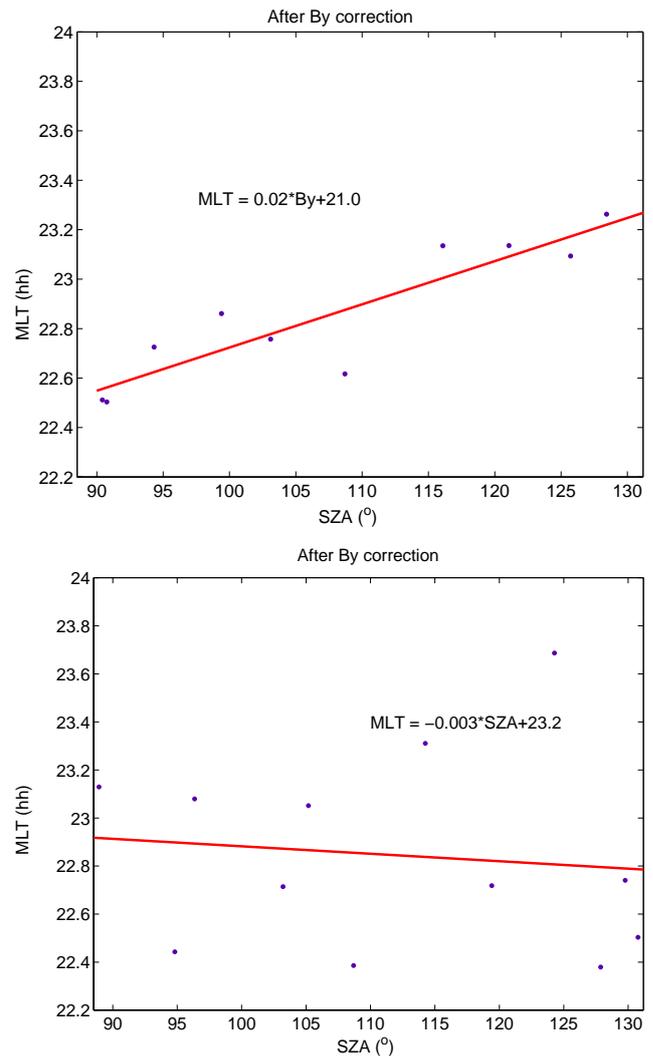


Fig. 10. Dependence of onset local time on solar zenith angle, SZA, after removal of IMF B_y effect.

the shift of onset MLT by IMF B_y . Since the onset latitude is closely related to the solar wind input, as was discussed in Sect. 4.1, the scaling factor in front of B_y becomes activity dependent.

In a recent study Østgaard et al. (2006) suggested that the azimuthal displacement of the onset location is closely related to the IMF clock angle. We have tested the efficiency of the IMF clock angle function presented by Østgaard et al. (2006) (Fig. 2) for correcting the MLT displacement shown in Fig. 9. The result is not convincing, since the obtained distribution is much less ordered than that of Fig. 10. For that reason we do not see any advantage in using the clock angle as a proxy for the MLT difference in the two hemispheres.

5 Summary

A detailed study of substorm onset locations has been performed. Based on a large set of substorm onset times and locations, obtained from both hemispheres, important dependence could be derived in a statistical approach. For several aspects it is the first time that statistically significant numerical factors can be given for the variation of break-up locations including the deviation from conjugate locations in both hemispheres. Important findings are:

1. A large majority of substorms is preceded by a southward oriented IMF, but about 30% of the events are preceded by northward IMF.
2. The latitude of the onset depends linearly on the square root of the merging electric field. A relation $|dMLat| = -5.2 E_m^{0.5}$ is equally valid for both hemispheres. In addition there seems to be a dependence of the onset latitude on season. Break-ups occur on average 2° more poleward during solstices than during equinoxes.
3. No dependence of onset latitude on solar illumination conditions was found. This answers one of the pending questions in Wang et al. (2005b).
4. The local time of a break-up depends on the IMF B_y component. Positive dawn-to-dusk fields shift the onsets in the Northern Hemisphere to earlier times and in the south to later local times. The scaling factor between IMF B_y and onset MLT depends on the latitude of the onset. As a consequence, the displacement of the onset locations from the conjugate points in the two hemispheres is larger during quiet periods than during active. For the data set investigated we obtained the average relation, $dMLT = 0.25 \text{ h/nT}$, between IMF B_y and onset MLT.
5. There seems to be an additional contribution to the hemispheric displacement coming from the illumination conditions. We have clear evidence from the Northern Hemisphere that substorms occur at earlier local times in sunlight than during darkness. The obtained relation is $dMLT = 1 \text{ min/deg}$, between the solar zenith angle and event MLT. The smaller sample number from the Southern Hemisphere does not support this tendency, but a sorting of all events in sunlit and dark indicates an equivalent dependency (earlier MLT for sunlit, cf. Fig. 5).

The reliability of the inter-hemispheric relation presented here should be verified on a case-by-case basis. For such a study the substorm onset catalogue could be used to select suitable events. There are several magnetometer networks in the Northern Hemisphere which can be used to provide in-situ substorm observations of South Pole events from the catalogue.

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