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# Simultaneous radar and spaced receiver VHF scintillation observations of ESF irregularities

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**Abstract.** Simultaneous observations of equatorial spread F (ESF) irregularities made on 10 nights during March–April 1998 and 1999, using an 18-MHz radar at Trivandrum (77° E, 8.5° N, dip 0.5° N) and two spaced receivers recording scintillations on a 251-MHz signal at Tirunelveli (77.8° E, 8.7° N, dip 0.4° N), have been used to study the evolution of Equatorial Spread F (ESF) irregularities. Case studies have been carried out on the day-to-day variability in ESF structure and dynamics, as observed by 18-MHz radar, and with spaced receiver measurements of average zonal drift  $V_o$  of the 251-MHz radio wave diffraction pattern on the ground, random velocity  $V_c$ , which is a measure of random changes in the characteristics of scintillation-producing irregularities, and maximum cross-correlation  $C_I$  of the spaced receivers signals. Results show that in the initial phase of plasma bubble development, the greater the maximum height of ESF irregularities responsible for the radar backscatter, the greater the decorrelation is of the spaced receiver scintillation signals, indicating greater turbulence. The relationship of the maximum spectral width derived from the radar observations and  $C_I$  also supports this result.

**Keywords.** Ionosphere (Equatorial ionosphere; Ionospheric irregularities; Plasma waves and instabilities)

## 1 Introduction

The inhomogeneities or irregularities in the nighttime F region of the equatorial ionosphere, termed as equatorial spread F (ESF), are known to be geomagnetic field-aligned and display a wide spectrum of scale sizes, ranging from a few centimeters to hundreds of kilometers (Kelley, 1989; Hysell, 2000 and the references therein). A hierarchy of plasma instabilities is believed to be the cause for the observed wide range of scale sizes, starting with a collisional shear instabil-

ity which is capable of producing large-scale plasma waves in the bottomside equatorial ionosphere at sunset, thus preconditioning the equatorial ionosphere for the onset of spread F (Hysell and Kudeki, 2004), due to the growth of the Generalized Rayleigh Taylor (GRT) instability, which may evolve into a full fledged plasma bubble extending to the topside, provided that the bottomside of the equatorial ionosphere is sufficiently high (Bhattacharyya, 2004). These two mechanisms produce irregularities with a scale in the 0.1 km to 20 km range, with an extension of the range to shorter scale lengths through secondary plasma instabilities (Haerendel, 1974; Chaturvedi and Kaw, 1976; Costa and Kelley, 1978). The GRT instability includes the destabilizing effects of the ambient electric field as well as the neutral wind and gravity (Sekar and Raghavarao, 1987; Sultan, 1996). The bottom-type scattering layers associated with ESF may be considered to arise due to primary waves excited by horizontal wind-driven instabilities (Hysell et al., 2004).

ESF irregularities produce amplitude and phase scintillations on transionospheric signals in the VHF and higher frequency ranges and show day-to-day variability, as well as variability with season, solar cycle, and magnetic activity (Basu et al., 1988; Aarons, 1993, 1997; Aarons et al., 2000; Bhattacharyya and Basu, 2002). Earlier studies have been mostly concerned with such variability in the occurrence pattern of ESF. However, day-to-day variability in the evolution of spatial structure in ESF irregularities, which is of prime concern as far as the impact of these irregularities on satellite-based communication and navigation systems, such as the GPS is concerned, has not received much attention. One of the problems encountered in studying the evolution of spatial structure in ESF irregularities is that for both scintillation and radar observations, except for the irregularities observed shortly after sunset, it is not known when the irregularities observed at a given time were initially generated, as their generally eastward drift may move them onto the signal path even several hours after their initial generation.

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**Table 1.** Radar parameters and other specifications used for the spread F experiments.

Frequency	18 MHz
Peak transmitted power	50 kW
Beam direction	Zenith
Beam width (one way half power full width)	6.3°
Pulse repetition frequency	100 Hz
Pulse width	100 $\mu$ s
Number of FFT points	256
No. of online spectral averaging	16
No. of range bins sampled	32
Height resolution	15 km
Velocity resolution	3.25 m s <sup>-1</sup>
Time resolution	2 min

There have been a few studies of the dynamics of ESF irregularities using scintillation observations (Bhattacharyya et al., 2001) and radar observations combined with scintillation observations (Rodrigues et al., 2004) which have shown the occurrence of large vertical drifts of the irregularities during the growth phase of equatorial plasma bubbles which could be associated with the perturbation electric field arising due to the GRT instability process. Here it is to be noted that although only the zonal drift,  $\mathbf{V}_o$ , of the ground scintillation pattern is calculated from spaced receiver scintillation observations, the orientation of the signal path is usually such that vertical drifts of the irregularities also contribute to  $\mathbf{V}_o$ , as will be discussed in the next section. The fluctuating perturbation electric fields associated with the growing plasma instability are expected to give rise to large variations in the irregularity drift during the growth phase of radar plumes, and hence larger spectral widths and greater decorrelation between scintillations recorded by spaced receivers. In the present paper, 18-MHz radar backscatter observations and scintillations recorded on a 251-MHz signal using spaced receivers are used to study this aspect of the nonlinear evolution of ESF irregularities.

## 2 Experimental details and data analysis

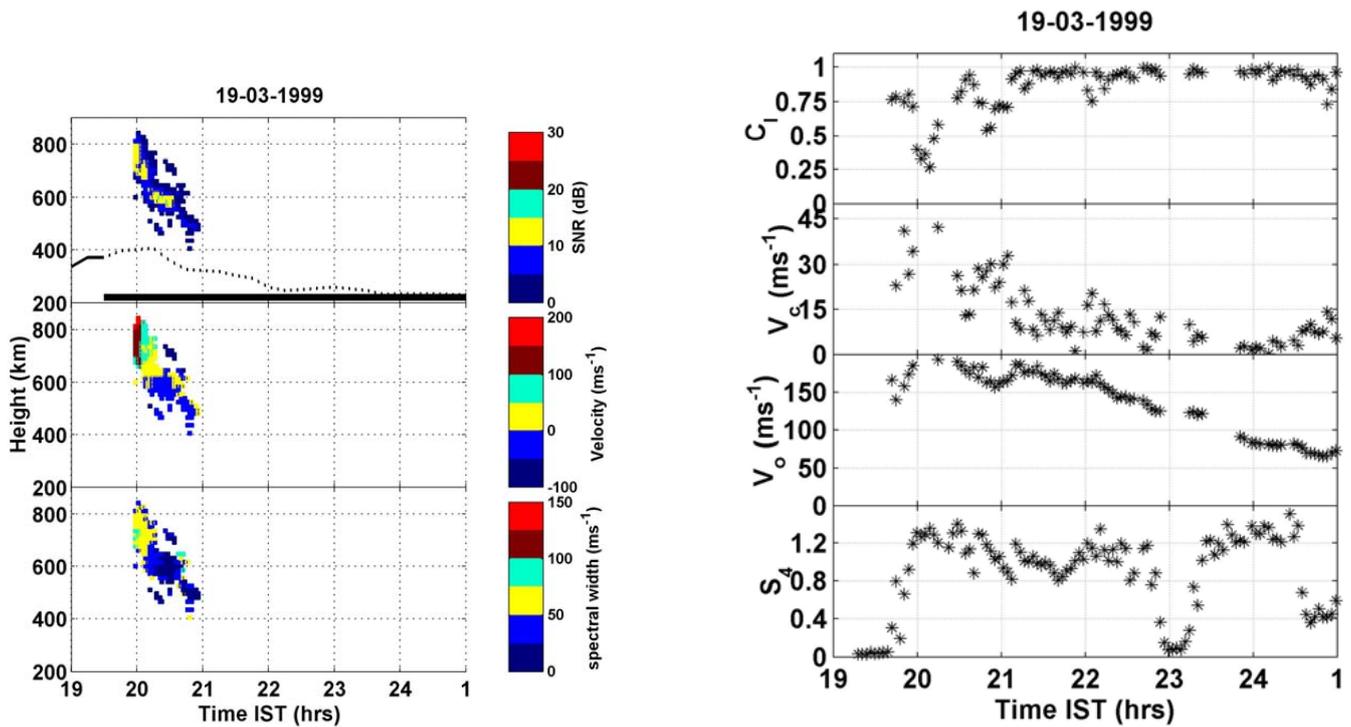
Simultaneous observations of ESF irregularities presented here were made using the 18-MHz radar located at Trivandrum (77° E, 8.5° N, dip 0.5° N), which pertain to 8.3-m scale size ESF irregularities, and spaced receivers recording scintillations on a 251-MHz signal at Tirunelveli (77.8° E, 8.7° N, dip 0.4° N). The zenith and azimuth angles for the signal path from the geostationary satellite FLEETSAT, located at 73° E, to Tirunelveli, are 11.4° and 105.6°, respectively. A calculation of the location of the ionospheric pierce

point (IPP) of the signal path at an altitude of 500 km shows that the scintillation receivers at Tirunelveli look at a volume that is approximately 50 km east of the volume seen by the radar at Trivandrum at an altitude of 500 km. It is expected that the ambient ionospheric conditions may not change substantially over this distance. Considering the movement of the sunset terminator there should be around 1.56 min delay in the onset of ESF over Trivandrum compared to that along the signal path to Tirunelveli. Thus, it is likely that if an 18-MHz radar was located beneath the IPP for the Tirunelveli–FLEETSAT signal path, it would have seen nearly the same pattern of echoes from ESF irregularities as is seen by the 18-MHz radar located at Trivandrum at a later time, after a delay of about 1.56 min.

The 18-MHz radar located at Trivandrum is a coherent pulsed Doppler radar and is operated with a maximum average power aperture product of  $1.25 \times 10^7$  Wm<sup>2</sup> (Janardhanan et al., 2001). Spectral characteristics of the signal backscattered by 8.3-m scale size ESF irregularities, which have been obtained using this radar, have been studied by Tiwari et al. (2004). The radar specifications and other important parameters used for ESF studies are given in Table 1. The three lower order (0th, 1st and 2nd) moments are estimated from the Doppler power spectra. Then the signal-to-noise ratio (SNR), line-of-sight irregularity drift and full spectral width respectively are derived from these moments, as described by Woodman (1985). These parameters have been used in the present study.

Two spaced receivers separated by a distance of 540 m, along an east-west baseline, have been used to record amplitude scintillations on a 251-MHz signal transmitted from the geostationary satellite “FLEETSAT”, and received at the equatorial station Tirunelveli. The zonal drift of the ground scintillation pattern,  $\mathbf{V}_o$ , has contributions from  $\mathbf{V}_E$ , the eastward drift and  $\mathbf{V}_Z$ , the vertical drift of the irregularities, which depends on the orientation of the signal path:  $\mathbf{V}_o = \mathbf{V}_E - \mathbf{V}_Z \tan \theta \sin \phi$ , where  $\theta$  is the zenith angle and  $\phi$  is the azimuth of the signal path measured eastward from the north. For the observations at Tirunelveli,  $\theta = 11.4^\circ$  and  $\phi = 105.6^\circ$ , yielding  $\mathbf{V}_o = \mathbf{V}_E - 0.2\mathbf{V}_Z$ . Hence, large upward and downward velocities of the irregularities, due to polarization electric fields associated with the GRT instability, give rise to large variations in  $\mathbf{V}_o$  (Bhattacharyya et al., 1989, 2001).

The scintillation data are sampled at 10 Hz. This data has been used to calculate the  $S_4$ -index, which is the standard deviation of normalized intensity, and the maximum cross-correlation  $C_I$  of the signals recorded by the two receivers, for each 3-min interval during a scintillation event. Further, full correlation analysis (Briggs, 1984) was used to calculate the average drift,  $\mathbf{V}_o$ , of the ground scintillation pattern transverse to the signal path, as well as the random velocity,  $\mathbf{V}_c$ , which is a measure of the random variability of the irregularities. For the estimations of  $\mathbf{V}_o$  and  $\mathbf{V}_c$ , data for only those three-minute intervals have been used for which  $C_I \geq 0.5$  and  $S_4 > 0.15$ , in order to reduce the effect of noise and to



**Fig. 1.** (a) Top, middle and bottom panels show the height-time-intensity (HTI), height-time-velocity (HTV) and height-time-width (HTW) maps of HF backscattered signal from spread F irregularities observed over Trivandrum on 19 March 1999. In the top panel, continuous and dashed lines represent the temporal variations of  $h'F$  in the absence and presence of ESF respectively and thick black horizontal line indicates the duration of ESF observed from the Ionosonde data (b) Four panels, from top to bottom, present the temporal variations of maximum cross-correlation  $C_1$  of the spaced receiver signals, random velocity  $V_c$ , average drift  $V_o$  of the ground scintillation pattern along the receiver baseline, and  $S_4$  index, respectively.

improve the validity of the assumptions upon which the full correlation analysis is based (Engavale et al., 2005).

### 3 Case studies

Simultaneous observations of ESF irregularities on 10 nights during March–April 1998 and 1999 made using 18-MHz radar at Trivandrum and spaced receivers recording 251-MHz scintillations at Tirunelveli have been used here to understand the evolution of spatial structure of ESF irregularities and their dynamics at different scale sizes. During the above periods, the  $A_p$  index varied in the range 3–16, which corresponds to magnetically quiet to moderate conditions.

The data here are classified into two categories, namely topside and bottomside, based on the height extent of the radar backscatter echoes. This enables us to illustrate different types of ESF structures and their association with the parameters estimated from spaced receiver scintillation measurements. It may be noted that the recorded scintillations are an integrated effect and as such, it is not possible to distinguish between topside and bottomside irregularities based on scintillation records alone. The base height of the equa-

torial F-region ( $h'F$ ) during the post-sunset period shows considerable day-to-day variability. Hence, the bottomside backscattered echoes may occur at different altitudes on different days. However, the classification here is also based on the vertical extent of the radar backscatter echoes: if the vertical extent of the echoes is less than 100 km, and the structure is essentially restricted to altitudes below 400 km, it is classified as a bottomside structure. Four case studies are shown here and later general features emerging from simultaneous observations on 10 nights are discussed. Out of the four cases presented here, three fall into the topside category, as the vertical extent of the radar backscatter echoes exceeds 100 km and extends above the 400-km altitude in these cases. During simultaneous observations on 10 nights, no example of an exclusively bottomside ESF structure was found. On 14 March 1999, a bottomside structure was observed for nearly 90 min, which was accompanied by a topside structure for a shorter duration, shortly after 20 IST (Indian standard time IST = UT+5.5 h). Study of these four cases reveals some of the different patterns of evolution of spatial structure in ESF irregularities that are part of the day-to-day variability of the ESF phenomenon. The apparent base height of the

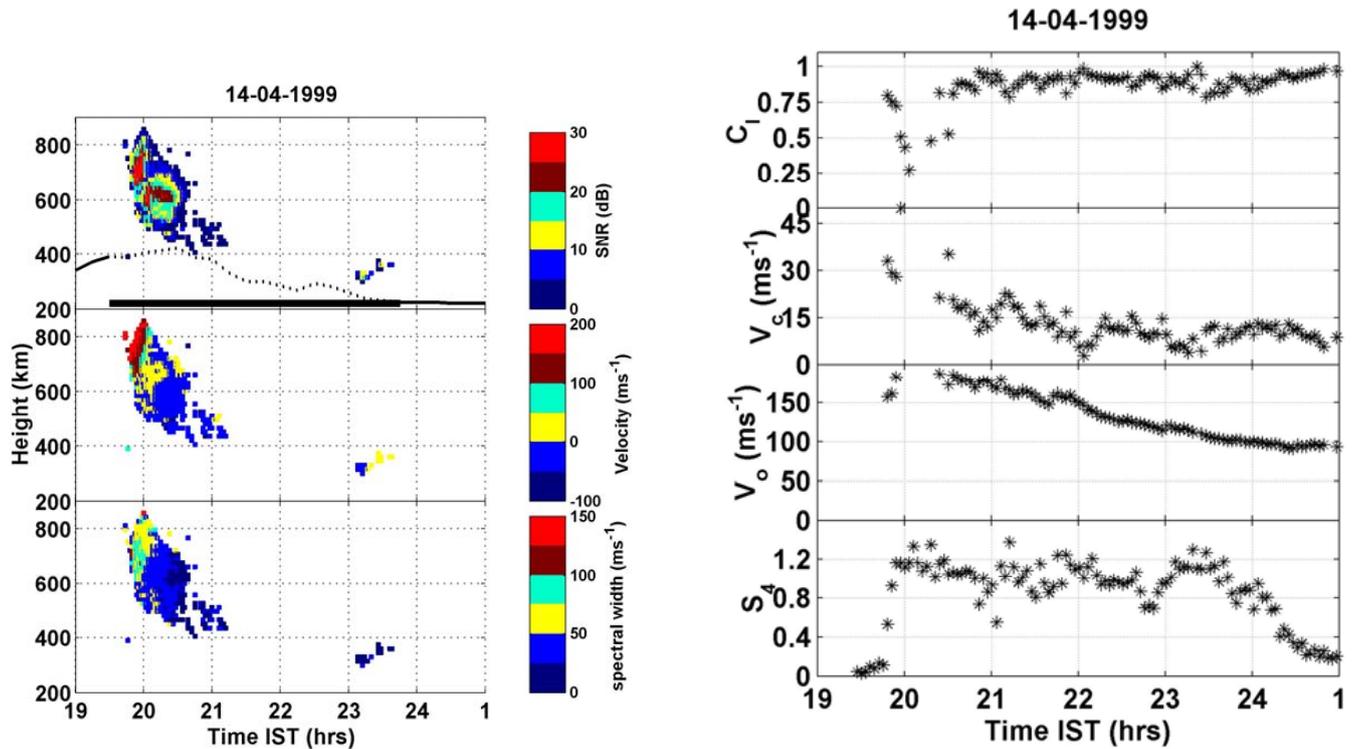


Fig. 2. (a) Same as Fig. 1a but for 14 April 1999. (b) Same as Fig. 1b but for 14 April 1999.

F-region ( $h'F$ ), as estimated from ionosonde data, obtained at Trivandrum are also indicated in each case as also the time duration of spread F seen on the Ionosonde data in each case.

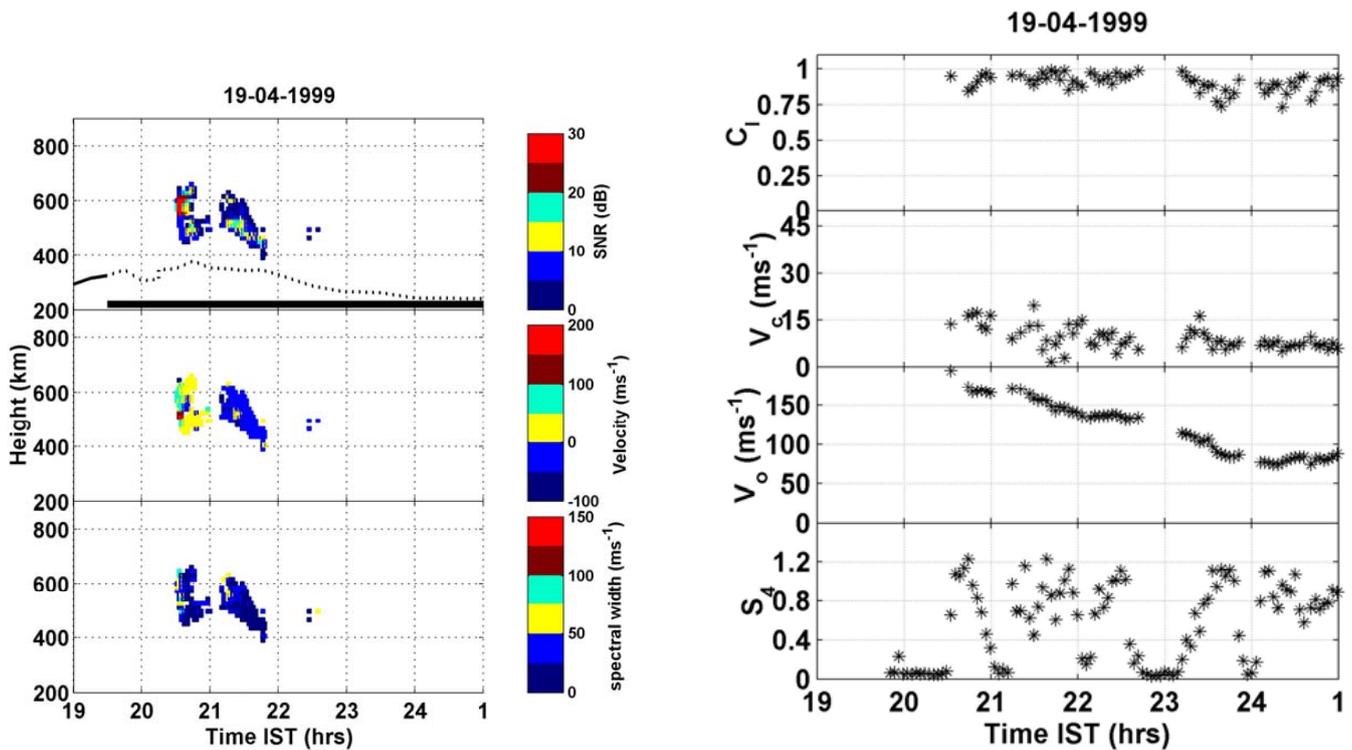
### 3.1 Topside ESF structures

#### (i) 19 March 1999 ( $A_p=6$ )

The top, middle and bottom panels of Fig. 1a present, respectively, radar measured height-time-intensity (HTI), height-time-velocity (HTV) and height-time-width (HTW) maps observed on 19 March 1999. Here “intensity” refers to the signal-to-noise-ratio (SNR) and width refers to the full spectral width, derived from the backscattered signals. On this day, ESF was observed from 19:40 to 20:40 IST and the HTI map was spread over a large altitude range of 400–820 km, extending well into the topside F region. The strongest SNR is found to be associated with the rising plume and is in the typical range of 10–15 dB above the noise level and mostly associated with upward velocities, where above  $100 \text{ m s}^{-1}$  in the topside is a clear manifestation of a rising plume structure, while the decay phase is associated with predominantly downward velocities with magnitudes typically about  $50 \text{ m s}^{-1}$ . The HTV plot shown here is indicative of the nature of updrafting and downdrafting of small plasma structures. Higher spectral widths are also found to be associated with the rising plume structure when the observed

irregularity drift exceeds  $50 \text{ m s}^{-1}$ . Continuous and dashed lines represent  $h'F$  (bottom side F layer height) as a function of time, as deduced from ionosonde data in the absence and presence of ESF respectively. The thick black line indicates the duration of ESF as seen in the ionosonde data on this night.

Four panels of Fig. 1b, starting from top, present the temporal variations of the maximum cross-correlation  $C_1$  of the spaced receiver scintillation records, random velocity  $V_c$ , average drift of the ground scintillation pattern along the baseline,  $V_o$ , and  $S_4$  index, respectively, calculated for each 3-min interval. The initial appearance of 251-MHz scintillations is seen earlier than the 18-MHz radar echoes. In the initial phase of plasma bubble development (topside ESF), as seen in the radar HTI map, the maximum cross-correlation  $C_1$  shows a sharp decrease to values below 0.5, which is attributed to the presence of perturbation electric fields associated with the GRT instability (Bhattacharyya et al., 2001).  $V_c$  has higher values in the early phase of the development of irregularities, with  $V_o$  having a larger variability during this period. Strong scintillations were observed in association with radar backscatter echoes. In the initial phase, the radar measured vertical irregularity drift velocity also has large values, with the corresponding spectral widths also being larger. It is interesting to note that although the backscattering structures were absent after



**Fig. 3.** (a) Same as Fig. 1a but for 19 April 1999. (b) Same as Fig. 1b but for 19 April 1999.

20:55 IST, a high level of scintillations was observed until 01:00 IST, similar to earlier observations reported elsewhere (Basu et al., 1978; Rodrigues et al., 2004). This is due to the decay of 8.3-m irregularities, while the longer scale length (>100 m) irregularities which formed earlier to the west of the observation point remained intact while they drifted eastward on to the signal path, to produce the recorded scintillations.

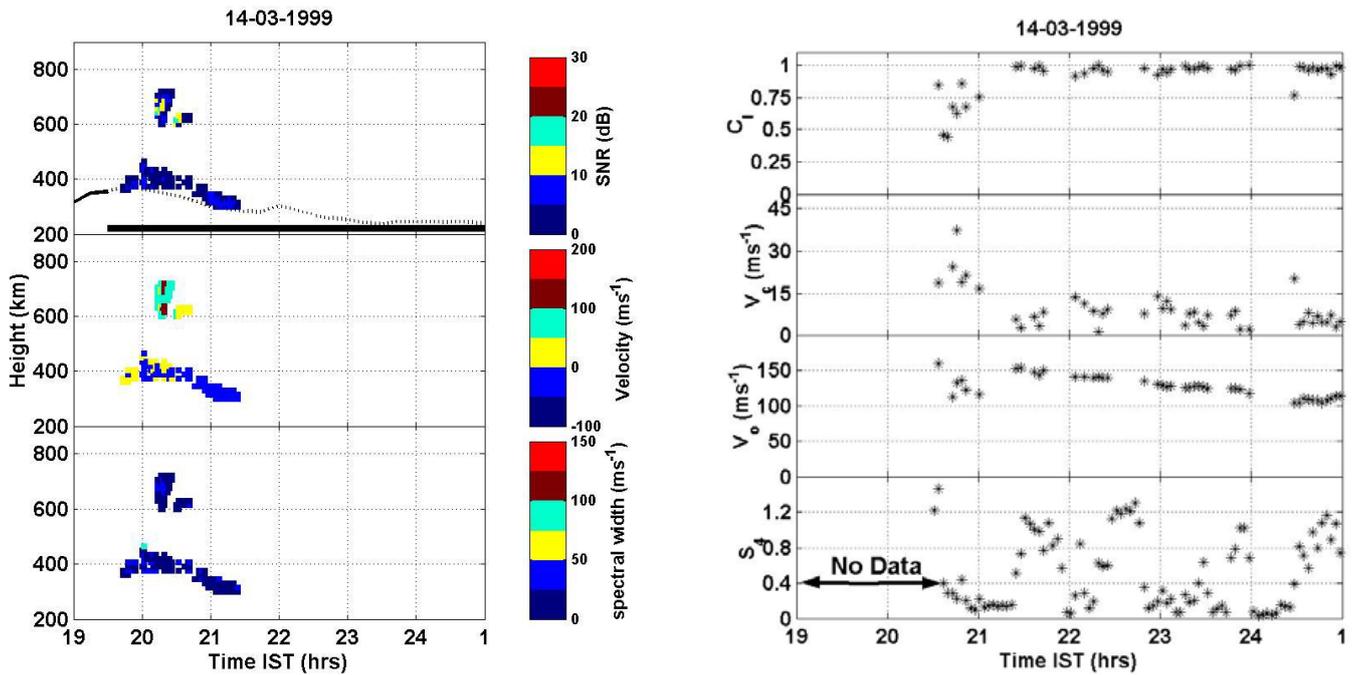
(ii) 14 April 1999 (Ap=7)

Figures 2a and b present the radar measured HTI, HTV and HTW and temporal variations of  $C_I$ ,  $V_c$ ,  $V_o$  and  $S_4$  for this day, similar to what has been presented above. On this day the onset of ESF irregularities was recorded at 19:42 IST and 18-MHz radar observed maximum SNR is found to be associated with the rising plumes and is in the typical range of 20–30 dB. Backscattering echoes disappeared at around 21:10 IST and again a very weak backscattering layer below 400 km appeared around 23:05 IST for a very brief period. On the whole, major features were similar to what has been observed on 19 March 1999. These observations suggest that the background ionospheric conditions were not very different on this day from that on 19 March 1999.

(iii) 19 April 1999 (Ap=12)

On this day the appearance of both radar echoes and VHF scintillations happened at a time later than in the previous two cases. HTI, HTV and HTW maps for this day have been shown in Fig. 3a. On this day the height extent of the radar echoes was confined to ~400–650 km and two backscatter structures were observed with the first one being stronger with maximum observed SNR in the range of 20–30 dB, and with predominantly upward velocities (50–100 m s<sup>-1</sup>) and observed spectral widths mostly in the range of 0–50 m s<sup>-1</sup>. Backscattering echoes disappeared for about 5–10 min and a second backscattering layer reappeared with weaker SNR values, dominated with downward velocity (~–50 m s<sup>-1</sup>) and small spectral widths varying in the range of 0–50 m s<sup>-1</sup>.

Figure 3b shows the temporal variations of  $C_I$ ,  $V_c$ ,  $V_o$ , and  $S_4$  index, respectively, for this day. On this day the observed maximum cross-correlation is always >0.75, with small values of  $V_c$  ( $\leq 15$  m s<sup>-1</sup>) and  $V_o$  also shows less variability compared to other events discussed above. At around the time of the disappearance of radar echoes, scintillations also disappeared briefly, but later, strong and weak scintillations were observed periodically after the disappearance of radar echoes, and this continued until 01:00 IST.



**Fig. 4.** (a) Same as Fig. 3a but for 14 March 1999 for bottomside ESF structure. (b) Same as Fig. 3b but for 14 March 1999 for bottomside ESF structure.

### 3.2 Bottomside ESF structures

14 March 1999 ( $A_p=15$ )

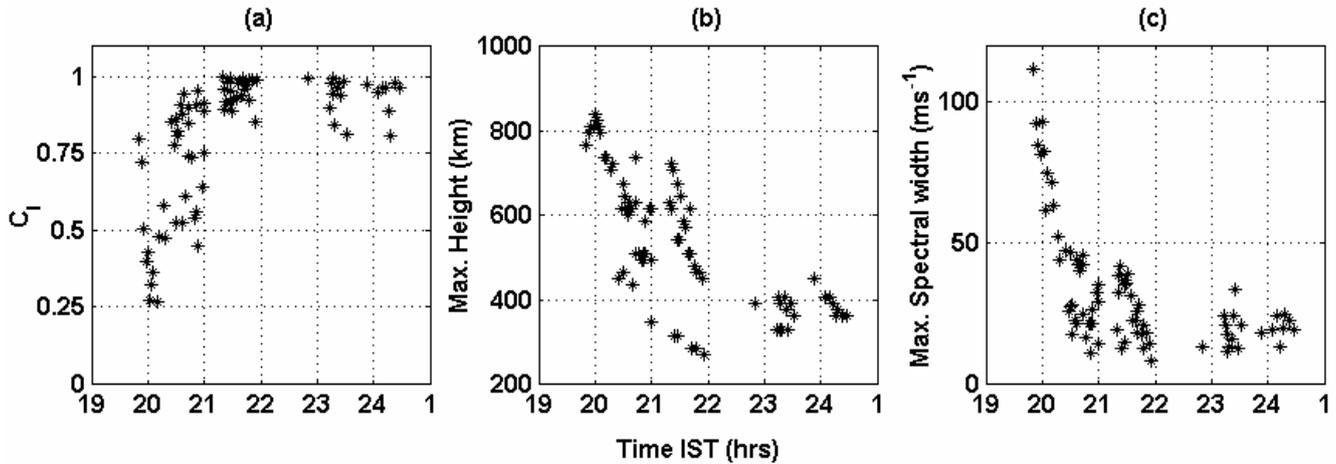
Figure 4a depicts the HTI, HTV and HTW maps for this day. The radar echoes were observed during 19:40–21:20 IST. Note that the layer does not expand significantly in height with time. The HTV map shows the upward flow in the bottomside F region with maximum velocities up to  $50 \text{ m s}^{-1}$  during the initial phase of the bottomside ESF structure. A rising plasma structure, detached from the bottomside structure, with an upward flow velocity in the range of  $50$  to  $100 \text{ m s}^{-1}$  and extending a little beyond  $700 \text{ km}$ , could be seen during 20:12–20:42 IST. The decay phase is predominantly associated with a downward velocity smaller in magnitude ( $30 \text{ m s}^{-1}$ ) at about  $300 \text{ km}$ . The thickness of the bottomside backscattering layer was of the order of  $100 \text{ km}$  with a weak SNR ( $0$ – $10 \text{ dB}$ ). Observed spectral widths on this day are mostly  $<50 \text{ m s}^{-1}$ .

Four panels of Fig. 4b, from top to bottom, present the temporal variations of  $C_I$ ,  $V_c$ ,  $V_o$ , and the  $S_4$  index, respectively. On this day no VHF scintillation data were available until 20:32 IST. From the available data, maximum cross-correlation  $C_I$  was found to fall to about  $0.5$  at 20:40 IST, with  $V_c$  showing higher values and  $V_o$  showing some variability around this time. Weak scintillations ( $S_4 < 0.25$ ) were associated with the bottomside ESF structures seen after 20:45 IST. Both strong and weak scintillations were ob-

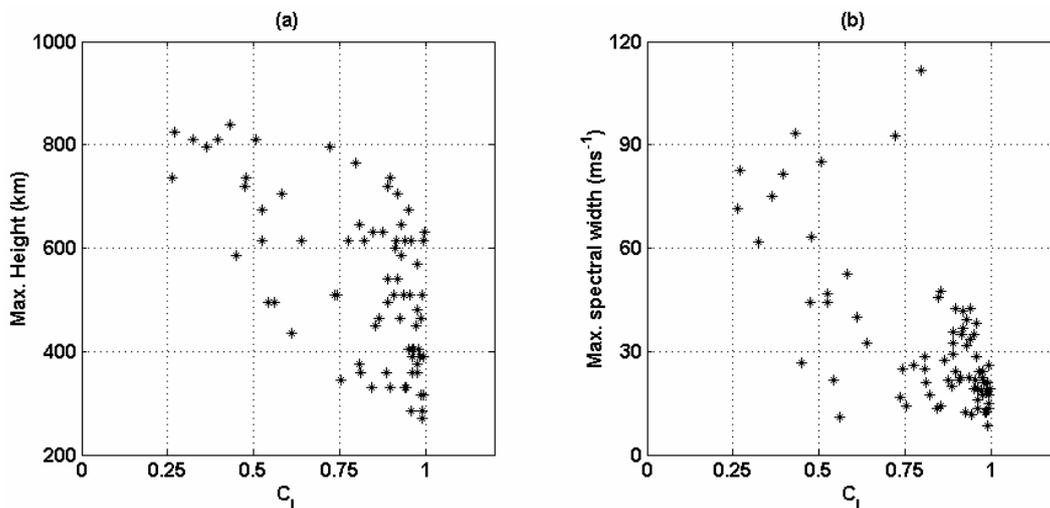
served periodically after the disappearance of radar echoes and continued until 01:00 IST.

### 3.3 General features

In order to compare the ESF related parameters estimated from 18-MHz radar and 251-MHz scintillation observations, the composite scatter plots of  $C_I$ , obtained from the scintillation observations; maximum height of 18-MHz radar echoes; and maximum spectral widths estimated from the radar observations; versus IST, for all 10 days of observations, are exhibited in Figs. 5a–c. For Figs. 6a–b, the radar derived parameters for the initial phase of bubble development have been time shifted, to take into account the time delay introduced by the movement of the sunset terminator. However, for  $C_I$  values pertaining to IST later than 21:00 IST, it is assumed that the irregularities, which produce the observed scintillation, were generated earlier to the west of the FLEETSAT-Tirunelveli signal path, and then drifted onto this signal path with an average eastward drift speed of  $150 \text{ m s}^{-1}$ . Hence, in Figs. 6a–b,  $C_I$  values have been plotted against radar derived parameters time shifted by  $5.44 \text{ min}$  to a later time, to match the time of the  $C_I$  values. From the temporal variations of  $C_I$ , as well as the radar measured maximum height of the radar echoes and maximum spectral width shown, respectively, in Figs. 5a–c, it can be seen very clearly that the rapid decrease in  $C_I$ , during the initial development of ESF, also has a counterpart in rapidly increasing



**Fig. 5.** Composite scatter plot of temporal variation of (a)  $C_I$ , (b) maximum height, and (c) maximum spectral width of the radar backscattered echoes.



**Fig. 6.** Composite scatter plot of (a) maximum height of the radar backscattered echoes versus  $C_I$  and (b) spectral width versus  $C_I$ .

spectral widths derived from radar echoes at this stage. The maximum height for the radar echoes also shows a similar pattern of variations with IST.

Figures 6a–b show the scatter plots of radar estimated maximum height versus  $C_I$ , and radar estimated maximum spectral width versus  $C_I$ , where, as mentioned above, the time for the radar parameters has been shifted to take into account (a) the movement of the sunset terminator, (b) the eastward drift of the irregularities. From Fig. 6a it is clear that for the values of  $C_I$  less than 0.5, the maximum height of the ESF irregularities is always more than 600 km. This supports the idea that during the period when the nonlinear evolution of the GRT-instability into plasma bubbles occurs resulting in topside structures, the spaced receiver signals are highly decorrelated and this fact can be used independently as an indicator of the scintillations being caused by freshly generated

plasma bubbles or by plumes. The plot of spectral width versus  $C_I$  (Fig. 6b) shows that values of  $C_I \leq 0.5$  are generally associated with spectral widths exceeding  $50 \text{ m s}^{-1}$ , indicating the role of increased turbulence in causing greater decorrelation of the spaced receiver scintillation signals. Lower values of spectral widths and higher values of  $C_I$  are generally associated with the decay phase of plasma bubbles which drift onto the signal path from the west or with the bottomside ESF structure.

#### 4 Discussion

The general features of parameters derived from 18-MHz radar backscatter and spaced receiver scintillations on a 251-MHz signal reveal that in the case of bottomside ESF

structures, generally the maximum cross correlation between the spaced receiver signals does not fall below 0.75, which is distinct from the case of the topside structures. A high degree of correlation found between spaced receiver records of scintillations have been shown to arise from the bottomside equatorial F-region irregularities layers, indicating the absence of random changes in these irregularities or their drift velocities, which would imply that the steady, nonlinear states reached are stable in this case (Bhattacharyya et al., 2001; Bhattacharyya, 2004). Some of the weaker echoes associated with bottomside structures have not been detected regularly using the 18-MHz radar system, due to low sensitivity of the system (Tiwari et al., 2004). These echoes have been observed regularly using the Jicamarca radar with SNR exceeding 30 dB (Woodman and La Hoz, 1976). Thus, the small plume structure seen detached from the bottomside structure in Fig. 4a may indeed have a neck attaching it to the bottomside structure which may not be seen due to low sensitivity of the radar.

The duration of the observed 251 MHz scintillations was generally seen to be longer than that of radar backscatter echoes, similar to observations reported elsewhere (Basu et al., 1978; Rodrigues et al., 2004). These observations indicate that in the initial phase, the kilometer- and meter-sized irregularities evolve and exist together because meter scale size irregularities are generated due to nonlinear evolution and growth of secondary plasma instabilities (Haerendel, 1974; Chaturvedi and Kaw, 1976; Costa and Kelley, 1978). In the decay phase, the meter-sized irregularities decay much more rapidly due to diffusion, whereas large-scale irregularities responsible for scintillations continue to retain their high spectral intensities (Basu et al., 1978).

The relation between  $C_I$  and the maximum height of the ESF structure shows that lower values of  $C_I$  ( $<0.5$ ) in the initial phase of irregularity generation can be used as a tracer for the evolution of plasma bubbles to topside, thus extending the usefulness of the inexpensive scintillation technique in studies of ESF irregularities.  $E \times B$  drifts, due to perturbation electric fields associated with the GRT instability, give rise to the large random fluctuations in the irregularity drift at this time.

## 5 Summary

Generally, it is not possible to determine the vertical extent of ESF irregularities from scintillation data, except in the case of weak scintillations, when the presence of Fresnel oscillations in the power spectrum of amplitude or phase scintillations indicates a relatively thin (thickness  $<100$  km) layer of irregularities (Yeh and Liu, 1982). In these special cases, the well-defined Fresnel minima also allow for the estimation of the average distance of the irregularities from the ground receiver, and hence of the average height of the irregularities, provided that the drift speed of the scintillation pattern on the

ground is available from spaced receiver measurements. The recently floated idea that, in the case of ESF irregularities, decorrelation of the scintillating signals recorded by spaced receivers is due to the presence of significant fluctuating perturbation electric fields associated with the GRT instability in the initial phase of its nonlinear evolution into a plasma bubble (Bhattacharyya et al., 2001), is supported by the following results derived from simultaneous observation of 18-MHz radar backscatter from a volume of ionosphere located about 50 km west of the ionospheric region, encountered by the signal from the FLEETSAT satellite to Tirunelveli.

1. In the initial phase of plasma bubble developments (topside ESF), as seen in the radar HTI map, the maximum cross-correlation  $C_I$  falls much below 1, due to large decorrelation of the 251-MHz signals recorded by the two spaced receivers.
2.  $V_c$  and the spectral width tend to have higher values in the early phase of the development of irregularities, because of the presence of perturbation electric fields associated with the GRT instability.
3. There is large variability in  $V_o$  at this stage in the development of the ESF irregularities. At the same time, radar measured irregularity vertical drift velocity has large values with the corresponding spectral widths being higher.
4. After about 21:00 IST when the perturbation electric field associated with the GRT instability is eroded, it is seen that the maximum value of  $C_I$  increases, and  $V_c$  becomes smaller, due to the decrease in the height of the irregularities as seen from the radar observations.
5. It is observed that plasma bubbles in the decay phase and bottomside ESF structures are generally associated with high  $C_I$  values.

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