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Absolute absorption cross-section and photolysis rate of I₂

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Abstract

Following recent observations of molecular iodine (I_2) in the coastal marine boundary layer (MBL) (Saiz-Lopez and Plane, 2004), it has become important to determine the absolute absorption cross-section of I_2 at reasonably high resolution, and also to evaluate the rate of photolysis of the molecule in the lower atmosphere. The absolute absorption cross-section (σ) of gaseous I_2 at room temperature and pressure (295 K, 760 Torr) was therefore measured between 182 and 750 nm using a Fourier Transform spectrometer at a resolution of 4 cm^{-1} (0.1 nm at $\lambda=500\text{ nm}$). The maximum absorption cross-section in the visible region was observed at $\lambda=533.0\text{ nm}$ to be $\sigma=(4.84\pm 0.60)\times 10^{-18}\text{ cm}^2\text{ molecule}^{-1}$. The spectrum is available as supplementary material accompanying this paper. The photo-dissociation rate constant (J) of gaseous I_2 was also measured directly in a solar simulator, yielding $J(I_2)=0.12\pm 0.03\text{ s}^{-1}$ for the lower troposphere. This agrees well with the value of $0.15\pm 0.03\text{ s}^{-1}$ calculated using the measured absorption cross-section, terrestrial solar flux for clear sky conditions and assuming a photo-dissociation yield of unity. A two-stream radiation transfer model was then used to determine the variation in photolysis rate with solar zenith angle (SZA), from which an analytic expression is derived for use in atmospheric models. Photolysis appears to be the dominant loss process for I_2 during daytime, and hence an important source of iodine atoms in the lower atmosphere.

1. Introduction

Over the past two decades, there has been a growing interest in the role that iodine chemistry plays in a number of atmospheric processes, in particular the iodine-catalysed destruction of O_3 in the troposphere (e.g. Jenkin et al., 1985; Chatfield and Crutzen, 1990; Davis et al., 1996; McFiggans et al., 2000). Furthermore, catalytic cycles involving chlorine and bromine appear not to be sufficient to explain O_3 depletion in the lower stratosphere at low latitudes, leading to speculation that iodine also plays

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a role in stratospheric chemistry (Solomon et al., 1994). Iodine may also be central to other atmospheric phenomena, including new particle formation in marine environments (O'Dowd et al., 2002a, b; Mäkelä et al., 2002) and the enhancement of halogen / inter-halogen release via uptake on sea-salt aerosol (Vogt et al., 1999; McFiggans et al., 2002).

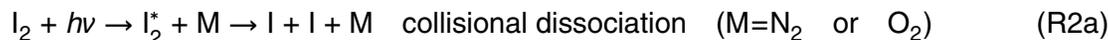
The primary source of iodine in the atmosphere has been postulated to be the evasion of biogenic iodocarbons from the sea surface (Carpenter, 2003a; Carpenter et al., 2003b). CH₃I, CH₂I₂, CH₂IBr and CH₂ICI have all been detected in coastal environments, with emissions showing a direct correlation with tidal height and high solar irradiance (Carpenter et al., 1999, 2001). In terms of the open ocean, evaporation of iodine from the water surface, via oxidation of iodide ions in seawater under solar irradiation, has also been suggested (Miyake and Tsunogai, 1963). An alternative mechanism for the liberation of iodine from seawater involves the reaction of iodide and iodate species with atmospheric O₃ (Garland and Curtis, 1981).

The recent detection of molecular iodine at Mace Head, Ireland by differential optical absorption spectroscopy (DOAS), suggests that I₂ may also make a significant contribution to the overall iodine budget in the atmosphere (Saiz-Lopez and Plane, 2004). I₂ was found at parts per trillion (ppt) levels in the coastal marine boundary layer (MBL), with a marked increase in concentration around low tide. This correlation between I₂ production and tidal height suggests an oxidative stress mechanism within exposed macro-algal beds as the probable source of I₂ in the coastal MBL.

In the troposphere, I atom production takes place through photo-dissociation of iodocarbons (RI/R'I₂)



Photolysis of I₂ in the atmosphere will also generate I atoms through the following channels,





where I_2^* is an electronically excited iodine molecule.

The subsequent reaction of I atoms with atmospheric O_3 leads to the formation of iodine monoxide (IO), which regenerates I atoms on photolysis,



or participates in iodine-catalysed O_3 destruction cycles through reactions with HO_2 (forming HOI), NO_2 (forming IONO₂) or halogen oxides (Jenkin et al., 1985; Davis et al., 1996; McFiggans et al., 2000). The IO self-reaction has also been identified as an important process leading to the homogeneous nucleation of MBL aerosol. New particle formation is thought to occur via the polymeric build-up of iodine oxide aerosol formed by reactions of the iodine dioxide (OIO) radical, one of the products from the self-reaction of IO (Bloss et al., 2001; Hoffmann et al., 2001; Jimenez et al., 2003).

The visible-infrared spectrum of I_2 has been the subject of numerous previous studies (Rabinowitch and Wood, 1936a; Goy and Pritchard, 1964; Ogryzlo and Thomas, 1965; Brewer and Tellinghuisen, 1972; Tellinghuisen, 1973). I_2 has a well-defined absorption structure in the visible region of the spectrum corresponding to the B←X electronic transition (Calvert and Pitts, 1966; Tellinghuisen, 1973). Photolysis in the visible region can proceed through two channels (Okabe, 1978):



where I represents a ground state atom $\text{I}(^2\text{P}_{3/2})$ and I^* denotes a meta-stable $\text{I}(^2\text{P}_{1/2})$ atom. Subsequent relaxation from the meta-stable state to the ground state occurs through quenching processes or radiative decay ($\tau \sim 50$ ms, Okabe, 1978):



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Notwithstanding this earlier work, a sufficiently high resolution absolute absorption cross-section for I₂ suitable for DOAS retrievals, and a measured rate of I₂ photolysis (appropriate for atmospheric modelling), do not appear to be available in the literature.

In this paper we will describe a measurement of the absolute absorption cross-section of I₂ from the vacuum ultra-violet to the near infra-red. The photolysis rate of I₂, determined directly in a solar simulator, will then be compared with values calculated using the measured absorption spectrum and single and two-stream radiation models.

2. Experimental

2.1. Cross-section determination

The I₂ spectrum was recorded using a Fourier Transform (FT) spectrometer (Bruker, Model IFS/66). Iodine crystals (Lancaster, 99.5%) were placed in an optical cell and allowed to equilibrate at room temperature (295 K) and pressure (760 Torr of air) before each measurement was made. The spectrum was recorded in three overlapping regions: 182–500 nm, GaP diode detector, Deuterium light source; 260–555 nm, GaP diode detector, Tungsten lamp; 500–1100 nm, Si diode, Tungsten lamp. A CaF₂ beam-splitter was used in each case. All measurements were made at a resolution of 4 cm⁻¹ (0.1 nm at λ=500 nm).

In order to calculate the absolute cross-section of I₂ from its absorption spectrum, an accurate measurement of the vapour pressure of I₂ is required. However, available data is sparse (see below), and so a new measurement was performed. Iodine crystals were placed in a 1/4" stainless steel U-tube, which was sealed at one end by a pair of Baratron pressure transducers, with 0–1000 Torr and 0–10 Torr ranges. The other end of the tube was attached to a vacuum line incorporating a diffusion pump, from which the tube could be isolated by a valve. The U-tube was immersed in water, in order to maintain it at a constant ambient temperature. The temperature of the water was

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monitored using a K-type thermocouple and was essentially constant (<0.4 K variation) over the duration of the experiment.

The vapour pressure of the iodine was measured by first evacuating the tube as far as possible, then isolating it from the vacuum line and allowing the pressure to rise until it reached a constant value. This procedure was repeated until a consistent final pressure was observed. The vapour pressure of I₂ at 295 K was found to be 0.225±0.020 Torr, equivalent to a vapour concentration of (7.4±0.7)×10¹⁵ molecules cm⁻³. This value is consistent with the vapour pressure of 0.30 Torr at 298 K from the JANAF Tables (Chase, 1998). In addition, our value compares very well with a pressure of 0.24 Torr evaluated from the measured data reported by Baxter et al. (1907), empirically fitted by Baxter and Grose (1915), and used by Shirley and Giauque (1959) in their measurements of thermodynamic properties of iodine.

2.2. Rate of photolysis of I₂

Figure 1 is a schematic diagram of the experimental set-up used. An excess of O₃ (>1×10¹⁶ molecules cm⁻³), produced from an O₂ flow through a corona-discharge, was used to ensure that the dominant loss process of I atoms was reaction with O₃ ($k_{295\text{K}}(\text{I}+\text{O}_3\rightarrow\text{IO}+\text{O}_2)=1.2\times 10^{-12}\text{ cm}^3\text{ molecule}^{-1}\text{ s}^{-1}$; De More et al., 1997). Under these conditions, I atom recombination in the gas phase and on the cell walls was negligible. I₂ vapour was generated by maintaining a steady low flow (5–10 standard cm³ min⁻¹) of N₂ across a few iodine crystals in a glass trap. The reaction took place in a glass cell (length=25 cm, volume≈150 cm³) onto which quartz windows were mounted at each end to allow for irradiation by the light source and measurement of the light intensity. Flows of N₂ (250–300 standard cm³ min⁻¹) were also passed across each cell window to prevent any build-up of solid product on the window surfaces. The total gas flow through the cell produced a gas residence time of 12.5±2.5 s.

Light from a “solar simulator” (Oriel Instruments, Model 91191), which employs a

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1000 W Xenon lamp with an air-mass optical filter to simulate the solar irradiance between 350 and 750 nm ($\lambda_{\max}=583$ nm), was directed through the cell. A condenser lens was used to ensure that the cell windows were illuminated completely with near-uniform light intensity. Neutral density filters were used to attenuate the light intensity, which was measured with a calibrated radiometer (Ramsden Scientific Instruments 550).

Downstream of the photolysis cell, the gas flowed through a 1 m-long cell fitted with quartz windows, which was used to monitor I₂ by absorption of the 546 nm emission line from a low pressure Hg lamp ($\sigma_{546\text{ nm}}=3.9\times 10^{-18}$ cm² molecule⁻¹, from this study). A monochromator (Optometrics, Model MC1-02, 1.2 nm resolution), with a side-on photomultiplier (Hamamatsu Photonics, Model H7710-12) and picoammeter allowed the I₂ concentration to be determined from its measured absorbance over the range 0.01 to 0.08, with a precision of $\pm 5\%$. Both the photolysis cell and the absorption cell were covered with black cloth to minimise the effects of the laboratory lights.

3. Results and discussion

3.1. I₂ absorption cross-section

The absorption cross-section of I₂ from 182 to 750 nm is shown in Fig. 2. The 182–499 nm region is dominated by the Cordes bands ($D^1\Sigma_u^+ \leftarrow X^1\Sigma_g^+$) between 182 and 200 nm. This spectral region is clearly irrelevant with respect to photolysis in the lower troposphere, where only solar radiation at $\lambda > 290$ nm is available. In the 500–750 nm region of the I₂ spectrum, the underlying continuum is overlapped with vibrational structure between 500 and 650 nm, which arises from the $B^3\Pi(0_u^+) \leftarrow X^1\Sigma_g^+$ transition (Calvert and Pitts, 1966; Okabe, 1978).

The last study to report I₂ spectral data (extinction coefficients) was that of Tellinghuisen (1973). Figure 3 shows the cross-section from our study at a resolution of 4 cm⁻¹ plotted against the values calculated from the data of Tellinghuisen measured at 2.6 nm resolution and reported at 10 nm wavelength intervals. The difference in res-

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olution makes direct comparison of values difficult. Nevertheless, the results show a relatively good agreement in the region from 570 to 630 nm, whereas below 570 nm, our values are significantly higher (27% higher at 500 nm). In the study of Tellinghuisen, measurements were made at room temperature (295–300 K) and with no added bath gas. In addition, the extinction coefficients from 510–620 nm were determined by extrapolating to zero pressure.

The effect of ambient pressure on the absorption spectrum of I₂ has been treated in a number of studies (Rabinowitch and Wood, 1936a; Goy and Pritchard, 1964; Ogryzlo and Thomas, 1965). These investigations demonstrated that as the total gas pressure is raised (at constant I₂ vapour pressure), absorption of I₂ in the structured region from 500 to 570 nm initially increases before levelling off at a pressure of about 500 Torr. Our cross-section value at and below the convergence limit of 499 nm agrees to within 10% of that calculated from Fig. 2 of the study of Rabinowitch and Wood, which was determined under comparable temperature-pressure conditions. This study, along with that of Ogryzlo and Thomas (1965), concluded that, as a result of pressure broadening in the structured region of the spectrum, the high pressure spectral profile of I₂ is the “true” profile in terms of the extinction coefficient (absorption cross-section) data. Consequently, the difference in experimental pressure conditions between our work and that of Tellinghuisen is likely to be responsible for the higher cross-section values determined in our study.

3.2. Photolysis rate measurements

Figure 4 illustrates the I₂ concentration measured downstream of the photolysis cell (see Fig. 1), as a function of time. The photolysis rate was determined from the following sequence of steps. Initially, [I₂] was measured with no O₃ in the cell and the photolysis lamp blocked. When the O₃ was then turned on, the dark (thermal) reaction of I₂ with O₃ ($k_{295\text{K}}=3.8\times 10^{-18}\text{ cm}^3\text{ molecule}^{-1}\text{ s}^{-1}$; Vikis and MacFarlane, 1985) proceeded until a (lower) steady-state [I₂] was reached (at 30–45 s). This concentration is termed [I₂]₀ below. The next step was to unblock the photolysis light (at 60 s) and

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irradiate the cell. The photo-oxidation of I₂ with O₃ was then monitored over a 3-min period, with [I₂] reaching a new steady state (at 75–90 s).

The observed changes in absorption at 546 nm in the downstream absorption cell are the result of both I₂ photolysis and extinction by particles generated from condensation of the iodine oxide products in the photolysis cell (Cox and Coker, 1983; Vikis and MacFarlane, 1985; Jimenez et al., 2003; Burkholder et al., 2004). In order to correct for particle extinction, measurements under dark and light reaction conditions were also taken at three other Hg lamp emission lines (365, 405 and 436 nm), where absorption by I₂ and O₃ is negligible. The absorption dependencies for the dark and light reactions were found to vary as λ⁻⁶ and λ⁻⁵, respectively. These data were extrapolated to fit the particle absorption at 546 nm and the initial absorption measurements (due to I₂+ particles) corrected accordingly. Particle absorption at 546 nm was found to be significant in both cases (30–35% of the total measured absorption for the dark reaction, and 80–90% when photolysis of I₂ occurred).

The rate at which I₂ is removed from the reaction cell as a result of photolysis (I₂+hν → I+I) is described by the first-order expression

$$\frac{d[I_2]}{dt} = -J(I_2) \times [I_2], \quad (1)$$

which integrates to give

$$[I_2]_t = [I_2]_0 \times \exp(-J(I_2) \times t), \quad (2)$$

where [I₂]_t is the concentration at time (t), [I₂]₀ is the initial concentration and J(I₂) is the first-order rate of photolysis.

J(I₂) values were calculated from Eq. (3) by taking [I₂]₀ as the dark reaction steady-state concentration, and [I₂]_t as the steady-state concentration after the I₂ had been irradiated for τ=12.5±2.5 s (the residence time in the photolysis cell):

$$J(I_2) = \ln\left(\frac{[I_2]_0}{[I_2]_t}\right) \times \frac{1}{\tau}. \quad (3)$$

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Figure 4 shows five data sets measured with an irradiance of 1350 W m^{-2} (~ 1 solar constant), from which a mean value for $J(\text{I}_2)$ of $0.14 \pm 0.04 \text{ s}^{-1}$ is obtained.

Measurements were made at different irradiances in order to establish the variation of photolysis rate with light intensity (Fig. 5). The data shows a linear fit with a slope of $(1.10 \pm 0.07) \times 10^{-4} \text{ s}^{-1} \text{ W}^{-1} \text{ m}^2$. A cloud free sky solar irradiance at the Earth's surface of 1100 W m^{-2} (assuming 20% attenuation by atmospheric absorption – Houghton, 2002), yields $J(\text{I}_2) = 0.12 \pm 0.03 \text{ s}^{-1}$ for an overhead sun.

The total rate of photolysis of I₂ can also be computed from the relation

$$\begin{aligned}
 J(\lambda) &= \int_{\lambda} F(\lambda) \times \sigma(\lambda) \times \phi(\lambda) d\lambda \\
 &= \int_{\lambda} j(\lambda) d\lambda,
 \end{aligned}
 \tag{4}$$

where $F(\lambda)$ denotes the incident actinic flux ($\text{photons cm}^{-2} \text{ s}^{-1} \text{ nm}^{-1}$), $\sigma(\lambda)$ is the absorption cross-section ($\text{cm}^2 \text{ molecule}^{-1}$) and $\phi(\lambda)$ is the quantum yield for photo-dissociation ($0 < \phi(\lambda) < 1$). $\phi(\lambda)$ for I₂ was set at a value of 1 for $290 < \lambda < 499 \text{ nm}$ and for $651 < \lambda < 750 \text{ nm}$. In the structured region of the spectrum ($500 < \lambda < 650 \text{ nm}$), competing fluorescence ($\text{I}_2^* + \text{M} \rightarrow \text{I}_2 + \text{M} + h\nu$) can reduce the photo-dissociation yield from unity (Okabe, 1978). At low ambient gas pressure, the quantum yield in this region was reported as a function of wavelength by Brewer and Tellinghuisen (1972) such that $\phi_{\text{min}} = 0.33$ at 590 nm. However at high pressure ($> 150 \text{ Torr}$), collisional quenching of excited I₂ molecules (R2a) leads exclusively to dissociation to I atoms (Rabinowitch and Wood, 1936b). Therefore for our calculations, $\phi(\lambda)$ was set at a value of 1 across the whole spectrum. Using the measured $\sigma(\lambda)$ from this study (Fig. 2) and extraterrestrial actinic flux $F_0(\lambda)$ from the SOLAR2000 irradiance model (Tobiska et al., 2000), corrected for atmospheric absorption, $j(\text{I}_2)$ values with respect to the tropospheric boundary layer were calculated at a resolution of 1 nm (Fig. 6). Integration of the area under this plot gives a total $J(\text{I}_2) = 0.15 \pm 0.03 \text{ s}^{-1}$ for an overhead sun. This estimate is in satisfactory agreement with the measured value above.

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A two-stream radiation scheme (Thompson, 1984) was used to determine the rate of photolysis of I₂ in the MBL as a function of solar zenith angle (SZA). Values corresponding to cloud-free conditions, were determined for early August (Julian Day 214) at Mace Head on the west coast of Ireland (53°20' N). This location and time-of-year correspond to the reported observations of I₂ (Saiz-Lopez and Plane, 2004). Much of the work into iodine oxide particle production has also been carried out at this site (O'Dowd et al., 2002b). Using our measured absorption cross-section, $J(\text{I}_2)=0.17 \text{ s}^{-1}$ at midday (SZA=36°) for a zero surface albedo. For a typical water surface albedo of 0.1, $J(\text{I}_2)$ increases to 0.20 s^{-1} .

Hough (1988) showed that for clear-sky conditions at an altitude of 0.5 km, the variation between the photolysis rate of a species and the SZA can be described by:

$$J(\text{SZA}) = L \cos(\text{SZA})^M \times \exp(-N \sec(\text{SZA})) \quad (5a)$$

$$\text{or for SZA} = 0, \quad J_{\max} = L \times \exp(-N), \quad (5b)$$

where L , M and N are parameters calculated by convoluting the absorption cross-section and quantum yield of a species with the solar actinic flux data, determined either by direct measurement using spectro-radiometry or calculated by radiative transfer modelling. Equation (5) has been used to generate J values for a number of volatile organic species (VOC's) found in the troposphere (Jenkin et al., 1997) and more recently with respect to halogen-containing molecules including I₂, Br₂ and Cl₂ (Kanaya et al., 2003). We have therefore fitted Eq. (5) to the clear-sky data obtained from the two-stream radiation code, yielding values of $L=0.223$, $M=0.002$ and $N=0.205$ for a zero surface albedo. Figure 7 illustrates that Eq. (5) provides a very satisfactory parameterisation of $J(\text{I}_2)$ for SZA<80°. For a surface albedo of 0.1, $L=0.252$, $M=0.002$ and $N=0.185$. Also shown in Fig. 7 is the variation in the photolysis rate calculated using the parameters given by Kanaya et al. (2003), which predicts $J(\text{I}_2)=0.15 \text{ s}^{-1}$ at a SZA of 36°. The differences observed in the two plots most likely result from our higher measured absorption cross-section, compared to the lower resolution, low pressure data of Tellinghuisen (1973) used by Kanaya et al.

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The photolysis rates were calculated for Cl₂ and Br₂ using the data from Kanaya et al. (2003). Table 1 lists these values along with other parameters which clearly show the trends from Cl₂→Br₂→I₂. Decreasing bond energies, increasing absorption cross-sections and the shift of peak absorption to the visible region of the solar spectrum lead to an increase in *J* values. Finally, Fig. 8 summarises the convoluted photolysis rate constants as a function of wavelength for a number of gas-phase species which release atomic iodine upon photolysis in the MBL. The integrated *J* values (290 nm<*λ*<750 nm) are (s⁻¹): 8.0×10⁻³ (CH₂I₂), 4.6×10⁻⁵ (CH₃I), 4.5×10⁻⁴ (CH₂ClI), 1.1×10⁻³ (CH₂BrI), 5.7×10⁻³ (HOI) and 0.15 (I₂).

4. Conclusions

A new measurement of the absolute absorption cross-section for I₂ over the wavelength range from 182 to 750 nm is reported for the conditions of 295 K and 760 Torr. We report a maximum absorption cross-section in the actinic wavelength range of $\sigma=(4.84\pm0.60)\times10^{-18}\text{ cm}^2\text{ molecule}^{-1}$ at $\lambda=533.0\text{ nm}$. The I₂ cross-section was measured at a resolution suitable for DOAS field measurements, and is provided as supplementary material to this paper.

The rapid photolysis of I₂ in the lower atmosphere (lifetime=5–10 s for an overhead sun) results from the strong absorption of visible wavelengths (400<*λ*<700 nm) which penetrate down to the troposphere. Photolysis is therefore almost certainly the major loss process for I₂ in the MBL during daytime, leading to the formation of iodine oxide aerosol and causing significant O₃ depletion in the MBL (Saiz-Lopez and Plane, 2004).

Supplementary material

The I₂ absorption cross-section from 182–750 nm reported in this study is available in spreadsheet format.

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Table 1. Trends in properties of the halogen diatomics.

Halogen (X ₂)	Bond Energy ^a (kJ mol ⁻¹)	(10 ²⁰)σ _{max} (cm ² molecule ⁻¹)	λ _{max} (nm)	J (s ⁻¹)
Cl ₂	243	26 ^b	330 ^b	0.0022 ^c
Br ₂	193	68 ^b	420 ^b	0.034 ^c
I ₂	151	484 ^d	533 ^d	0.18 ^e

^a at 298 K (CRC Handbook of Chemistry & Physics, 2001).

^b at 298 K, from the IUPAC gas kinetic database.

^c values calculated for SZA=0 using parameters given in Kanaya et al. (2003).

^d measured in this study.

^e from this study for SZA=0.

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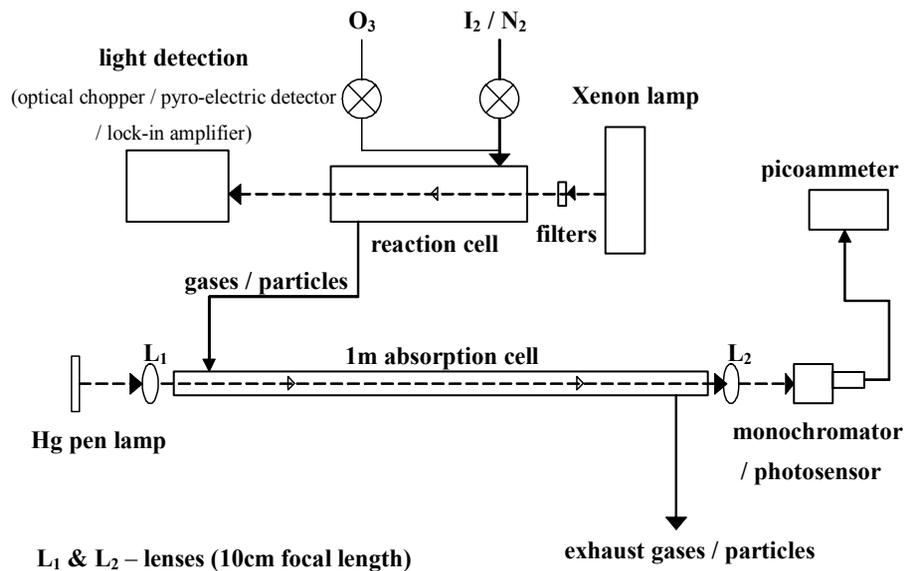


Fig. 1. Schematic of the optical absorption experiment used to measure the rate of photolysis of I_2 .

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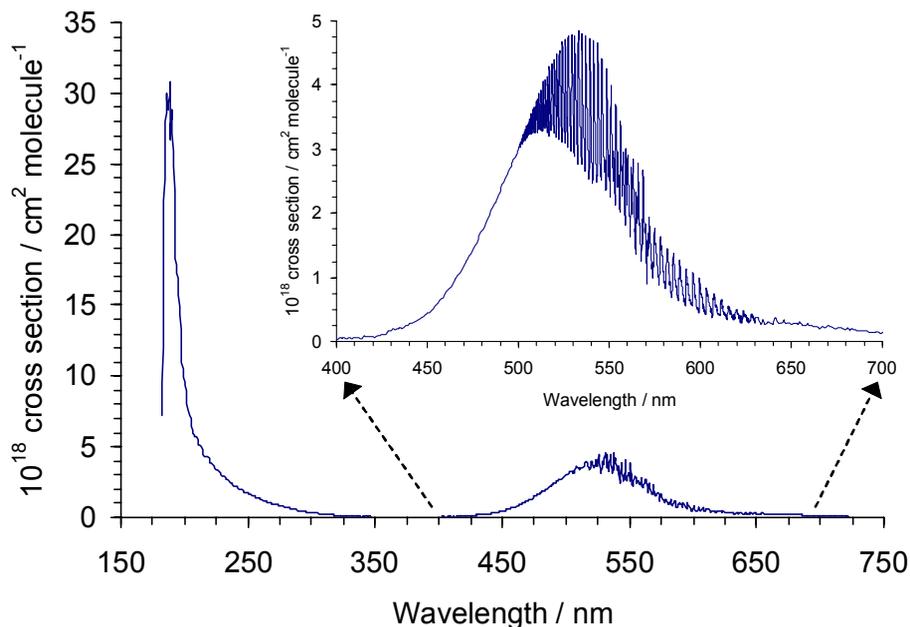


Fig. 2. Measured I₂ absorption cross-section profile (full range – 1 nm interval) with visible region (400–700 nm) expanded at higher resolution (4 cm⁻¹) to show the detailed vibrational structure from 500–650 nm.

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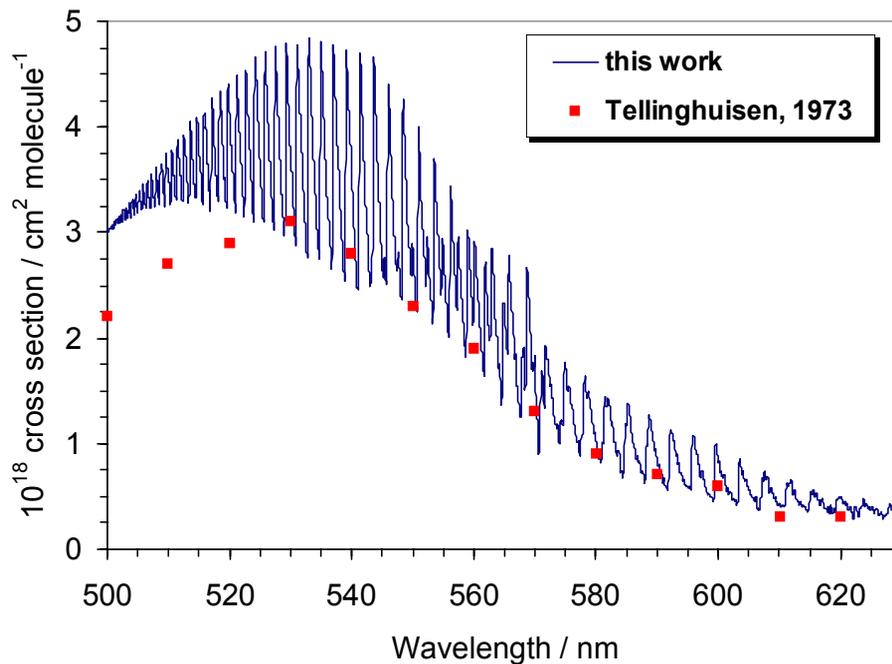


Fig. 3. Comparison of the high resolution I₂ spectrum (500–630 nm) from this study with the values calculated using the data from Tellinghuisen (1973).

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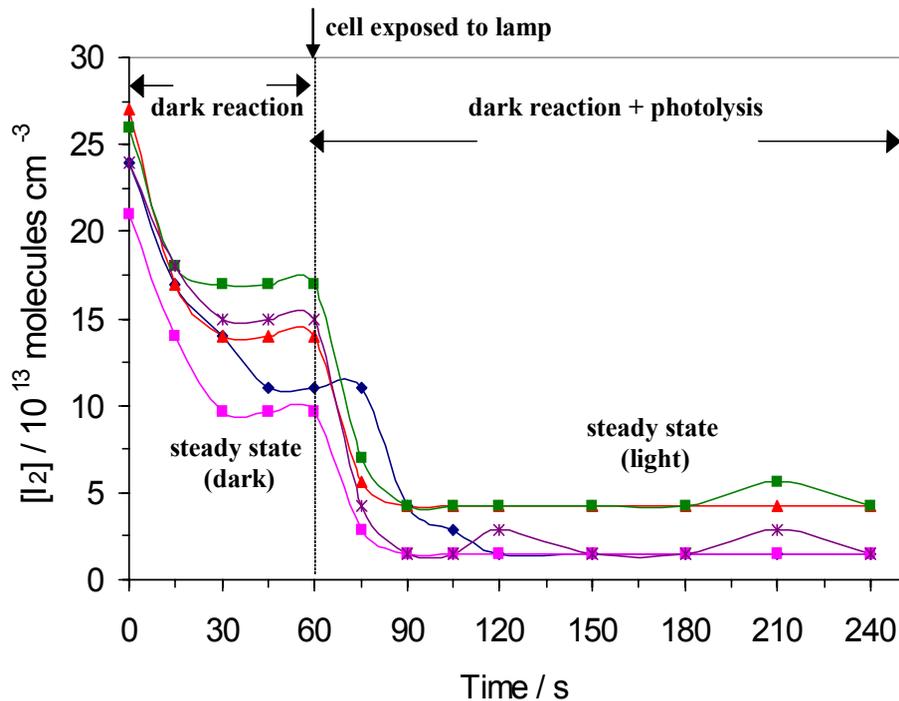


Fig. 4. Data sets for I₂+O₃ photo-oxidation at a cell irradiance of 1350 Wm⁻². The different symbols refer to individual experimental runs under the same conditions.

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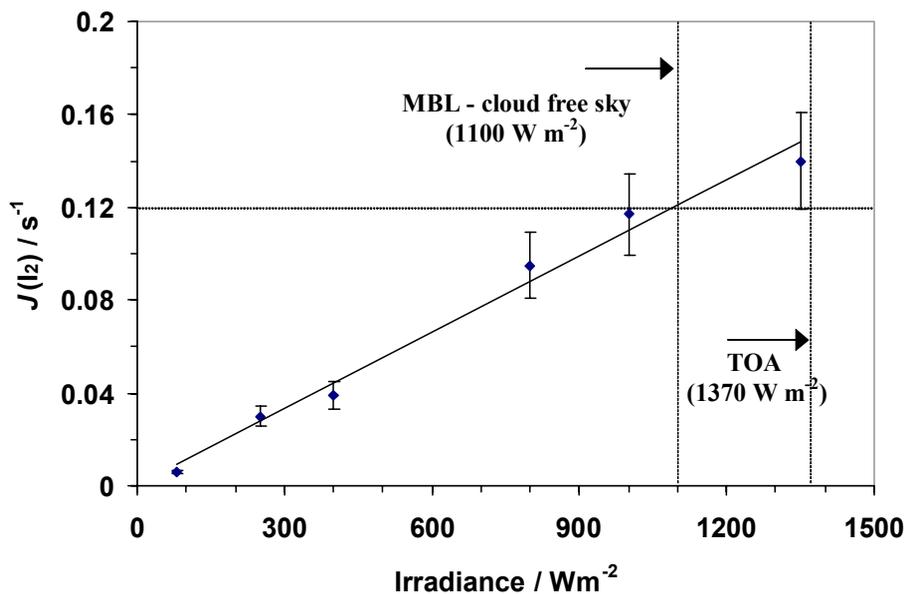


Fig. 5. Variation of the laboratory measured rate of I_2 photolysis with light irradiance. Data points shown correspond to the mean values calculated from 5 experiments run for each light irradiance value, all other reaction parameters being identical (error bars represent standard deviation in individual measurements). TOA (top of atmosphere) irradiance = 1 solar constant.

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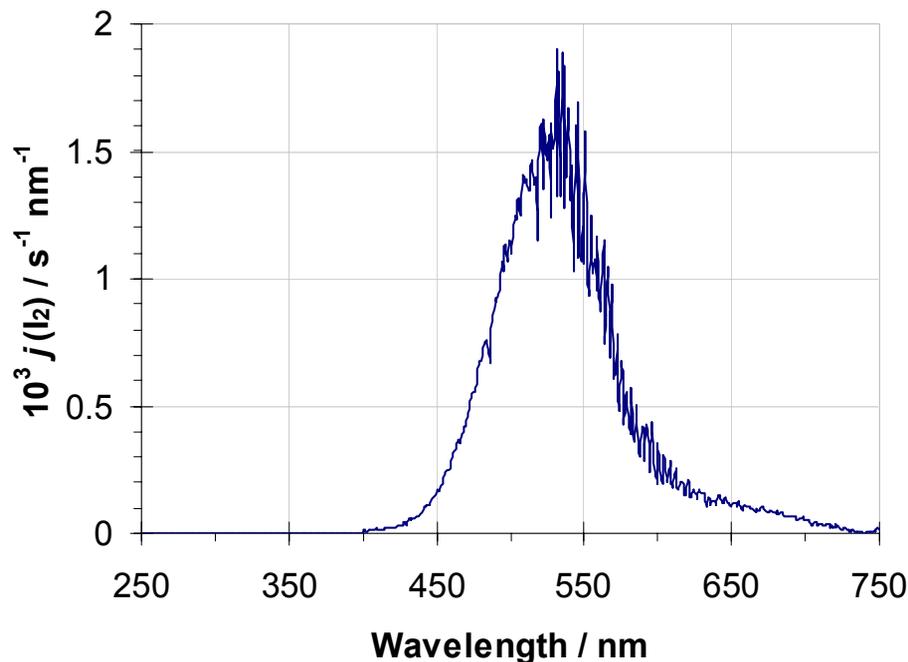


Fig. 6. Wavelength dependence of the I₂ photolysis rate in the lower atmosphere, calculated at 1 nm resolution. Values were convoluted using the measured I₂ cross-section, solar flux data from Tobiska et al. (2000) (attenuated by atmospheric absorption) and a photo-dissociation yield of unity.

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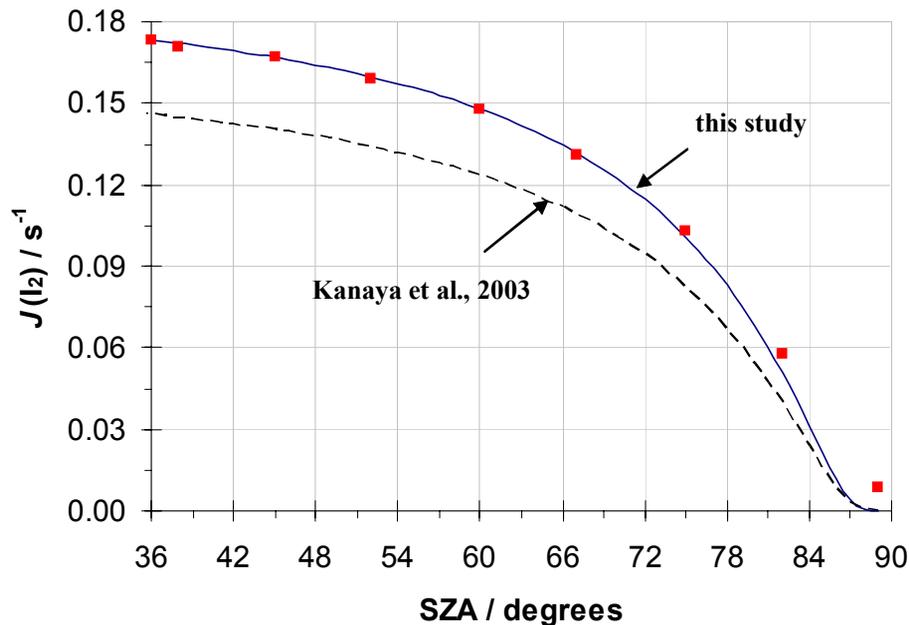


Fig. 7. Variation of $J(I_2)$ with SZA. Values were calculated for a mid-latitude location (53° N) in early August under cloud free conditions. The discrete data points are from the two-stream radiation model (Thompson, 1984), the solid line is the fit to these points using Eq. (5) with $L=0.223$, $M=0.002$ and $N=0.205$ (see text) and the dashed line is the data calculated for $L=0.1912$, $M=0.0019$ and $N=0.2185$ (Kanaya et al., 2003).

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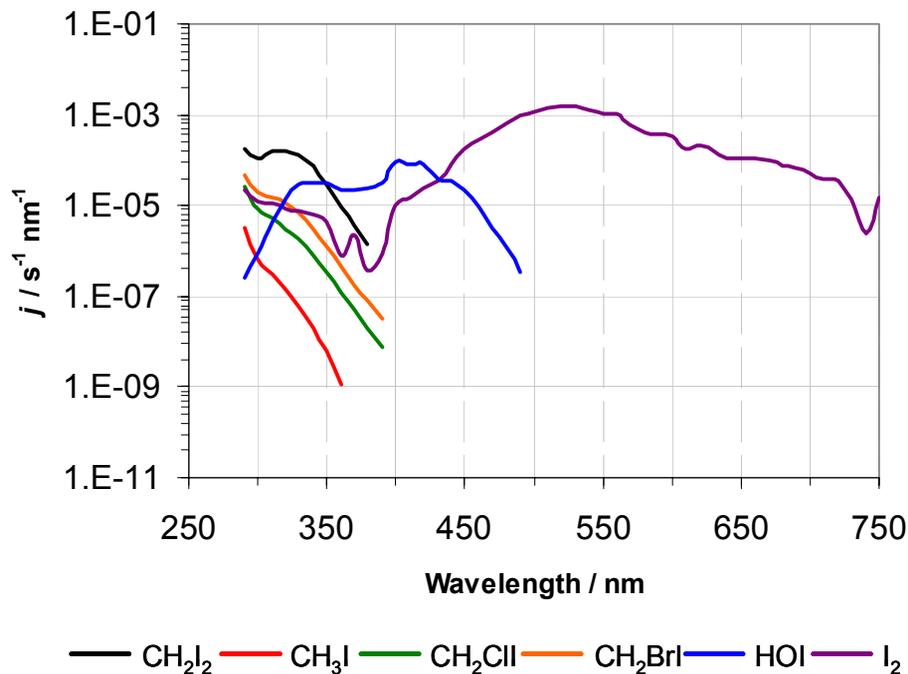


Fig. 8. Convolved tropospheric photolysis rate variation (clear sky, overhead sun) for a number of iodine-containing species (absorption cross-sections at 10 nm intervals). Cross-section data was taken from the IUPAC gas kinetic database (with the exception of I₂ for which our measured values were used), the photo-dissociation quantum yields were set to unity for all species, and direct actinic fluxes calculated from top of atmosphere data (corrected for atmospheric absorption). Integration across the wavelength range gives the photolysis rate constant (J) for each species (see text for values).

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